



Moderate Increase in TCO₂ Enhances Photosynthesis of Seagrass *Zostera japonica*, but Not *Zostera marina*: Implications for Acidification Mitigation

Cale A. Miller^{1*}, Sylvia Yang² and Brooke A. Love¹

¹ Department of Environmental Science, Huxley College of the Environment, Western Washington University, Bellingham, WA, United States, ² Shannon Point Marine Center, Western Washington University, Anacortes, WA, United States

OPEN ACCESS

Edited by:

Bernardo Duarte,
Mare - Marine and Environmental
Sciences Centre, Portugal

Reviewed by:

Rodrigo Riera,
Atlantic Environmental Marine Center
(CIMA SL), Spain
Iris Eline Hendriks,
University of the Balearic Islands,
Spain

*Correspondence:

Cale A. Miller
camiller16@alaska.edu

† Present Address:

Cale A. Miller,
Department of Marine Biology, College
of Fisheries and Ocean Sciences
Fairbanks, AK, United States

Specialty section:

This article was submitted to
Marine Pollution,
a section of the journal
Frontiers in Marine Science

Received: 06 December 2016

Accepted: 06 July 2017

Published: 20 July 2017

Citation:

Miller CA, Yang S and Love BA (2017)
Moderate Increase in TCO₂ Enhances
Photosynthesis of Seagrass *Zostera
japonica*, but Not *Zostera marina*:
Implications for Acidification
Mitigation. *Front. Mar. Sci.* 4:228.
doi: 10.3389/fmars.2017.00228

Photosynthesis and respiration are vital biological processes that shape the diurnal variability of carbonate chemistry in nearshore waters, presumably ameliorating (daytime) or exacerbating (nighttime) short-term acidification events, which are expected to increase in severity with ocean acidification (OA). Biogenic habitats such as seagrass beds have the capacity to reduce CO₂ concentration and potentially provide refugia from OA. Further, some seagrasses have been shown to increase their photosynthetic rate in response to enriched total CO₂ (TCO₂). Therefore, the ability of seagrass to mitigate OA may increase as concentrations of TCO₂ increase. In this study, we exposed native *Zostera marina* and non-native *Zostera japonica* seagrasses from Padilla Bay, WA (USA) to various levels of irradiance and TCO₂. Our results indicate that the average maximum net photosynthetic rate (P_{max}) for *Z. japonica* as a function of irradiance and TCO₂ was 3x greater than *Z. marina* when standardized to chlorophyll ($360 \pm 33 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ and $113 \pm 10 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$, respectively). Additionally, *Z. japonica* increased its P_{max} ~50% when TCO₂ increased from ~1,770 to 2,051 $\mu\text{mol TCO}_2 \text{ kg}^{-1}$. In contrast, *Z. marina* did not display an increase in P_{max} with higher TCO₂, possibly due to the variance of photosynthetic rates at saturating irradiance within TCO₂ treatments (coefficient of variation: 30–60%) relative to the range of TCO₂ tested. Our results suggest that *Z. japonica* can affect the OA mitigation potential of seagrass beds, and its contribution may increase relative to *Z. marina* as oceanic TCO₂ rises. Further, we extended our empirical results to incorporate various biomass to water volume ratios in order to conceptualize how these additional attributes affect changes in carbonate chemistry. Estimates show that the change in TCO₂ via photosynthetic carbon uptake as modeled in this study can produce positive diurnal changes in pH and aragonite saturation state that are on the same order of magnitude as those estimated for whole seagrass systems. Based on our results, we predict that seagrasses *Z. marina* and *Z. japonica* both have the potential to produce short-term changes in carbonate chemistry, thus offsetting anthropogenic acidification when irradiance is saturating.

Keywords: *Zostera marina*, *Zostera japonica*, seagrass, ocean acidification, photosynthetic potential, mitigation, TCO₂

INTRODUCTION

The uptake of CO₂ from anthropogenic fossil fuel emissions by the global oceans is shifting the acid-base balance of the carbonate system in a process known as ocean acidification (OA). Increasing CO₂ concentration outpaces the natural buffering capacity of seawater and increases the total CO₂ (TCO₂), which is the sum of all forms of carbonic acid and its conjugate bases (Doney et al., 2009; Hönlisch et al., 2012). The dissolution of anthropogenic CO₂ in nearshore waters interacts with a host of other processes that drive the dynamics of nearshore carbonate chemistry, such as biological metabolism, riverine discharge and associated organic matter composition, tidal pumping, upwelling, nutrient input, and eutrophication (Feely et al., 2008, 2010; Cai, 2011; Duarte et al., 2013; Waldbusser and Salisbury, 2014; Wallace et al., 2014). The synergy of these factors induces high variability to the carbonate system, and results in periodic and episodic decreases in pH and aragonite saturation state (Ω_{ar}) that are more extreme than the ~ 0.4 pH and ~ 1.5 Ω_{ar} decreases predicted for global ocean averages by the year 2100 (Ciais et al., 2013; Duarte et al., 2013; Waldbusser and Salisbury, 2014). Among these drivers, photosynthesis and respiration are the dominant processes controlling coastal ocean carbonate chemistry (Gattuso et al., 1998; Sunda and Cai, 2012; Waldbusser and Salisbury, 2014). The diurnal variability of biological photosynthesis and respiration can, therefore, either dampen or amplify the magnitude of extreme carbonate chemistry events, which may occur from episodic influxes of fresh water or upwelling. Gaining a better understanding of how these biological signals modify and potentially ameliorate acidification is imperative, particularly when the effects of acidification can impact the economic and social stability of coastal human communities that are dependent on ocean resources (Ekstrom et al., 2015).

