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# An investigation on the effects of increasing maximum wind speed of tropical cyclones on the return periods of water levels in the sea area of the Yangtze River Delta

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This paper investigates the impact of increasing maximum wind speed of tropical cyclones on the return periods of water levels in the sea area of the Yangtze River Delta. To conduct this study, a series of numerical experiments are performed using historical tropical cyclones that impacted the Yangtze River Delta from 1949 to 2019. The aim is to analyze the effects of global climate change on extreme water levels and the corresponding return periods. To obtain the historical water levels in the sea areas of the Yangtze River Delta, a storm surge model is driven by the selected tropical cyclones. The simulated astronomical tidal levels during the same period are also used. The extreme water levels of different return periods are then calculated. The maximum wind speeds of the selected tropical cyclones are increased by 11% according to the expected amount of increase under global climate change. The extreme water levels of different return periods under this scenario are calculated with the same procedure. The results of the study show that the impact of increasing maximum wind speed of tropical cyclones on the increases of extreme water levels and the decrease of return periods is more significant in the inner area of the estuaries than in the outer areas. Moreover, the responses of the extreme water levels and the corresponding return periods in the Yangtze River Estuary and the Hangzhou Bay show different characteristics. The results of this study provide significant reference value for the management of future coastal disaster prevention and mitigation in the Yangtze River Delta. Furthermore, the methodology used in this study can be applied in other estuaries to investigate the potential impacts of changes in climate and hydrology factors on extreme water levels and the corresponding return periods.

### KEYWORDS

return period, extreme water levels, Yangtze River Delta, tropical cyclone, maximum wind speed

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## **1** Introduction

Under the background of global climate change, the maximum wind speed of tropical cyclones and extreme high tide level in coastal areas are likely to increase (IPCC, 2012), which pose a great threat to coastal and estuary areas (e.g., Pan et al., 2013; Pan et al., 2020; Li et al., 2022) and also impact the present return period system based on historical hydrological data. Although some soft shore protection technologies can be used to resist the coastal disasters to some degree (e.g., Temmerman et al., 2013; Pan et al., 2022a), it is necessary to understand the impacts of global climate change on the return period system for the purpose of future coastal disaster prevention and mitigation.

The Yangtze River Delta (YRD) is formed by two estuaries, the Yangtze River Estuary and the Hangzhou Bay (i.e., the Qiantangjiang River Estuary), as seen in Figure 1. It is one of the economic centers of China with dense population and developed commodity economy, and it is often affected by storm surges induced by tropical cyclones. Studies are conducted on the impacts of global climate change on the storm surge induced by tropical cyclones in the Yangtze River Delta (e.g., Wang et al., 2012; Zhao et al., 2014; Feng et al., 2018; Pan and Liu, 2019; Shen et al., 2019) and other estuaries (e.g., Karim and Mimura, 2008; Robins et al., 2016; Yin et al., 2017; Li et al., 2020; Yuan et al., 2022; Zhou et al., 2022). Most of these studies focus on the impacts of climate change on the distribution of extreme storm surge levels in the estuaries. With these studies, a portray of the changes in distribution of extreme storm surge level under climate change can be gained. However, the sea dike risk due to climate change is determined by both the changes in extreme storm surge levels and the original sea dike crest levels, which are determined by historical data and different alone the coast of the estuaries. Therefore, it is necessary to investigate the changes in return periods of water levels due to climate change, to provide reference for future planning of coastal and offshore projects such as sea levees (Pan et al., 2015a; Pan et al., 2015b), harbours (Gao et al., 2020; Wang et al., 2020; Gao et al., 2021), offshore wind power fields (Guan et al., 2019; He et al.,



FIGURE 1 The study area: the Yangtze River Delta.

2019), artificial beaches (Pan et al., 2017; Pan et al., 2018; Li et al., 2021; Li et al., 2022), etc.

