



# Response to elevated CO<sub>2</sub> in the temperate C3 grass *Festuca arundinaceae* across a wide range of soils

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Soils vary widely in mineral nutrient availability and physical characteristics, but the influence of this variability on plant responses to elevated CO<sub>2</sub> remains poorly understood. As a first approximation of the effect of global soil variability on plant growth response to CO<sub>2</sub>, we evaluated the effect of CO<sub>2</sub> on tall fescue (*Festuca arundinacea*) grown in soils representing 10 of the 12 global soil orders plus a high-fertility control. Plants were grown in small pots in continuously stirred reactor tanks in a greenhouse. Elevated CO<sub>2</sub> (800 ppm) increased plant biomass in the high-fertility control and in two of the more fertile soils. Elevated CO<sub>2</sub> had variable effects on foliar mineral concentration—nitrogen was not altered by elevated CO<sub>2</sub>, and phosphorus and potassium were only affected by CO<sub>2</sub> in a small number of soils. While leaf photosynthesis was stimulated by elevated CO<sub>2</sub> in six soils, canopy photosynthesis was not stimulated. Four principle components were identified; the first was associated with foliar minerals and soil clay, and the second with soil acidity and foliar manganese concentration. The third principle component was associated with gas exchange, and the fourth with plant biomass and soil minerals. Soils in which tall fescue did not respond to elevated CO<sub>2</sub> account for 83% of global land area. These results show that variation in soil physical and chemical properties have important implications for plant responses to global change, and highlight the need to consider soil variability in models of vegetation response to global change.

**Keywords:** soil taxonomy, soil orders, elevated CO<sub>2</sub>, *Festuca arundinaceae*, tall fescue

## INTRODUCTION

Substantial attention has been given to the effects of elevated CO<sub>2</sub> concentration on plant growth and physiology (Körner, 2006), reflecting concern about the performance of both cultivated and wild plants in future climates characterized by elevated CO<sub>2</sub> (IPCC, 2007a). Studies of the responses of crops and ecosystems to elevated CO<sub>2</sub> (Körner, 2006; Ziska and Bunce, 2006, 2007) often report increased growth and use efficiency of nitrogen and water when nutrient availability is optimal or near optimal (Woodward et al., 1991; Bunce, 2004). Though the majority of studies consider elevated CO<sub>2</sub> in isolation, plant responses to elevated CO<sub>2</sub> may be affected by other environmental factors, including soil properties (Diaz et al., 1993; Bassirad et al., 2001; Spinnler et al., 2002; Lynch and St. Clair, 2004, 2010; Fay et al., 2009, 2012a).

When suboptimal nutrient availability has been considered (generally as deficiency of either N or P), a commonly observed response is that a limited supply of N or P leads to a reduction in photosynthetic rates and foliar N concentrations and increased concentrations of non-structural carbohydrates (NSC) (van Noordwijk et al., 1998; Gifford et al., 2000; Gifford, 2004). Few studies have considered multiple nutrient stresses (deficiencies and/or toxicities) in conjunction with elevated CO<sub>2</sub> (Körner, 2006), and our understanding of the mechanisms behind the interaction of soil characteristics with elevated CO<sub>2</sub> is far from

complete (Lynch and St. Clair, 2004). In contrast to research with crops and model plants, forestry and ecological research has considered the effects of elevated CO<sub>2</sub> in natural soils without amendments or fertilizers. Such studies generally indicate multiple and complex limitations, mostly of edaphic origin, that trees face under elevated CO<sub>2</sub> (Bucher-Wallin et al., 2000; Bassirad et al., 2001; Egli et al., 2001; Poorter and Perez-Soba, 2001; Spinnler et al., 2002).

Natural soils vary widely across terrestrial ecosystems; the USDA soil taxonomy system addresses this variability by classifying soils into 12 orders based on factors related to soil formation (Wilding, 2000; Table S1). These are further divided into 64 suborders and additional sub-categories based on climatic and edaphic modifiers (Soil Survey Staff, 1999). One purpose of such taxonomic systems is to optimize land use over the range of soils so as to maximize productivity and sustainability (Driessen and Konijn, 1992). The variability described by soil taxonomy may also be useful in understanding climate change effects, such as the effect of elevated CO<sub>2</sub> on plant growth, but there has been little effort made to test whether, and to what extent, the effects of elevated CO<sub>2</sub> on plant growth depend on soil taxonomic variation.

A literature search performed on combinations of "Elevated CO<sub>2</sub>" and edaphic and soil taxonomy terms produced a relatively small number of hits (Table S2 in Supplemental Materials)—

only 57 unique publications were returned by searching “Elevated CO<sub>2</sub> AND soil type” and “Elevated CO<sub>2</sub> AND (any USDA soil order)” on Web of Science. Further analysis of topics for the 57 unique publications highlight the scarcity of consideration of the range of possible soil impacts on CO<sub>2</sub> responses (Tables S3, S4 in Supplemental Materials). Only 21 of these publications report results from an experiment with two or more soils and elevated CO<sub>2</sub>, and the maximum number of soils considered was three. Furthermore, these 21 publications report on only 6 unique experiments, and 16 of the 21 reports and 3 of the 6 unique experiments relate to woody plants. Although several authors have noted the importance of soils in determining plant responses to elevated CO<sub>2</sub> (Bassirirad et al., 2001; Poorter and Perez-Soba, 2001; Spinnler et al., 2003; Fay et al., 2009, 2012a), it appears that no attempt has been made to test the extent to which plant responses to elevated CO<sub>2</sub> vary across the natural variability of soils.

As noted above, work on woody plants may be more advanced than work in non-woody plants in this area. Watanabe et al. (2013) reported no CO<sub>2</sub> × soil interaction on photosynthetic traits of hybrid *Larix* grown for two seasons on a fertile forest soil and an infertile volcanic ash soil. In contrast, Spinnler et al. (2002) found that while in *Picea* there was no CO<sub>2</sub> × soil interaction on biomass, in *Fagus* elevated CO<sub>2</sub> only stimulated growth in a more fertile calcareous soil, and actually suppressed growth on an acidic soil. In the same system it was found that responses to elevated CO<sub>2</sub> may differ between root and shoot (Sonnleitner et al., 2001). Such differences may have ecosystem consequences; another report on the same system showed that the acidic soil increased its carbon content to a much greater degree than the calcareous soil, even though it supported much less biomass (Hagedorn et al., 2003). Interestingly, while biomass responses of *Picea* and *Fagus* in this system were strongest in the early years of this 4 year study, plant water relations still responded to elevated CO<sub>2</sub> in a soil-dependent manner (Bucher-Wallin et al., 2000). In contrast, another study showed no effect of elevated CO<sub>2</sub> on stomatal conductance or leaf hydraulic conductivity in *Betula* or *Quercus* grown on two contrasting soil, though responses to CO<sub>2</sub> differed between sun and shade leaves (Eguchi et al., 2008).

The best examples for non-woody plants are a series of reports on experiments carried out with monoliths of three soils that were exposed to a CO<sub>2</sub> gradient from sub- to super-ambient. In this system early growth of *Panicum virgatum* was enhanced by elevated CO<sub>2</sub> but not regrowth after clipping. An interaction of Soil × CO<sub>2</sub> was seen for soil moisture but not for annual net primary productivity (Fay et al., 2012b). Another study using this system with constructed prairie plant communities found that aboveground biomass response to CO<sub>2</sub> was greatest on soils with greater plant available water (a Mollisol and an Alfisol) and was reduced on a heavy clay soil (a Vertisol) with lower plant available water (Fay et al., 2012a). Another report on this system showed that in the Mollisol forbs responded more strongly to elevated CO<sub>2</sub> than grasses, though grasses were stimulated by increasing CO<sub>2</sub> in all three soils (Polley et al., 2012).