Many marine calcifiers and a variety of other marine species will be negatively affected by OA as decreases in pH, calcium carbonate saturation state and the substrate-to-inhibitor ratio— $[\text{HCO}_3^-]/[\text{H}^+]$ —have been shown to inhibit calcification, growth, and acid-base regulation (Pörtner, 2008; Kroeker et al., 2013; Thomsen et al., 2015; Waldbusser et al., 2015; Fassbender et al., 2016). Conversely, some autotrophic organisms may directly benefit from the increase in seawater CO₂ associated with acidification, which can stimulate photosynthesis and increase growth (reviewed in Kroeker et al., 2010, 2013; Koch et al., 2013). For example, photosynthetic rates of many seagrass species have been shown to increase with TCO₂ (Beer and Koch, 1996; Thom, 1996; Zimmerman et al., 1997; Invers et al., 2001; Cox et al., 2016; Ow et al., 2016). In addition, some studies have concluded that the carbon uptake by seagrass beds may mitigate acidification on specific spatial and temporal scales when irradiance is high (Manzello et al., 2012; Unsworth et al., 2012; Hendriks et al., 2014).

OA mitigation by seagrass arises from the removal of CO₂ from seawater through photosynthesis, which raises pH and reduces TCO₂ thus minimizing acidification on short timescales (Marbá et al., 2006; Unsworth et al., 2012). This direct action of OA mitigation by seagrass is, however, reversed during

times of dark respiration, which transiently increases TCO₂. For most seagrass systems, periods of high photosynthetic activity are robust enough to reduce TCO₂ in spite of the enhanced remineralization of allochthonous organic matter, which accumulates within seagrass beds and fuels the production of TCO₂ via respiration (Koch et al., 2006; Duarte et al., 2013; Hendriks et al., 2014; Baumann et al., 2015). Since most seagrass systems tend to be net autotrophic, or at least seasonally net autotrophic, they can sequester copious amounts of carbon on seasonal timescales that correspond to changes in above-ground biomass, while carbon can be sequestered on longer timescales by below-ground retention in the sediment (Duarte et al., 2005, 2010; Chung et al., 2011; McLeod et al., 2011; Fourqurean et al., 2012; Unsworth et al., 2012; Marbá et al., 2015; Poppe, 2016). Estimated values of organic carbon retained in seagrass sediments to a depth of 1 m range from 9.1 to 625 mg ha⁻¹, with actual measurements extending to 829 mg ha⁻¹ (Fourqurean et al., 2012). Despite the importance of carbon storage, it is the instantaneous photosynthetic carbon uptake that drives OA mitigation potential on short timescales.

Presumably, extreme acidification events in nearshore coastal waters will increase in frequency, duration, and magnitude as a result of anthropogenic OA (Harris et al., 2013; Hauri et al., 2013). For early life stage calcifiers undergoing rapid growth and development, sensitivity to OA is heightened and driven by the duration and intensity of exposure (Kurihara, 2008; Talmage and Gobler, 2009; Hettlinger et al., 2012; Waldbusser et al., 2015). Therefore, an increase in acidification severity will likely result in chemical conditions that periodically surpass physiological thresholds of resident organisms already living near their tolerance limits (Grantham et al., 2004; Waldbusser and Salisbury, 2014). Instantaneous carbon uptake rates by seagrasses, however, may be able to expand periods of favorable carbonate chemistry for these sensitive species, or dampen the episodic extremes of acidification on hourly timescales, thus lessening the exposure to extreme acidification.

In the U.S. Pacific Northwest (PNW), two *Zostera* seagrass species are among the seagrasses commonly found in soft-sediment habitats in the Salish Sea: native *Zostera marina* L. and non-native *Zostera japonica* Ascher & Graebner (Harrison and Bigley, 1982). The native seagrass *Z. marina* has specifically been identified by Washington State as a biological means to ameliorate acidification (Washington State Blue Ribbon Panel on Ocean Acidification Ocean Acidification., 2012). In addition, both *Z. marina* and *Z. japonica* have strong habitat-associations with organisms vulnerable to OA, such as bivalves (Ferraro and Cole, 2012; Mach et al., 2014; Dumbauld and McCoy, 2015). The non-native *Z. japonica* has colonized previously unvegetated mudflats and is found in the mid to upper intertidal zone, whereas *Z. marina* has a distribution extending from the lower intertidal to shallow subtidal region; species overlap between *Z. marina* and *Z. japonica* can occur in the lower intertidal zone on flat shorelines (Harrison, 1982; Thom, 1990; Kaldy, 2006; Ruesink et al., 2010). The increasing presence and distribution of *Z. japonica* in the PNW warrants an inclusion of this non-native species when examining the potential of seagrass in the Salish Sea to mitigate OA.

Research has shown that seagrass carbon uptake rates are species-specific and vary in response to altered carbonate chemistry (Campbell and Fourqurean, 2013; Koch et al., 2013). Thus far, *Z. marina* has been the focus of many studies examining its photosynthetic response to increases in TCO_2 (Beer and Koch, 1996; Thom, 1996; Zimmerman et al., 1997; Koch et al., 2013); however, this is not the case for *Z. japonica*. There have not been any published studies to our knowledge examining *Z. japonica* response to TCO_2 . Only a handful of studies have examined *Z. japonica* photosynthesis and physiology for populations established in the PNW (see Shafer et al., 2011; Shafer and Kaldy, 2014; Kaldy et al., 2015). Shafer and Kaldy (2014) found that *Z. japonica* local to the central Oregon coast has a substantially greater photosynthetic rate than *Z. marina* under the same chemical conditions. While this study provides insight into the relative photosynthetic differences between the two species, a comprehensive understanding of variable TCO_2 and irradiance response is needed in order to determine how this difference drives OA mitigation potential. In addition, it is important to examine if the local adaptations that exist throughout a species' distribution (Backman, 1991 and references therein; Shafer et al., 2011) result in photosynthetic differences that are dissimilar between central Oregon and Salish Sea populations.

In this study, we conducted a series of laboratory experiments to better understand how *Z. marina* and *Z. japonica* may alter carbonate chemistry under conditions of elevated TCO_2 and variation in light intensity, as would occur over a diurnal period. We aimed to determine how the photosynthetic carbon uptake of *Z. marina* and *Z. japonica* shift the carbonate system on short timescales potentially counteracting acidification. Specifically, we conducted one experiment per species to (1) examine the differences in species' photosynthetic rates (i.e., carbon uptake rates), (2) quantify the response of photosynthetic rates to differing levels of TCO_2 , and (3) use those results to estimate how the photosynthetic carbon uptake could induce hourly changes on the carbonate system over a diurnal cycle at various biomass to water volume ratios.

