



The Importance of Mesozooplankton Diel Vertical Migration for Sustaining a Mesopelagic Food Web

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We used extensive ecological and biogeochemical measurements obtained from quasi-Lagrangian experiments during two California Current Ecosystem Long-Term Ecosystem Research cruises to analyze carbon fluxes between the epipelagic and mesopelagic zones using a linear inverse ecosystem model (LIEM). Measurement constraints on the model include ¹⁴C primary productivity, dilution-based microzooplankton grazing rates, gut pigment-based mesozooplankton grazing rates (on multiple zooplankton size classes), ²³⁴Th:²³⁸U disequilibrium and sediment trap measured carbon export, and metabolic requirements of micronekton, zooplankton, and bacteria. A likelihood approach (Markov Chain Monte Carlo) was used to estimate the resulting flow uncertainties from a sample of potential flux networks. Results highlight the importance of mesozooplankton active transport (i.e., diel vertical migration) in supplying the carbon demand of mesopelagic organisms and sequestering carbon dioxide from the atmosphere. In nine water parcels ranging from a coastal bloom to offshore oligotrophic conditions, mesozooplankton active transport accounted for 18–84% (median: 42%) of the total carbon transfer to the mesopelagic, with gravitational settling of POC (12–55%; median: 37%), and subduction (2–32%; median: 14%) providing the majority of the remainder. Vertically migrating zooplankton contributed to downward carbon flux through respiration and excretion at depth and via mortality losses to predatory zooplankton and mesopelagic fish (e.g., myctophids and gonostomatids). Sensitivity analyses showed that the results of the LIEM were robust to changes in nekton metabolic demand, rates of bacterial production, and mesozooplankton gross growth efficiency. This analysis suggests that prior estimates of zooplankton active transport based on conservative estimates of standard (rather than active) metabolism are likely too low.

Keywords: biological carbon pump, export production, DVM, LIEM, active transport, inverse model, carbon export, ecosystem model

INTRODUCTION

Although mesopelagic food webs are believed to depend entirely on productivity generated in the euphotic zone, reconciling mesopelagic metabolic demand with estimates of export has been challenging (del Giorgio and Duarte, 2002; Steinberg et al., 2008; Burd et al., 2010; Henson et al., 2011; Hannides et al., 2015). Due to large uncertainties in rate measurements for

meso- and bathypelagic organisms as well as low sampling resolution, steady-state budgets must either report wide ranges or otherwise exclude some processes, such as mortality and defecation of diel vertical migrators at depth. Even among recent studies, global carbon export budgets have been highly variable (Boyd and Trull, 2007; Henson et al., 2011, 2015; Laws et al., 2011; Siegel et al., 2014). Compounding this issue, several analyses have reported carbon demands by mesopelagic bacteria alone that exceed calculated carbon export (Ducklow and Harris, 1993; Burd et al., 2010), sometimes by an order of magnitude (Steinberg et al., 2008). This apparent imbalance between carbon supply to the mesopelagic and estimated metabolic demand suggests either that export estimates fail to capture important dynamics or that metabolic calculations are highly biased (Burd et al., 2010).

Some work has demonstrated that diel vertical migrators are important for net transfer of organic carbon from the euphotic zone to the mesopelagic, a transfer not measured with traditional carbon export methods (Morales, 1999; Steinberg et al., 2000). Since export by mesozooplankton is not captured by sediment traps or radioisotope disequilibria methods, we must rely on net tows coupled to assumptions about *in situ* respiration rates, or on indirect modeling syntheses. For example, using remote sensing fields and a size-structured ecosystem model, Archibald et al. (2019) found that global zooplankton diel vertical migration (DVM) can increase export production by 14% annually. This is consistent with previous modeling exercises based on zooplankton behavior (Bianchi et al., 2013) and community size structure (Aumont et al., 2018). Zooplankton behavior models argue that for DVM to be evolutionarily advantageous (Cohen and Forward, 2009), the energy expenditure should be offset by a commensurate reduction in predation risk. Using this modeled-behavior approach, Hansen and Visser (2016) found that 16–30% mid-latitude export production in the North Atlantic was likely due to DVM mesozooplankton. Each of these models note sensitivities to zooplankton biomass and the fraction of the zooplankton population that undergoes DVM, which are ecosystem metrics that are difficult to generalize.

Linear inverse ecosystem models (LIEM) have been shown to be a versatile and robust framework for integrating a wide range of ecosystem data (Vézina et al., 1988; Gontikaki et al., 2011; van Oevelen et al., 2012; Sailley et al., 2013; Stukel et al., 2018b). A LIEM combines an ecosystem network with observations and generalized constraints to determine possible energy flows through the ecosystem. Unlike a forward model (e.g., an NPZ model; Franks, 2002), the relationships between organisms are not prescribed by functional responses of model state variables (e.g., assuming a Monod functional form controls phytoplankton nutrient uptake responses or an Ivlev grazing formulation). Instead, the model includes all possible combinations of fluxes that are compatible with the assumed model structure and input constraints. The most likely ecosystem structure is then retrieved based on a random walk through the solution space (van den Meersche et al., 2009). This inverted approach has the advantage of not requiring *a priori* assumptions of functional ecological responses but instead relies on an assumed basic ecosystem structure (i.e., which functional groups should be included and

who eats whom) and many independent constraints on the food web.

The California Current Ecosystem (CCE) is an eastern boundary current upwelling biome with extensive temporal and spatial variability. As a result of high mesozooplankton biomass and strong DVM (Stukel et al., 2013; Powell and Ohman, 2015; Ohman and Romagnan, 2016), we expect a substantial contribution to export production by diel vertical migrators and a commensurately important role in satisfying the mesopelagic carbon demand. Stukel et al. (2013) suggested that active transport could be responsible for 1.8–29% of total export in the CCE. However, their study focused only on active transport fluxes due to zooplankton respiration and only included basal metabolism. To more thoroughly investigate the potential importance of active transport, we designed a two-layer LIEM, which includes non-living organic matter, primary producers, zooplankton, and planktivorous nekton organized into two layers: an epipelagic and a mesopelagic ecosystem. Using extensive data from two cruises of the CCE Long-Term Ecological Research (LTER) Program in the southern California Current region, our LIEM data synthesis suggests that active transport of carbon from the epipelagic down to depth is a significant mechanism supporting the mesopelagic carbon demand. Although previous studies have indicated that active transport may be responsible for 10–30% of total carbon flux (Yeber et al., 2005; Bianchi et al., 2013; Hansen and Visser, 2016; Aumont et al., 2018; Archibald et al., 2019), our LIEM suggests that 20–80% of carbon export in the CCE can be attributed to mesozooplankton DVM.

MATERIALS AND METHODS

Ecosystem Data

The data presented here (Appendix A) were collected during two cruises of the California Current Ecosystem Long Term Ecological Research (CCE LTER) program (P0704 in April 2007; P0810 in Oct. 2008). On these cruises, *in situ* drift arrays were used for quasi-Lagrangian tracking of water parcels for periods of 3–5 days (Landry et al., 2009, 2012), while the water column was repeatedly sampled for the following variables: CTD-derived physical data, phytoplankton diversity and biomass (flow cytometry, epifluorescence microscopy, and pigment analyses, Taylor et al., 2012), primary production ($H^{14}CO_3$ -uptake, Morrow et al., 2018), mesozooplankton biomass and community analyses (paired day-night bongo and Multiple Opening and Closing Net with Environmental Sampling System, MOCNESS net tows, Ohman et al., 2012; Powell and Ohman, 2012), microzooplankton biomass (epifluorescence microscopy), microzooplankton grazing (dilution method, Landry et al., 2009), mesozooplankton grazing (gut pigment methods, Landry et al., 2009), meso- and epipelagic micronekton biomass and metabolic demands (see section Phytoplankton, Bacteria, and Protist Constraints; Oozeki net tows, multi-frequency EK60 echosounder, and individual-based metabolic model, Davison et al., 2013, 2015), bacterial production (3H -leucine uptake, Samo et al., 2012), and gravitational particle export (sediment traps and ^{234}Th : ^{238}U disequilibrium, Stukel et al.,

2013). The use of a quasi-Lagrangian sampling framework also allowed us to assess net rates of change of phytoplankton biomass. Bulk rates and associated errors for the 3–5 day cycles were calculated by averaging vertically integrated rates or biomasses for each experimental cycle. The data and detailed methods can be found on the CCE LTER Datazoo website (<http://oceaninformatics.ucsd.edu/datazoo/data/ccelter/datasets>) and/or in published manuscripts cited above.

The quasi-Lagrangian experiments (hereafter “cycles” of repeated measurements in the same water parcel) spanned much of the physical, chemical, and ecological variability of the CCE domain (**Table 1**, **Figure 1**) which allowed us to classify cycles according to nutrient conditions, the primary driver of ecosystem variability within the CCE (Landry et al., 2012). Cycle classification was defined as: nutrient-limited cycles which were conducted in off-shore, low nutrient regions (P0704-2, P0810-2, P0810-6); transition region cycles which were characterized by low surface nutrient concentrations and intermediate NPP and biomass (P0810-1, P0810-3, P0810-4); and upwelling cycles in which surface nutrient concentrations and phytoplankton growth rates were highest (P0704-1, P0704-4, P0810-5; **Table 1**).

Phytoplankton, Bacteria, and Protist Constraints

Daily *in situ* primary productivity measurements using $H^{14}CO_3^-$ uptake (^{14}CPP) were conducted at 6–8 depths spanning the euphotic zone using 4 L incubations subsampled in triplicate (Morrow et al., 2018). A 250 mL dark bottle was used to correct for non-photosynthetic ^{14}C uptake. Contemporaneously, *in situ* dilution experiments, using the two-treatment approach of Landry et al. (2008), were conducted to measure protistan zooplankton grazing rates and chlorophyll-*a* growth rates (Landry et al., 2009). Chlorophyll to carbon ratios were determined by the ratio of vertically integrated chlorophyll-*a* growth rates and ^{14}CPP . Euphotic zone primary production and protistan zooplankton grazing rates were vertically integrated and averaged by cycle.

Rates of 3H -leucine incorporation into bacteria were measured in triplicate at multiple depths during each cycle (Samo et al., 2012). Each profile was vertically integrated and

then averaged by cycle in order to determine production rates of epipelagic bacteria. Additionally, upper and lower bounds for mesopelagic bacterial production were calculated by integrating bacterial production attenuation curves and scaling by the epipelagic bacterial production (Equation 1).

$$\text{Mesopelagic BP} = BP_{100} \int_{100}^{450} \left(\frac{z}{z_0}\right)^{-\alpha} dz \quad (1)$$

where BP_{100} is the measured BP rate at 100 m and α (BP attenuation factor) = 1.47 (Yokokawa et al., 2013) for the lower limit and $\alpha = 0$ (i.e., no attenuation) for the upper limit.

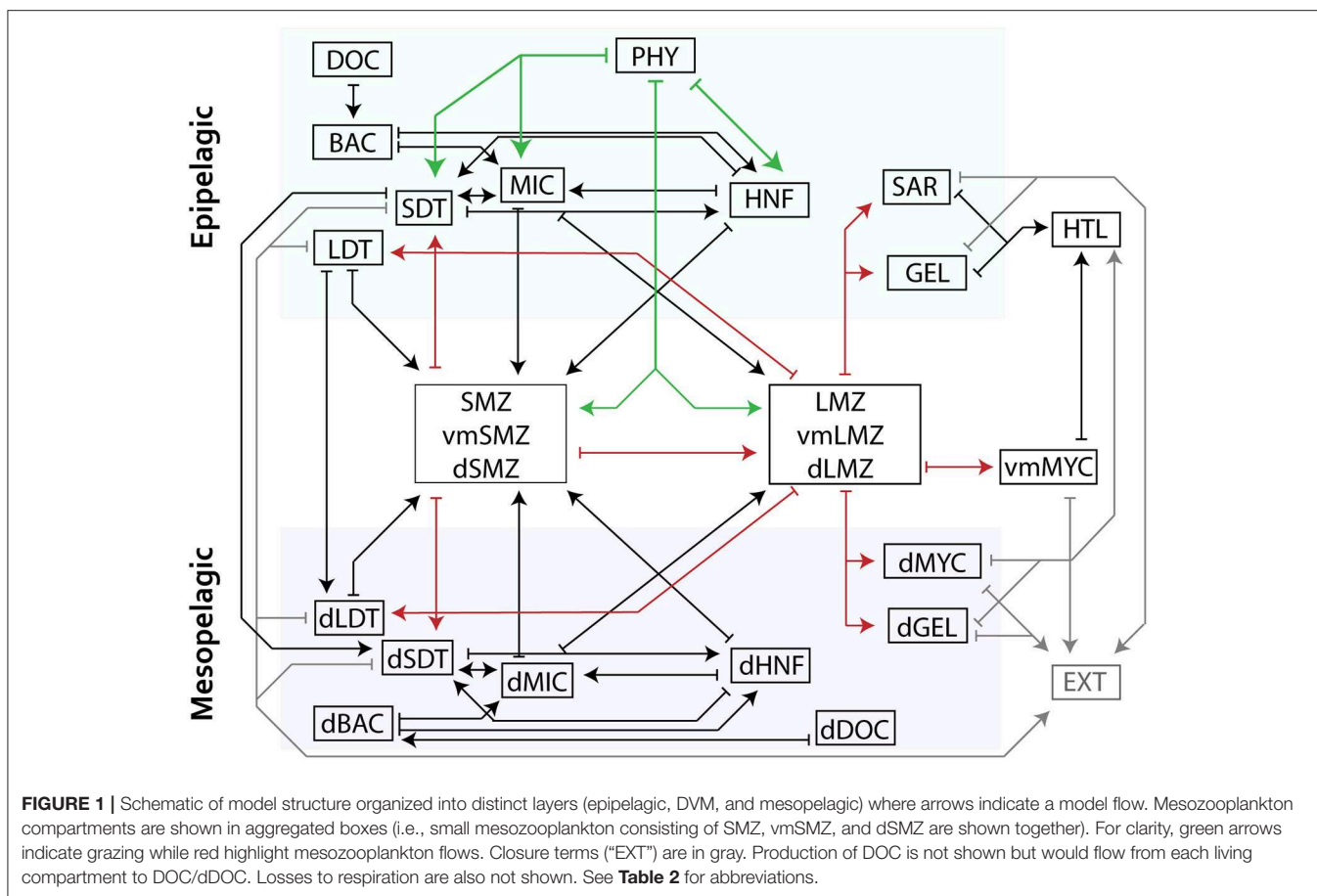
Mesozooplankton and Nekton Constraints

Data for the mesozooplankton constraints comes primarily from day-night paired oblique bongo net tows through the epipelagic (for grazing rates) or day-night paired 202 μm mesh MOCNESS tows taken at 9 depth horizons spanning the upper 450 m (for biomass and metabolism estimates). MOCNESS samples were analyzed by ZooScan digital scanner (Gorsky et al., 2010; Ohman et al., 2012), vignettes provisionally classified using machine learning methods, then 100% manually validated. Organisms were sorted (Stukel et al., 2013) into groups including euphausiids, nauplii, copepods, appendicularians, siphonophores, and “other crustaceans”. For this study, we separated the mesozooplankton community into two size classes (i.e. <1 and >1 mm ESD) of grazers and one compartment for gelatinous predators (siphonophores). We also partitioned the large and small mesozooplankton into non-vertically migrating epipelagic residents, vertical-migrators, or mesopelagic resident communities. Biomass estimates of non-migrating epipelagic mesozooplankton were calculated from day time net tows in the upper 100 m, while the non-migrating, mesopelagic biomass was calculated based on nighttime mesopelagic (100–450 m) net tows (Stukel et al., 2013). We note that epipelagic estimates are likely conservative due to net avoidance. Biomass estimates for the DVM mesozooplankton were calculated by averaging the difference in the night and day epipelagic biomass estimates with the difference in the day and night mesopelagic biomass estimates. This approach was used in order to be the most consistent with both the epipelagic and mesopelagic biomass estimates for non-vertically migrating biomass. For a list of abbreviations used for all model compartments, see **Table 2**.