Grasses (Poaceae) include cereal crops that provide over half of the calories and protein consumed by humans (Cordain, 1999) and are the principle vegetation of the 30% of global land area

occupied by natural grasslands (Bartholome and Belward, 2005; Lambin and Geist, 2006). Grasses represent a large, variable group of plants that are successful in many environments and that have evolved several mechanisms to adapt to extreme soil conditions (Marschner, 1998).

In order to begin to characterize the effects of the wide range of soils on plant responses to elevated CO<sub>2</sub> we grew tall fescue (*Festuca arundinaceae* Schreb.), a temperate C3 grass, in elevated and ambient CO<sub>2</sub> on 13 different soils, representing ten of the twelve soil orders, and a high fertility control, and assessed plant growth, mineral acquisition, gas exchange, and non-structural carbohydrate accumulation.

## MATERIALS AND METHODS

### EXPERIMENTAL SETUP

Tall fescue was grown in eight Continuous Stirred Tank Reactors (CSTRs; mylar covered cylindrical steel frames approximately 2 m in diameter and 2 m tall with a continually rotating stirring paddle near the top to ensure even mixing of the atmosphere; Heck et al., 1975) in a greenhouse at the Pennsylvania State University (40°85'N, 77°83'W). The CSTRs were covered in transparent mylar and fitted with a positive pressure ventilation system that provided an airflow of 1 L per minute to each CSTR. Each CSTR was equipped with an external overhead 1000 watt HID Lamps for supplemental light; maximum light intensity at plant level averaged 350 μmol PAR s<sup>-1</sup> m<sup>-2</sup> (This relatively low lighting intensity reflects both the attenuation of solar radiation through both greenhouse roof and the mylar covering of the CSTRs and the difficulty of controlling heat load from the HID lighting supplied to each CSTR).

The eight CSTRs were grouped into four pairs, with one of each pair receiving near ambient (400 ppm CO<sub>2</sub>, which was near the ambient level in the greenhouse) and the other receiving elevated (800 ppm CO<sub>2</sub>, corresponding to the IPCC's “worst case” A1F1 scenario for mid; IPCC, 2007b). Elevated CO<sub>2</sub> was maintained by bleeding 99.8% dry CO<sub>2</sub> from a pressurized tank via a needle valve into a manifold from which four valves controlled the flow of CO<sub>2</sub> to each of the elevated CO<sub>2</sub> CSTRs. These valves were adjusted daily to maintain the target CO<sub>2</sub> concentration. CO<sub>2</sub> concentration (measured with a Li-Cor 6262 infrared gas analyzer connected to a multiplexing pump), temperature, photosynthetically active radiation (PAR), and relative humidity for each CSTR were recorded every 16 min. CO<sub>2</sub> concentrations were relatively stable, with mean values (±1 standard deviation of 790 ± 14 ppm for Elevated CO<sub>2</sub> and 399 ± 11 ppm for near ambient CO<sub>2</sub>).

### PLANTING

Soil samples representing 10 taxonomic orders (Table 1) were obtained from Puerto Rico, Ecuador and the U.S. (Alaska) during 2005 and 2006. In each location we collected from areas with no known history of fertilizer use. Soils were air dried and transported to University Park, PA., where they were kept in refrigerated storage (6°C) until the experiment began in January 2007, when the soils were sieved (2 mm) to exclude gravel and organic debris, and eight pots (400 ml volume) were filled with each soil type. Pots were also prepared with a high-fertility

**Table 1 | Properties of soils used in the study.**

Soil ID	Order (suborder)	Origin	pH	Total N (%)	P (ppm)	K (ppm)	Mg (ppm)	Ca (ppm)	Zn (ppm)	Cu (ppm)	S (ppm)	Clay%	Silt%	Sand%
ALF	Alfisol (Udalf)	PR	6.4	0.33	7	210	147	2065	3.6	4.4	46	15.3	22.3	62.3
AND	Andisol (Aquand)	EC	5.4	0.47	12	148	73	439	3.7	5.9	29.3	1.06	28.4	70.5
ARD	Aridisol	EC	7.9	0.09	71	182	633	2582	4.2	9.4	36.2	8.06	41.5	50.4
INC1	Inceptisol (Tropent)	EC	5.6	tr	7	86	168	768	3.5	3.1	19.6	0.45	5.46	94.1
INC2	Inceptisol (Udept)	PR	4.7	0.05	1	84	17	120	0.3	3.8	216	NA	NA	NA
GEL	Gelisol	AK	7.7	0.19	9	50	164	5490	NA	NA	NA	12.1	59.7	28.2
MOL	Mollisol (Ustol)	PR	7.5	0.16	200	689	497	5001	4.5	10.2	22.4	30.1	31.4	38.5
OXI1	Oxisol (Ustox)	PR	7.7	0.21	11	280	124	2987	1.7	5.2	17.7	26.0	29.2	44.8
OXI2	Oxisol (Udox)	PR	5.2	0.15	1	44	57	189	0.7	1.4	285	4.1	11.6	84.3
OXI3	Oxisol	EC	5.8	0.26	9	28	52	522	1.4	2	18.9	13.0	29.1	57.9
SPO	Spodosol (Orthod)	PR	7.3	0.19	20	33	125	3566	19.4	13.6	29.5	1.56	11.1	87.3
ULT	Ultisol (Humult)	PR	5.4	0.29	14	545	225	1248	2.6	6.4	44.3	20.8	33.2	46.1
VRT	Vertisol (Ustert)	PR	6.8	0.15	10	150	1737	5016	1.4	8	23.3	22.9	29.7	47.4

PR, Puerto Rico; EC, Ecuador; AK, Alaska (United States); NA, Not available; tr, trace values.

control treatment (“CTR”), consisting of a standard horticultural medium based on peat and vermiculite (Sunshine Mix #3, Sun Gro Horticulture, Bellevue, WA) amended with a complete slow-release fertilizer (Osmocote 14-14-14 in a rate of approximately 3 g per pot; Scotts Miracle-Gro, Marysville, OH).

Approximately 20 seeds of tall fescue (cultivar Kentucky 31; SeedLand, Inc. Wellsborn, FL) were broadcast on the surface of each pot on Jan 23, 2007. The pots were covered in clear plastic until germination, after which pots were thinned to 15 plants pot<sup>-1</sup>. 14 pots (the 13 soils plus the fertile control) were randomized within each of the eight CSTRs, for a total of 112 pots. Pots were irrigated manually with distilled water every day.

#### DATA COLLECTION

The presence of the endophyte *Neotyphodium sp.* was assessed from a sample of 100 seeds and two growing tillers per pot (from two replicates), using a commercial immunoblot detection kit (Agrinostics Ltd. Co., Watkinsville, GA).