In this paper, the effects of increasing maximum wind speed of tropical cyclones on the return periods of water levels in the sea area of the Yangtze River Delta are investigated via a serious of numerical experiments. Although global climate change comes with a serious of effects, e.g., sea level rise, increasing maximum wind speed of tropical cyclones, and changes in the tracks of tropical cyclones, this study only focus on the effects of increasing maximum wind speed of tropical cyclones. Under a hypothesis that the river run-off, astronomical tide, mean sea level, and tracks of tropical cyclones are unchanged, while the maximum wind speeds of tropical cyclones increase in a possible scenario (IPCC, 2012), the effects of increasing maximum wind speed of tropical cyclones on the return periods of water levels in the sea area of the Yangtze River Delta are investigated via a numerical model. Based on numerical simulations under different scenarios, the distributions of increase of water levels and decreases of return periods are analyzed and discussed.

# 2 Methodology

A storm surge model of the Yangtze River Estuary is set up based on the ADCIRC model (e.g., Dietrich et al., 2011) and verified via a comparison between the simulated and measured data. With the storm surge model, a two-step numerical study scheme is designed to investigate the possible impact of increasing maximum wind speed of tropical cyclones on the return periods of water levels in the sea area of the Yangtze River Delta. The numerical study scheme is designed based on the hypothesis that the river run-off, astronomical tide, mean sea level, and tracks of tropical cyclones are unchanged, while the maximum wind speeds of tropical cyclones increase in a possible scenario. In the first step, all the tropical cyclones that impact the Yangtze River Delta during 1949 to 2019 are picked out and used to drive the storm surge model. The astronomical tides during 1949 to 2019 are also simulated to provide the maximum annual astronomical tidal levels. With the simulated annual tidal levels (may from either storm surges or astronomical tides), the water levels of different return period are calculated in the sea area of the Yangtze River Delta. In the second step, all the processes in step 1 are repeated except that the tropical cyclones are strengthened according to IPCC (2012), and the water levels of different return periods are calculated accordingly. The new-calculated water levels of different return periods are then compared to the old ones and the impacts of increasing maximum wind speed of tropical cyclones on the spatial distribution of return periods of water levels in the sea area of the Yangtze River Delta are discussed.

## 2.1 Storm surge model

The storm surge model is set up with ADCIRC model. The ADCIRC model is a shallow-water circulation model that solves for water levels and currents at a range of scales (e.g., Westerink et al.,

2008; Dietrich et al., 2011). The ADCIRC model has been widely used in the numerical studies on storm surge processes around the world (e.g., Parker et al., 2019; Zhang et al., 2020; Shankar and Behera, 2021; Li et al., 2022; Pan et al., 2022b; Wang et al., 2022).

The computational grid and stations used in model verification and data analysis are given in Figure 2. The computational grid covers the sea area of the East China Sea, part of south China sea, and a small part of Pacific Ocean. The total number of computational nodes and elements are 60338 and 112455, respectively. The ocean boundary conditions are provided by the TPXO 7.2 database. In the verification cases, the upper boundary of the Yangtze River is adapted from the measurement at Datong (DT) station, while in the cases for numerical experiments the monthlyaveraged run-off discharges are used. Because the runoff of the Qiantang River has little influence on the flow field in Hangzhou Bay, the monthly-averaged runoff of Qiantang River is used for the upper boundary of the Qiantang River in both verification cases and the cases for numerical experiments. A modified Holland model (Holland, 1980; Mattocks and Forbes, 2008; Pan et al., 2016) is used to generate the wind fields according to tropical cyclone tracks. The equation of the wind speed is given as

$$v_{rot} = \sqrt{\frac{B}{\rho_a} \left(\frac{R_{max}}{r}\right)^B (P_n - P_c) exp\left(-\left(\frac{R_{max}}{r}\right)^B\right) + \left(\frac{rf}{2}\right)^2} - \frac{rf}{2} \quad (1)$$

$$B = \frac{((v_r - v_{mc})/W_{PBL})^2}{P_n - P_c}$$
(2)

where  $v_{rot}$  is the velocity of the rotating storm;  $P_n$  is the background surface pressure;  $P_c$  is the central surface pressure;  $\rho_a$  is the density of air; *f* is the Coriolis force; *B* is the hurricane shape parameter, which controls the eye diameter and steepness of the tangential velocity gradient;  $v_r$  is the maximum wind speed;  $v_{mc}$  is the moving speed of the storm;  $W_{PBL}$  is a wind reduction factor, defined at the gradient wind flow level above the influence of the planetary boundary layer (Powell et al., 2003).

