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Harnessing MEMS sensors and statistics to unravel rock fracture

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This study focuses on obtaining differences in rock fracture surface morphology under various loading directions and speeds to infer rock damage mechanics by using micro-electro-mechanical system (MEMS) sensors, which can measure stress, strain, and displacement during loading accurately, providing detailed data for understanding the rock fracture mechanism for physics-informed statistics. Statistical variables analyze directional angle samples of the normal vector central line. The deviation normal distribution coefficient (DNDC) for rock fracture surface normal vectors is defined by the kurtosis coefficient. Brazilian splitting tests calculate the DNDC for Brazilian disk fracture surfaces. The variation in the DNDC with a measurement scale distinguishes morphological differences. Three results are obtained: the DNDC has a scale effect; loading the specimen in another direction before compression causes internal damage; and different loading speeds do not significantly change the DNDC. This research holds promise for a better understanding of rock fractures.

KEYWORDS

micro-electromechanical system, physics-informed statistics, rock mechanics, rock fracture surface, kurtosis coefficient, normal distribution

1 Introduction

Rock deformation and fracture commence with changes in nano/micro-scale morphology. Micro-electro-mechanical system (MEMS) sensors [1, 2] can be used to measure stress, strain, and displacement during the loading process, and they have ultra-sensitivity [3, 4]. This provides more detailed data on the behavior of rock and enables a more comprehensive understanding of the rock fracture mechanism. Additionally, MEMS sensors can be used to monitor the internal structure of rock in real-time, detecting any micro-cracks or damage that may occur before macroscopic fractures become visible [5]. This could assist in predicting rock failure and taking preventive measures to ensure safety in engineering applications.

Statistical methods can help determine the probability of rock fracture based on the sensor readings. For example, by analyzing the variance in stress measurements over time, it is possible to predict when a rock is approaching its breaking point. This combination of MEMS sensors and statistics produces a physics-informed statistical method.

In the process of Earth's crustal movement, such as earthquakes, landslides, and mudslides, most processes are accompanied by rock deformation and fracture. For example, in construction projects such as mining, civil engineering, slopes, bridges, and tunnels, these processes are accompanied by rock deformation and fracture. Rock, as a geological mineral of building materials, is increasingly receiving widespread attention from industry professionals, thus forming the discipline of rock mechanics. Rock mechanics mainly studies the stress and fracture conditions of rocks from a mechanical perspective. Nowadays, more and more researchers are starting from the morphology of rock fracture

surfaces to infer the mechanical mechanism of rock stress until fracture, also known as micro- and macro-rock mechanics.

In the study of rock fracture morphology, experts have proposed a series of methods, such as Mandelbrot's theory [6–10] and the two-scale fractal geometry [11–15]. Due to the highly irregular and complex fractal characteristics of rock fracture surfaces, researchers have proposed that fractal dimension can be used to characterize the roughness of rock fracture surfaces, and the morphology of rock fracture surfaces can be quantified by roughness [16–19]. Heping Xie, Jingan Wang, and others accurately characterized the relationship between the roughness of the rock fracture surface and its singular indexes using multi-fractal spectra and drew a graph of function changes [20]. Hongwei Zhou, H. P. Xie, and others used power spectral density and cumulative power spectral density to describe the anisotropy of the profile line on rock fracture surfaces [21]. T. Belem, F. Homand-Etienne M. Souley discovered the joint roughness coefficient of rock joint, which divides the roughness of rock joint surfaces into 10 levels [22]. V. Rasouli and J.P.Harrison proposed using the Reimann statistical method to estimate the roughness of a rock fracture section with linear profiles [23, 24]. M. Borri-Brunetto, A. Carpinteri, S. Invernizzi, and M. Paggi researched the micro-variation of rough surfaces under cyclic tangential loading [25]. M. Borri Brunetto, B. Chiaia, S. studied the multi-fractal properties of rough rock interfaces under pressure and the porosity of the contact domain [26]. A. Carpinteri, B. Chiaia, S. Invernizzi studied the direct fractal measure of fracture surfaces [27]. A. Carpinteri, Chiaia B, Invernizzi S. studied the three-dimensional fractal analysis of the microstructure morphology of concrete [28].

This paper proposes a new experimental method and uses mathematical statistics to study the morphological differences of rock fracture surfaces under different loading directions and loading rates and attempts to infer the mechanical mechanism and mechanism of rock fracture. This physics-informed statistics is an emerging and powerful approach that combines principles from physics and statistics.

2 Experimental design and operation

The experimental procedure is as follows: first, granite from Beishan, Gansu, was selected as the rock material as this material is relatively uniform. The ZS100 rock drilling machine used in engineering takes cylindrical rock cores with a diameter of 50 mm and a height of 120 mm and then cuts them into Brazilian disks with a diameter of 50 mm and a thickness of 20 mm using a cutting machine. The disks were polished into standard Brazilian disk specimens with a smooth surface using a polishing machine. The purpose of polishing was to ensure that the rock has a relatively uniform stress process as much as possible. Second, a mechanical test system (MTS) machine was used in the laboratory to load the disk vertically at different rates (as shown in Figure 1), with loading speeds of 0.01 mm/min, 0.1 mm/min, and 1 mm/min until the disk was fractured vertically. Moreover, the following experiments were taken. First, the MTS machine loaded the disk to 13 KN at a speed of 0.01 mm/min, which was approximately equal to 2/3 of the rock fracture threshold. Then, the upper and lower loading jaws of the MTS machine were released, the disk was rotated clockwise around the center by 30°, and it was



FIGURE 1
MTS machine loads the Brazilian disk.