Leaf photosynthesis ( $A_{\max}$ , area basis) of a young, fully expanded leaf in one plant per pot was determined on weeks 8 and 10 at mid-day with a LI-6400 portable photosynthesis system (Li-Cor, Lincoln, NE) with leaf temperature set to 20°C. Measurements were made at 600, 800, and 1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  PAR, after measurements at 200–1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  PAR on a subset of plants showed that maximum photosynthetic rate

occurred in the 600–1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  range. Also at week 10 the net pot CO<sub>2</sub> exchange (including soil) was measured with a Li-6200 Infrared Gas Analyzer system (Li-Cor, Lincoln, NE) using a 12 liter chamber in which the whole plant and pot were enclosed for 2 min. During these measurements temperature ranged from 25 to 28°C, and light (PAR) ranged from 320 to 380  $\mu\text{mol m}^{-2} \text{s}^{-1}$ , reflecting the ambient growing conditions.

Above-ground tissue was harvested following CO<sub>2</sub> exchange measurement. Immediately following shoot excision, the same method was used to measure the amount of respiration from the roots and soil. Canopy CER was estimated as the difference between the net pot CO<sub>2</sub> exchange and root + soil CO<sub>2</sub> exchange.

Excised shoots were divided into three samples. The first sample (~5 g fresh weight) was frozen at –80°C. These were later processed to quantify ethanol soluble sugars and starch from approximately 100 mg of ground tissue with the enzyme-coupled colorimetric method described by Hendrix (1993).

All remaining above-ground tissue was oven dried at 60°C for 48 h to determine total dry weight. A sample of about 0.2 g of dried and ground leaf tissue was digested in a microwave (Miller, 1998). The diluted (250:1) extract was analyzed in a Varian Induced Couple Mass Spectrophotometer (Varian Inc., Palo Alto CA) to determine the content of: phosphorus, potassium, calcium, magnesium, manganese, iron, copper, boron, aluminum, zinc and sodium. A second sample of dried and ground tissue

was analyzed using a Perkin-Elmer EA2400 elemental analyzer with combustion and reduction columns to determine carbon and nitrogen content.

**STATISTICAL ANALYSIS**

The data were analyzed with a split-plot design. The four pairs of CSTRs were treated as blocks, with CO<sub>2</sub> as the main plot factor and soil the subplot factor. Data were analyzed in R (R Core Team, 2014) using *lme* (Pinheiro et al., 2014), to fit linear mixed effects models with block and CSTR as random effects (The model was “response ~ CO<sub>2</sub> + Soil + Soil:CO<sub>2</sub>, random = ~1|Block/CSTR”), and residuals were checked to ensure that regression assumptions were not violated. For A<sub>max</sub>, the mixed effect model included all saturating levels of PAR with PAR as a random effect (The model was “A<sub>max</sub> ~ CO<sub>2</sub> + Soil + Soil:CO<sub>2</sub>, random = ~1|Block/CSTR/PAR”). Pair-wise Tukey comparisons for the effect of CO<sub>2</sub> in each soil were obtained using *multcomp* (Hothorn et al., 2008).

Given the large number of variables measured (31) Principal component analysis (PCA) was used to characterize the principal sources of variability in the data. PCA was carried out for all observations with growth, photosynthesis, soil analysis and leaf mineral concentrations as variables using *prcomp* (R Core Team, 2014).

**RESULTS**

About 40% of the seeds and 95% of all tillers tested positive for the *Neotyphodium* endophyte, and *Neotyphodium* colonization did not differ among soils or CO<sub>2</sub> treatments.

Of the soils used in this experiment, only INC2 did not support any germination of tall fescue. Plants grown in GEL, HIS and OXI2 exhibited severe reductions (>97%) in biomass compared to CTR, and did not produce sufficient tissue for all analyses to be completed. They were therefore eliminated from most of the subsequent analyses. The remaining soils produced 50–80% less shoot biomass than CTR, indicating significant differences among soils (Table 2). These differences were especially notable under elevated CO<sub>2</sub>, where biomass of CTR increased by about 60% (Figure 1A). Elevated CO<sub>2</sub> significantly increased plant biomass in the CTR, ULT and ALF (by approximately 20–40%), but did not increase biomass in the other soils (Figure 1A).

In the soils used in this experiment elevated CO<sub>2</sub> did not significantly change leaf nitrogen concentration (*p* = 0.323; Figure 1B, Table 2), though there were differences in leaf N associated with the soils (*p* < 0.001), as would be expected. Furthermore, the C:N ratio was not altered by CO<sub>2</sub> or the CO<sub>2</sub> × soil interaction (*p* > 0.300, data not shown), though it did differ among the soils (*p* < 0.001, Table 2).

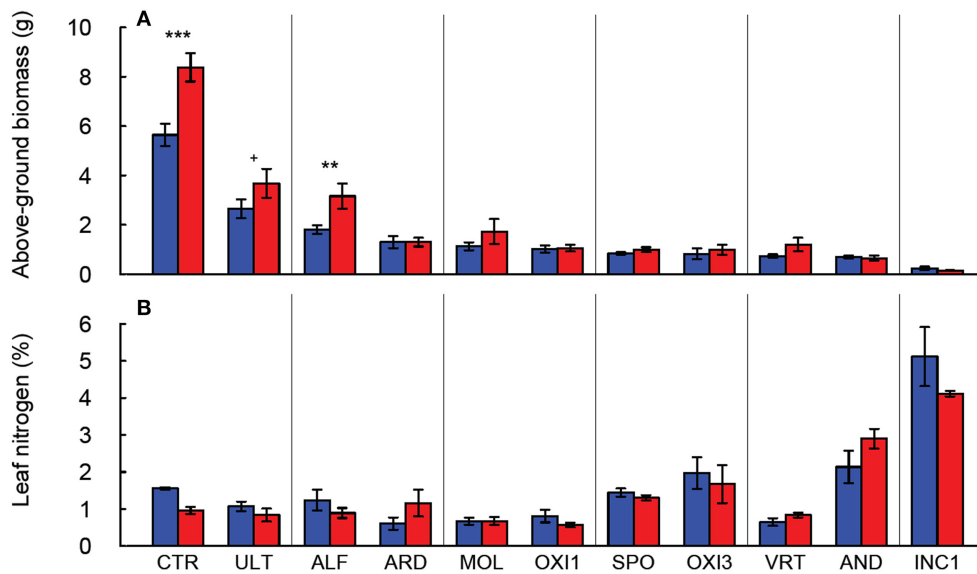
Leaf phosphorus concentration declined approximately 30% with elevated CO<sub>2</sub> in CTR, while in the remaining soils phosphorus concentration changed only marginally (*p* = 0.062; Figure 2A). Leaf potassium concentration was reduced by 40% under elevated CO<sub>2</sub> in ALF soil, and increased by 100% in INC1, but was not affected in the remaining soils (*p* > 0.650; Figure 2B). Leaf K concentration was dramatically lower in SPO than in CTR, but the remaining soils produced leaf K values within about ±20% of CTR (Figure 2B).