The storm surge model is verified by simulating three storm surges that impact the sea area of the Yangtze River Delta significantly. The three simulated storm surges are induced by the tropical cyclones Winnie (9711), Fung-wong (1416) and Lekima (1909) whose tracks are illustrated in Figure 3. Comparisons of simulated and measured water levels at different stations during the three storm surges are plotted in Figure 4. As seen, good agreements can be found between model simulations and observations of water level, indicating that the storm surge model provides satisfactory descriptions of the water levels.

## 2.2 Numerical study scheme

The tropical cyclones that impact the sea area of the Yangtze River Delta from 1949 to 2019 are firstly picked out. To do this, a circle is made with the center located LCG and the radius of 500 km. All tropical cyclones that moved in this circle area during 1949 to 2019 are considered to have significant impact on the sea area of the





The tracks of all selected tropical cyclones that impact the sea area of the Yangtze River Delta and the three tropical cyclones used for model verification.

Yangtze River Delta and a total of 154 tropical cyclones are selected, as seen in Figure 3. The tropical cyclone track data are from the CMA (China Meteorological Administration) Tropical Cyclone Best Track Dataset (Ying et al., 2014; Lu et al., 2021; data available in tcdata.typhoon.org.cn). All the selected tropical cyclones are used to drive the storm surge model to get the water level distributions under the effects of historical tropical cyclones (the cases are referred to as historical cases hereinafter). In addition, the astronomical tides from 1949 to 2019 are also simulated to provide the maximum annual astronomical tidal levels, in case of the annual maximum tidal levels of some years are induced by the highest astronomical tide rather than storm surge (the cases are referred to as astronomical cases hereinafter).

Then the tropical cyclones are enhanced with increased maximum wind speeds and decreased central pressures. According to the special report of IPCC (2012) on extreme events and disasters, the mean maximum wind speed of tropical cyclones may increase 2% to 11% by the end of the 21st century. Based on this, the maximum wind speeds of all selected tropical cyclones are increased according to the upper limit (11%), and the central pressures are deceased accordingly with the empirical relationship

$$V_c = 1.176(1010 - P_c)^{0.87} + 6.5$$
(3)

where  $V_c$  is the maximum wind speed, m/s;  $P_c$  is the central pressure of a tropical cyclone, hPa. The empirical relationship is extracted according to the relationships between the maximum wind speed and the central pressure of the selected 154 tropical cyclones. It can be seen that the units on the two sides of Equation (1) are not the same. However, because the central pressure of a tropical cyclone has much less influence on the storm surge than the maximum wind speed (usually it is considered that the pressure-induced surge accounts for about 5% of the storm surge), the errors in the center pressure have little influence on the simulated results and a generally reasonable decrease of the central pressure meets the requirement of the numerical study scheme.

The enhanced tropical cyclones are then used to drive the storm surge model to get the water level distributions under the effects of increased maximum wind speed of tropical cyclones (the cases are referred to as enhanced cases hereinafter). The annual maximum tidal levels obtained based on the results of historical cases + the astronomical cases (namely historical scenario) can be labeled as Historical Annual Maximum Tidal Levels (HAMTL), while annual maximum tidal levels obtained based on the results of enhanced cases + the astronomical cases (namely enhance scenario) can be labeled as Enhanced Annual Maximum Tidal Levels (EAMTL).

The two set of annual maximum tidal levels can be used to extract return period information. A theoretical distribution should be chosen firstly. For two reasons, 1) this study focuses on the effects of enhanced tropical cyclones, and 2) the analysis is on all spatial points in the sea area of the Yangtze River Delta for the purpose of spatial comparison, the theoretical distribution needs to meet two requirements: 1) the theoretical distribution can be fitted by only part of the data (in this study the upper part is used to reflect more feature of the extreme high water levels), and 2) the curve-fitting need to be straightforward without manual adjustment. According to the two requirements, the two-parameter Weibull distribution is chosen rather than the Gumbel Probability Distribution (which cannot be fitted by part of the data) and Pearson III Probability Distribution (which has tuning parameters that need manual adjustment). The probability distribution function of twoparameter Weibull distribution is given by

$$P(X \le x) = 1 - exp\left(-\left(\frac{x}{a}\right)^b\right) \tag{4}$$

where P is the probability of the studied variable X being less than or equal to x, a is the scale factor, and b is the shape factor. Because the two-parameter Weibull distribution starts with zero, the X is the difference between the annual maximum tidal levels and the minimum value of the annual maximum tidal levels at each spatial point.