Minimum respiration estimates for each mesozooplankton group were calculated using published temperature-length-basal respiration relationships (Ikeda et al., 2001). Oxygen consumption was converted to carbon units using the scale factor $9.88 \times 10^{-3} \text{ mg C d}^{-1} (\mu\text{L O}_2 \text{ h}^{-1})^{-1}$. Mesozooplankton grazing on phytoplankton was calculated from gut pigment contents of oblique bongo net tow tows (202 μm mesh, $D = 0.71 \text{ m}$) and estimated gut passage rates (Dam and Peterson, 1988). Carbon-based grazing rates were then calculated from chlorophyll (Chl) consumption, and C:Chl ratios computed as the ratio of NPP to chlorophyll-specific growth rates obtained from the dilution experiments. Mesozooplankton grazing rates were size fractionated as above. Mesozooplankton gut contents

TABLE 1 | Overview of conditions for each cycle along with the attributed classifications: upwelling, transition region, and nutrient limited.

Cycle	Classification	Surface Chl ($\mu\text{g Chl a L}^{-1}$)	^{14}C Primary productivity ($\text{mg C m}^{-2} \text{ d}^{-1}$)	Mesozooplankton biomass (mg C m^{-2})
P0704-1	Upwelling	1.35	1,233	2,695
P0704-2	Nutrient limited	0.22	587	391
P0704-4	Upwelling	0.99	2,314	1,715
P0810-1	Transition region	0.45	554	740
P0810-2	Nutrient limited	0.20	484	528
P0810-3	Transition region	0.72	892	923
P0810-4	Transition region	1.05	674	832
P0810-5	Upwelling	1.47	1,672	1,098
P0810-6	Nutrient limited	0.22	325	628



samples were improperly frozen for P0810-5, P0810-6, and most of P0810-4. In order to provide estimates for these grazing rates, average grazing rates from the cycle with the same classification were used (e.g., P0810-5 was an upwelling cycle so grazing rates were averaged from the other upwelling cycles). Conservative uncertainty estimates were set to be 2x the error calculated by propagation of error. This higher level of uncertainty is a reasonable compromise given the data limitations. For additional details on gut pigment processing, see Landry et al. (2009).

Nekton biomass was estimated based on catches made by a 5 m² Matsuda-Oozeki-Hu net trawl (Davison et al., 2013). For each station, epipelagic net tows were conducted at night after the ascent of the deep scattering layer. Preserved specimens from each net tow were identified to species and measured. Fish were classified as either non-vertical migrating or vertically migrating based on species. An individual based model was then used to determine metabolic rates and requirements for each nekton population: resident epipelagic, diel vertical migrant, and resident mesopelagic (Davison et al., 2013).

Export Production

VERTEX-style sediment traps consisting of 8–12 tubes per depth were deployed and recovered at the start and end of each cycle (Knauer et al., 1979; Stukel et al., 2013). Tubes were filled with a hypersaline, poisoned brine solution. Upon recovery >200- μ m swimming mesozooplankton taxa were manually removed

during inspection under a stereomicroscope. Samples for C and N or C:²³⁴Th ratios were filtered through pre-combusted glass fiber and quartz filters, respectively, prior to analysis on a CHN elemental analyzer or a RISO beta multi-counter.

²³⁴Th:²³⁸U disequilibrium measurements were made at 12 depths spanning the upper 200 m at the start and end of each cycle using standard small-volume procedures (Benitez-Nelson et al., 2001; Pike et al., 2005). Thorium-234 export rates were then computed using a 1-box steady state model (Savoye et al., 2006). The C:²³⁴Th ratio measured from sediment trap particles was used to convert to carbon export. For additional details, see Stukel et al. (2019).

Subduction of POC provides an alternative mechanism for the export of organic matter to the mesopelagic, that is not measured by either sediment traps or ²³⁴Th profiles, which only record gravitational settling of particles. A three-dimensional particle advection model was used to determine a range of possible subduction rates (Stukel et al., 2018c). The maximum and minimum estimates of particle subduction were used as bounds on two size-fractionated subduction flows within the LIEM.

Linear Inverse Model

We developed a LIEM for the CCE to investigate mechanisms of epipelagic-mesopelagic coupling. The LIEM consists of 140 flows (i.e., ecosystem fluxes, **Supplemental Table 2**)

TABLE 2 | Names and abbreviations of all model compartments.

	Epipelagic abbreviation	Name	Mesopelagic abbreviation	
Organisms	PHY	Phytoplankton		
	HNF	Heterotrophic Nanoflagellates	dHNF	
	MIC	Microzooplankton	dMIC	
	SMZ, vmSMZ	Small Mesozooplankton	dSMZ, vmSMZ	
	LMZ, vmLMZ	Large Mesozooplankton	dLMZ, vmLMZ	
	SAR		Sardines and other planktivorous fish	
			Non-DVM Myctophids	dMYC
	GEL	Gelatinous Predators	dGEL	
	vmMYC	Vertically Migrating Myctophids	vmMYC	
	POC& DOC	BAC	Bacteria	dBAC
SDT		Small Detritus	dSDT	
LDT		Large Detritus	dLDT	
	DOC	Dissolved Organic Matter	dDOC	
Closures	HTL	Higher Trophic Levels	HTL	
	RES	Respiration	dRES	
	EXT	Fecal Matter & External	EXT	

An abbreviation in the left column indicates inclusion in the epipelagic, while an abbreviation in the right column indicates inclusion in the mesopelagic. Each abbreviation is a distinct compartment in the LIEM with the prefix “vm” signifying vertical migration and “d” signifying the mesopelagic.

and 24 compartments (i.e., standing stocks; **Table 2**) organized into two layers: the surface epipelagic and a deeper mesopelagic ecosystem (defined as 100–450 m depth to match with *in situ* measurements). The epipelagic and mesopelagic ecosystems consist of 73 flows and 64 flows, respectively, with four explicit flows (particle sinking and subduction) and three implicit flows (active transport) linking the two layers (**Figure 1**). Three vertically migrating compartments (small and large mesozooplankton and nekton) connect the epipelagic and mesopelagic through a transfer associated with DVM (i.e., respiration, excretion, and mortality). Constraints consist of 24 mass balance equations, 18 approximate equations (i.e., *in situ* rate measurements) and 133 inequalities, which are provided in the **Supplement Model Spreadsheet**.

The 18 approximate equations are ecosystem observations, which can be directly compared to flows within the model (**Appendix A**). These equations are net primary productivity (NPP), phytoplankton biomass net rate of change, protistan grazing, size-fractionated grazing rates (<1 and >1-mm) for epipelagic resident and DVM mesozooplankton, sediment trap and ²³⁴Th-based export fluxes, bacterial production, and mesopelagic fish respiration, mortality and fecal pellet production rates. The model was provided an estimated value and associated uncertainty for each measurement.

Respiration, mesopelagic export, nekton fecal pellets, and losses to higher trophic levels were included as closure terms. Within the model, every organism loses carbon to respiration, DOC excretion, and defecation or mortality to detritus/fecal pellets. Grazing was allowed between organisms whose ecological roles and size ranges permit grazing (e.g., small mesozooplankton graze on nano- and microplankton; sardines consume only >1-mm mesozooplankton). Mass balance was required for each compartment. All compartments were assumed to be at steady state except for PHY, for which changes in biomass were measured (via Chl-*a* proxy) during each cycle and incorporated into the model. This flexibility was essential to capture the bloom phase of the ecosystem since dramatic shifts in Chl-*a* were observed during some cycles.

Inequality Constraints

The formulas used in the inequality constraints are provided in the **Supplement**. Upper and lower limit estimates of POC subduction from the epipelagic to the mesopelagic layer were taken from Stukel et al. (2018c), and minimum fecal pellet fluxes were assigned based on the assumption that recognizable fecal pellets in sediment trap material represented a lower limit on total fecal pellet flux. Minimum and maximum Gross Growth Efficiencies (GGE) were assigned according to previously accepted literature values: 10–40% GGE for protistan zooplankton (HNF & MIC) and gelatinous predators (Straile, 1997); 10–30% for mesozooplankton (Anderson et al., 2018); and 5–30% for bacteria (del Giorgio and Cole, 1998). The Absorption Efficiencies (AE) for all heterotrophs were limited to 50–90% (Conover, 1966).

Minimum respiration requirements were considered as both active respiration and basal respiration. Active respiration was set as a fraction of ingestion, and basal respiration was set as a function of biomass and temperature. Valid solutions fulfilled both criteria. Diel vertical migrator biomass, as determined from MOCNESS net tows, was used to calculate a minimum respiration based on temperature. DOC excretion was required to be >10% of ingestion (or 2% of NPP for phytoplankton) and less than respiration (or 35% of NPP). All inequality constraints are listed in **Supplemental Table 1**.

Model Solution

Because the LIEM is under-constrained, infinite possible solutions satisfy the equality, and inequality constraints. To choose mean solutions and determine uncertainties within the possible solution space, we use a Markov Chain Monte Carlo (MCMC) sampling method (Kones et al., 2009; van den Meersche et al., 2009; van Oevelen et al., 2010), which has been shown to reconstruct unmeasured flows more accurately than the L₂ minimum norm approach (Stukel et al., 2012, 2018a; Saint-béat et al., 2013). Implementation details are provided in the **Supplement Detailed Methods**.

As a metric for discussing model results with respect to the approximation equations (i.e., the observations), we use the model-observation misfit relative to the model uncertainty: $\Sigma = (X_{model} - X_{obs})/\sigma_{obs}$. Here X_{model} is the model prediction, X_{obs} is the observed value, and σ_{obs} is the standard deviation of the

observed value. The square of this quantity (Σ^2) is summed over all approximate equations yielding the solution cost function, and thus Σ is a proxy for disagreement between the LIEM and observations. Unless otherwise stated, LIEM solutions are given as ranges based on the mean solutions for each cycle as well as the median value for all cycles. Displaying data in this way allows us to highlight inter-cycle variability. For value and uncertainty in all rate constraints, see **Appendix A**.

Analyses and Model Comparisons

Indirect Analysis

An indirect analysis permits investigation of the contributions of carbon between any two compartments through indirect linkages. By taking the normalized matrix of flows between compartments (G) and the identity matrix (I), the matrix $(I-G)^{-1}$ provides all the indirect flows data (Kroes, 1977). In this way the contribution of the surface compartments to the deep ones can be ascertained even when no direct flows exist. For example, if the food chain were A B C, an indirect analysis would reveal that 100% of the flows to C go through A.

Independent DVM Estimates

A model to predict the export flux due to zooplankton DVM was recently published by Archibald et al. (2019), which adds a diel vertical migration module to the Siegel et al. (2014) ecosystem model. The Archibald et al. model parameterizes the export production based on NPP, size-fractionated grazing (i.e., protists and mesozooplankton), and the proportion of DVM mesozooplankton. The export production attributed to vertically migrators who defecate at depth is a function of total grazing, the gut clearance rate, and the proportion of zooplankton undergoing DVM (Equation 2).

$$Export_{twilight} = p_{DVM} \cdot (1 - f_{fec}) \cdot (m_{fec} \cdot G_m + n_{fec} \cdot G_n) \quad (2)$$

where p_{DVM} is the fraction of mesozooplankton that undergo DVM, and f_{fec} is the fraction of fecal pellets produced by diel vertical migrators in the euphotic zone. m_{fec} and n_{fec} are the proportions of grazing that are exported by mesozooplankton and protistan zooplankton, respectively. G_m and G_n are the grazing rates for mesozooplankton and protistan zooplankton, respectively.

The respiration conducted by vertically migrating zooplankton can be calculated based on the metabolic efficiency, fraction of mesozooplankton undergoing DVM, and their grazing rate (Equation 3).

$$Export_{resp} = p_{met} \cdot p_{dvm} \cdot f_{met} \cdot \left[(1 - m_{fec}) \cdot G_m + \frac{n_{fec}}{m_{fec}} (1 - m_{fec}) \cdot G_n \right] \quad (3)$$

where p_{met} is the temperature dependent metabolic rate with ΔT , the temperature difference between the mesopelagic and epipelagic and $p_{met} = 2^{(\Delta T/10)} / (2^{(\Delta T/10)} + 1)$. f_{met} is the metabolic efficiency of the zooplankton, assumed to be 0.50. We calculated active transport from Equations 3 and 4 following Archibald, but using the CCE-optimized parameter set that

Stukel et al. (2015) determined for the Siegel et al. (2014) model. The fraction of mesozooplankton undergoing DVM (p_{dvm}) was calculated as described in Section Mesozooplankton and Nekton Constraints. Fecal pellet production for meso- and microzooplankton were set to $m_{fec} = 0.3$ and $n_{fec} = 0.06$ (Archibald et al., 2019), respectively.

Since the Archibald et al. model does not include mortality at depth as export and excludes any mesopelagic ingestion or excretion, the total export flux is the sum of Equations 2 and 3. To compare with the LIEM presented here, a modified LIEM active transport flux will be calculated using the total active transport for mesozooplankton and subtracting mesopelagic mortality.