MATERIALS AND METHODS

Sample Site and Collection

Padilla Bay, Washington, is a tidally dominated estuary in the Salish Sea, and is a part of the National Estuarine Research Reserve System (48°31'14.1"N, 122°35'24.4"W). The *Z. marina* and *Z. japonica* meadows in Padilla Bay constitute a submerged and emergent total area of ~4,000 ha, where *Z. marina* accounts for ~3,000 of the total area (Bulthuis, 2013). The higher intertidal region is dominated by *Z. japonica*, which is morphologically different than *Z. marina* and has a leaf surface area and mass that is ~5x less than that of *Z. marina*. *Z. japonica* constitutes ~22% of the seagrass biomass in Padilla Bay, and the spread of the non-native species has increased the total areal extent of seagrass acreage. In the intertidal where *Z. japonica* overlaps with *Z. marina*, a 250 m transect was marked for shoot collection, which occurred every 25 m (48°29'36.6"N, 122°29'8.5"W). Two weeks prior to shoot collection, two HOBO Pendant Temperature/Light 64k data loggers were attached to the transect marking poles

~1 m above the sediment (above the height of the canopy) to capture *in situ* irradiance at the sample site which was used to generate a realistic but idealized light field to drive model visualizations. HOBO data (in units of lux) were converted to daylight photon flux density (PFD) as described by Thimijan and Heins (1983). The measurements from the two sensors were averaged and smoothed, and this curve was scaled to the maximum value recorded by land-based PAR measurements at the Padilla Bay farm station. It is important to note that the HOBO light sensors only measure in the planar flux and not the spherical or scalar flux, which is a more comprehensive irradiance flux; therefore, the sensor data are used to define the shape of the curve, while the scaling should bring the overall range of irradiance values within realistic bounds.

Approximately 200 shoots of each species were collected from Padilla Bay along the entire 250 m transect. Healthy-looking adult *Z. marina* shoots with intact rhizomes were collected by hand during low tide (0.8 m MLLW) on August 16, 2015 from Padilla Bay, placed in a cooler, and transported to Shannon Point Marine Center in Anacortes, Washington, within 1 h of collection. Ten days later on August 26, 2015, adult *Z. japonica* shoots with intact rhizomes were collected in the same manner at low tide (~0.0 MLLW) for a second experiment. One hundred of the most healthy-looking shoots (i.e., shoots without visible damage and well-preserved rhizomes) were then haphazardly selected, rinsed with seawater, and dispersed among four separate 40-l acrylic flow through tanks under low PFD (~50 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$) on a 12:12 L:D cycle for ~48 h before experimentation.

Experimental Design

We used a 5 × 5 factorial design that targeted 25 treatment CO_2 and light combinations: estimated PFD levels 0, 40, 200, 500, and 750 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$, and $p\text{CO}_2$ of 140, 250, 400, 650, and 900 μatm , where the 400 $p\text{CO}_2$ treatment is the approximate ambient condition for the open ocean. Treatment $p\text{CO}_2$ values were selected because they correspond to an observed range of diurnal *in situ* variability of the carbonate system in Padilla Bay (Love et al., 2016). Even though $p\text{CO}_2$ was the variable used to prepare treatments by gas equilibration, TCO_2 is a more appropriate metric for our experiment given that CO_2 and HCO_3^- are presumably both utilized for photosynthesis; therefore, we identify treatment levels by initial TCO_2 rather than $p\text{CO}_2$ henceforth. In order to achieve the full factorial design, we used closed experimental incubation vials housing leaf segments rather than whole shoots in large open aquaria. Each treatment combination had quadruplicate replication and duplicate blanks for a total of 100 leaf-segment and 50 blank vials. Blanks had the same PFD exposure and initial TCO_2 but lacked leaf segments, and were used to account for any changes in seawater chemistry induced by microbial activity.

During each experiment, 150 incubation vials (20 ml borosilicate scintillation vials with a polyethylene cone-shaped liner) were placed in five clear acrylic water bath trays (56.5 × 7.62 × 3.81 cm) fitted with flow-through seawater for temperature control. Each water bath tray was mounted to a single 1.5 cm clear acrylic sheet, and placed directly above an individual bulb in a light fixture housing five T5 high output

54W 6,500 K Spectralux bulbs. Incubation vials were nearly fully submerged (water line stopped at cap) when placed in water bath trays. Vinyl mesh wraps were constructed and fitted to incubation vials to attenuate light. Mesh covers either had one, three, or six layers, which provided a PFD range from ~ 40 to $500 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$.

Treatment Water and Leaf Segment Preparation

To obtain a spectrum of TCO_2 concentrations, seawater pumped from Guemes Channel was collected, $0.2 \mu\text{m}$ filtered, sterilized via autoclave, and distributed into four 20 L polycarbonate carboys. Four mixtures of pure CO_2 (research grade-5) and compressed ambient air stripped of CO_2 using a regenerative molecular sieve adsorber (Twin Tower Engineering, CAS2-11) were generated using four sets of Sierra SmartTrak mass flow controllers. Treatment water was bubbled for at least 72 h with gas mixtures that were first passed through humidifiers (one liter Nalgene bottles filled halfway with tap water). Carboys were held in an incubator, which maintained water temperature at approximately 12.5°C . Further details are described in the equilibration system portions of Love et al. (2017). A fifth CO_2 treatment was produced by mixing, equilibrated, identical volumes of the two lowest CO_2 mixtures.

Preparation of seagrass tissue used for incubations began approximately 20 h before each experiment. This began by selecting the middle section of the second youngest leaf in a shoot, and wiping the entire leaf clean of epiphytes. To test similar leaf surface area for both species, a 2 cm leaf segment for *Z. marina* and 4 cm leaf segment for *Z. japonica* was excised. While the middle section of the leaf was selected for, variation among shoot length of the second youngest leaf existed. Each leaf segment was then cut in two, and one half randomly frozen for chlorophyll extraction and the other placed into the incubation vial filled with 5 ml of filtered and sterilized seawater. Leaf segment stocked vials were then held in a low light incubator at $\sim 12^\circ\text{C}$ for 16–20 h until experimentation.