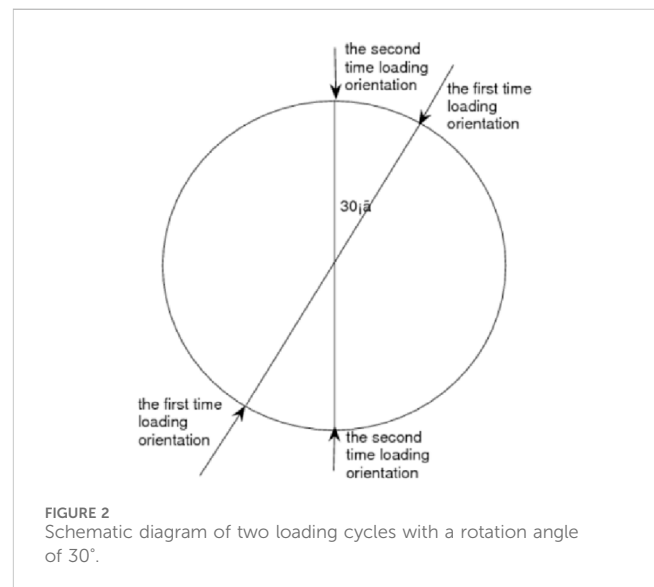
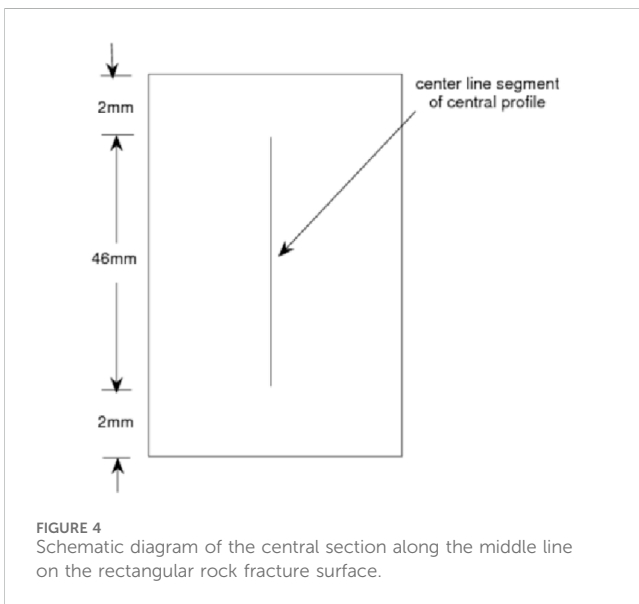
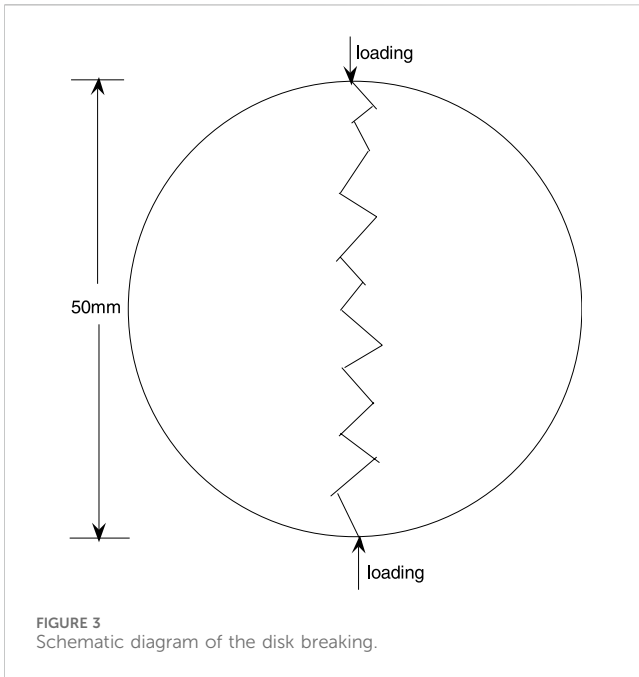


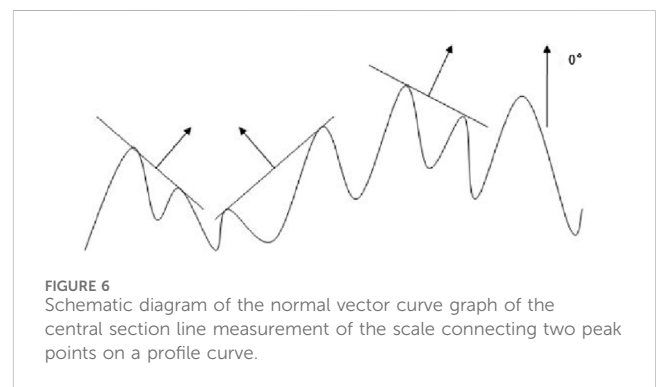
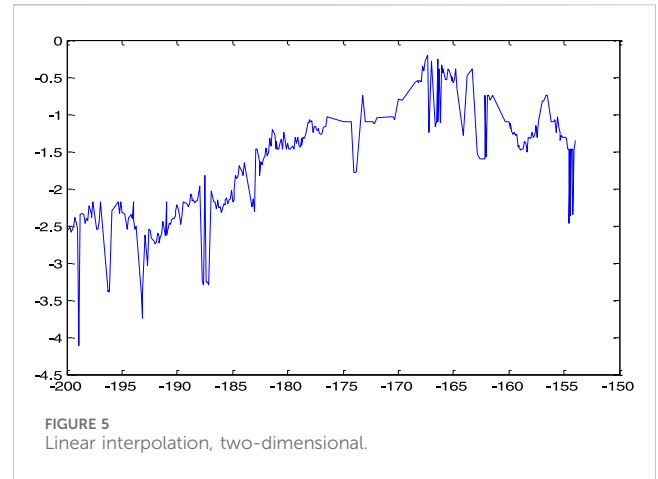
FIGURE 2
Schematic diagram of two loading cycles with a rotation angle of 30°.