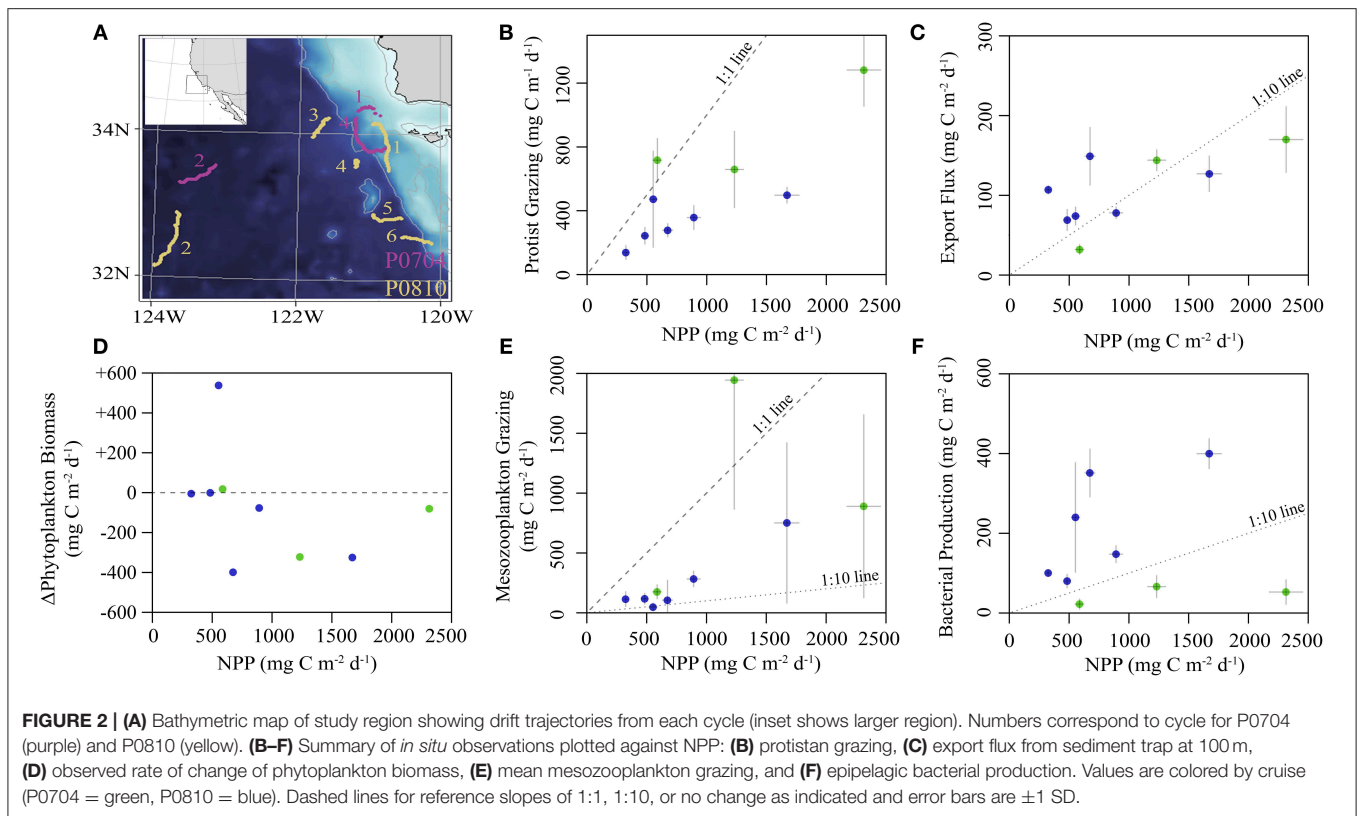
RESULTS

In situ Ecosystem Observations

The locations for each study site were chosen to maximize the range of environmental conditions (**Figure 2**). Sea surface chlorophyll *a* (Chl *a*) varied from 0.2 to 1.5 mg Chl *a* m⁻³ with vertically-integrated primary productivity varying from 325 to 2,314 mg C m⁻² d⁻¹. Productivity and biomass typically declined with distance from the Point Conception upwelling center. Most cycles were in water masses with steady or declining phytoplankton biomass (**Figure 2D**), with the exception of P0810-1. Sediment trap-derived carbon export at 100 m depth varied from 32 to 170 mg C m⁻² d⁻¹ (**Figure 2C**), with observed e-ratios (i.e., sediment trap export/¹⁴CPP) ranging from 5 to 33%. Standing stock of zooplankton correlated positively with NPP and export (Spearman correlations of 0.36 and 0.40, respectively). Protistan zooplankton were responsible for grazing ~50% of NPP (**Figure 2B**) while mesozooplankton grazed, on average, ~30% of NPP with one exception (**Figure 2E**). The proportion of mesozooplankton biomass exhibiting DVM behavior ranged from 35 to 86% (median: 58%). Epipelagic bacterial production rates did not correlate with NPP but ranged from 22 to 400 mg C m⁻² d⁻¹ (**Figure 2F**), with the three lowest rates observed during the P0704 Cruise.

Model-Observation Mismatch

The LIEM solutions consistently show general agreement with all *in situ* observations except for modeled NPP, which is elevated by 18–56% (median: 22%) from ¹⁴CPP estimates (**Figure 3A**), or 3.0–9.3 Σ (median: 3.6 Σ). This degree of misfit corresponds to 18–82% (median: 46%) of the total model-observation misfit. Model agreement with the sediment trap was high (–33–25%; **Figure 3E**) with a modeled e-ratio (i.e., sediment trap export/NPP) of 5–35% (median: 14%), which compares well to the observed e-ratio of 5–33% (median: 11%). Modeled protistan grazing rates and mesozooplankton grazing rates were reasonably close to observations (**Figure 3B**). Modeled microzooplankton (MIC) grazing was lower than observed for cycles P0704-2 (–2.8 Σ) but agreed reasonably well (–1.5– +0.1 Σ) for the other cycles (**Supplemental Figure 1**). For P0704-1, mesozooplankton grazing rates were lower than observations for SMZ (–1.8 Σ), total non-DVM grazing (–1.8 Σ) and for vmSMZ grazing (–1.6 Σ). During the course of this cycle, phytoplankton biomass declined (–322 mg C m⁻² d⁻¹) and had high zooplankton



grazing rates compared to the other cycles. This water parcel may have been in a declining bloom stage where observed grazing rates were unsustainable. Model-data agreement among the seven nekton-related observations (e.g., **Figure 3F**) was satisfactory ($|\Sigma| < 1$) except for P0810-1, which showed reduced vertically migrating nekton activity relative to estimates (vmMYC epipelagic respiration: -1.5Σ , vmMYC mesopelagic respiration: -1.7Σ , and vmMYC mesopelagic mortality: -1.1Σ). This cycle was along the edge of an anti-cyclonic eddy, where lateral gradients were likely high.

Epipelagic Ecosystem Model

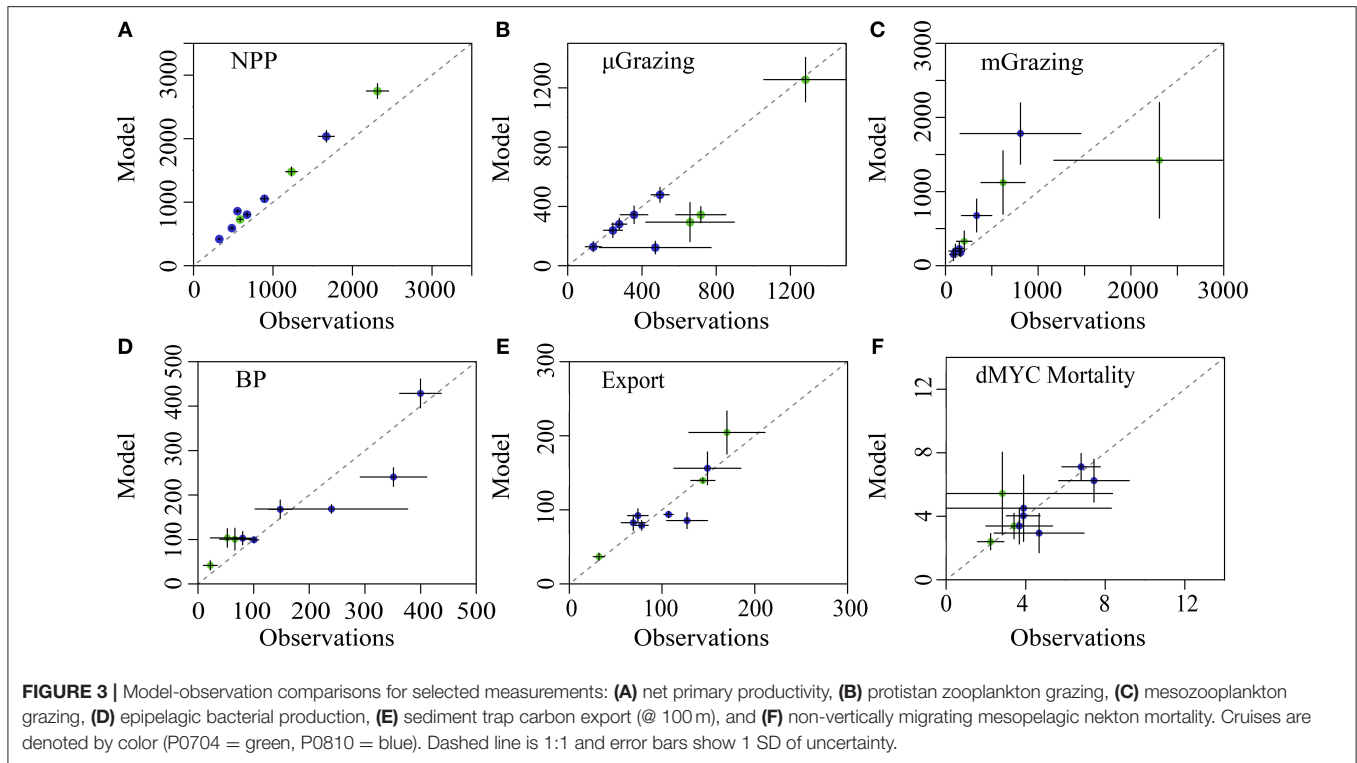
According to the LIEM, phytoplankton respired 18–39% (median: 30%) of GPP, lost 14–26% (median: 18%) as DOC, lost 2–42% (median: 6%) to non-grazer mortality and the remaining 5–54% (median: 45%) was grazed by zooplankton. Modeled NPP ranged from $421 \text{ mg C m}^{-2} \text{ d}^{-1}$ to $2,750 \text{ mg C m}^{-2} \text{ d}^{-1}$ (median: $861 \text{ mg C m}^{-2} \text{ d}^{-1}$). The LIEM suggested that protists and mesozooplankton had relatively similar grazing impacts on phytoplankton across all cycles, although the proportional role was greater for mesozooplankton in coastal regions and greater for protists under oligotrophic conditions. Between 14 and 47% (median: 33%) of NPP was grazed by protistan zooplankton (MIC + HNF) and 18–96% (median: 45%) by mesozooplankton (SMZ + vmSMZ + LMZ + vmLMZ). We note that protistan grazing rates normalized to NPP was higher than observations while protistan grazing generally

matched the observations (**Figure 3; Supplemental Figure 1**). 57–82% (median: 74%) of mesozooplankton grazing was by small mesozooplankton (SMZ + vmSMZ). Vertically migrating mesozooplankton were responsible for 52–89% (median: 63%) of total mesozooplankton grazing, 58–85% (median: 77%) of which was done by vmSMZ (i.e., vmSMZ grazing/total vm grazing).

Mortality relative to ingestion for mesozooplankton was similar for the different epipelagic mesozooplankton (i.e., SMZ, LMZ, vmSMZ, and vmLMZ): SMZ: 24–25%, vmSMZ: 23–25%, LMZ: 22–25%, and vmLMZ: 24–27%, as was fecal pellet production (between 30 and 40% of ingestion).

Overall, 19–44% (median: 29%) of NPP was transferred from the epipelagic to the mesopelagic with 3–8% (median: 5%) of NPP leaving the epipelagic through higher trophic levels (SAR + vmMYC). Gravitational settling and subduction of POC accounted for 12–55% (median: 37%) and 2–32% (median: 14%) of epipelagic export, respectively, while 18–84% (median: 41%) was through active transport of DVM mesozooplankton (vmSMZ + vmLMZ). Vertically migrating myctophids (vmMYC) transferred 2–6% (median: 4%) of total export. Section New Production, Export and DVM provides a more detailed description of export production.

The gross growth efficiencies (GGE) for each type of organism are shown in **Figure 4A**. Overall, BAC GGE was 7–29% (median: 25%) with an upper bound set to 30%. Notably, BAC GGE differed based on cruise, with P0704 cycles ranging between 8 and 13% and P0810 ranging between 23 and 29%. MIC



GGE was 35–38% (median: 37%), and HNF GGE ranged from 32 to 35% (median: 33%), which is slightly higher than typical estimates of protistan zooplankton GGE (Straille, 1997) although reported variability is high (Steinberg and Landry, 2017). GGEs for epipelagic mesozooplankton were consistently above 20%.

Trophic Level and Diets

Trophic levels for each organism (**Figure 4B**) were calculated by assuming that primary productivity, detritus and DOC were at trophic level 1. Trophic level indices were not affected by the overall cycle productivity (i.e., NPP), time of year, or by nutrient regime. The trophic level of small epipelagic mesozooplankton (SMZ) ranged from 2.2 to 2.5 (median: 2.2) and large mesozooplankton (LMZ) ranged from 2.2 to 2.9 (median: 2.6). The SAR trophic level was 3.3–3.8 (median: 3.5), and vmMYC was similar at 3.3–4.0 (median: 3.8). Modeling these higher trophic levels is important for structuring the ecosystem, and the nekton trophic levels found here are consistent with findings from ^{15}N amino acid studies (Choy et al., 2015).

The modeled mesozooplankton ingestion can be classified into four distinct dietary types: (1) Herbivory = phytoplankton diet, (2) Protistivory = protistan zooplankton diet, (3) Detritivory = detrital diet (i.e., SDT or LDT), and (4) Carnivory = mesozooplankton diet. Using this partitioning, the relative contributions of each dietary component were assessed for large, and small vertically migrating mesozooplankton compartments (**Figure 5**). The largest proportion of the diet for resident epipelagic mesozooplankton (i.e., SMZ & LMZ) was balanced between herbivory (19–57% median: 40%) and protistivory

(26–59% median: 40%). Detritivory was 9–21% (median: 13%) of total diet. Inter-cycle variability in carnivory was low for resident epipelagic mesozooplankton and contributed 6–8% (median: 6%) of their diet.