**Table 2 | Summary of analysis of variance results of biomass, mineral content, and gas exchange parameters for *Festuca arundinacea* var.**

		<b>Biomass</b>		<b>A<sub>max</sub></b>		<b>Shoot CER</b>		<b>Pot Resp.</b>	
Source	df	F-value	df	F-value	df	F-value	F-value	F-value	F-value
CO <sub>2</sub>	1/3	6.98*	1/3	7.71+	1/3	1.01		7.29+	
Soil	13/77	117***	10/208	15.5***	10/60	10.8***		19.1***	
CO <sub>2</sub> × Soil	13/77	5.50***	10/208	6.57***	10/60	0.903		2.24*	
		<b>N</b>		<b>C:N</b>		<b>Starch</b>		<b>Sucrose</b>	
CO <sub>2</sub>	1/3	2.12	1/3	1.13	1/3	15.8*		3.01	
Soil	12/59	25.5***	10/55	7.50***	9/54	2.05+		2.79**	
CO <sub>2</sub> × Soil	12/59	1.32	10/55	0.696	9/54	1.25		0.711	
		<b>P</b>		<b>K</b>		<b>Ca</b>		<b>Mg</b>	
Source	df	F-value	F-value	F-value	F-value	F-value	F-value	F-value	F-value
CO <sub>2</sub>	1/3	8.43+	4.51	18.9	15.8*	11.6*			
Soil	10/56	55.1***	20.7***	72.6***	26.7***	126***			
CO <sub>2</sub> × Soil	10/56	3.26**	3.70***	3.08**	1.78+	1.75+			
		<b>Al</b>		<b>Fe</b>		<b>B</b>		<b>Na</b>	
CO <sub>2</sub>	1/3	0.00111	0.0181	0.00137	3.10	7.75+			
Soil	10/56	4.68**	16.4***	13.5***	23.4***	104***			
CO <sub>2</sub> × Soil	10/56	0.423	1.77+	0.624	1.72+	3.60**			

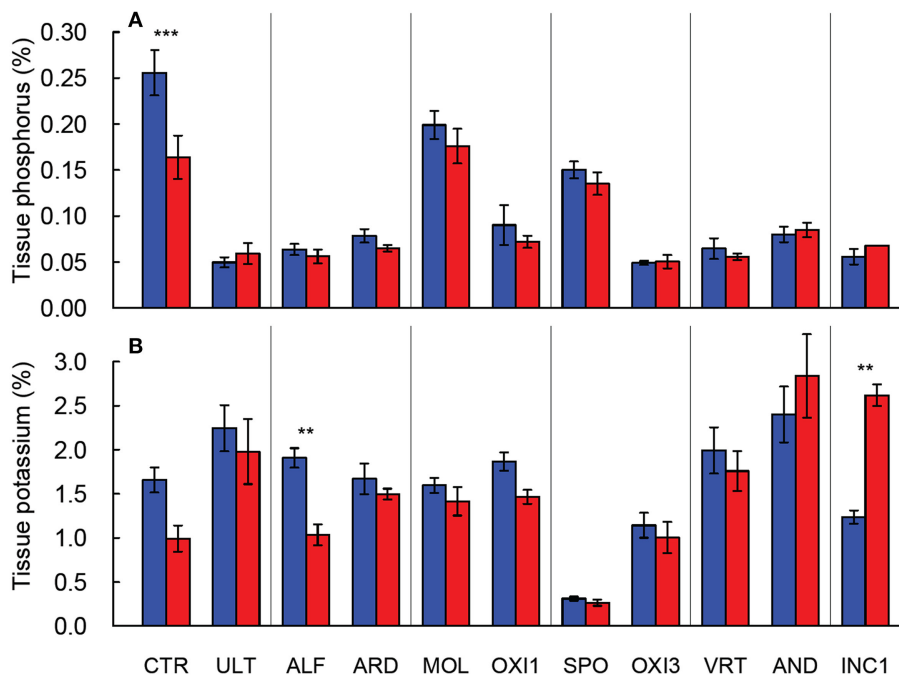
Kentucky 31 grown in ambient (400 ppm) and elevated (800 ppm) CO<sub>2</sub> in 13 different soils plus a high-fertility control.\*

\* Not all soils produced sufficient biomass for all analyses, leading to differences in degrees of freedom between analyses. Degrees of freedom (df) printed to the left of each variable and listed as numerator df/denominator df. Pot Resp. = Root + Soil CO<sub>2</sub> exchange rate. Significance denoted as: + *p* = 0.1–0.05; \* *p* = 0.05–0.01; \*\* *p* = 0.01–0.001; \*\*\* *p* < 0.001.



**FIGURE 1 | Production of above-ground biomass (A) and leaf nitrogen content (B) of *Festuca arundinacea* grown in 12 different soils and a high-fertility control under elevated (800 ppm; red) and ambient (400 ppm; blue) atmospheric CO<sub>2</sub>.**

Mean of four replicates ± one standard error shown. Significance indicated based on pair-wise Tukey comparisons. The soils are ordered by decreasing biomass in ambient CO<sub>2</sub>. +*p* = 0.1–0.05; \*\**p* = 0.01–0.001; \*\*\**p* < 0.001.



**FIGURE 2 | Leaf P (A) and K (B) concentration in shoots of *Festuca arundinacea* grown in 10 soils and a high fertility control under elevated (800 ppm; red) and ambient (400 ppm; blue) CO<sub>2</sub>. Mean of four replicates**

± one standard error shown; lack of error bar indicates missing data points. The soils are ordered by decreasing biomass in ambient CO<sub>2</sub>, as in Figure 1A. \*\**p* = 0.01–0.001; \*\*\**p* < 0.001.

The ANOVA analysis for leaf mineral concentrations found highly significant differences (*p* < 0.001) for the soil effect for all leaf nutrient concentrations measured. However, the results for CO<sub>2</sub> and the soil × CO<sub>2</sub> interaction were less consistent. For

example the soil × CO<sub>2</sub> interaction was significant for P, K, Ca, and Zn, and only marginally significant for Mg, Mn, and Fe. (Table 2). Elevated CO<sub>2</sub> increased content of calcium in INC1 and SPO (*p* < 0.001), of magnesium in CTR (*p* = 0.038) and

OXI3 ( $p = 0.003$ ), of manganese in OXI3 ( $p < 0.001$ ); of sodium in SPO ( $p < 0.001$ ); and of zinc in CTR ( $p < 0.001$ ) and SPO ( $p = 0.033$ ).

Elevated CO<sub>2</sub> increased leaf photosynthesis ( $A_{\max}$ ,  $\mu\text{mol CO}_2 \text{ s}^{-1} \text{ m}^{-2}$  leaf area) in most of the soils, but not in CTR, ALF, MOL, and OXI3, reflecting a significant soil  $\times$  CO<sub>2</sub> interaction ( $p < 0.001$ ; **Table 2**, **Figure 3A**). Among the soils  $A_{\max}$  in ambient CO<sub>2</sub> was highest in CTR and AND in ambient and elevated CO<sub>2</sub> respectively and lowest in INC1 (both CO<sub>2</sub> levels; **Figure 3A**; **Table 2**).

Aboveground carbon exchange rate (Shoot CER;  $\text{nmol CO}_2 \text{ s}^{-1} \text{ mg}^{-1}$  DW canopy) did not differ significantly between elevated and ambient CO<sub>2</sub> overall ( $p = 0.389$ ), but varied between soils ( $p < 0.001$ ), and there was no interaction between CO<sub>2</sub> and soil (**Table 2**, **Figure 3B**). Root + soil CO<sub>2</sub> exchange rate ( $\text{nmol CO}_2 \text{ s}^{-1}$ ) increased in CTR, ALF, ARD, and MOL, but not in the others (Figure S1), leading to a significant Soil  $\times$  CO<sub>2</sub> interaction ( $p = 0.027$ ; **Table 2**).

Combining CO<sub>2</sub>, biomass production, mineral content and photosynthesis variables we created a matrix of 29 variables and 76 observations after records with missing data were excluded (mostly soils in which insufficient biomass was produced for all analyses to be completed; HIS, GEL, INC2, and OXI2) were excluded. Since % Sand, % Silt, and % Clay sum to 100%, we excluded % Sand from this analysis. Similarly, since CEC is reflective of soil Ca and Mg, we excluded CEC.