loaded to fracture (as shown in Figure 2). Three sample specimens were prepared for each of the experiments for a total of 12 specimens. From the principle of rock mechanics, this experiment involved indirectly stretching the Brazilian disk specimen horizontally toward both sides until the disk fractured vertically from the middle (as shown in Figure 3). During the loading process, due to the brittle nature of the rock, the stress on the edges of the rock was uneven, and some specimens may have a small amount of rock debris falling off the edges. Therefore, relatively speaking, the stress on the inner layer of the rock was most uniform. In view of this, the 2-mm equidistant ends of the rectangular section were removed separately (as shown in Figure 4), and the center of the rock rectangular section was kept 46 mm long. This way, the measured data are most representative. Finally, a laser scanner was used to scan the fracture surface of the rock, with a spacing of 0.1 mm. A total of 461 sets of coordinate data were extracted. Before scanning, contrast imaging agents can be sprayed onto the fracture surface to prevent reflection of certain points on the fracture surface.



3 Data acquisition, analysis, and the purpose of researching

The coordinate data on a central section line were extracted along the loading direction on the rock fracture surface from the data obtained from the above scanning for research and analysis (as shown in Figure 4). In order to reduce errors, a linear interpolation method was used to fit the approximate two-dimensional curve image of this section line (as shown in Figure 5) because the error between the curve fitted by the linear interpolation method and the real curve was within the allowable range [29–31]. Two “peaks” with similar curve contours were connected using a straight line segment (hereafter referred to as the scale), the vector perpendicular to this



scale and pointing upward was taken as the normal vector (as shown in Figure 6), and the angle at which the normal vector deviates from the vertical upward vector was measured [23]. Here, it was agreed that the angle of the vector pointing vertically upward was zero degrees, the angle of the vector deviating to the left was negative, and the angle of the vector deviating to the right was positive. These statistical degree data are referred to as the direction angle of the normal vector of the center section line on the fracture surface of each specimen (herein referred to as the direction angle of the specimen). The distribution of directional angles of each specimen at different scales is researched below. Through computer program operation, it was shown that the measurement scale from 1.1 to 0.7 mm is a relatively effective transition range for the directional angle distribution of most specimens from accepting normal distribution to rejecting normal distribution. Therefore, the above measurements were carried out at scales of 1.1, 1.0, 0.9, 0.8, and 0.7 mm to obtain sample data on the directional angle of each specimen. Finally, the sample data obtained by statistical methods were processed to calculate the statistical measures of the directional angles of each specimen at different scales (herein referred to as the statistical measures of the specimens). The above design ideas aimed to achieve three objectives: first, to compare the variation patterns of corresponding statistical quantities of the same specimen at different measurement scales; second, to compare corresponding statistics between specimens that were loaded to 2/3 of the fracture strength threshold at a loading rate of 0.01 mm/min and then rotated at a 30°

angle before being crushed and specimens that were directly crushed without rotation; and finally, to reveal the effect and degree of the loading rate on the corresponding statistics of each specimen.

4 Using a statistical method to process the data on rock fracture surfaces

On one hand, we calculate the following statistics. Let X_1, X_2, \dots, X_n be a set of samples. Denote

$$\bar{X} = \frac{1}{n} \sum_{i=1}^n X_i \quad (1)$$

as the sample mean (herein referred to as the mean), which reflects the central tendency of the sample data. Denote

$$s^2 = \frac{1}{n-1} \sum_{i=1}^n (X_i - \bar{X})^2 \quad (2)$$

as sample variance (herein referred to as variance). Denote

$$s = \sqrt{\frac{1}{n-1} \sum_{i=1}^n (X_i - \bar{X})^2} \quad (3)$$

as sample standard variance (herein referred to as standard variance). Variance and standard deviation are statistical measures used to describe the deviation trend of sample data from the mean. Denote

$$g_1 = \frac{1}{s^3} \sum_{i=1}^n (X_i - \bar{X})^3 \quad (4)$$

as the skewness coefficient. Denote

$$g_2 = \frac{1}{s^4} \sum_{i=1}^n (X_i - \bar{X})^4 \quad (5)$$

as the kurtosis coefficient.

Skewness and kurtosis are two statistical measures of describing the shape of sample data. The skewness coefficient reflects the symmetry of the distribution of the sample data. It is called right skewed when $g_1 > 0$, which indicates that the data to the right of the mean are more scattered than the data to the left of the mean. Conversely, it is called left skewed when $g_1 < 0$, which indicates that the data to the left of the mean are more scattered than the data to the right of the mean. When g_1 approaches 0, it is called unbiased, and the distribution can be considered symmetric. Kurtosis reflects the normal distribution of the samples, where the kurtosis of the normal distribution is 3. When $g_2 > 3$, it indicates that there is a significant amount of data in the sample that are far from the mean. The distributional density function curve has a flatter shape than that of normal distribution. Conversely, when $g_2 < 3$, it indicates that there are a fewer data points in the sample that are far from the mean. The distributional density function curve has a steeper shape than that of normal distribution. Moreover, in the distribution function graph, the red dashed straight line represents the standard normal distribution function graph. The closer the distribution function graph is to the red dashed line, the closer the data distribution is to a normal distribution. On the contrary, if the distribution function graph deviates further from the red dashed line, it indicates that the data distribution deviates further from the normal distribution.