Immediately before each experiment, stocked vials were emptied of filtered seawater, rinsed and filled with treatment seawater. Vials were overflowed to eliminate headspace, and a 0.5-mm glass bead was inserted as a stirring mechanism. Mesh coverings (no mesh, 1 layer, 3 layer, 6 layer, and black electrical tape for opaque) were placed around vials to achieve varied levels of PFD. Vials were haphazardly placed into water bath trays, and incubated for 90 min until termination of the experiment. Vials were physically inverted by hand, three times, every 5 min to stir water inside and minimize the development of a TCO_2 poor boundary layer around leaf segments. Temperature in water bath trays was continuously monitored and recorded every 20 min with a Fluke 1523 reference thermometer from the beginning of the experiment until all vials were terminated. PFD was measured at each vial location with a QSL-101 PAR irradiance sensor (Biospherical Instruments Inc.) by setting vials with mesh coverings in their respective locations after the conclusion of the experiment. Due to the discontinuity of PFD along a bulb (i.e., vials positioned at the center of a bulb received more light

than those at either end), light measurements were treated as continuous data for all statistical analyses.

Leaf Segment and Carbonate Chemistry Analysis

Photosynthetic rates were determined from leaf segment incubations of *Z. marina* and *Z. japonica* by measuring treatment TCO_2 concentration before and after incubation. While photosynthetic rates are normalized to TCO_2 , the O_2 convention is followed. That is, positive photosynthetic rates are a negative flux of TCO_2 from the medium, while respiration is a positive flux of TCO_2 to the medium: this is opposite when O_2 normalized. Initial TCO_2 samples were collected in triplicate 20 ml scintillation vials, poisoned with $10 \mu\text{l}$ of saturated HgCl_2 , capped, wrapped with parafilm to minimize any potential gas leakage, and refrigerated at 2°C until analysis. Initial total alkalinity (TA) samples were collected in triplicate 350 ml amber glass bottles with polyurethane-lined crimp-sealed metal caps and poisoned with $30 \mu\text{l}$ of saturated HgCl_2 . Salinity was measured from each TA sample with a refractometer before poisoning occurred, and again for each treatment vial when samples were processed.

Incubations were terminated by removing vials, one at a time haphazardly across treatments, but ordinarily by replicate assignment—this allowed for an equal experimental termination time across treatments. Leaf segments were removed from the vial, marked for incubation time, and stored in an empty vial for dry-weight measurement. The experimental vial was immediately poisoned and stored in the same manner as initial condition samples. TCO_2 samples were analyzed within 5 days of each experiment using an Apollo SciTech AS-C3 dissolved inorganic carbon analyzer. TA samples were titrated within 30 days of the experiment using the open-cell method as in Dickson et al. (2007) with a Metrohm 888 Titrando. Certified reference material was used to construct a five-point standardization curve for TCO_2 and to verify accuracy of TA open-cell titration (Batch 144, A.G., Dickson, Scripps Institute of Oceanography). All other carbonate chemistry parameters were calculated using CO_2SYS (Pierrot et al., 2006) with K_1 and K_2 equilibrium constants from Mehrbach et al. (1973) and refit by Dickson and Millero (1987). Dry weight was recorded after rinsing leaf segments 3x with deionized water and drying at 55°C for at least 24 h. Frozen leaf segments were prepped for chlorophyll extraction by sonicating for 30 s in a 10 ml 90% acetone solution. Segments were then refrozen at -20°C for 24 h and centrifuged for 5 min directly before chlorophyll measurement. Extract was measured with a Trilogy fluorometer (Turner designs), acidified with 0.1 N HCl, and measured again. Chlorophyll and phaeopigment concentrations were calculated following the methods described by Lorenzen (1966).

Statistical Methods and Photosynthetic Response

Predicted photosynthetic rates for both species were determined by modeling the empirical data separately for each experiment as a function of continuous PFD, and a combined effect of

continuous PFD and TCO_2 in one integrated, iterative (600 iterations) model, using the non-linear and linear curving fitting functions and curve fitting tools in the MathWorks software Matlab (V. 2015b). Species comparison was determined from the robustness of the model fit and whether or not 95% confidence intervals overlapped for each predicted photosynthetic parameter. Photosynthetic rates were normalized to chlorophyll (a/b) rather than dry-weight in order to account for variability of pigment concentration that occurs between and along leaves (Enríquez et al., 2002). Following the methods described in Jassby and Platt (1976), the net photosynthetic rate (P_{net}) was calculated as:

$$P_{\text{net}} = P_{\text{max}} \tanh\left(\frac{\alpha E}{P_{\text{max}}}\right) + R_d \quad (1)$$

where P_{max} is the maximum photosynthetic rate ($\mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$), E is the PFD ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$), R_d is the dark respiration rate ($\mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$), and α (photosynthetic efficiency) is the initial slope of the photosynthetic-irradiance curve ($\mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1} (\mu\text{mol photons m}^{-2} \text{ s}^{-1})^{-1}$). In order to integrate the empirical instantaneous photosynthetic rates (i.e., carbon uptake) to hourly changes in the carbonate system, TCO_2 was integrated into the standard photosynthesis-irradiance curve (Equation 1) as a linear effect on P_{max} . A vector of initial TCO_2 concentration was applied to the P_{max} term, thus allowing P_{max} to vary with TCO_2 :

$$P_{\text{max}}^{\text{TC}} = P_{\text{base}} + \beta * \text{TCO}_{2i} \quad (2)$$

where $P_{\text{max}}^{\text{TC}}$ is the TCO_2 dependent P_{max} , β is the slope of the P_{max} response to TCO_2 , P_{base} is the intercept, and TCO_{2i} is the initial TCO_2 concentration. The integrated TCO_2 -irradiance model was, thus, a combination of the above equations:

$$P_{\text{net}} = P_{\text{max}}^{\text{TC}} \tanh\left(\frac{\alpha E}{P_{\text{max}}^{\text{TC}}}\right) + R_d \quad (3)$$