Some samples of the Weibull curve-fitting are given in Figure 5. The curves are fitted with the upper 25% of the annual maximum tidal levels to have a better reflection of the distribution of largest water levels. Using a different theoretical distribution or portion of data might yields different results in detailed values, but the general spatial patterns should be similar. As seen, the Weibull distribution fits well with both the HAMTL and EAMTL, and significant differences can be found between the best fit curves of HAMTL and EAMTL. It can be noticed that in Figure 5D) (station 4) the curve-fitting is moderate, with the middle and lower part of the data have a deviation from the curve. A possible explanation is that the station 4 is close to open sea area and it also influenced little by the Zhoushan islands compared to station 5, so the middle and lower part of the data are more induced by the astronomical tide rather than the storm surge. For all the spatial points in the sea area of the Yangtze River Delta, the Weibull fit is conducted to both the HAMTL and EAMTL. With the best fit curves, the water levels of different return periods under historical and enhanced scenarios can be obtained as the data base for further analysis.



FIGURE 4

(1909). The (A-F) respectively correspond to the water levels at YL, WS, ZP, KP and LCG during tropical cyclone Winnie (9711). The (F) and (G) respectively correspond to the water levels at YL and XLJ during tropical cyclone Fung-wong (1416). The (H) and (I) respectively correspond to the water levels at KP and ZP during tropical cyclone Lekima (1909). Note: because the measured data are from different sources, different stations are verified in the three storm surge processes.

# **3** Results and discussion

The differences in extreme water levels (50-year, 100-year, and 200-year return period) between historical and enhanced scenarios are calculated by subtracting the values from the historical scenarios from those from the enhanced scenarios. In this way, the impacts of the increasing maximum wind speed of tropical cyclones on the extreme water levels can be reflected straightforwardly. The distribution of the difference in extreme water levels are illustrated in Figure 6. As seen, under the impact of increased



maximum wind speed, the increase patterns of extreme water levels are majorly controlled by the estuary topography. Generally, the extreme water levels increase more significantly in the inner area than the outer area of the estuary. The increases of the extreme water levels of different return periods increase along the estuary from out sea to the upstream, which applies to both the Yangtze River Estuary and the Hangzhou Bay. A comparison between the two estuaries indicates that the extreme water levels of the Hangzhou Bay are more sensitive to the increase of maximum wind speed of tropical cyclones. At the upstream end of the Yangtze River Estuary (i.e. near the XLJ station in Figure 2) the increase of 200-year water level is about 0.15 m, while at the upstream end of the Hangzhou Bay (i.e. near the station 6 in Figure 2) the value is about 0.4 m. The reason might be the squeezing effect of the hornshape of the Hangzhou Bay to the storm surge. Noticing that the hypothesis of the numerical study scheme is that the river run-off, astronomical tide, mean sea level, and tracks of tropical cyclones are unchanged, while the maximum wind speeds of tropical cyclones are increased, strictly speaking the changes in extreme water levels are due to the maximum wind speeds of tropical cyclones rather than the climate change. However, the patterns of spatial distributions are of reference significance to the management of future coastal disaster prevention and mitigation.