On the other hand, a hypothesis test is conducted on the normal distribution of the input direction angle data. The angle data here are subjected to Jarque–Bera testing, with a default significance level. In each data table, “ P ” represents the probability value of accepting the null hypothesis, and “ $Jbstat$ ” is the value of the test statistic. $CV = 5.9915$ is the threshold for rejecting the null hypothesis. “ H ” represents the testing result, and if H is equal to 0, it can be assumed that the data of the angle belong to the null hypothesis of normal distribution. If H is equal to 1, it cannot be assumed that the angle data follow the null hypothesis of normal distribution. In addition, there are also the same judgment methods consistent with the previously mentioned judgments. If $P < \alpha$, the null hypothesis of normal distribution can be rejected. Similarly, if $Jbstat > CV$, the null hypothesis of normal distribution can also be rejected [32, 33].

- (a) Process 1: Under the condition of a loading speed of 0.01 mm/min, first, the specimen is loaded to 13K N; the upper and lower jaws of the MTS machine are then released; and the specimen is then rotated clockwise at a 30° angle along the center of the disk, continuing to load it vertically until it fractures. Process 2: The loading speed remains at 0.01 mm/min, but the specimen did not rotate and is directly loaded until it fractures. The commonalities and differences in the statistical and directional angle distributions of each specimen in processes 1 and 2 are compared (reference Table 1).

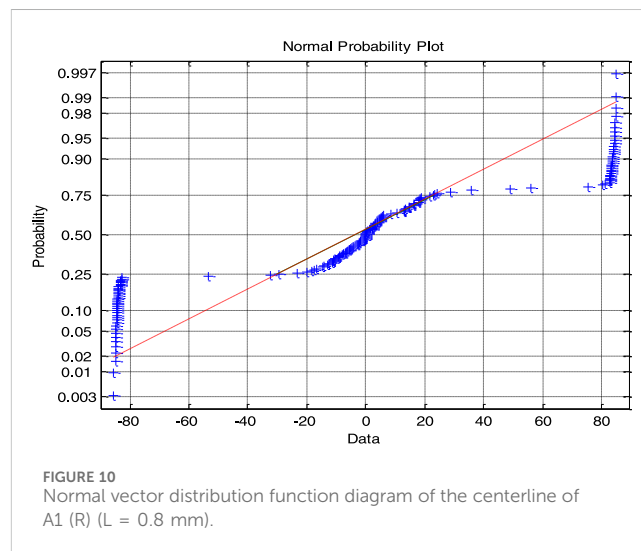
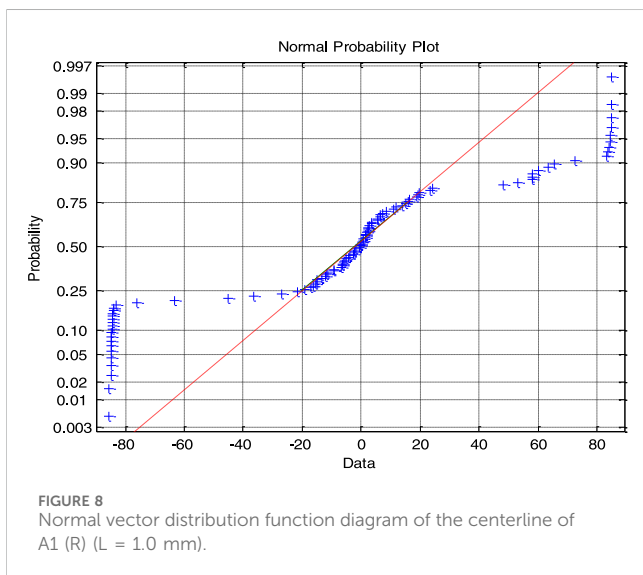
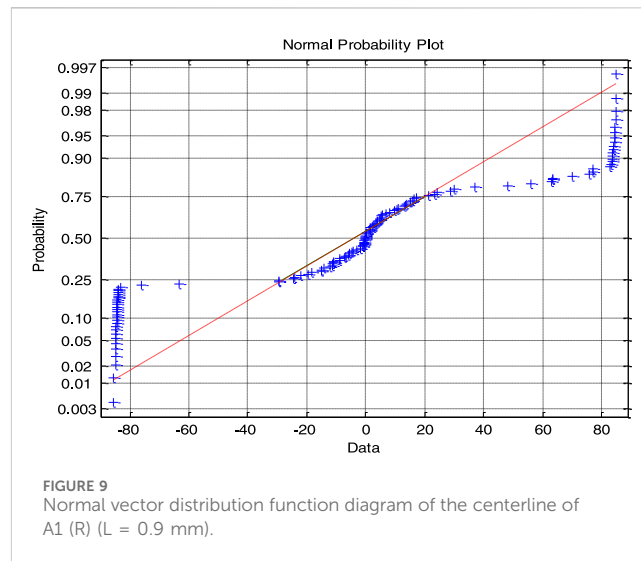
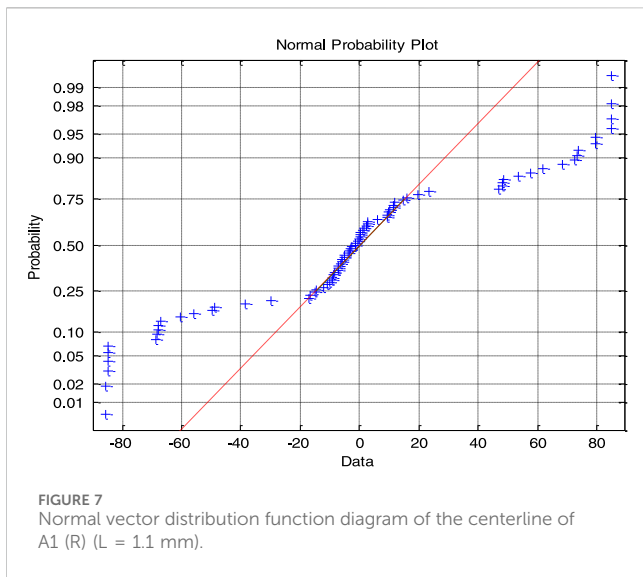
Table 1 shows that the measurement scale range is from 1.1 mm to 0.7 mm, and the commonality of the statistics between specimens rotated 30° and specimens not rotated is that as the measurement scale decreases, the distribution of directional angles of the specimens deviates more and more from the normal distribution. Therefore, if the measurement scale range is from 0.7 mm to 0.8 mm, the null hypothesis of normal distribution is completely denied. Taking sample specimen A1 (R) as an example, its probability distribution function diagrams are shown in Figures 7–11 within the scale range of 1.1 mm–0.7 mm. The figure shows that as the scale decreases, most of the points on the image deviate more from the “red dashed line,” indicating that the distribution of the normal vector deviates more from the normal distribution. Moreover, the variance and standard deviation further increase with the decrease in the measurement scale, indicating that as the scale decreases, the degree of deviation of the angle data from the mean increases. However, from the skewness coefficient being close to 0, it can be concluded that the degree of data dispersion on both sides of the mean is comparable, and other specimens also have similar commonalities. The difference is in statistical characteristics between specimens rotated 30° and those not rotated. From the variation patterns of H , P , $Jbstat$, and kurtosis coefficient values, it can be concluded that the distribution of directional angles for the three specimens rotated 30° can be accepted as a normal distribution when the measurement scale ranges from 1.1 mm to 0.9 mm, while the distribution of directional angles for the specimens not rotated can only be accepted as a normal distribution when the measurement scale is 1.1 mm. At the same scale, the average variance and standard deviation of the three specimens rotated at a 30° angle are smaller than those of the three specimens not rotated, indicating that the deviation of the direction angle of the specimens rotated is smaller than that of the specimens not rotated. The