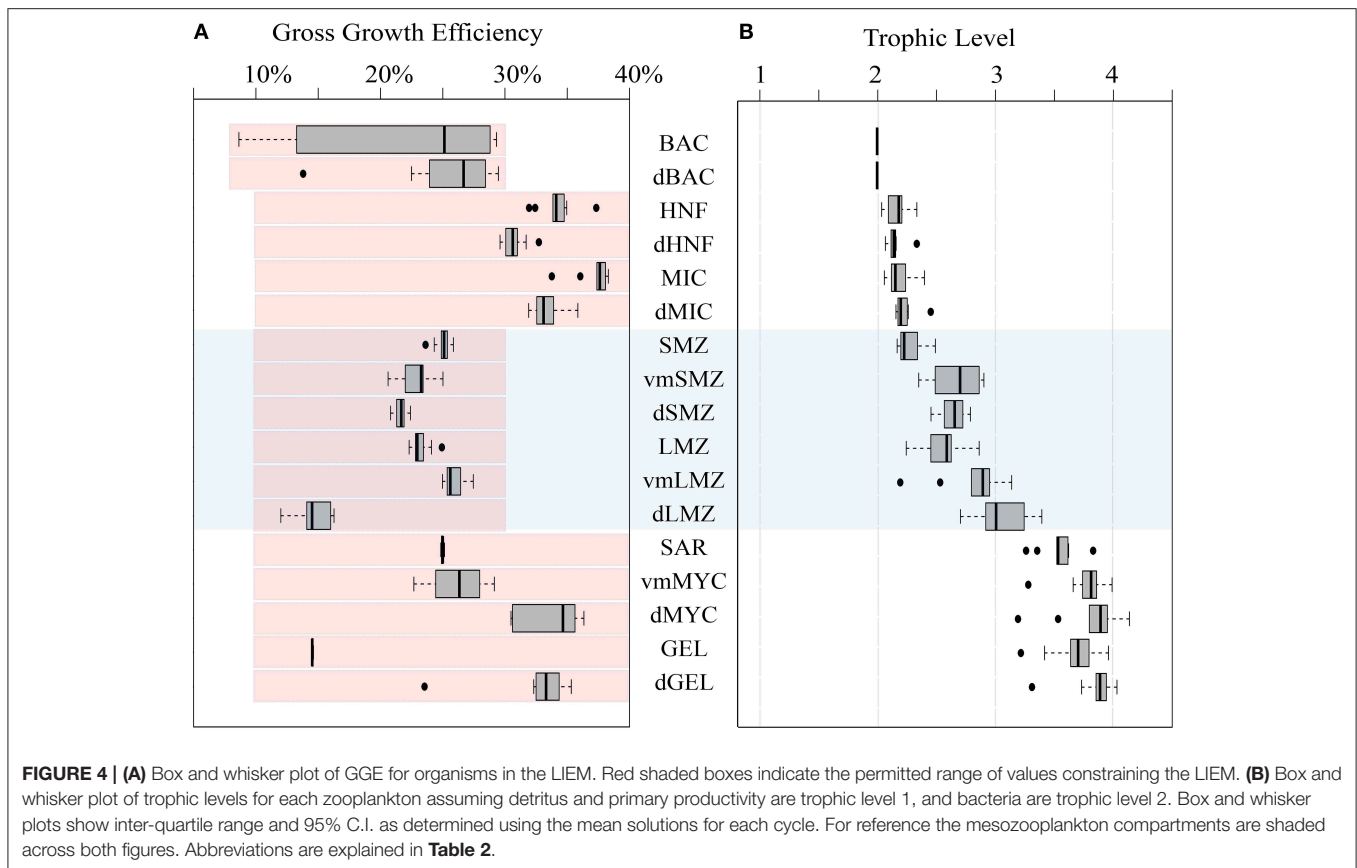
Comparing the LIEM solutions between the nutrient limited and upwelling cycles, we found that large mesozooplankton (LMZ) grazing increased from 9–16% (median: 13%) in the nutrient limited cycles to 22–65% under upwelling conditions (median: 30%) of NPP. However, the overall diets of the mesozooplankton did not systematically change with nutrient condition.

New Production, Export and DVM

Total export ranged from 163 to 707 mg C m⁻² d⁻¹ (median: 282 mg C m⁻² d⁻¹), with distinctly elevated values associated with upwelling cycles (**Figure 6A**). The fraction of export attributed to mesozooplankton DVM (vmSMZ + vmLMZ) covaried with nutrient regime: mesozooplankton active transport contributed 14–37% of total export under nutrient limited conditions and 44–84% under upwelling conditions (**Figure 6B**). There was no significant relationship ($p < 0.1$) between the total export efficiency (i.e., total export/NPP) and NPP (**Figure 6C**).

For vmSMZ, 77–80% (median: 80%) of their respiration took place in the epipelagic, along with 67–87% (median: 85%) of their DOC excretion. This is consistent with the suggestion that mesozooplankton respiration and excretion are elevated in the warmer epipelagic waters (Ikeda, 1985), where activity is highest.

The fate of active export flux is important for understanding the ecological impact of this carbon supply. Within the



mesopelagic, mesozooplankton respired 11–104 mg C m⁻² d⁻¹ (median: 33 mg C m⁻² d⁻¹) and excreted 7–116 mg C m⁻² d⁻¹ (median: 20 mg C m⁻² d⁻¹; **Figure 7A**). Predation on vertically migrating mesozooplankton accounted for a loss of 23–352 mg C m⁻² d⁻¹ (median: 59 mg C m⁻² d⁻¹) in the mesopelagic. Mesozooplankton fecal pellet production in the mesopelagic was 8–29 mg C m⁻² d⁻¹ (median: 13 mg C m⁻² d⁻¹). Resident mesopelagic mesozooplankton were the dominant mortality term for the vertically migrating mesozooplankton (**Figure 7B**).

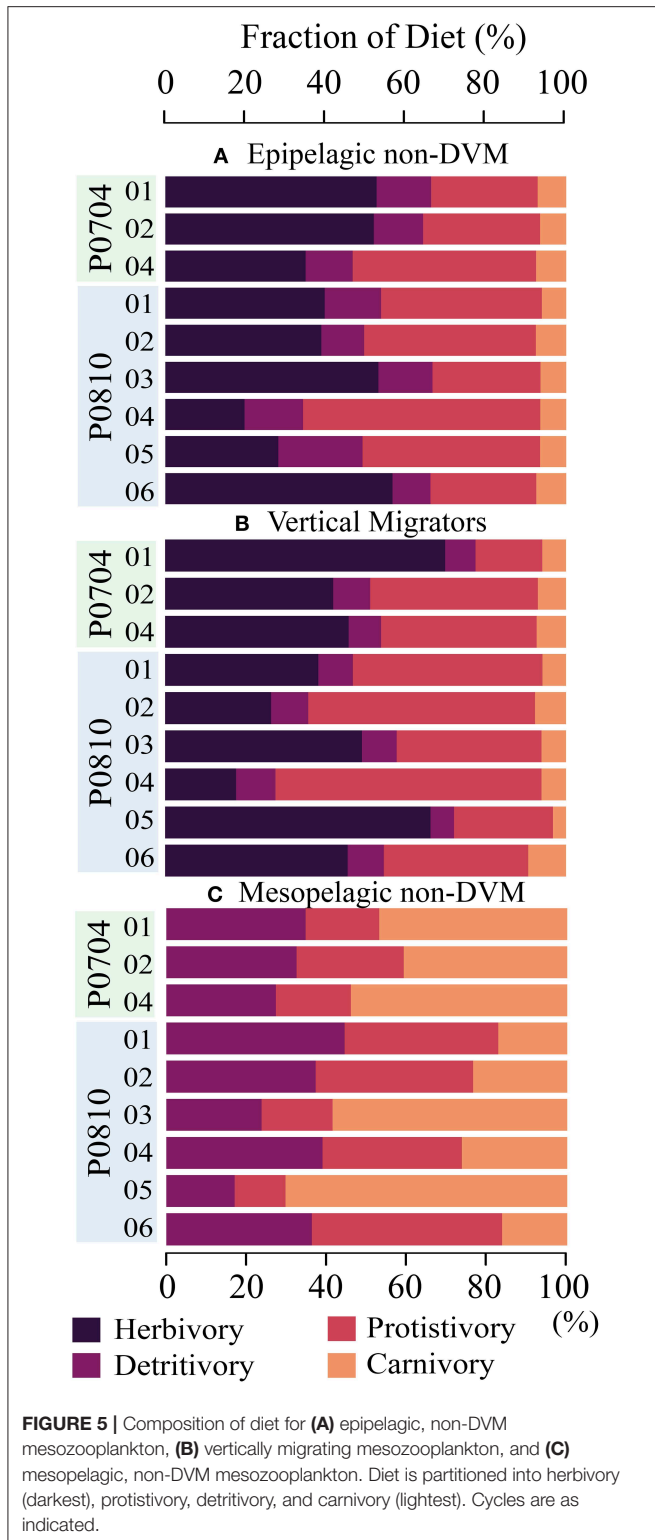
Mesopelagic Ecosystem

Deep bacteria (dBAC) made up 6–30% (median: 11%) of the mesopelagic protistan zooplankton diet with the remainder supplied by detritus/fecal pellets. Mesopelagic mesozooplankton (i.e., dSMZ & dLMZ) had a more variable diet than the epipelagic mesozooplankton (**Figure 5**), with detritivory ranging from 17 to 43% (median: 39%) of their diet, protistivory at 14–51% (median: 30%) and carnivory at 10–68% (median: 33%).

Systematic increases in trophic level between the epipelagic and mesopelagic resident zooplankton and nekton were observed (**Figure 4**). The trophic level of epipelagic microzooplankton (MIC) was 2.0–2.3 while dMIC was 2.3–2.5. Similar increases between the epipelagic and mesopelagic were observed for mesozooplankton, where SMZ had a trophic level of 2.2–2.5 (median: 2.2) dSMZ had a trophic level of 2.5–2.8 (median: 2.6).

Likewise, dLMZ trophic levels were elevated by ~0.4 relative to LMZ. The trophic level of dMYC (3.2–4.1) was more variable than the other micronekton (e.g., vmMYC: 3.5–4.0), illustrating a greater variability in diet.

Mesopelagic respiration is a useful diagnostic loss term for determining which organisms are responsible for the mesopelagic carbon demand (**Supplemental Figure 2**). Mesopelagic bacteria accounted for the largest proportion of mesopelagic respiration (31–41% median: 34%). High respiration of mesopelagic bacteria was found despite relatively high GGE for these organisms (median 26%, **Figure 4A**). Mesopelagic protistan zooplankton and resident mesozooplankton were responsible for 14–30% (median: 25%) and 14–24% (median: 15%), respectively. Resident gelatinous predators and myctophids are responsible for 4–8% of mesopelagic respiration combined. The proportion of export due to active transport covaried with resident mesopelagic respiration (**Figure 8A**), illustrating the coupling between active transport and mesopelagic activity in the LIEM. The effect of higher active transport relative to total export can be shown with an indirect analysis where the relative contribution of carbon from epipelagic detritus (i.e., a passive transport proxy) and vertically migrating mesozooplankton (i.e., an active transport proxy) in the diet of each organism can be measured. Indirect flux analyses show that a higher proportion of the carbon consumed by mesopelagic bacteria, protists, and mesozooplankton originated from passive rather than active



transport (Figure 8B). However, mesopelagic nekton (dMYC) were predominantly supported by carbon derived from active transport pathways.

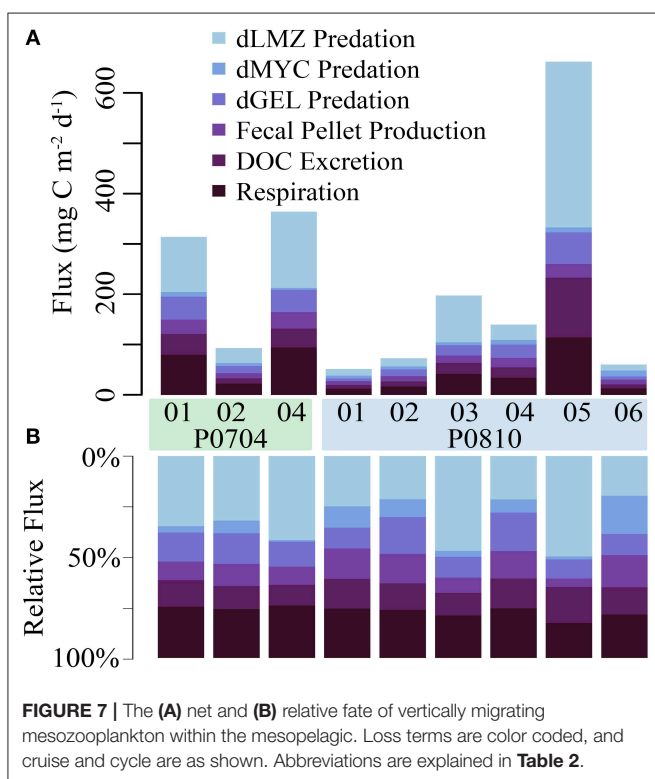
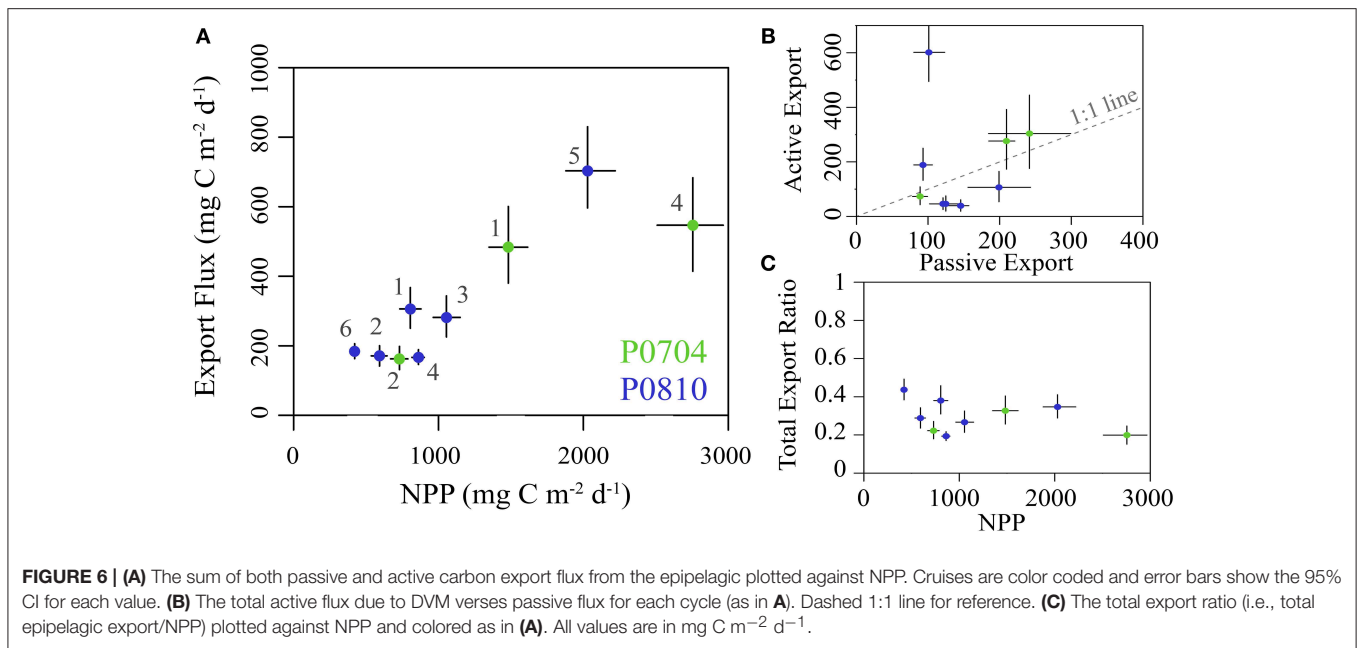
DISCUSSION