where P_{net} is now a function of irradiance and TCO_2 . The estimated mean P_{max} and photosynthetic efficiency (α) from the TCO_2 -irradiance model (Equation 3) output were used to calculate the saturation irradiance:

$$E_k = \frac{P_{\text{max}}^{\text{TC}}}{\alpha} \quad (4)$$

where E_k is the saturation irradiance and $P_{\text{max}}^{\text{TC}}$ and α the outputs from the TCO_2 -irradiance model (Equation 3). Since all TCO_2 -irradiance model predictions of P_{max} are TCO_2 dependent ($P_{\text{max}}^{\text{TC}}$), for simplicity, we will simply refer to them as P_{max} rather than $P_{\text{max}}^{\text{TC}}$. P_{max} -values for both species at every initial TCO_2 treatment were statistically compared based on the predicted standard error and confidence bounds from the model output. In addition, P_{max} -values as a function of initial CO_2 and HCO_3^- were observed as both carbonate species are generally utilized for photosynthesis.

Model Evaluation: Photosynthesis and Carbon Uptake

The TCO_2 -irradiance model was conceived by applying realistic PFD values representing a diurnal cycle and the experimental TCO_2 treatment values as model input parameters (Equation 3), producing hourly photosynthesis estimates and the concomitant uptake of TCO_2 . A Gaussian function was fit to field collected PFD measurements resulting in a generalized light cycle. The TCO_2 values for the ambient treatment from both experiments were averaged ($1,964 \mu\text{mol kg}^{-1}$) and utilized as a baseline concentration from which hourly photosynthetic carbon uptake could be used to calculate changes in TCO_2 at each hourly step over a diurnal period. Assuming TA constant (also averaged from both experiments— $2,134 \mu\text{mol kg}^{-1}$), the relative changes in pH and Ω_{ar} could then be calculated based on the change in TCO_2 as driven by the estimated photosynthetic carbon uptake of each species. The calculated differences were always determined from the baseline TCO_2 for each time step. This is in accordance with the initial condition of each of our experimental TCO_2 and light conditions, and is analogous to the situation when residence time in a seagrass bed is small. In the scenario where reduction in TCO_2 is compounded, concentrations in the conceived model would be reduced to minimal values rapidly upon initiation of photosynthesis and the effects of both species-specific rates and initial TCO_2 would disappear under the general carbon limitation.

Estimated Changes in Carbonate Chemistry

The experimental results were determined at a fixed ratio of biomass to water volume, constraining the ability to compare the laboratory results to other studies. Therefore, the model predicted results were examined at different biomass to water volume ratios, which would be similar to variation in seagrass density and water depth in the field. Predicted P_{max} results were applied to a simple box model consisting of a 1 m area, with seagrass at a moderate density, and carbonate chemistry predictions driven by saturating irradiance, TCO_2 and water depth. The specific changes induced on the carbonate system by both species' maximum photosynthetic potential were extrapolated from vial volume (20 ml) to different volumes of static water corresponding to various depths over a 1 m² patch of seagrass with a biomass of 100 gDW m⁻². It should be noted that we don't take into account light attenuation at these depths, but at mid-day, irradiance should still be saturating. Estimates for both species from various locations in Padilla Bay reported in Bulthuis (2013), show a range from ~60–200 gDW m⁻² with a mean of ~100 for both species. The change in TCO_2 was calculated assuming a well-mixed water column, the average mg chl gDW⁻¹ (1.88 ± 0.89 and 0.67 ± 0.33 for *Z. marina* and *Z. japonica*, respectively) for each species, and using projected P_{max} values, as:

$$\Delta \text{TCO}_2 = \text{TCO}_{2i} - \left(C_b * \left(\frac{b_i}{a} \right) * \left(\frac{1}{d} \right) * f * r \right) \quad (5)$$

where TCO_{2i} is the initial TCO_2 , C_b is mg chl gDW⁻¹, b_i is the gDW over a m² (a), d is depth (m), r is P_{max} ($\mu\text{mol TCO}_2 \text{ mg}$

$\text{chl}^{-1} \text{h}^{-1}$), and f is the unit conversion $\text{m}^3 \text{1,000 L}^{-1}$. Changes in TCO_2 from photosynthetic carbon uptake by each species were then used to determine the changes in carbonate chemistry parameters while assuming TA constant (averaged from both experiments— $2,134 \mu\text{mol kg}^{-1}$).

RESULTS

Incubation Conditions

Experimental conditions for the leaf segment incubations were similar between experiments (Table 1), which were run at nearly the exact same time of day in the mid-morning (<2 min apart). Temperature averaged $13.2 \pm 0.4^\circ\text{C}$, for the *Z. marina* experiment, and $13.4 \pm 0.1^\circ\text{C}$ for the *Z. japonica* experiment. Photon flux density (PFD) had a greater variance for the high light treatments than for the lower, more heavily shaded treatments. Salinity ranged from 30.5 to 32 ppt for the *Z. japonica* experiment (Table 1). This variation was likely a result of unequal effectiveness by the humidifiers during bubbling, or from differential evaporation in the autoclaved bottles used to fill carboys. The $p\text{CO}_2$ of the seawater treatments did not indicate complete equilibration with the gas mixtures, which ranged from 100 to $1,200 \mu\text{atm}$ whereas seawater treatment conditions only ranged from 138 to $918 \mu\text{atm}$ (Table 1). However, the treatment values did produce a large range in TCO_2 and were similar across both experiments. The high TCO_2 treatment had the greatest difference between experiments, which was $55 \mu\text{mol kg}^{-1}$ higher for the *Z. marina* experiment than for the *Z. japonica* experiment (Table 1). It should be noted that sample size by treatment TCO_2 varied from 20 to 17 due to either the malfunction of the dissolved inorganic carbon analyzer or loss of treatment seawater stored in the incubation vial.