Principal components analysis yielded four principal components (PC) which explained about 65% of the variability in the data (**Figure 4**). PC1 (29% of variability) was most strongly influenced by foliar mineral concentrations (Zn, Cu, Mg, Ca, and Na) and soil clay content (minerals declining with increasing clay). PC2 (19% of variability) was most influenced by soil copper, soil calcium, soil pH foliar P and foliar Mn concentrations (all others decrease when foliar Mn increases). PC3 (9% of variability) was most strongly influenced by gas exchange (leaf photosynthesis

and stomatal conductance) and foliar Al, P, Fe and C (photosynthesis increasing with decrease in foliar minerals). PC4 (8% of variability) was most strongly influenced by above-ground biomass, sucrose, soil S, soil K, and Mg (with Mg decreasing when others increase; **Table S5**).

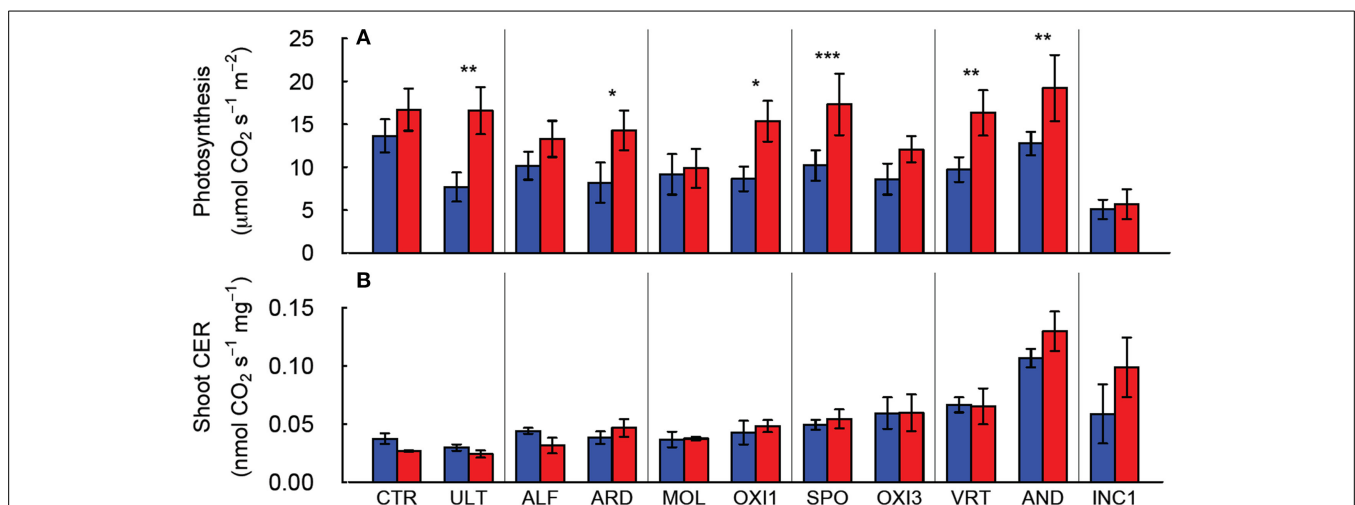
High and low molecular weight NSC (starch and sucrose respectively) responded differently in this experiment (**Table 2**). Starch concentrations increased under elevated CO<sub>2</sub> ( $p = 0.029$ ), but sucrose concentrations were not affected by CO<sub>2</sub> ( $p = 0.181$ ). Levels of NSC were influenced by the different soils ( $p = 0.051$  and 0.009 for starch and sucrose respectively; **Table 2**). However, the effect of elevated CO<sub>2</sub> on NSC did not depend on soil type ( $p = 0.288$  and 0.697 for starch and sucrose respectively; **Table 2**, **Figure 5**).

Maximum photosynthetic rate ( $A_{\max}$ ) did not show a strong correlation with foliar N (**Figure 6A**). Furthermore, the relationship between  $A_{\max}$  and foliar N differed between soils ( $p = 0.0008$ ) and marginally with CO<sub>2</sub> ( $p = 0.079$ ). A graphical analysis (based on non-overlap of SE ellipses, **Figure 6A**) suggests that for OXI1, ULT, SPO, CTR, and AND  $A_{\max}$  may not be strongly related to foliar N. The tukey test of pair-wise differences confirms this for OXI1 and ULT ( $p = 0.006$  and 0.062).

Analysis of N:P ratios (**Figure 6B**, **Table 2**) shows a strong effect of soil ( $p < 0.001$ ) but no effect of CO<sub>2</sub> ( $p = 0.410$ ), or their interaction ( $p = 0.131$ ). Three rough groupings of soils are differentiated here. CTR and MOL, with a low N:P ratio, INC1, AND, and OXI3, with a high N:P ratio, and ARD, ALF, OXI1, SPO, ULT, and VRT with intermediate values.

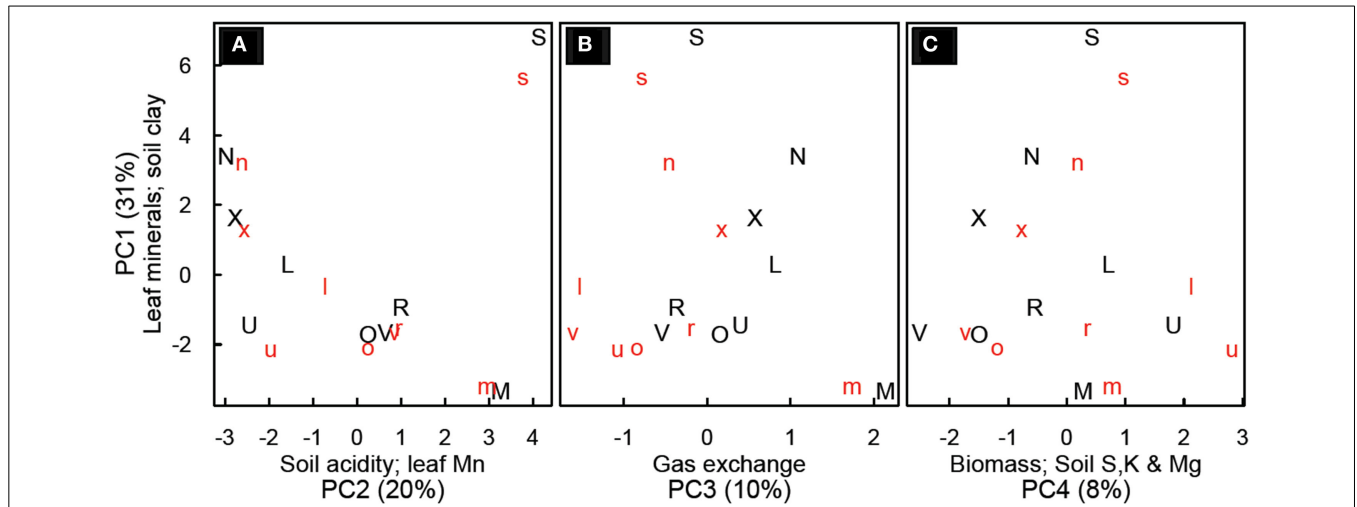
## DISCUSSION

In this experiment we evaluated the effect of elevated CO<sub>2</sub> on the growth and physiology of *Festuca arundinacea* encountering different chemical and physical soil characteristics presented by soils from 9 of the 12 soil orders, spanning the global range of soil