Diel Vertical Migration and Active Transport in the CCE

In contrast to common assumption about the processes driving the biological pump, our results suggest that active transport may be as, if not more, important than sinking particle flux. We found that active transport (mesozooplankton and fish combined) was responsible for 39–606 mg C m⁻² d⁻¹ (median: 107 mg C m⁻² d⁻¹), corresponding to 21–86% of total export to the mesopelagic, while sinking particles contributed 14–79%. This finding is not directly forced by an *a priori* assumption of the importance of active transport. Indeed, we placed no direct constraint on the amount of mesozooplankton mortality in the mesopelagic, and the minimum constraints on basal metabolism by zooplankton in the mesopelagic (Stukel et al., 2013) implied that active transport could have been as low as 2–40% of sinking flux (median: 18%). Nevertheless, the importance of active transport was a robust result of the inverse analyses. For P0810-6, the cycle with the lowest relative contribution of active transport to total export (21%), the total flux was 184 ± 23 mg C m⁻² d⁻¹ (95% CI) and active transport was 39 ± 21 mg C m⁻² d⁻¹ (95% CI). This cycle was oligotrophic and had the lowest ¹⁴CPP measurements of any cycle on the two cruises. In contrast, cycle P0810-5 had the highest relative contribution of active transport (86 ± 4% of total export at the 95% CI). P0810-5 was on the coastal (i.e., high biomass) side of a strong frontal feature with high rates of primary productivity and large standing stocks of zooplankton.

Although these rates of active transport are higher than reported in many studies, they are fully consistent with mesozooplankton community dynamics in the CCE. The model suggests that total epipelagic mesozooplankton consumption on phytoplankton, protists, detritus, and other mesozooplankton ranged from 361 to 2,966 mg C m⁻² d⁻¹ (median: 1,006 mg C m⁻² d⁻¹). Vertically stratified day-night net tows showed that 35–86% (median: 57%) of the mesozooplankton community was vertically migrating to depth each day and that most of these vertical migrants were copepods and euphausiids (Stukel et al., 2013). Our model results indicate that only 20–23% of respiration and 16–34% of excretion by vertical migrants occurred at depth. None of these assumptions are particularly aggressive. Furthermore, our results (Figure 9) are consistent with estimates of DVM in the zooplankton derived from the model of Archibald et al. (2019), if specific dynamics of the CCE are taken into account (e.g., zooplankton consume nearly all of NPP, Landry et al., 2009; microphytoplankton are negligible contributors to sinking flux, Stukel et al., 2013). Our estimates of the total export ratio 19–44% are also consistent with typical *f*-ratio estimates (new production to total export) in our study region, which varied from 0.23 to 0.40 (Krause et al., 2015). Our results thus do not arise from unusual parameterizations but instead may reflect the fact that estimates of active export using standard metabolism calculated from Ikeda (1985) and Ikeda et al. (2001) may be conservative underestimates.

Our results also reflect realistic coupling between the epipelagic and mesopelagic communities. Model results



suggested that the carbon demand was equal to <1–4% (median: 1.1%) of NPP for mesopelagic fish, 1–7% (median: 3%) of NPP for predatory gelatinous zooplankton, 8–22% (median: 14%) of NPP for resident mesopelagic zooplankton, and 6–19% (median: 11%) of NPP for mesopelagic bacteria. These mesopelagic carbon demands must be met by carbon flux from the surface layer, the

most likely sources of which are sinking particle flux (which we experimentally measured using two independent approaches) and active transport. While it is possible that both sediment traps and ^{238}U - ^{234}Th disequilibrium underestimated sinking carbon flux, the inverse analysis offers compelling evidence that active transport is more likely to support mesopelagic fish and gelatinous predator communities. Although sinking particles can efficiently support bacterial production (as they are likely directly colonized by particle-attached bacteria), many fish and gelatinous zooplankton are predators that feed more on living organisms than on the sinking fecal pellets that typically dominate particle flux in the CCE. For these planktivorous organisms, sustaining their metabolism through a food chain supported by sinking particles would likely require one (if not more) trophic levels to separate them from the export source, depending on whether the sinking particles are consumed by filter- or flux-feeding zooplankton or by microbes (Stukel et al., 2019). Thus, sustaining the high carbon demand of mesopelagic myctophids with sinking particles requires substantially more total carbon flux than does sustaining it via active transport of the myctophids' prey.

Mesopelagic sources of mortality have implications for the fitness of vertical migrators. It is often assumed that DVM is ecologically advantageous when the costs associated with not feeding during the day and actively swimming to depth are offset by the benefits of reduced predation pressure and/or reduced metabolism at colder mesopelagic temperatures (Bianchi et al., 2013; Hansen and Visser, 2016; Morozov and Kuzenkov, 2016). Our model suggests that mortality normalized to ingestion is similar across all mesozooplankton compartments and across a wide range of ecosystem states (SMZ: 24–26%, LMZ: 22–25%, vmSMZ: 21–25%, vmLMZ: 25–27%, dSMZ: 21–23%, dLMZ: 19–23%). Even though vmSMZ experience similar predation to SMZ and dSMZ, approximately half of the predation on vertically

migrating zooplankton takes place in the mesopelagic, thereby transferring carbon to depth despite the fact that their excretion and respiration occur primarily in the epipelagic.

The comparable mortality experienced by vertically-migrating mesozooplankton in the mesozooplankton may seem counterintuitive in light of extensive research suggesting that the adaptive advantage of DVM may be to reduce predation (Ohman and Romagnan, 2016; Bandara et al., 2018). However, in the CCE, it is not particularly surprising when the large abundances of myctophids, gonostomatids, and other mesopelagic fish are considered. Davison et al. (2013) and Davison et al. (2015) demonstrated high biomass of these fish comprising both vertically-migrating and mesopelagic resident communities. Mesozooplankton may thus face as high, if not higher, predator abundance at mesopelagic depths than in the epipelagic, although colder temperatures and reduced irradiance may diminish predation rates at depth. DVM may remain advantageous as a lifestyle because if these organisms were present at the surface during the day then they might experience substantially greater predation than in the mesopelagic.

Sensitivity Analysis and Ecological Connections

The ecosystems generated in the 9 model runs were as varied as the cruise measurements: including observations from dynamic coastal blooms to quiescent oligotrophic communities. All 9 cycles had significantly elevated NPP compared to the observed ^{14}C PP (Figure 3; Supplemental Figure 1) with 95% CI from the MCMC random walk. Whether this result can be considered a model bias or is derived from possible systematic differences between ^{14}C PP and true net primary production (Peterson, 1980; Lefevre et al., 1997; Marra, 2009; Milligan et al., 2015) is not known. Compared to shorter ^{14}C labeling experiments (e.g. dawn-to-dusk, 8 h, pulse-chase), the 24 h incubations used here are generally thought to measure NPP rates directly (Milligan et al., 2015); however, long-term incubations are susceptible to biases introduced by heterotrophic processes and DOC excretion (Laws et al., 2000; Dickson et al., 2001). Since rapid consumption of net primary productivity by grazers, cell lysis, and excretion of DOC (all of which are explicitly included in the LIEM) will reduce the apparent ^{14}C -bicarbonate uptake rates, ^{14}C PP rates may be biased low, especially when turnover times are short. In fact, when comparing dilution-based growth rates with 24 h ^{14}C PP incubations in the equatorial Pacific, Landry et al. (2011) found that ^{14}C PP estimates needed to be adjusted upwards by 29% on average. An alternative explanation may stem from a bias in the MCMC approach used. Since the random walk is strictly required to yield solutions where flows are positive through the “mirror” algorithm, the region of permitted solutions is non-symmetric and may favor the broader solution-space of high NPP solutions (as noted in Stukel et al., 2012). A thorough investigation into the potential biases of the ^{14}C PP method of the MCMC solution algorithm are beyond the scope of this study, but the impact of a potential bias in modeled NPP are discussed below.

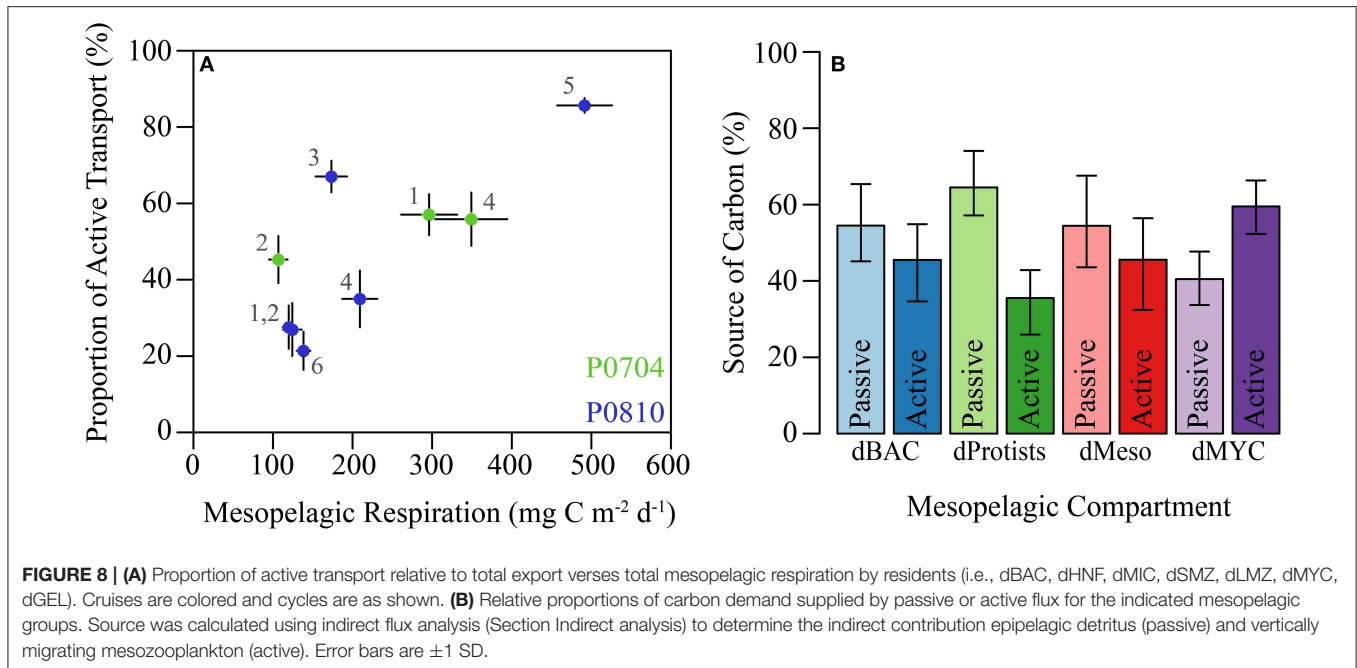
To test the model's sensitivity to the misfit with ^{14}C PP and to confirm that our results were not driven by a potential bias in

the model, the LIEM was rerun assuming that ^{14}C PP uncertainty was 1/10th of the actual estimated uncertainty (i.e., 0.6% relative uncertainty). The model-observation misfit increased by nearly 2.5x with vmSMZ and SMZ grazing rates, myctophid metabolic estimates, and sediment trap export all reduced by $\sim 2 \Sigma$ relative to the standard model run. This result shows that the model needed to increase NPP in the standard model run in order to match the observed mesozooplankton grazing rate and myctophid metabolic requirements. However, the proportion of export resulting from active transport remained relatively unchanged. It varied from 106 to 641 mg C m⁻² d⁻¹ across the cycles (compared to 162 to 707 mg C m⁻² d⁻¹ in the standard model run). This suggests that our primary conclusions about export flux were not contingent on elevated model NPP.

Because bacterial activity in the mesopelagic was not measured, we set a high upper and low minimum bounds for bacterial production. For the minimum bound on mesopelagic BP, we chose an attenuation coefficient of $\alpha = 1.47$ (Yokokawa et al., 2013). This resulted in model-determined mesopelagic bacterial carbon demand that may have been lower than true *in situ* values. Other reported values for the attenuation of BP in the mesopelagic include slopes of $\alpha = 1.15$ (Tanaka and Rassoulzadegan, 2004) and 1.03 (Gasol et al., 2009), which would result in 25 and 36% higher estimates of mesopelagic BP, respectively. When the minimum mesopelagic bacterial production estimates were halved ($\alpha = 0.64$; Equation 1), the model responded by increasing NPP by +2% (inter-cycle median) and total export flux by 11%. Since passive particle flux is constrained by observations, passive flux increased by 0–12% (median: 4%) while active transport by mesozooplankton increased by 0–56% (median: 26%). Active transport by nekton was also elevated (0–14%, median: 10%). Model-observation misfit increased by an average of 17% with notable changes in NPP (+0.42 Σ), sediment trap flux (+0.34 Σ) and Thorium-234 flux (+0.22 Σ).

The standard model results were also robust to changes in other observations. When the nekton metabolic estimates were halved, export by vmMYC was reduced by 51% (inter-cycle median), a change of < 5 mg C m⁻² d⁻¹, while other forms of export were unchanged. Increasing the upper limit of mesozooplankton GGE from 30 to 40% led to a $\sim 20\%$ increase in mesozooplankton active transport and no change in nekton-derived active flux or passive flux.