Species-Specific Photosynthetic Response to Irradiance and TCO_2

All photosynthetic parameters for *Z. japonica* were best predicted by the TCO_2 -irradiance model (Equation 3), which had a more robust fit (RMSE = $150 \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1}$) than the standard photosynthesis-irradiance model (Equation

1), RMSE = $161 \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1}$ (Table 2). There was no difference in the predictive power between the TCO_2 -irradiance and the standard photosynthesis-irradiance models for *Z. marina*, both had an RMSE of $43 \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1}$. Based on the TCO_2 -irradiance model relating the photosynthesis-irradiance curve to TCO_2 , the mean maximum photosynthetic rate (P_{max}) for *Z. japonica* was 3x greater than *Z. marina*, and photosynthetic efficiency (α) was 4.5x greater on a per chlorophyll basis (Table 2, Figure 1). The non-linear fit for both seagrasses in Figure 1 is the mean fit over all initial TCO_2 , which displays higher variability for *Z. japonica* compared

TABLE 2 | Predicted photosynthetic parameters from the TCO_2 -irradiance (TCO_2 -IRR) model (Equation 3) and the standard photosynthesis-irradiance (P vs. E) model (Equation 1).

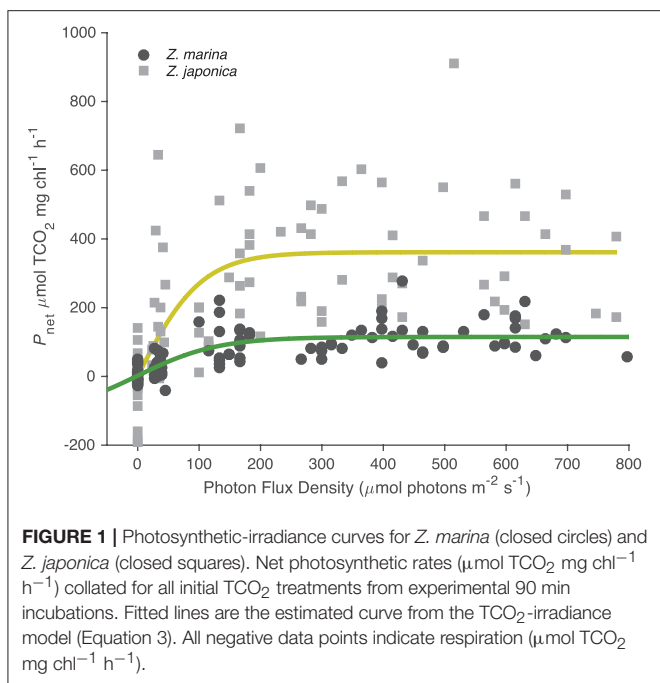
| Species | Model | Parameter | Estimate | SE | tstat | p-value |
|--------------------|-----------------------|------------------|----------|-------|--------|---------|
| <i>Z. marina</i> | TCO ₂ -IRR | P _{max} | 113 | 9.91 | 2.02 | 0.0466 |
| | | α | 0.819 | 0.211 | 3.85 | <0.001 |
| | | R _d | 2.92 | 69.68 | 0.337 | 0.737 |
| | | β | -0.054 | 0.055 | -0.986 | 0.327 |
| <i>Z. marina</i> | P vs. E | P _{max} | 113 | 10.7 | 10.5 | <0.001 |
| | | α | 0.842 | 0.211 | 3.99 | <0.001 |
| | | R _d | 2.13 | 8.77 | 0.242 | 0.809 |
| <i>Z. japonica</i> | TCO ₂ -IRR | P _{max} | 360 | 33.0 | 2.98 | 0.003 |
| | | α | 3.70 | 1.14 | 3.24 | 0.002 |
| | | R _d | 0.180 | 31.2 | 0.005 | 0.995 |
| | | β | 0.841 | 0.219 | 3.84 | <0.001 |
| <i>Z. japonica</i> | P vs. E | P _{max} | 360 | 40.3 | 8.92 | <0.001 |
| | | α | 3.47 | 1.11 | 3.13 | 0.002 |
| | | R _d | 1.42 | 33.7 | 0.042 | 0.966 |

P_{max} is the maximum photosynthetic rate, α (photosynthetic efficiency) is the initial slope of the photosynthesis-irradiance curve, and β is the slope of linear relationship between P_{max} and initial TCO_2 concentration. Total sample size for *Z. marina* and *Z. japonica*: $n = 98$ and 93 , respectively. The TCO_2 -irradiance model fit for *Z. marina* produced a RMSE of $43 \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1}$ and 150 for *Z. japonica*. Units: P_{max} and $R_d = \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1}$; $\alpha = \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1} (\mu\text{mol photons } \text{m}^{-2} \text{ s}^{-1})^{-1}$; $\beta = \mu\text{mol TCO}_2 \text{mg chl}^{-1} \text{h}^{-1} (\mu\text{mol TCO}_2 \text{kg}^{-1})^{-1}$.

TABLE 1 | Initial conditions for both *Z. marina* and *Z. japonica* experiments.

| Experiment | $p\text{CO}_2$ (μatm) treatment | TCO_2 ($\mu\text{mol kg}^{-1}$) | TA ($\mu\text{mol kg}^{-1}$) | $p\text{CO}_2$ (μatm) | pH (total) | Ω_{ar} | Temp. $^\circ\text{C}$ | Sal. ppt |
|--------------------|--|--|--------------------------------|------------------------------------|-----------------|----------------------|------------------------|----------|
| <i>Z. marina</i> | 140 | 1779 ± 2.7 | 2157 ± 15 | 140 ± 4.4 | 8.40 ± 0.01 | 4.0 ± 0.0 | 14.2 ± 0.8 | 32 |
| | 250 | 1873 ± 4.5 | 2157 ± 6.3 | 225 ± 4.0 | 8.24 ± 0.01 | 3.0 ± 0.1 | 14.4 ± 0.6 | 32 |
| | 400 | 1972 ± 1.8 | 2141 ± 7.4 | 421 ± 6.6 | 8.01 ± 0.01 | 1.9 ± 0.0 | 14.0 ± 0.2 | 32 |
| | 650 | 2083 ± 2.1 | 2197 ± 5.5 | 652 ± 18 | 7.85 ± 0.01 | 1.4 ± 0.0 | 14.4 ± 0.3 | 32 |
| | 900 | 2106 ± 1.5 | 2180 ± 1.2 | 863 ± 15 | 7.73 ± 0.01 | 1.1 ± 0.0 | 14.1 ± 0.3 | 32 |
| <i>Z. japonica</i> | 140 | 1770 ± 6.1 | 2147 ± 6.2 | 138 ± 1.8 | 8.41 ± 0.00 | 3.9 ± 0.1 | 14.0 ± 0.5 | 32 |
| | 250 | 1868 ± 1.1 | 2147 ± 2.1 | 214 ± 4.8 | 8.26 ± 0.01 | 3.0 ± 0.0 | 13.6 ± 0.4 | 31 |
| | 400 | 1956 ± 1.5 | 2127 ± 1.9 | 393 ± 8.4 | 8.03 ± 0.01 | 1.9 ± 0.0 | 13.6 ± 0.4 | 31 |
| | 650 | 2018 ± 0.8 | 2113 ± 6.2 | 684 ± 9.6 | 7.81 ± 0.01 | 1.2 ± 0.0 | 13.5 ± 0.3 | 32 |
| | 900 | 2051 ± 1.7 | 2103 ± 3.2 | 918 ± 15 | 7.70 ± 0.01 | 0.9 ± 0.1 | 13.2 ± 0.1 | 30.5 |