Except for the increases of the extreme water levels, the decrease of return periods of extreme water levels due to increasing maximum wind speed is also a concerned issue and of great importance to the management of future coastal disaster prevention and mitigation. To calculate the decreases of return periods, the extreme water levels of certain return periods (i.e., 50, 100 and 200 years) are calculated firstly, and then the corresponding return periods of the extreme water levels of certain return periods are calculated based on the simulated results of the enhanced scenarios. The decreases of return periods are then



calculated by subtracting the new-calculated return periods (based on the enhanced scenarios) from the given ones (based on the historical scenarios) and plotted in Figure 7. It can be seen that the decrease of return periods and increase patterns of extreme water levels have generally similar distributions and some differences in the two estuaries. Under increased maximum wind speed, the return periods decrease more rapidly in the upstream of the estuaries generally. In the sea area out of the estuaries the return periods decrease about 50% (e.g., 25 years for 50-year return period case, 50 years for 100-year return period case, 110 years for 200-year return period case), while in the upper stream ends of the estuaries the return period case, 80 years for 100-year return period case, 170 years for 200-year return period case). A comparison between the two estuaries indicates that, on the contrary of the increase of extreme water levels, the decrease of return periods in the Yangtze River Estuary is more sensitive to increasing maximum wind speed than those of the Hangzhou Bay, the reason of which might be that the extreme water levels are higher in the Hangzhou Bay than those in the Yangtze River Estuary under the same return period. As discussed in preceding paragraph, due to the hypothesis of the study, the changes in extreme water levels are due to the maximum wind speeds of tropical cyclones rather than the climate change. but the patterns of spatial distributions are of reference significance to the management of future coastal disaster prevention and mitigation.

It should also be noticed that all the 154 historical tropical cyclones are strengthened according to the upper limit (11%) of the IPCC (2012) estimation, so the quantitative values of the increases of extreme water levels and the decrease of return periods might be overestimated, but the patterns of spatial distributions are of reference significance.



From the points of view of the changes in both extreme water levels and return periods, the inner areas of the estuaries will be under greater pressures from storm surges compared to the open sea areas, which applies to both the Yangtze River Estuary (rivertide dominated) and the Hangzhou Bay (tide dominated). Hence attentions show be paid on the disaster prevention of the inner areas of the estuaries under the background of increasing tropical cyclones.

# 4 Conclusions

Under the hypothesis that the river run-off, astronomical tide, mean sea level, and tracks of tropical cyclones that affect the Yangtze River Delta are unchanged, while the maximum wind speeds of tropical cyclones increase 11%, the impacts of increasing maximum wind speeds of tropical cyclones on the return periods of water levels in the sea area of the Yangtze River Delta are investigated *via* a series of numerical experiments. The following conclusions can be drawn:

(1) Under the scenario of increased maximum wind speeds of tropical cyclones, the extreme water levels under certain return periods (i.e., 50, 100 and 200 years) increase more significantly in the inner area than the outer area of the estuary. The maximum increases of extreme water level under 50, 100 and 200 years return period are about 0.3, 0.4 and 0.5 m, which appear at the upstream end of the Hangzhou Bay. The extreme water levels in the Hangzhou Bay seems more sensitive to the increase of maximum wind speeds of tropical cyclones, which might be due to the squeezing effect of the horn-shape of the Hangzhou Bay.

(2) Under the scenario of increased maximum wind speeds of tropical cyclones, the decrease of return periods ranges about 10 to

35 years for 50-year return period, 30 to 80 years for 100-year return period, and 70 to 170 years for 200-year return period in the Yangtze River Delta and the sea area out of it. The general patten of the decrease of return periods is similar to that of the increase pattern of extreme water levels, i.e., the changes in the inner areas of the estuaries are more significant than those in the sea areas out of the estuaries. However, the decrease of return periods is less significant in the Hangzhou Bay than that in the Yangtze River Estuaries, which might be due to that the extreme water levels are higher in the Hangzhou Bay themselves.

It should be mentioned that the historical tropical cyclones are strengthened according to the upper limit (11%) of the IPCC (2012) estimation, so the quantitative values of the increases of extreme water levels and the decrease of return periods might be overestimated, but the patterns of spatial distributions are of reference significance to the management of future coastal disaster prevention and mitigation. The methodology can also be used in other estuaries to investigate the potential impacts of changes in climate and hydrology factors on extreme water levels and the corresponding return periods.

## Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

## Author contributions

Conceptualization and research plan: YP, XC. Numerical Experiments: WL, JT, PY. Data analysis and writing: JT, PY, YC.

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# **Conflict of interest**

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