Zooplankton basal respiration rates have been shown to be suppressed under low-oxygen conditions (Ekau et al., 2010; Seibel, 2011), such as those seen in the midwater oxygen minimum zones (OMZ) often encountered below the productive Eastern boundary current upwelling biomes (Chavez and Messié, 2009; Bettencourt et al., 2015). During our study periods, water-column dissolved oxygen concentrations fell below 44.7 μM (0.5 ml/l), indicating hypoxic conditions (Helly and Levin, 2004; Gilly et al., 2013), between 271 and 470 m water depth (Supplemental Figure 3). Notably, most of the zooplankton captured in our study were found at depths shallower than 300 m. While there are questions remaining with regard to how mesozooplankton respiration rates would be affected by the intermediate oxygen depletion observed in

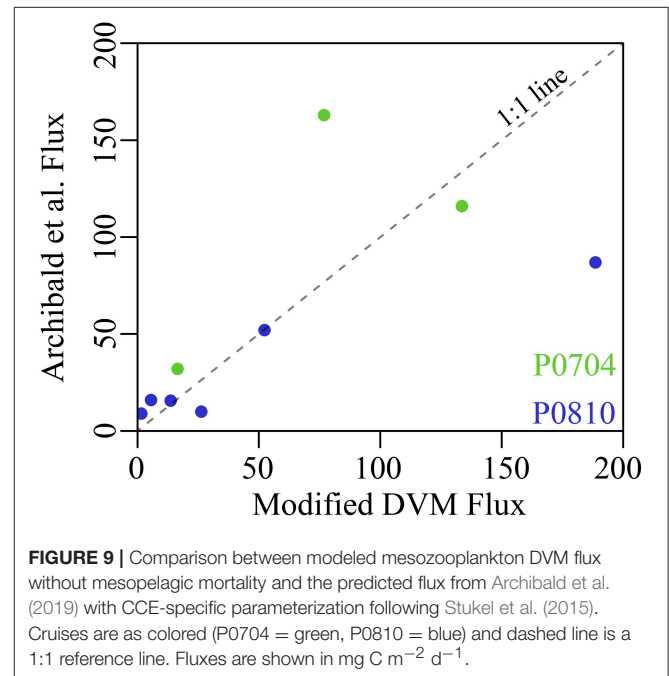


our study (Teuber et al., 2013; Kiko et al., 2016), the model results are largely insensitive to a possible reduction in basal metabolism. In particular, the respiration rates of the mesopelagic organisms in the model were consistently higher than the basal metabolic constraint placed on them (typically $> 2x$). Including ecological implications of the OMZ directly would be a valuable contribution to the field that necessitates a more depth-resolved model due to the importance of vertical gradients in oxygen and temperature.

Linear Inverse Models

LIEMs are powerful tools for assimilating diverse *in situ* measurements and constraints with a food web perspective. The use of a two-layer model (Jackson and Eldridge, 1992) is particularly powerful because it allows information from the mesopelagic to constrain epipelagic food web flows and vice versa. Compared to most previously published LIEMs, the model presented here includes many more *in situ* rate measurements, made possible by the suite of contemporaneous rate measurements made during quasi-Lagrangian experiments. When constrained by fewer rate measurements (Dubois et al., 2012; van Oevelen et al., 2012; Sailley et al., 2013), the LIEM solution relies more heavily on greater than/less than constraints derived from biomass measurements, leading to correspondingly higher uncertainty. This highlights a need for studies that simultaneously quantify the activity of many different plankton functional groups.

Since a LIEM is fundamentally a data-regression technique, our results are emergent from (A) our observations, (B) the assumptions used (e.g., GGE), and (C) the ecosystem structure of the model. Thus, we believe the resulting model solutions to be descriptive of the dominant *in situ* processes in the CCE LTER study region. However, it is important to note



that there were large uncertainties associated with some model flows, and that these could be quantified using the MCMC approach (**Supplemental Table 2**). We thus highly recommend the MCMC approach (Kones et al., 2009; van den Meersche et al., 2009), which has been shown to be robust in its ability to recover ecosystem rates relative to the L-2 minimum norm (Stukel et al., 2012; Saint-béat et al., 2013). Even more important is its ability to generate confidence intervals that realistically

represent the uncertainties in model outputs with respect to both measurements and under-determinacy of the model. For instance, for cycle P0810-6, we found that the 95% confidence interval for HNF ingestion of detritus was 5–127 mg C m⁻² d⁻¹, providing no real knowledge of whether or not this connection was an important part of the ecosystem. However, for Cycle P0810-5, we found that mesopelagic mesozooplankton predation on small vertical migrators was 233–423 mg C m⁻² d⁻¹ (95% CI), indicating a higher degree of confidence that this flow was substantial at this location. Investigation of the confidence intervals can thus inform which conclusions can be considered robust. Developing even better-resolved ecosystem models likely requires incorporation of more diverse measurement types, such as ¹⁵N isotopic data (Stukel et al., 2018a).

The Biological Carbon Pump and Mesopelagic Flux Attenuation

Reports of active transport by vertically migrating biota have long suggested that these organisms can transport a globally significant amount of carbon to depth. However, most early studies suggested that active transport was substantially less important than passive flux of sinking particles (Morales, 1999; Davison et al., 2013; Steinberg and Landry, 2017). At the oligotrophic BATS station off Bermuda, Dam et al. (1995) found that respiration by mesozooplankton augmented the passive carbon flux at 150 m by 18–70%. Also at BATS, Steinberg et al. (2000) reported a significant vertical transfer of nitrogen by zooplankton, including dissolved organic nitrogen (DON). In fact, vertical migrators were found to perform 15–66% of the total nitrogen transport. Hansen and Visser (2016) estimated that across the North Atlantic active transport by mesozooplankton may constitute 27% of total export out of the surface mixed layer. In addition to zooplankton, vertical migrations by micronekton can also lead to significant export fluxes (Angel and Pugh, 2010; Davison et al., 2013; Hernandez-Leon et al., in review). Using biomass estimates and metabolic relationships, Davison et al. (2013) found micronekton contributions of 22–24 mg C m⁻² d⁻¹ (or 15–17% of estimated passive export) in the northeast Pacific at 150 m water depth. In the North Pacific Subtropical Gyre, Al-Mutairi and Landry (2001) estimated that active transport due to zooplankton respiration was responsible for carbon flux equal to 18% of passive flux at 150 m. Using a conservative approach (Longhurst et al., 1990), estimated that active export by zooplankton DVM was 13–58% that of passive flux at 150 m when accounting for respiration alone in subtropical waters, which is similar to our results where the LIEM suggests that mesozooplankton respiration at depth is 9–113% (median: 34%) that of passive export at 100 m. Global modeling estimates have indicated that active transport may be responsible for 14% (Archibald et al., 2019) or 15 to 40% (Bianchi et al., 2013) increases in carbon export out of the euphotic zone relative to sinking particles alone. More recent results have suggested increased importance for active transport, potentially rivaling that of passive flux. In the Costa Rica Dome, a region with high mesozooplankton biomass like the CCE, Stukel et al. (2018b) identified active transport by zooplankton DVM as responsible for 21–45% of total euphotic zone export. Hernández-León et al. (in review) found that active transport

was equal to one quarter of passive flux in oligotrophic regions, but was 2-fold higher than passive flux in eutrophic areas of the tropical and subtropical Atlantic. Our results that total active transport (zooplankton and nekton) may be responsible for 18–84% (median: 42%) of total carbon export at 100 m in the CCE are thus somewhat higher than found in most studies, but consistent with recently published values for high zooplankton biomass regions. Furthermore, our results are in line with other biogeochemical and ecological expectations (e.g., mesopelagic carbon demand, euphotic zone new production, mesozooplankton energy partitions). We thus suggest that active transport in high biomass regions may be more important, in fact, than some previous studies suggest, and we recommend focused research to investigate the potentially conservative assumptions made in previous studies that rely on standard (rather than active) estimates of zooplankton metabolic rates.

Within the mesopelagic, zooplankton also play an important biogeochemical role in the attenuation of particle flux (Steinberg et al., 2008; Buesseler and Boyd, 2009; Stukel et al., 2019) and in effecting elemental cycling (Kiko et al., in review; Robinson et al., 2010). Our results suggest that mesozooplankton detritivory accounted for the consumption of 57–71% of sinking particles from the epipelagic, with bacterially mediated remineralization of the majority of the remainder (i.e., mesopelagic export efficiency is < 10%). Notably, 3 of the 4 cycles with the lowest proportion of detritivory and the largest proportion of carnivory in the resident mesopelagic zooplankton were during upwelling cycles. This is opposite to the findings of Wilson et al. (2010), who observed increases in fatty-acid biomarkers associated with carnivory at station Aloha relative to K2 and attributed the increase to the lower primary productivity at station Aloha. Our result that zooplankton rely more heavily on carnivory in the mesopelagic agrees with fecal pellet characteristic analyses and fatty acid biomarkers measured by Wilson et al. (2008) and Wilson et al. (2010), respectively. However, given the advective nature of an eastern boundary current and frequency of non-steady state conditions, it is difficult to generalize from our results to the rest of the Pacific. Clearly additional studies are necessary.

CONCLUSIONS

The LIEM used here incorporated numerous *in situ* measurements made during quasi-Lagrangian experiments in the CCE in order to constrain carbon flows through the ecosystem. These observations were made in water parcels spanning a wide range of conditions from highly productive upwelling regions to an oligotrophic offshore domain and consistently found that active transport of carbon by mesozooplankton was important to supplying the mesopelagic carbon demand. The model suggests that, relative to total export, gravitational settling contributes 12–55% (median: 37%) and subduction contributes 2–32% (median: 14%) of carbon flux. This finding has implications for the interpretation of sediment trap and ²³⁴Th disequilibrium measurements and for helping to reconcile the long-studied imbalance in the mesopelagic carbon budget. The LIEM also highlights the central importance of zooplankton in marine food webs and

biogeochemistry. Excretion by vertical migrants is important for meeting bacterial carbon demand, while predation on vertical migrants supports mesopelagic resident fish communities. Our analysis comprises a unique, fully resolved phytoplankton-to-fish coupled food web of the epipelagic and mesopelagic ocean. Nevertheless, substantial uncertainties remain, and targeted studies are necessary to validate the suggested relationships *in situ* and to test their applicability across the global ocean.

DATA AVAILABILITY

All datasets generated for this study are included in the manuscript and/or the **Supplementary Files**.

AUTHOR CONTRIBUTIONS

ML was responsible for cruise design and protistan zooplankton data. MO was responsible for mesozooplankton data. RG was responsible for phytoplankton data. PD was responsible for myctophid data. MS was responsible for particle export data. TK and MS designed the model. TK wrote the manuscript. All authors contributed to editing the manuscript.

REFERENCES

- Al-Mutairi, H., and Landry, M. R. (2001). Active export of carbon and nitrogen at station ALOHA by diel migrant zooplankton. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 48, 2083–2103. doi: 10.1016/S0967-0645(00)00174-0
- Anderson, T. R., Martin, A. P., Lampitt, R. S., Trueman, C. N., Henson, S. A., and Mayor, D. J. (2018). Quantifying carbon fluxes from primary production to mesopelagic fish using a simple food web model. *ICES J. Mar. Sci.* 76, 690–701. doi: 10.1093/icesjms/fsx234
- Angel, M. V., and Pugh, P. R. (2010). Quantification of diel vertical migration by micronektonic taxa in the northeast Atlantic. *Hydrobiologia* 440, 161–179. doi: 10.1023/A:1004115010030
- Archibald, K. M., Siegel, D. A., and Doney, S. C. (2019). Modeling the impact of zooplankton diel vertical migration on the carbon export flux of the biological pump. *Glob. Biogeochem. Cycles* 33, 181–199. doi: 10.1029/2018GB005983
- Aumont, O., Maury, O., Lefort, S., and Bopp, L. (2018). Evaluating the potential impacts of the diurnal vertical migration by marine organisms on marine biogeochemistry. *Glob. Biogeochem. Cycles* 32, 1622–1643. doi: 10.1029/2018GB005886
- Bandara, K., Varpe, Ø., Ji, R., and Eiane, K. (2018). A high-resolution modeling study on diel and seasonal vertical migrations of high-latitude copepods. *Ecol. Model.* 368, 357–376. doi: 10.1016/j.ecolmodel.2017.12.010
- Benitez-Nelson, C. R., Buesseler, K. O., Van Der Loeff, M. R., Andrews, J., Ball, L., Crossin, G., et al. (2001). Testing a new small-volume technique for determining ²³⁴Th in seawater. *J. Radioanal. Nucl. Chem.* 248, 795–799. doi: 10.1023/A:1010621618652
- Bettencourt, J. H., López, C., García, E. H., Montes, I., Sudre, J., Dewitte, B., et al. (2015). Boundaries of the peruvian oxygen minimum zone shaped by coherent mesoscale dynamics. *Nat. Geosci.* 8, 937–940. doi: 10.1038/ngeo2570
- Bianchi, D., Stock, C., Galbraith, E. D., and Sarmiento, J. L. (2013). Diel vertical migration: ecological controls and impacts on the biological pump in a one-dimensional ocean model. *Glob. Biogeochem. Cycles* 27, 478–491. doi: 10.1002/gbc.20031
- Boyd, P. W., and Trull, T. W. (2007). Understanding the export of biogenic particles in oceanic waters: is there consensus? *Prog. Oceanogr.* 72, 276–312. doi: 10.1016/j.pocan.2006.10.007

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fmars.2019.00508/full#supplementary-material>