Measured mean values and standard deviation of TA, TCO_2 , temperature, salinity, and calculated values pH (total), and aragonite saturation state (Ω_{ar}).



to *Z. marina* indicating a greater response by *Z. japonica* to TCO_2 . The model produced a good fit for *Z. japonica* and for *Z. marina* for two out of the three photosynthetic parameters, however, the mix of positive and negative TCO_2 fluxes in the respiration vials resulted in a positive value rather than a negative value predicted by the model for *Z. marina* and *Z. japonica* respiration (intercept) (Figure 1, Table 2). Across all treatments, the mean leaf segment mg chl:gDW^{-1} ratio—used to normalize the model results—was $\sim 3\times$ greater for *Z. marina* (1.88 ± 0.89) than for *Z. japonica* (0.67 ± 0.33). Model predicted mean P_{max} for *Z. japonica* was $360 \pm 33 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ with an α of $3.70 \pm 1.1 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1} (\mu\text{mol photons m}^{-2} \text{ s}^{-1})^{-1}$, and where considered significant with standard error (SE) (Table 2). Model prediction for *Z. marina* P_{max} was $113 \pm 10 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ with an α of $0.819 \pm 0.21 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1} (\mu\text{mol photons m}^{-2} \text{ s}^{-1})^{-1}$, and were also predicted as significant (Table 2). In addition, the point at which each species experiences a saturating irradiance (E_k) was also different. *Z. japonica* displayed a lower E_k at 97 (propagated SE = 31) $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$, which was $\sim 70\%$ of *Z. marina*'s E_k of 138 (propagated SE = 37) $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$. The maximum photosynthetic rates—which were used to fit the P_{max} values—for *Z. japonica* appeared to respond positively with increasing TCO_2 (Figure 2). *Z. marina*, however, did not display any positive increase in photosynthetic rate with increasing TCO_2 . It is important to note that Figure 2A strictly shows one half of the TCO_2 -irradiance model (photosynthetic rate to TCO_2) and does not directly reflect the interaction between TCO_2 and light, as is really described by the model.

The TCO_2 -irradiance model predicted a positive increase in P_{max} to increasing carbon availability by *Z. japonica* over the entire range of TCO_2 tested (Figure 3). The TCO_2 -irradiance

model predicted a significant and positive β term for *Z. japonica*, whereas β for *Z. marina* was non-significant (where β is the slope of the P_{max} to TCO_2 relationship). Due to the variance of maximum photosynthetic rates and the resulting non-significant effect of TCO_2 on *Z. marina*, we are unable to properly analyze any trends in *Z. marina* P_{max} with TCO_2 , HCO_3^- , and CO_2 . According to the model, *Z. japonica* P_{max} increased by $\sim 50\%$ from the lowest to highest initial TCO_2 treatment (Figure 3). Across all initial TCO_2 treatments, *Z. japonica* had a significantly higher P_{max} than *Z. marina* (i.e., non-overlapping 95% CI) (Table 3).

Model Predictions for Diurnal Variation of the Carbonate System

Under realistic natural light conditions, both species are at saturating irradiances for about 6 h during a typical summer day. The resulting maximum predicted photosynthetic rates for *Z. japonica*, are $535 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ when initial TCO_2 is set to $2051 \mu\text{mol kg}^{-1}$ (highest TCO_2 treatment), $387 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ when TCO_2 is set at $1964 \mu\text{mol kg}^{-1}$ (ambient treatment) and only $145 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ when TCO_2 is set at $1770 \mu\text{mol kg}^{-1}$ (lowest TCO_2 treatment) (Figure 4). The predicted maximum photosynthetic rate for *Z. marina* is $116 \mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ under these saturating light conditions, and no sensitivity to TCO_2 was detected. Due to the poor respiration data, estimates of respiration were excluded during times of zero irradiance. When light was saturating, model estimates predict that the carbon drawdown based on a P_{max} of $387 \text{ TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$ for *Z. japonica* would result in a maximum increased change in pH and omega of 0.65 and 3.8 units $\text{mg chl}^{-1} \text{ h}^{-1}$ compared to the 0.25 and 1.1 unit increase by *Z. marina*. The estimated changes are, however, based on the total carbon draw down by the leaf segment in a 20 ml vial and in isolation from *in situ* bed conditions, which would be driven by additional considerations.

Model (Equation 5) estimated hourly changes in pH, Ω_{ar} , and the substrate-to-inhibitor ratio ($[\text{HCO}_3^-]/[\text{H}^+]$) for biomass to water volume ratios similar to a typical seagrass bed under 1 m of water were greater when mediated by *Z. japonica* than for *Z. marina*, with identical initial TCO_2 (Figure 5). Under these conditions, *Z. japonica* induced an hourly rate of change for pH and Ω_{ar} that was approximately 20% greater than *Z. marina* (Figure 5). Both seagrass species induced positive changes in carbonate chemistry parameters. The greatest change in pH (0.49 h^{-1}) and Ω_{ar} (1.61 h^{-1}) occurred at a shallow depth of 0.2 m with high TCO_2 ($2051 \mu\text{mol kg}^{-1}$) and was mediated by *Z. japonica* (Figure 5). The greatest change in the substrate-to-inhibitor ratio was not determined by high carbon uptake rates via photosynthesis, but rather by the low concentration of H^+ at low TCO_2 .