TABLE 1 Statistical values of specimens rotated 30° and not rotated ("A" represents a loading speed of 0.01 mm/min. "1, 2, and 3" represent the specimen number. "R" represents specimens rotated 30°).

Test statistics Rock label	Scale (unit: mm)	H	P	J_{bstat}	Variance	Standard variance	Skewness coefficient	Kurtosis coefficient
A1 (R)	1.1	0	0.7657	0.5338	2,076.8	45.5719	-0.0115	2.6684
	1.0	0	0.4795	1.4698	2,525.5	50.2547	-0.0661	2.4697
	0.9	0	0.1533	3.7513	3,002.5	54.7946	-0.0330	2.1823
	0.8	1	0.0494	6.0165	3,292.5	57.3803	-0.0216	2.0827
	0.7	1	0.0360	6.6476	3,260.2	57.0980	0.0229	2.0682
A2 (R)	1.1	0	0.7673	0.5297	1,600.3	40.0043	-0.0782	2.9409
	1.0	0	0.9466	0.1098	2,144.7	46.3107	-0.0524	2.9205
	0.9	0	0.4888	1.4316	2,631.2	51.2956	-0.0371	2.4404
	0.8	0	0.0855	4.9174	3,241.8	56.9364	-0.0917	2.0747
	0.7	1	0.0392	6.4790	3,336.6	57.7632	0.0845	1.9986
A3 (R)	1.1	0	0.6456	0.8752	2,401.9	49.0096	-0.0305	2.5642
	1.0	0	0.5844	1.0744	2,543.8	50.4359	-0.0866	2.5637
	0.9	0	0.1023	4.5602	3,203.7	56.6016	-0.0586	2.1050
	0.8	1	0.0290	7.0822	3,373.4	58.0812	-0.0289	1.9665
	0.7	1	0.0110	9.0147	3,514.0	59.2791	0.0322	1.9209
A1	1.1	0	0.0938	4.7328	3,319.2	57.6121	0.0312	1.9323
	1.0	1	0.0080	9.6580	3,966.1	62.9768	-0.0327	1.6389
	0.9	1	0.0016	12.9376	4,191.0	64.7377	0.0085	1.5917
	0.8	1	4.1940e-005	20.1586	4,599.9	67.8225	-0.0072	1.4554
	0.7	1	1.3472e-006	27.0349	4,795.5	69.2498	-0.0039	1.4097
A2	1.1	0	0.0948	4.7119	3,452.4	58.7568	0.0234	1.9397
	1.0	1	0.0230	7.5465	3,853.2	62.0739	0.0230	1.7713
	0.9	1	0.0086	9.5047	3,911.7	62.5435	-0.0040	1.7348
	0.8	1	0.0051	10.5384	3,908.4	62.5174	0.0250	1.7420
	0.7	1	0.0039	11.0936	3,842.6	61.9888	0.0398	1.7682
A3	1.1	0	0.1329	4.0369	3,180.4	56.3954	-0.0088	1.9903
	1.0	1	0.0140	8.5375	3,818.1	61.7909	-0.0772	1.7362
	0.9	1	0.0056	10.3812	3,877.1	62.2662	-0.0884	1.7290
	0.8	1	9.8485e-004	13.8460	4,109.3	64.1038	-0.0660	1.6351
	0.7	1	3.1029e-004	16.1560	4,166.6	64.5488	-0.0503	1.6135

difference between the statistical values of the specimens in the above two experimental processes fully demonstrates the difference in statistical characteristics between specimens with a 30° rotation angle and specimens without a rotation angle. From the changes in the values of H , P , J_{bstat} , and kurtosis coefficient values, it can be concluded that the distribution of directional angles for the three

specimens with a 30° rotation angle can be accepted as a normal distribution when the measurement scale ranges from 1.1 mm to 0.9 mm, while the distribution of directional angles for specimens without a rotation angle can only be accepted as a normal distribution when the measurement scale is 1.1 mm. At the same scale, the average variance and standard deviation of the three



specimens rotated at a 30° angle are smaller than the average variance and standard deviation of the three specimens not rotated, indicating that the deviation of the direction angle of the specimens rotated is smaller than that of the specimens not rotated. The difference between the statistical values of the specimens in the above two experimental processes fully demonstrates that the 13-KN pressure experienced before the turning angle caused some damage to the interior of the rock specimen, thereby changing the directional angle distribution of the specimen.