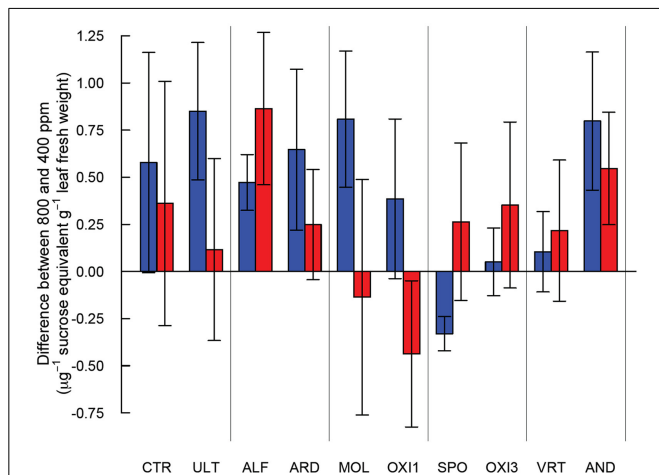
**FIGURE 3 | Leaf photosynthesis of individual leaves (A) and net carbon exchange rate (CER) for the shoot (B) of *Festuca arundinacea* grown in 10 soils, plus a high-fertility control under elevated (800 ppm; red) and**

**ambient (400 ppm; blue) atmospheric CO<sub>2</sub>.** Mean of four replicates  $\pm$  one standard error shown. The soils are ordered by decreasing biomass in ambient CO<sub>2</sub>, as in **Figure 1A**. \* $p = 0.05$ –0.01; \*\* $p = 0.01$ –0.001; \*\*\* $p < 0.001$ .

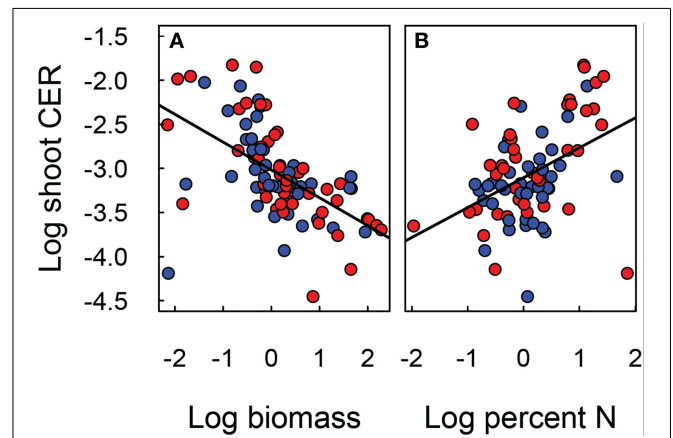


**FIGURE 4 | Principal component analysis for *Festuca arundinacea* grown in 10 soils, plus a high-fertility control at two CO<sub>2</sub> levels.** Principal component (PC) 1 vs. PC2 (A); PC1 vs. PC3 (B) and PC 1 vs. PC4 (C). Lower-case red letters represent elevated CO<sub>2</sub> (800 ppm) and upper-case black letters ambient CO<sub>2</sub> (400 ppm). Axes are labeled with the principal

component, the variables most strongly related to that principle component, and (in parentheses) the percent of variability explained by that principle component. The symbols represent the soils as follows: L, Alfisol; N, Andisol; R, Aridisol; M, Mollisol; O, Oxisol1; X, Oxisol3; S, Spodosol; U, Ultisol; V, Vertisol.



**FIGURE 5 | Differences in starch (blue) and sucrose (red) content between *Festuca arundinacea* grown under elevated and ambient atmospheric CO<sub>2</sub> in 10 soils, plus a high-fertility control.** The soils are ordered by decreasing biomass in ambient CO<sub>2</sub> as in Figure 1A, and those which produced insufficient biomass for analysis are not shown.



**FIGURE 6 | Relationship of shoot CER to shoot biomass (A) and nitrogen content (B) of *Festuca arundinacea* grown in 10 soils, plus a high-fertility control under elevated (800 ppm; red) and ambient (400 ppm; blue) atmospheric CO<sub>2</sub>.** The solid line indicates N:P ratio of 13 and the dotted lines ratios of 9 and 19. The solid line illustrates the linear regression model fit to the data in each panel.

variability. The effect of elevated CO<sub>2</sub> on growth, photosynthesis, and leaf chemistry depended on the soil in which the plants were grown. Since we provided adequate irrigation, we assume that the responses we observed mostly reflect the ability of *Festuca* to acquire nutrients from the different soils under contrasting CO<sub>2</sub> regimes.

Of the 13 soils used in this experiment four soils either failed to permit germination or to produce sufficient tissue for all of our assays. These are extreme soils for which *Festuca*, despite its wide range of adaptability (Malinowski and Belesky, 2000; Rahman and Saiga, 2007), is apparently not well adapted. These soils

include the two soils with the lowest phosphorus values (INC2 and OXI2) and the lowest pH of all the soils (INC2).

Although *Neotyphodium* is restricted to the aerial parts of the plant, its presence has been related to increased tolerance of fescue to several edaphic stresses, including P and Ca, (Malinowski et al., 2000) though the effect of endophyte infection may depend on plant genotype and soil (Rahman and Saiga, 2007). We found no effect of soil or CO<sub>2</sub> on colonization of *Neotyphodium* in fescue; we therefore do not expect *Neotyphodium* colonization to favor specific soil treatments (Malinowski and Belesky, 2000).

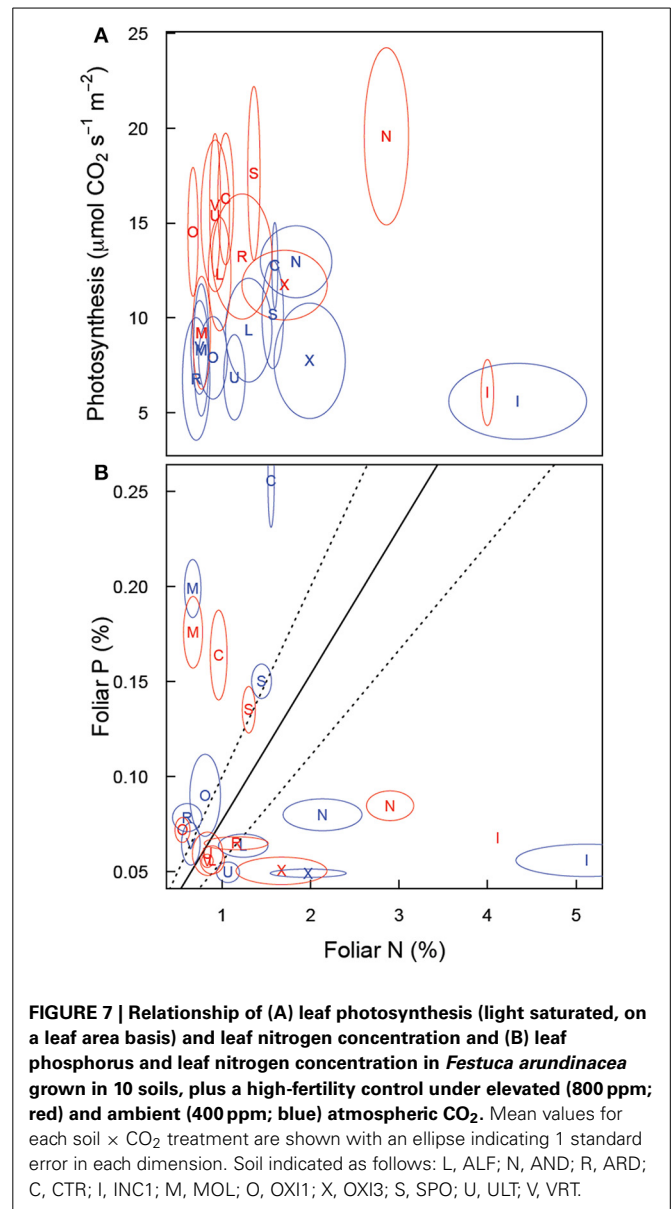
We found large differences between soils in the effects of elevated CO<sub>2</sub> on *Festuca*. First, there was a large increase in biomass