- Buesseler, K. O., and Boyd, P. W. (2009). Shedding light on processes that control particle export and flux attenuation in the twilight zone of the open ocean. *Limnol. Oceanogr.* 54, 1210–1232. doi: 10.4319/lo.2009.54.4.1210
- Burd, A. B., Hansell, D. A., Steinberg, D. K., Anderson, T. R., Aristegui, J., Baltar, F., et al. (2010). Assessing the apparent imbalance between geochemical and biochemical indicators of meso- and bathypelagic biological activity: What the @#! is wrong with present calculations of carbon budgets? *Deep Sea Res. Part II Top. Stud. Oceanogr.* 57, 1557–1571. doi: 10.1016/j.dsr.2.2010.02.022
- Chavez, F. P., and Messié, M. (2009). A comparison of Eastern boundary upwelling ecosystems. *Prog. Oceanogr.* 83, 80–96. doi: 10.1016/j.pocan.2009.07.032
- Choy, C. A., Popp, B. N., Hannides, C. C. S., and Drazen, J. C. (2015). Trophic structure and food resources of epipelagic and mesopelagic fishes in the north pacific subtropical gyre ecosystem inferred from nitrogen isotopic compositions. *Limnol. Oceanogr.* 60, 1156–1171. doi: 10.1002/lno.10085
- Cohen, J. H., and Forward, R. B. (2009). Zooplankton diel vertical migration — a review of proximate control. *Ocean. Mar. Biol. Annu. Rev.* 47, 77–109. doi: 10.1201/9781420094220.ch2
- Conover, R. J. (1966). Assimilation of organic by zooplankton. *Limnol. Oceanogr.* 11, 338–345. doi: 10.4319/lo.1966.11.3.0338
- Dam, H. G., and Peterson, W. T. (1988). The effect of temperature on the gut clearance rate constant of planktonic copepods. *J. Exp. Mar. Biol. Ecol.* 123, 1–14. doi: 10.1016/0022-0981(88)90105-0
- Dam, H. G., Roman, M. R., and Youngbluth, M. J. (1995). Downward export of respiratory carbon and dissolved inorganic nitrogen by diel-migrant mesozooplankton at the JGOFS Bermuda time-series station. *Deep-Sea Res. Part I* 42, 1187–1197. doi: 10.1016/0967-0637(95)00048-B
- Davison, P., Lara-Lopez, A., and Anthony Koslow, J. (2015). Mesopelagic fish biomass in the southern California current ecosystem. *Deep-Sea Res. Part II Top. Stud. Oceanogr.* 112, 129–142. doi: 10.1016/j.dsr.2.2014.10.007
- Davison, P. C., Checkley, D. M., Koslow, J. A., and Barlow, J. (2013). Carbon export mediated by mesopelagic fishes in the northeast Pacific Ocean. *Prog. Oceanogr.* 116, 14–30. doi: 10.1016/j.pocan.2013.05.013
- del Giorgio, P. A., and Cole, J. J. (1998). Bacterial growth efficiency in natural aquatic systems. *Annu. Rev. Ecol. Syst.* 29, 503–541. doi: 10.1146/annurev.ecolsys.29.1.503
- del Giorgio, P. A., and Duarte, C. M. (2002). Respiration in the open ocean. *Nature* 420, 379–384. doi: 10.1038/nature01165

- Dickson, M.-L., Orchardo, J., Barber, R. T., Marra, J., McCarthy, J. J., and Sambrotto, R. N. (2001). Production and respiration rates in the Arabian Sea during the 1995 northeast and southwest monsoons. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 48, 1199–1230. doi: 10.1016/S0967-0645(00)00136-3
- Dubois, S., Del Amo, Y., Grami, B., Jude, F., Marquis, E., David, V., et al. (2012). Network analysis of the planktonic food web during the spring bloom in a semi enclosed lagoon (Arcachon, SW France). *Acta Oecologica.* 40, 40–50. doi: 10.1016/j.actao.2012.02.002
- Ducklow, H. W., and Harris, R. P. (1993). Introduction to the JGOFS North Atlantic bloom experiment. *Deep-Sea Res. Part II* 40, 1–8. doi: 10.1016/0967-0645(93)90003-6
- Ekau, W., Auel, H., Pörtner, H.-O., and Gilbert, D. (2010). Impacts of hypoxia on the structure and processes in pelagic communities (zooplankton, macro-invertebrates and fish). *Biogeosciences* 7, 1669–1699. doi: 10.5194/bg-7-1669-2010
- Franks, P. J. S. (2002). NPZ models of plankton dynamics: their construction, coupling to physics, and application. *J. Oceanogr.* 58, 379–387. doi: 10.1023/A:1015874028196
- Gasol, J. M., Alonso-Sáez, L., Vaqué, D., Baltar, F., Calleja, M. L., Duarte, C. M., et al. (2009). Mesopelagic prokaryotic bulk and single-cell heterotrophic activity and community composition in the NW Africa-Canary Islands coastal-transition zone. *Prog. Oceanogr.* 83, 189–196. doi: 10.1016/j.pocean.2009.07.014
- Gilly, W. F., Beman, J. M., Litvin, S. Y., and Robison, B. H. (2013). Oceanographic and biological effects of shoaling of the oxygen minimum zone. *Annu. Rev. Mar. Sci.* 5, 393–420. doi: 10.1146/annurev-marine-120710-100849
- Gontikaki, E., van Oevelen, D., Soetaert, K., and Witte, U. (2011). Food web flows through a sub-arctic deep-sea benthic community. *Prog. Oceanogr.* 91, 245–259. doi: 10.1016/j.pocean.2010.12.014
- Gorsky, G., Ohman, M. D., Picheral, M., Gasparini, S., Stemmann, L., Romagnan, J.-B., et al. (2010). Digital zooplankton image analysis using the ZooScan integrated system. *J. Plankton Res.* 32, 285–303. doi: 10.1093/plankt/fbp124
- Hannides, C. C. S., Drazen, J. C., and Popp, B. N. (2015). Mesopelagic zooplankton metabolic demand in the North Pacific Subtropical Gyre. *Limnol. Oceanogr.* 60, 419–428. doi: 10.1002/lno.10032
- Hansen, A. N., and Visser, A. W. (2016). Carbon export by vertically migrating zooplankton: an optimal behavior model. *Limnol. Oceanogr.* 61, 701–710. doi: 10.1002/lno.10249
- Helly, J. J., and Levin, L. A. (2004). Global distribution of naturally occurring marine hypoxia on continental margins. *Deep Sea Res. Part Oceanogr. Res. Pap.* 51, 1159–1168. doi: 10.1016/j.dsr.2004.03.009
- Henson, S. A., Sanders, R., Madsen, E., Morris, P. J., Le Moigne, F., and Quartly, G. D. (2011). A reduced estimate of the strength of the ocean's biological carbon pump. *Geophys. Res. Lett.* 38, 10–14. doi: 10.1029/2011GL046735
- Henson, S. A., Yool, A., and Sanders, R. (2015). Variability in efficiency of particulate organic carbon export: a model study. *Glob. Biogeochem. Cycles* 29, 33–45. doi: 10.1002/2014GB004965
- Ikeda, T. (1985). Metabolic rates of epipelagic marine zooplankton as a function of body mass and temperature. *Mar. Biol.* 85, 1–11. doi: 10.1007/BF00396409
- Ikeda, T., Kanno, Y., Ozaki, K., and Shinada, A. (2001). Metabolic rates of epipelagic marine copepods as a function of body mass and temperature. *Mar. Biol.* 139, 587–596. doi: 10.1007/s002270100608
- Jackson, G. A., and Eldridge, P. M. (1992). Food web analysis of a planktonic system off Southern California. *Prog. Oceanogr.* 30, 223–251. doi: 10.1016/0079-6611(92)90014-Q
- Kiko, R., Hauss, H., Buchholz, F., and Melzner, F. (2016). Ammonium excretion and oxygen respiration of tropical copepods and euphausiids exposed to oxygen minimum zone conditions. *Biogeosciences* 13, 2241–2255. doi: 10.5194/bg-13-2241-2016
- Knauer, G. A., Martin, J. H., and Bruland, K. W. (1979). Fluxes of particulate carbon, nitrogen, and phosphorus in the upper water column of the northeast Pacific. *Deep Sea Res. Part Oceanogr. Res. Pap.* 26, 97–108. doi: 10.1016/0198-0149(79)90089-X
- Kones, J. K., Soetaert, K., van Oevelen, D., and Owino, J. O. (2009). Are network indices robust indicators of food web functioning? a monte carlo approach. *Ecol. Model.* 220, 370–382. doi: 10.1016/j.ecolmodel.2008.10.012
- Krause, J. W., Brzezinski, M. A., Goericke, R., Landry, M. R., Ohman, M. D., Stukel, M. R., et al. (2015). Variability in diatom contributions to biomass, organic matter production and export across a frontal gradient in the California Current Ecosystem. *J. Geophys. Res. Oceans* 120, 1032–1047. doi: 10.1002/2014JC010472
- Kroes, H. W. (1977). The niche structure of ecosystems. *J. Theor. Biol.* 65, 317–326. doi: 10.1016/0022-5193(77)90327-7
- Landry, M. R., Brown, S. L., Rii, Y. M., Selph, K. E., Bidigare, R. R., Yang, E. J., et al. (2008). Depth-stratified phytoplankton dynamics in cyclone opal, a subtropical mesoscale eddy. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 55, 1348–1359. doi: 10.1016/j.dsr2.2008.02.001
- Landry, M. R., Ohman, M. D., Goericke, R., Stukel, M. R., Barbeau, K. A., Bundy, R., et al. (2012). Pelagic community responses to a deep-water front in the California Current Ecosystem: overview of the a-front study. *J. Plankton Res.* 34, 739–748. doi: 10.1093/plankt/fbs025
- Landry, M. R., Ohman, M. D., Goericke, R., Stukel, M. R., and Tsyrklevich, K. (2009). Lagrangian studies of phytoplankton growth and grazing relationships in a coastal upwelling ecosystem off Southern California. *Prog. Oceanogr.* 83, 208–216. doi: 10.1016/j.pocean.2009.07.026
- Landry, M. R., Selph, K. E., Taylor, A. G., Décima, M., Balch, W. M., and Bidigare, R. R. (2011). Phytoplankton growth, grazing and production balances in the HNLC equatorial Pacific. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 58, 524–535. doi: 10.1016/j.dsr2.2010.08.011
- Laws, E. A., D'Sa, E., and Naik, P. (2011). Simple equations to estimate ratios of new or export production to total production from satellite-derived estimates of sea surface temperature and primary production. *Limnol. Oceanogr. Methods* 9, 593–601. doi: 10.4319/lom.2011.9.593
- Laws, E. A., Landry, M. R., Barber, R. T., Campbell, L., Dickson, M.-L., and Marra, J. (2000). Carbon cycling in primary production bottle incubations: inferences from grazing experiments and photosynthetic studies using ¹⁴C and ¹⁸O in the Arabian Sea. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 47, 1339–1352. doi: 10.1016/S0967-0645(99)00146-0
- Lefevre, D., Minas, H. J., Minas, M., Robinson, C., Le, B., Williams, P. J., et al. (1997). Review of gross community production, primary production, net community production and dark community respiration in the Gulf of Lions. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 44, 801–832. doi: 10.1016/S0967-0645(96)00091-4
- Longhurst, A. R., Bedo, A. W., Harrison, W. G., Head, E. J. H., and Sameoto, D. D. (1990). Vertical flux of respiratory carbon by oceanic diel migrant biota. *Deep Sea Res. Part Oceanogr. Res. Pap.* 37, 685–694. doi: 10.1016/0198-0149(90)90098-G
- Marra, J. (2009). Net and gross productivity: weighing in with ¹⁴C. *Aquat. Microb. Ecol.* 56, 123–131. doi: 10.3354/ame01306
- Milligan, A. J., Halsey, K. H., and Behrenfeld, M. J. (2015). Horizons: advancing interpretations of ¹⁴C-uptake measurements in the context of phytoplankton physiology and ecology. *J. Plankton Res.* 37, 692–698. doi: 10.1093/plankt/fbv051
- Morales, C. E. (1999). Carbon and nitrogen fluxes in the oceans: the contribution by zooplankton migrants to active transport in the North Atlantic during the joint global ocean flux study. *J. Plankton Res.* 21, 1799–1808. doi: 10.1093/plankt/21.9.1799
- Morozov, A. Y., and Kuzenkov, O. A. (2016). Towards developing a general framework for modelling vertical migration in zooplankton. *J. Theor. Biol.* 405, 17–28. doi: 10.1016/j.jtbi.2016.01.011
- Morrow, R. M., Ohman, M. D., Goericke, R., Kelly, T. B., Stephens, B. M., and Stukel, M. R. (2018). CCE V: primary production, mesozooplankton grazing, and the biological pump in the California Current Ecosystem: variability and response to El Niño. *Deep Sea Res. Part Oceanogr. Res. Pap.* 140, 52–62. doi: 10.1016/j.dsr.2018.07.012
- Ohman, M. D., Powell, J. R., Picheral, M., and Jensen, D. W. (2012). Mesozooplankton and particulate matter responses to a deep-water frontal system in the southern California Current System. *J. Plankton Res.* 34, 815–827. doi: 10.1093/plankt/fbs028
- Ohman, M. D., and Romagnan, J. B. (2016). Nonlinear effects of body size and optical attenuation on diel vertical migration by zooplankton. *Limnol. Oceanogr.* 61, 765–770. doi: 10.1002/lno.10251
- Peterson, B. J. (1980). Productivity and the ¹⁴C-CO₂ method: a history of the productivity problem. *Ecology* 11, 359–385. doi: 10.1146/annurev.es.11.110180.002043
- Pike, S. M., Buesseler, K. O., Andrews, J., and Savoye, N. (2005). Quantification of ²³⁴Th recovery in small volume sea water samples by inductively