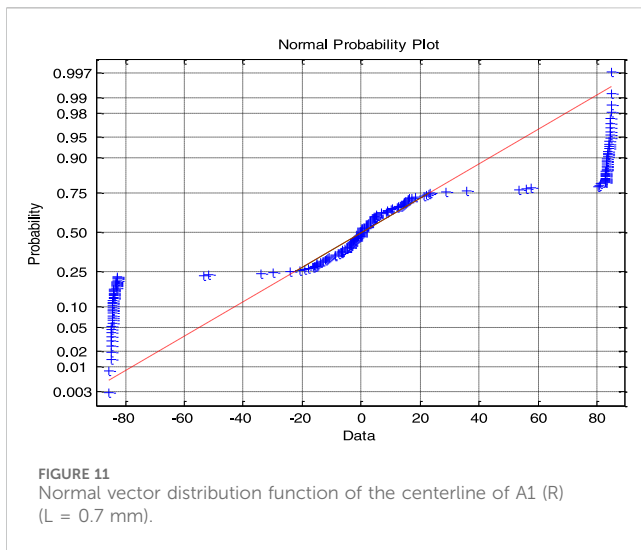
- (b) The differences in directional angle distribution between specimens with loading speeds 0.01 mm/m and 0.1 mm/m are shown in Tables 1, 2.

Compared with the statistical values of specimens A1, A2, and A3 in Table 1, the difference between them is that for the first velocity A (0.01 mm/m), the three specimens can accept a normal

distribution at a measurement scale of 1.1 mm, while for the second velocity B (0.1 mm/min), the variation patterns of H , P , J_{stat} , and kurtosis coefficient values indicate that the three specimens cannot accept a normal distribution at a measurement scale ranging from 1.1 mm to 0.7 mm. This indicates that these two loading speeds also caused a difference in the distribution of directional angles of the specimens, thus indicating that these two loading speeds caused a difference in the fracture of the specimens.

- (c) The relationship and differences between the statistical values of the directional angle samples of each specimen under loading rate 0.01 mm/m and 1 mm/m are compared (refer to Tables 1–3).

After comparison, the statistical value of the directional angle data on the specimens under the third velocity C is similar to the corresponding statistical value under the first velocity A. Both are



accepted as normal distributions at a measurement scale of 1.1 mm. However, overall, at a measurement scale of 1.1 mm, the average value of the statistical value P of the directional angles of the three specimens under the third velocity C is larger than that under the first velocity A. The average value of the test statistic J_{bstat} is smaller than the average value of J_{bstat} of the three specimens at the first

velocity A, and of course, all of them are less than the measurement threshold $CV = 5.9915$.

The mean value of the kurtosis coefficient is also higher than that of the first velocity A, indicating that under the condition of loading speed of 1 mm/min, the probability of the distribution of the specimen direction angle accepting the null hypothesis of normal distribution is greater than that under the condition of velocity of 0.01 mm/m. This indicates that there are differences in the fracture morphology of granite caused by these two different loading rates. On the other hand, under the conditions of a loading rate of 0.1 mm/m and a loading rate of 1 mm/m, comparing Tables 2, 3 showed that the faster the loading rate, the more deviated the degree distribution of the directional angle from the normal distribution. This is because under the condition of a loading rate of 1 mm/m, when the measurement scale is 1.1 mm, the distribution can accept the null hypothesis of normal distribution. However, when the loading rate is 0.1 mm/m, the null hypothesis of normal distribution cannot be accepted within the range of 1.1 mm–0.7 mm.

5 Skewness coefficient

Based on the above analysis, for a kurtosis coefficient of 3, the sample of the direction angle of the normal vector of the center section on the fracture surface of the specimen follows a standard

TABLE 2 Under the condition of loading speed 0.1 mm/m, statistical values of each specimen (“B” represents a loading speed of 0.01 mm/min. “1, 2, and 3” represent the specimen number).

Test statistics Rock label	Scale (unit: mm)	H	P	J_{bstat}	Variance	Standard variance	Skewness coefficient	Kurtosis coefficient
B1	1.1	1	0.0403	6.4232	3,597.0	59.9751	-0.0033	1.8392
	1.0	1	0.0090	9.4292	3,964.9	62.9676	-0.0076	1.7309
	0.9	1	0.0023	12.1815	4,081.8	63.8889	0.0259	1.6617
	0.8	1	0.0011	13.5585	4,078.0	63.8594	0.0460	1.6844
	0.7	1	4.0387e-004	15.6288	4,095.3	63.9944	0.0144	1.6638
B2	1.1	1	0.0218	7.6550	3,800.6	61.6490	0.0409	1.7693
	1.0	1	0.0015	13.0320	4,386.0	66.2267	-0.0131	1.6091
	0.9	1	4.9083e-004	15.2388	4,444.2	66.6648	-0.0030	1.5819
	0.8	1	1.6742e-004	17.3901	4,494.2	67.0391	-0.0105	1.5804
	0.7	1	1.6711e-004	17.3938	4,328.3	65.7901	-0.0182	1.5567
B3	1.1	1	0.0420	6.3414	3,671.8	60.5955	0.0662	1.8380
	1.0	1	0.0031	11.5568	4,266.8	65.3206	0.0595	1.6349
	0.9	1	0.0013	13.2818	4,198.4	64.7947	0.0808	1.6002
	0.8	1	3.6955e-004	15.8065	4,351.3	65.9645	0.0400	1.5949
	0.7	1	1.0420e-004	18.3384	4,378.0	66.1661	0.0332	1.5755

TABLE 3 Statistical values of directional angle data for each specimen under the condition of loading speed 1 mm/m ("C" represents a loading speed of 1 mm/m. "1, 2, and 3" represent specimen numbers).