under elevated CO<sub>2</sub> on soils where biomass was greater under ambient CO<sub>2</sub>, in contrast to the lack of stimulation seen in soils that supported less biomass production in ambient CO<sub>2</sub>. To some extent, this may simply indicate that soils where *Festuca* grows well support a greater increase in biomass in elevated CO<sub>2</sub>; i.e., there may be a correlation between biomass increase in elevated CO<sub>2</sub> and biomass in ambient CO<sub>2</sub>. This relationship is significant, but only explains 44% of the variability in CO<sub>2</sub> response ( $p < 0.001$ ,  $R^2 = 0.44$ ). This suggests that other factors may also be important.

Our results generally agree with previous reports of the lack of response to increased CO<sub>2</sub> under nutrient-limited conditions (Poorter and Perez-Soba, 2001; Ziska and Bunce, 2006). Differences in leaf elemental concentration highlight a second important response; some soils have inherent low levels of N, P and K, and plants do not accumulate these elements to sufficient levels, and therefore may experience limitation by these elements. For example plants grown in Spodosols had very low foliar K (Figure 2B), suggesting K limitation as a possible cause for the low biomass production (Figure 1A) in this soil. In some cases elevated CO<sub>2</sub> led to the accumulation of non-limiting elements. For example under elevated CO<sub>2</sub> plants in AND and INC1 had higher levels of foliar N than CTR or any other more fertile soil; plant growth in these soils was apparently not limited by nitrogen, but the very low biomass of these plant points at some limitation (Figure 1).

The carbon exchange rate for whole shoot and leaf photosynthesis ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ leaf s}^{-1}$ ) also showed contrasting results (Figure 3). While leaf photosynthesis ( $A_{\text{max}}$ ) showed the expected increase in response to elevated CO<sub>2</sub> in six of the soils, shoot CER was not altered by elevated CO<sub>2</sub>. The difference between leaf and canopy photosynthetic responses to elevated CO<sub>2</sub> was not a simple product of changes in biomass as shoot CER was normalized by biomass. There was a significant negative relationship (log-linear) between shoot CER and shoot biomass (Figure 6A,  $p < 0.001$ ,  $R^2 = 0.30$ ). However, shoot CER was positively related (log-linear) with leaf N (Figure 6B,  $p < 0.001$ ,  $R^2 = 0.17$ ). There are reports of growth dilution of leaf N by elevated CO<sub>2</sub> (Luo et al., 1994; Idso and Idso, 2001; Taub and Wang, 2008; Wieser et al., 2008), but since we saw no difference in leaf N (Figure 1B) or C:N ratio (not shown) with CO<sub>2</sub> there is no evidence of growth dilution. While some of the difference between leaf and canopy level responses may be explained by the lower light levels for the measurement of CER relative to that used for the leaf photosynthesis (300 vs. 1000  $\mu\text{mol PAR s}^{-1} \text{ m}^{-2}$ ), the relationship of leaf N concentration and canopy CER suggests that there are fundamental differences in photosynthetic N use between plants grown in different soils.

In a review of research on N:P ratios, Güsewell (2004) reported that N:P ratios near 13 were typical for plants grown in their native conditions. In our study, only ALF, ULT and VRT under elevated CO<sub>2</sub> showed values near this (Figure 7B). A range of 9–19 for N:P ratios was also reported for a range of plants in a range of vegetative communities (Güsewell, 2004), with a range of 10–14 for graminoids. As shown in Figure 7B, few of our plants fell within this range (ALF, ARD, and ULT in elevated CO<sub>2</sub>, and VRT in both atmospheres). Without more detail on the



soil the reported N:P ratios represent it is difficult to interpret these results—it is possible that the range of N:P ratios reported does not represent the range of soil variability we are testing here. Alternatively, the wide range of N:P ratios we report here may indicate that fescue is not well adapted to some of these soils. However, correlation between the divergence of observed N:P ratio from the “optimum” value of 13 and biomass was very low ( $r = -0.25$ ) suggests that divergence of N:P ratio from some optimum is not strongly related to biomass.

The progressive nitrogen limitation hypothesis (Luo et al., 2004) suggests that increased plant biomass (and hence soil organic matter) stimulated by elevated CO<sub>2</sub> can immobilize sufficient N to lead to increasing N limitation with elevated CO<sub>2</sub>. We saw little support for this in this study. Foliar N had limited influence on  $A_{\text{max}}$  (Figure 7A). Furthermore, biomass was stimulated in only 3 soils (Figure 1A) and root respiration was only



modestly increased in four soils (of which two showed above-ground biomass increase also). So in 6 of the 11 soils for which we could fully analyze results there was no stimulation of biomass or below-ground respiration. We note however that the progressive nitrogen limitation hypothesis is likely to operate more strongly on ecosystem spatial scales and over multiple seasons, as a key mechanism for progressive limitation is that a greater proportion of N ends up in plant tissue and in soil organic matter. In this study, the elevated CO<sub>2</sub> treatment was confined to 10 weeks, which is very unlikely to be a sufficient time-span for progressive limitation to occur to a notable extent.

A substantial number of studies have considered the effect of elevated CO<sub>2</sub> on root:shoot allocation. We did not harvest roots in this study because the relatively small sizes of plants in this study mean that roots will mostly be rather fine. This fact, combined with the wide range of soil textures (Table 1) would not have yielded reliable data, as the recovery rate for roots would have varied rather strongly with soil texture.

An increase in non-structural carbohydrates (NSC) as been reported in plants exposed to elevated CO<sub>2</sub> (Poorter and Perez-Soba, 2001; Ziska and Bunce, 2006, 2007). The pattern of accumulation of NSC we observed suggests that carbohydrate physiology may be strongly influenced by differences in soils, though the variability in NSC was rather high (Figure 5). Accumulation of NSC in elevated CO<sub>2</sub> in the high-fertility control was not significant, while in ALF and AND there was accumulation of both high- and low-weight NSC. In ULT, ARD, and MOL only high weight NSC accumulated. In contrast, for SPO high weight NSC were reduced in elevated CO<sub>2</sub>. The lack of accumulation in SPO, OXI3 and VRT suggests active use of carbon in the plant, possibly in high metabolic demand processes such as mineral acquisition. Alternatively, carbon losses could take place through respiration or rhizo-deposition (Nguyen, 2003). These contrasting responses highlight the ways in which carbohydrate assimilation and metabolism may be influenced by soil conditions.