- coupled plasma-mass spectrometry. *J. Radioanal. Nucl. Chem.* 263, 355–360. doi: 10.1007/s10967-005-0062-9
- Powell, J. R., and Ohman, M. D. (2012). Use of glider-class acoustic doppler profilers for estimating zooplankton biomass. *J. Plankton Res.* 34, 563–568. doi: 10.1093/plankt/fbs023
- Powell, J. R., and Ohman, M. D. (2015). Changes in zooplankton habitat, behavior, and acoustic scattering characteristics across glider-resolved fronts in the Southern California Current System. *Prog. Oceanogr.* 134, 77–92. doi: 10.1016/j.pocean.2014.12.011
- Robinson, C., Steinberg, D. K., Anderson, T. R., Aristegui, J., Carlson, C. A., Frost, J. R., et al. (2010). Mesopelagic zone ecology and biogeochemistry – a synthesis. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 57, 1504–1518. doi: 10.1016/j.dsr2.2010.02.018
- Sailley, S. F., Ducklow, H. W., Moeller, H. V., Fraser, W. R., Schofield, O. M., Steinberg, D. K., et al. (2013). Carbon fluxes and pelagic ecosystem dynamics near two western Antarctic Peninsula adélie penguin colonies: an inverse model approach. *Mar. Ecol. Prog. Ser.* 492, 253–272. doi: 10.3354/meps10534
- Saint-béat, B., Vézina, A. F., Asmus, R., Asmus, H., and Niquil, N. (2013). The mean function provides robustness to linear inverse modelling flow estimation in food webs: a comparison of functions derived from statistics and ecological theories. *Ecol. Model.* 258, 53–64. doi: 10.1016/j.ecolmodel.2013.01.023
- Samo, T. J., Pedler, B. E., Ball, G. I., Pasulka, A. L., Taylor, A. G., Aluwihare, L. I., et al. (2012). Microbial distribution and activity across a water mass frontal zone in the California Current Ecosystem. *J. Plankton Res.* 34, 802–814. doi: 10.1093/plankt/fbs048
- Savoye, N., Benitez-Nelson, C., Burd, A. B., Cochran, J. K., Charette, M., Buesseler, K. O., et al. (2006). ²³⁴Th sorption and export models in the water column: a review. *Mar. Chem.* 100, 234–249. doi: 10.1016/j.marchem.2005.10.014
- Seibel, B. A. (2011). Critical oxygen levels and metabolic suppression in oceanic oxygen minimum zones. *J. Exp. Biol.* 214, 326–336. doi: 10.1242/jeb.049171
- Siegel, D. A., Buesseler, K. O., Doney, S. C., Sailley, S. F., Behrenfeld, M. J., and Boyd, P. W. (2014). Global assessment of ocean carbon export by combining satellite observations and food-web models. *Glob. Biogeochem. Cycles* 28, 181–196. doi: 10.1002/2013GB004743
- Steinberg, D. K., Carlson, C. A., Bates, N. R., Goldthwait, S. A., Madin, L. P., and Michaels, A. F. (2000). Zooplankton vertical migration and the active transport of dissolved organic and inorganic carbon in the Sargasso Sea. *Deep Sea Res. Part Oceanogr. Res. Pap.* 47, 137–158. doi: 10.1016/S0967-0637(99)00052-7
- Steinberg, D. K., and Landry, M. R. (2017). Zooplankton and the ocean carbon cycle. *Annu. Rev. Mar. Sci.* 9, 413–444. doi: 10.1146/annurev-marine-010814-015924
- Steinberg, D. K., Van Mooy, B. A. S., Buesseler, K. O., Boyd, P. W., Kobari, T., and Karl, D. M. (2008). Bacterial vs. zooplankton control of sinking particle flux in the ocean's twilight zone. *Limnol. Oceanogr.* 53, 1327–1338. doi: 10.4319/lo.2008.53.4.1327
- Straile, D. (1997). Gross growth efficiencies of protozoan and metazoan zooplankton and their dependence on food concentration, predator-prey weight ratio, and taxonomic group. *Limnol. Oceanogr.* 42, 1375–1385. doi: 10.4319/lo.1997.42.6.1375
- Stukel, M. R., Décima, M., and Kelly, T. B. (2018a). A new approach for incorporating ¹⁵N isotopic data into linear inverse ecosystem models with markov chain monte carlo sampling. *PLoS ONE* 13:e0199123. doi: 10.1371/journal.pone.0199123
- Stukel, M. R., Décima, M., Landry, M. R., and Selph, K. E. (2018b). Nitrogen and isotope flows through the costa rica dome upwelling ecosystem: the crucial mesozooplankton role in export flux. *Glob. Biogeochem. Cycles* 32, 1815–1832. doi: 10.1029/2018GB005968
- Stukel, M. R., Kahru, M., Benitez-Nelson, C. R., Décima, M., Goericke, R., Landry, M. R., et al. (2015). Using lagrangian-based process studies to test satellite algorithms of vertical carbon flux in the eastern North Pacific Ocean. *J. Geophys. Res. Oceans* 120, 7208–7222. doi: 10.1002/2015JC011264
- Stukel, M. R., Kelly, T. B., Aluwihare, L. I., Barbeau, K. A., Goericke, R., Krause, J. W., et al. (2019). The Carbon:²³⁴Thorium ratios of sinking particles in the California current ecosystem I: relationships with plankton ecosystem dynamics. *Mar. Chem.* 212, 1–15. doi: 10.1016/j.marchem.2019.01.003
- Stukel, M. R., Landry, M. R., Ohman, M. D., Goericke, R., Samo, T., and Benitez-nelson, C. R. (2012). Do inverse ecosystem models accurately reconstruct plankton trophic flows? comparing two solution methods using field data from the California current. *J. Mar. Syst.* 91, 20–33. doi: 10.1016/j.jmarsys.2011.09.004
- Stukel, M. R., Ohman, M. D., Kelly, T. B., and Biard, T. (2019). The roles of filter-feeding and flux-feeding zooplankton as gatekeepers of particle flux into the mesopelagic ocean. *Front. Mar. Sci.* 6:397. doi: 10.3389/fmars.2019.00397
- Stukel, M. R., Ohman, M. D., Benitez-Nelson, C. R. C., and Landry, M. M. R. (2013). Contributions of mesozooplankton to vertical carbon export in a coastal upwelling system. *Mar. Ecol. Prog. Ser.* 491, 47–65. doi: 10.3354/meps10453
- Stukel, M. R., Song, H., Goericke, R., and Miller, A. J. (2018c). The role of subduction and gravitational sinking in particle export, carbon sequestration, and the remineralization length scale in the California Current Ecosystem. *Limnol. Oceanogr.* 63, 363–383. doi: 10.1002/lno.10636
- Tanaka, T., and Rassoulzadegan, F. (2004). Vertical and seasonal variations of bacterial abundance and production in the mesopelagic layer of the NW Mediterranean Sea: bottom-up and top-down controls. *Deep Sea Res. Part Oceanogr. Res. Pap.* 51, 531–544. doi: 10.1016/j.dsr.2003.12.001
- Taylor, A. G., Goericke, R., Landry, M. R., Selph, K. E., Wick, D. A., and Roadman, M. J. (2012). Sharp gradients in phytoplankton community structure across a frontal zone in the California Current ecosystem. *J. Plankton Res.* 34, 778–789. doi: 10.1093/plankt/fbs036
- Teuber, L., Kiko, R., Séguin, F., and Auel, H. (2013). Respiration rates of tropical Atlantic copepods in relation to the oxygen minimum zone. *J. Exp. Mar. Biol. Ecol.* 448, 28–36. doi: 10.1016/j.jembe.2013.06.012
- van den Meersche, K., Soetaert, K., van Oevelen, D., Meersche, K., Van den, Soetaert, K., and Van Oevelen, D. (2009). xsample(): an R function for sampling linear inverse problems. *J. Stat. Softw. Code Snippets* 30, 1–15. doi: 10.18637/jss.v030.c01
- van Oevelen, D., Soetaert, K., and Heip, C. (2012). Carbon flows in the benthic food web of the porcupine abyssal plain: the (un)importance of labile detritus in supporting microbial and faunal carbon demands. *Limnol. Oceanogr.* 57, 645–664. doi: 10.4319/lo.2012.57.2.0645
- van Oevelen, D., van den Meersche, K., Meysman, F. J. R., Soetaert, K., Middelburg, J. J., and Vézina, A. F. (2010). Quantifying food web flows using linear inverse models. *Ecosystems* 13, 32–45. doi: 10.1007/s10021-009-9297-6
- Vézina, A., Platt, T., Vezina, A. F., and Platt, T. (1988). Food web dynamics in the ocean. I. Best-estimates of flow networks using inverse methods. *Mar. Ecol. Prog. Ser.* 42, 269–287. doi: 10.3354/meps042269
- Wilson, S. E., Steinberg, D. K., and Buesseler, K. O. (2008). Changes in fecal pellet characteristics with depth as indicators of zooplankton repackaging of particles in the mesopelagic zone of the subtropical and subarctic North Pacific Ocean. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 55, 1636–1647. doi: 10.1016/j.dsr2.2008.04.019
- Wilson, S. E., Steinberg, D. K., Chu, F.-L. E., and Bishop, J. K. B. (2010). Feeding ecology of mesopelagic zooplankton of the subtropical and subarctic North Pacific Ocean determined with fatty acid biomarkers. *Deep Sea Res. Part Oceanogr. Res. Pap.* 57, 1278–1294. doi: 10.1016/j.dsr.2010.07.005
- Yebra, L., Almeida, C., and Hernández-León, S. (2005). Vertical distribution of zooplankton and active flux across an anticyclonic eddy in the Canary Island waters. *Deep Sea Res. Part Oceanogr. Res. Pap.* 52, 69–83. doi: 10.1016/j.dsr.2004.08.010
- Yokokawa, T., Yang, Y., Motegi, C., and Nagata, T. (2013). Large-scale geographical variation in prokaryotic abundance and production in meso- and bathypelagic zones of the central Pacific and Southern ocean. *Limnol. Oceanogr.* 58, 61–73. doi: 10.4319/lo.2013.58.1.0061

Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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APPENDIX A

TABLE A1 | Measurement constraints used in the LIEM.

Quantity	P07.1		P0704-2		P0704-4		P0810-1		P0810-2		P0810-3		P0810-4		P0810-5		P0810-6	
	μ	σ	μ	σ	μ	σ	μ	σ	μ	σ	μ	σ	μ	σ	μ	σ	μ	σ
1 C-14PP	1233	74	587	35	2314	139	554	33	484	29	893	54	674	40	1672	100	325	19
2 Delta PHY	-322	209	18	253	-80	602	538	1680	-1	50	-76	166	-399	4771	-325	173	-5	62
3 Microzooplankton Grazing	659	238	717	135	1282	225	472	301	243	51	357	74	277	41	498	48	138	43
4 SMZ Grazing	2249	1132	123	37	856	652	21	9	90	26	199	53	48*	21*	836*	688*	85*	37*
5 LMZ Grazing	93	87	42	18	123	55	18	12	14	9	66	26	21*	15*	51*	55*	19*	14*
6 vmSMZ Grazing	2092	907	137	44	669	499	50	30	110	34	249	113	72*	50*	753*	538*	101*	47*
7 vmlMZ Grazing	232	171	42	18	132	42	19	11	20	10	51	25	19*	14*	95*	101*	22*	16*
8 SMZ + LMZ Grazing	2342	1077	166	57	979	763	35	31	104	44	265	65	67*	38*	887*	668*	104*	61*
9 vmlMZ + vmSMZ Grazing	2306	1137	203	82	622	238	86	42	165	11	336	163	107*	70*	809*	650*	150*	41*
10 Sed Trap @ 100m	144	13	32	6	170	41	74	11	69	13	78	7	149	36	127	22	107	5
11 Thorium @ 100m	77	11	32	10	121	45	32	60	51	29	18	22	46	38	53	13	54	8
12 Fecal Pellet Flux @ 100m	135		4		54		4		8		5		35		14		4	
13 Minimum Subduction	39		47		25		11		33		13		26		10		31	
14 Maximum Subduction	79		55		45		33		47		15		51		19		59	
15 Epi Bacterial Prod (0-100m)	66	27	22	12	53	30	240	137	80	17	148	21	351	60	400	37	101	8
16 Min dBAC BP	16		5		8		16		13		14		27		102		21	
17 Max dBAC BP	52		16		27		52		41		45		89		336		69	
18 Deep NM Resp (dMYC)	2.8	5.5	3.4	1.0	2.2	0.7	4.6	2.2	3.6	1.7	3.8	0.8	3.8	4.4	6.7	0.9	7.3	1.8
19 Deep NM Poop (dMYC)	1.2	2.4	1.5	0.5	1.0	0.3	2.1	1.0	1.6	0.7	1.7	0.4	1.7	1.9	3.0	0.4	3.3	0.8
20 Deep NM Mort (dMYC)	2.8	5.5	3.4	1.1	2.2	0.7	4.7	2.3	3.7	1.7	3.9	0.8	3.9	4.4	6.8	1.0	7.4	1.8
21 Epi VM Resp (vmMYC)	7.5	14.7	2.8	2.2	12.2	4.8	12.0	2.8	2.8	0.6	14.5	4.1	8.0	4.8	8.9	1.0	5.5	2.0
22 Deep VM Resp (vmMYC)	9.5	18.6	3.5	2.8	15.5	6.0	14.9	3.4	3.0	0.7	16.9	5.0	9.7	6.0	10.6	0.8	6.2	2.1
23 Deep VM Poop (vmMYC)	5.1	10.0	1.9	1.5	8.3	3.2	8.1	1.9	1.8	0.4	9.5	2.8	5.3	3.3	5.9	0.5	3.5	1.3
24 Deep VM Mort (vmMYC)	6.0	11.7	2.2	1.8	9.8	3.8	9.5	2.2	2.1	0.5	11.1	3.2	6.3	3.8	6.9	0.6	4.1	1.5

Values given show the mean (μ) and 1 SD (σ) for each cycle except for min/max constraints which are blank. Marked values (*) were assumed values calculated from cycles of the same classification (see section Mesozooplankton and Nekton Constraints). All values are given in $mg\ C\ m^{-2}\ d^{-1}$.

TABLE A2 | Mesozooplankton biomass and minimum respiration estimates used in the LIEM.

Quantity	P0704-1	P0704-2	P0704-4	P0810-1	P0810-2	P0810-3	P0810-4	P0810-5	P0810-6
Epipeleagic Min Resp (SAR)	8.0	8.6	8.0	0.0	0.0	0.0	0.0	8.9	0.0
SMZ Min Resp	24.4	5.1	20.6	4.8	26.3	18.0	36.3	31.6	27.3
vmSMZ Min Resp	5.3	0.4	5.5	16.4	3.3	8.6	0.0	21.5	0.0
dSMZ Min Resp	2.6	2.6	5.7	9.4	2.5	7.1	1.1	14.5	3.3
LMZ Min Resp	7.0	3.3	11.9	5.5	10.4	5.7	7.8	3.0	5.8
vmLMZ Min Resp	19.9	1.5	63.2	24.7	14.9	50.9	11.1	64.2	11.8
dLMZ Min Resp	3.2	5.4	6.7	20.4	7.6	13.3	4.9	31.3	13.5
GEL Min Resp	0.2	0.1	0.1	0.3	0.7	0.2	0.0	0.0	0.7
dGEL Min Respiration	0.4	0.1	0.5	0.4	1.1	0.3	0.1	0.4	0.8
SMZ Biomass	654	106	511	106	833	509	1249	1891	1346
vmSMZ Biomass	104	33	115	722	127	282	0	1120	0
dSMZ Biomass	499	280	715	968	307	834	111	1430	381
LMZ Biomass	424	173	478	971	586	1126	867	938	1763
vmLMZ Biomass	557	347	5175	993	1557	9859	1681	6122	1671
dLMZ Biomass	5928	2839	5702	7180	1818	3579	1395	7134	3942

Respiration is given in $mg\ C\ m^{-2}\ d^{-1}$ and biomass in $mg\ C\ m^{-2}$.