Test statistics Rock label	Scale (unit: mm)	H	P	J_{bstat}	Variance	Standard variance	Skewness coefficient	Kurtosis coefficient
C1	1.1	0	0.3197	2.2808	2,675.8	51.7279	0.0166	2.1944
	1.0	1	0.0180	8.0355	3,854.7	62.0863	-0.0015	1.7296
	0.9	1	0.0015	13.0695	4,350.0	65.9544	0.0095	1.5675
	0.8	1	2.3934e-004	16.6753	4,489.2	67.0013	0.0057	1.5654
	0.7	1	2.1869e-004	16.8557	4,352.0	65.9697	-0.0452	1.5217
C2	1.1	0	0.2297	2.9417	2,964.3	54.4451	-0.0197	2.1967
	1.0	1	0.0380	6.5400	3,498.3	59.1464	0.0033	1.9184
	0.9	1	0.0124	8.7872	3,681.3	60.6736	0.0004	1.8348
	0.8	1	0.0018	12.6248	3,966.9	62.9837	-0.0132	1.7226
	0.7	1	8.4586e-004	14.1503	3,914.9	62.5692	-0.0384	1.7184
C3	1.1	0	0.0892	4.8343	3,442.3	58.6713	0.0589	1.9578
	1.0	1	0.0204	7.7876	3,822.7	61.8278	0.0821	1.9416
	0.9	1	0.0161	8.2527	3,779.7	61.4790	0.0492	1.8866
	0.8	1	0.0166	8.1959	3,690.2	60.7471	0.0418	1.8530
	0.7	1	0.0218	7.6503	3,539.6	59.4947	0.0281	1.8357

normal distribution (referred to as the sample of the direction angle of the specimen follows the standard normal distribution). Therefore, from the kurtosis coefficient of the direction angle of each specimen at different scales, the deviation normal distribution coefficient (DNDC) that varies with scale can be defined. The formula is defined as follows:

$$D(j) = (|K_j - 3|)/3, \quad (6)$$

where $D(j)$ is the skewness coefficient, K_j is the kurtosis coefficient, and the subscript j is equal to 1.1, 1.0, 0.9, 0.8, and 0.7 and its unit is millimeters (mm). The skewness coefficient reflects the relative error of the deviation of the direction angle of the specimen from the normal distribution at different scales from the perspective of the kurtosis coefficient. In particular, when K_j is equal to 3, the skewness coefficient $D(j)$ is equal to 0. The distribution of the sample strictly follows normal distribution, which is called unbiased.

The content below is an analysis of the difference in the skewness coefficient of the normal vector of the central section line on the fracture surface of the specimen under different loading methods (referred to as the skewness coefficient of the specimen).

- (i) The relationship between the skewness coefficient and scale of each specimen is revealed.
- (ii) The difference in the skewness coefficient between specimens rotated 30° and those not rotated at a loading rate 0.01 mm/m is revealed.

- (iii) The differences in the skewness coefficients of specimens under different loading rates are revealed (the data are given in Table 4).

The overall conclusions could be obtained from Table 4. First, for each specimen, its skewness coefficient increases as the scale decreases (refer to Figures 12, 13). Second, for two kinds of specimens at loading speed 0.01 mm/m and the specimens rotated 30° , the variation range of their skewness coefficient is concentrated from 0.1105 to 0.3597, and as the scale decreases, differences in their skewness coefficient become larger and larger. However, for the not-rotated-specimens, their skewness coefficient is concentrated between 0.3534 and 0.5301; so, from Figure 12 and Figure 13, overall, the skewness coefficient of specimens without rotating angles is greater than that of specimens with rotating 30° angles, which shows that when the specimen is loaded to about $2/3$ of the fracture threshold, the inside of the rock is indeed damaged. As mentioned above, the skewness coefficients of specimens with a loading rate of 0.01 mm/m are concentrated between 0.3534 and 0.5301, and its range is 0.1767. For the specimens of loading rate 0.1 mm/m, their skewness coefficients range from 0.3869 to 0.4811, with a range of 0.0942. For the specimens of loading rate 1 mm/m, their skewness coefficients are concentrated from 0.2678 to 0.4928, with a range of 0.225. From the range of skewness coefficients, the following conclusion can be obtained. Neither the faster the loading rate of the specimen, the greater the range of skewness coefficients, nor the slower the loading speed of the specimen, the larger the range of skewness coefficients (refer to Figure 13). The variable

TABLE 4 Skewness coefficient values between specimens rotating 30° and not rotated ("A" represents loading speed 0.01 mm/m. "1, 2, and 3" represent specimen numbers. "R" represents the specimens of rotating 30°).