The four principal components we found in this study improve our understanding of the inter-relationships among mineral content in soils, foliar concentration of minerals, photosynthesis and biomass (Figure 4). In the soils we sampled, foliar levels of Zn, Cu, Mg, Ca, and Na tended to be higher in soils with lower clay content (PC1). In general cations such as Mg and Ca are more available in soils with greater cation exchange capacity, which tend to be soils with higher clay and organic matter content. However, here we see a tendency for lower soil clay to be associated with higher values of these minerals in leaves, suggesting that soil availability of these minerals is not the strongest determinant of their foliar concentration. Spodosols were differentiated on PC1, likely due to the high concentration of calcium and sodium observed in leaves and the low clay content. The high Na content in SPO may have reduced K availability and altered carbohydrate physiology as reflected in the distinct NSC patterns in SPO discussed above. In most soils we observed a decline in the values of PC1 with elevated CO<sub>2</sub>, suggesting a trend of dilution of minerals as was observed clearly with CTR, ALF and ULT. The risk of mineral dilution and the consequent loss of food and forage quality has been mentioned by others (Idso and Idso, 2001; Wieser et al.,

2008); our findings suggest that this may affect plants on some soils differently than on others.

PC2 was most influenced by soil minerals (Cu, Ca, Zn), soil pH, foliar P, and foliar Mn (with the opposite sign). This is not surprising, as soil pH governs soil mineral availability, and it is well known that in acidic soils low foliar concentrations of P and high concentration of Mn can inhibit growth (Marschner, 1998). On PC3 photosynthesis and stomatal conductance are opposite in sign to foliar Al, P, Fe, and C. This grouping may indicate limitation of growth and photosynthesis by something other than P; in such conditions foliar P might be less correlated to photosynthetic responses. On PC4, the loading of biomass, low molecular weight NSC (sucrose), and foliar C indicate that growth is favored under conditions that favor sucrose, rather than starch, accumulation in leaves. Starch accumulation in leaves is one symptom of severe P deficiency (Marschner, 1998).

The fact that the first four PCs only captured 65% of the variation in the data indicates that the relationships between photosynthesis, leaf mineral content, and soil physical and chemical properties is complex and highly dimensional—a small number of variables will not adequately describe the range of differences seen. The strongest loadings for CO<sub>2</sub> were on PC6 (4.7% of the variation) and PC12 (2% of the variation), and CO<sub>2</sub> was also loaded on PC5 (5.4% of variation). This suggests that variability associated with CO<sub>2</sub> was relatively low in this data set compared with that associated with plant responses to diverse soils. This suggests that more work is needed with highly diverse soils to better map the potential responses of plants to global change variables.

Soil texture and its influence on plant available water has been suggested as a mechanism that mediates differing responses to elevated CO<sub>2</sub> (Fay et al., 2012a). The interaction of soil texture and elevated CO<sub>2</sub> via water availability is an important mechanism that requires further investigation. Our methodology in this study did not test these responses as all plants were well watered. In natural systems where water availability is limiting, responses to elevated CO<sub>2</sub> could be larger than what we observed. However, as noted by Lynch and St. Clair (2004), toxicity of metals such as Mn can be strongly controlled by soil moisture, so it is also possible that increasing soil water could have negative effects on plant growth. Given the importance of this interaction, a more complete exploration of this interaction is clearly needed. In order to avoid artifacts introduced by sieving or mixing soils, such a test would best be achieved using soil monoliths and a method of providing experimental units with the same total water over the growing season.

The differences among soils in the response to elevated CO<sub>2</sub> suggest some caution in predicting plant responses to elevated CO<sub>2</sub> based on the world-wide network of free-air carbon enrichment (FACE) sites without considering how the FACE sites reflect the global diversity of soils. Such caution has been suggested by others who have noted that the distribution of soils limited by acidity (von Uexkull and Mutert, 1995) and phosphorus deficiency (Sanchez, 1976, 1981; Fairhurst et al., 1999; Jaramillo, 2011) and how these contrast with the geographic concentration of the free-air concentration enrichment (FACE) studies in countries in zones free of these edaphic limitations (Schimel, 2006).

**Table 3 | Percent area of each of the soil orders in six continents.**

Order	S. America	Africa	Asia	Oceania	N. America	Europe
Alfisols	<b>10.2</b>	<b>12.3</b>	<b>4.90</b>	<b>14.8</b>	<b>10.4</b>	<b>25.4</b>
Andisols	1.39	0.16	0.57	0.88	4.53	0.46
Aridisols	8.22	14.0	11.8	35.1	7.89	0.85
Entisols	15.0	41.8	11.3	26.3	6.89	6.27
Gelisols	0.49	0.00	27.4	0.00	13.3	5.84
Histosols	0.25	0.06	1.58	0.02	2.39	2.68
Inceptisols	11.8	6.36	26.1	4.19	25.5	19.0
Molisols	6.43	0.35	7.75	1.65	11.0	14.5
Oxisols	30.9	13.9	0.21	1.15	0.56	0.00
Spodosols	0.16	0.00	0.71	0.93	11.0	24.2
Ultisols	<b>14.2</b>	<b>7.35</b>	<b>6.48</b>	<b>3.56</b>	<b>4.30</b>	<b>0.04</b>
Vertisols	0.93	3.64	1.30	11.5	2.25	0.77

Soils in which biomass of *Festuca* was stimulated by elevated CO<sub>2</sub> are indicated in bold.

These differences also raise the possibility that models derived from FACE studies on high fertility sites could be overestimating any positive “silver-lining” effect of climate change on food production (Reilly and Schimmelpennig, 1999; Long et al., 2006; Leakey et al., 2012).

In this study elevated CO<sub>2</sub> increased biomass of *Festuca* in only ULT, ALF, and in the high-fertilty control. ULT and ALF are only present in relatively small areas of the world (Table 3), accounting for 17–25% of land area, depending on continent. Others have noted that CO<sub>2</sub> enrichment studies have predominantly reflected temperate biomes, which may respond differently than do tropical or arctic biomes (Leakey et al., 2012). Since *Festuca* did not germinate on the Gelisols we cannot speculate on the possible response of plants growing in regions from the tundra and other areas with permafrost soils. These frigid zones are the ones that could experience faster and greater impact of the expected temperature increase due to global change (IPCC, 2007a,b). While we do not claim that the samples utilized in this experiment represent the range of characteristics within each soil order, the soil orders that produced plants with small biomass and with reductions in their carbon assimilation under elevated CO<sub>2</sub>, occupy about 80% of the agricultural area in the world (including grasslands). This highlights the urgent need to better understand the real effects of climate change on plant growth across a representative range of soils, and its implications for food production and ecosystem management.

While we acknowledge the limitations in our study in the use of only one plant species, small pots, and the lack of water stress treatments, the differences between soils in the response to elevated CO<sub>2</sub> and the importance of soil variables in explaining these differences suggests that a more nuanced consideration of the consequences of soil variability on plant responses to global change may be warranted. The need for studies that address multiple climate change variables and their interactions, particularly with soil, has been noted by others (Lynch and St. Clair, 2004, 2010; Körner, 2006). This study suggests that the variability among soils might be quite large. This points to the need for a substantial research effort to better characterize the

interactions of the wide range of soil variability with global change variables. Such an effort should: (1) include not only elevated CO<sub>2</sub>, but water deficit and nitrogen deposition as well; (2) include multiple representatives for each soil order; (3) use intact soil monoliths to assess soils in a more natural context; (4) include more than one species of plant, to address the issue of plant adaptation to specific soils. Furthermore, the methodological challenge of accurately estimating root biomass and length for plants grown in widely differing soils needs to be solved.

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## SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <http://www.frontiersin.org/journal/10.3389/fpls.2015.00095/abstract>

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