Scale DNDC Rock	1.1 mm	1.0 mm	0.9 mm	0.8 mm	0.7 mm
A1 (R)	0.1105	0.1768	0.2726	0.3058	0.3106
A2 (R)	0.0197	0.0265	0.1865	0.3084	0.3338
A3 (R)	0.1453	0.1454	0.2983	0.3445	0.3597
A1	0.3559	0.4537	0.4694	0.5149	0.5301
A2	0.3534	0.4096	0.4217	0.4193	0.4106
A3	0.3366	0.4213	0.4237	0.4550	0.4622
B1	0.3869	0.4230	0.4461	0.4385	0.4454
B2	0.4102	0.4636	0.4727	0.4732	0.4811
B3	0.3873	0.4550	0.4666	0.4684	0.4748
C1	0.2685	0.4235	0.4775	0.4782	0.4928
C2	0.2678	0.3605	0.3884	0.4258	0.4272
C3	0.3474	0.3528	0.3711	0.3823	0.3881

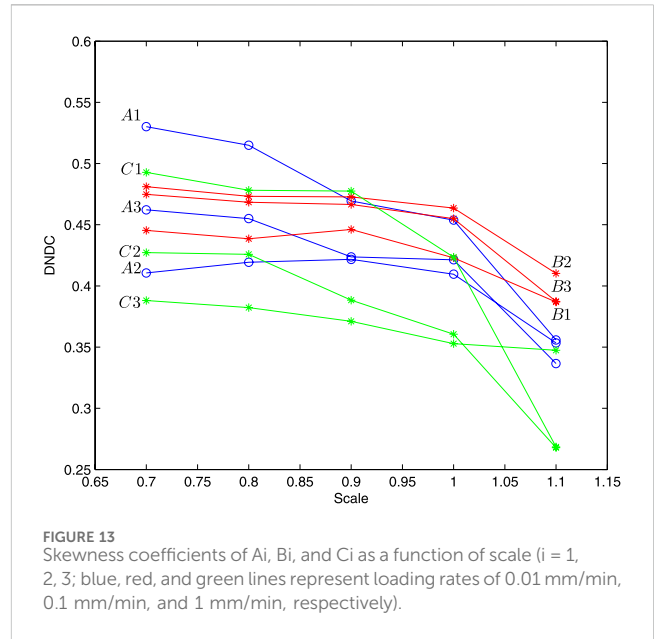


FIGURE 13 Skewness coefficients of Ai, Bi, and Ci as a function of scale (i = 1, 2, 3; blue, red, and green lines represent loading rates of 0.01 mm/min, 0.1 mm/min, and 1 mm/min, respectively).

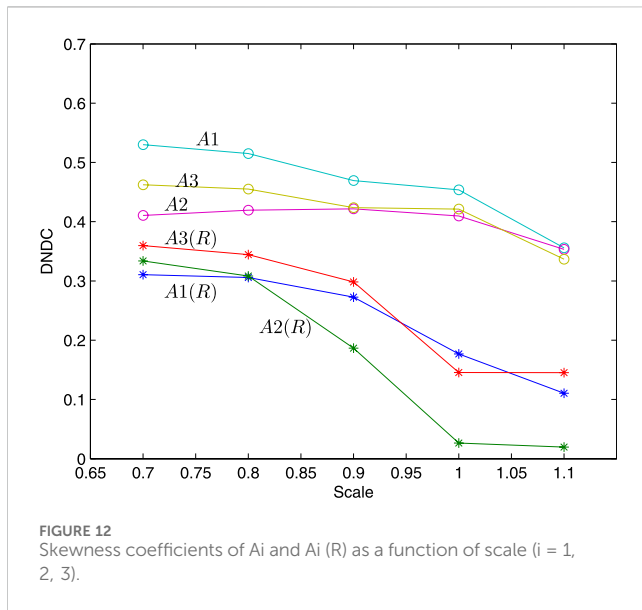


FIGURE 12 Skewness coefficients of Ai and Ai (R) as a function of scale (i = 1, 2, 3).

curves in Figure 12 and Figure 13 show that there is no significant change in the skewness coefficient under the above three loading rates. It indicates that the effects of the three loading rates on the rock fracture morphology of the specimens are not significantly different.

6 Conclusion and outlook

MEMS sensors play a crucial role in understanding rock fracture. Coupled with statistical analysis, they offer a powerful approach to studying this complex phenomenon.

MEMS sensors can be strategically placed on or within rocks to measure parameters such as stress, strain, and vibration. These real-time measurements provide valuable data on the behavior of rocks under different conditions. Meanwhile, statistics comes into play by analyzing these data to identify patterns and trends.

In practical applications, this approach can be used in mining and construction industries to assess the stability of rock formations and plan safe operations. It can also aid in geologic research to better understand the processes that lead to rock fractures. By leveraging MEMS sensors and statistics, we can gain a deeper understanding of rock fracture and take proactive measures to ensure safety and stability.

The following conclusion can be obtained from the above rock testing and normal hypothesis testing analysis. First, the distribution status of normal vectors of the central profile on rock fracture surfaces has scale effects from the variable regularity of statistics and skewness coefficient. It indicates that the smaller the measurement scale, the more the distribution deviates from the normal distribution. Second, under the loading speed 0.01 mm and the same scale, there is obviously a difference in the directional angle statistics and skewness coefficient of specimens rotating 30° and not rotating 30° angles. Overall, it is more likely that the distribution of the direction angle of the specimen rotated by 30° is accepted as a normal distribution than that of the specimen without rotation, indicating that the 13-KN pressure applied before rotation caused some damage to the interior of the specimen. Finally, the differences in the above three loading rates in rock fracture surfaces morphology are not obvious. If conditions permit, larger loading rates can be used to analyze the differences in rock fracture surface morphology and identify its variation regularity.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary Material; further inquiries can be directed to the corresponding author.

Author contributions

XP: data curation, funding acquisition, writing—original draft, writing—review and editing, and formal analysis. GD: investigation, software, and writing—review and editing.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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