DISCUSSION

Species-specific difference in TCO_2 -dependent photosynthetic carbon uptake is a critical component that determines the capacity of seagrass to remove dissolved CO_2 from seawater and

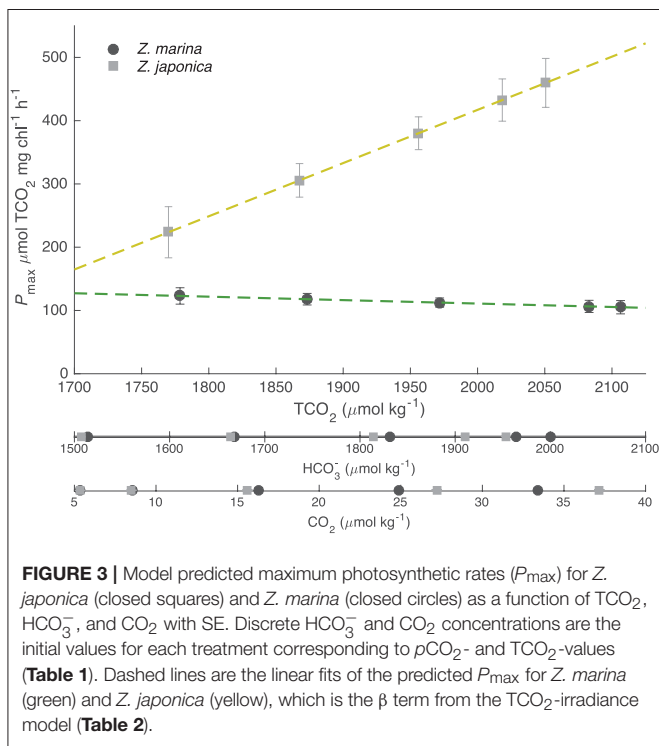
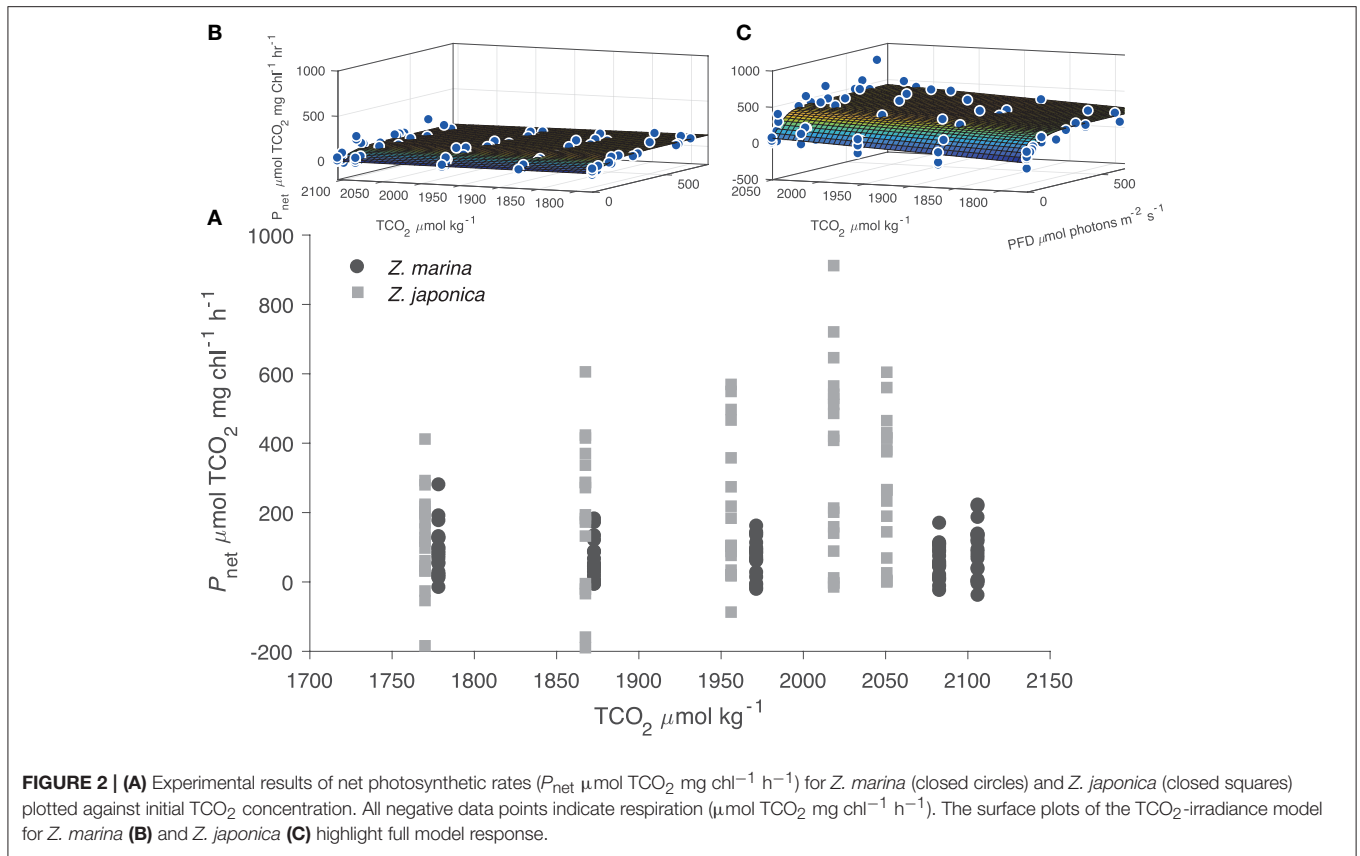
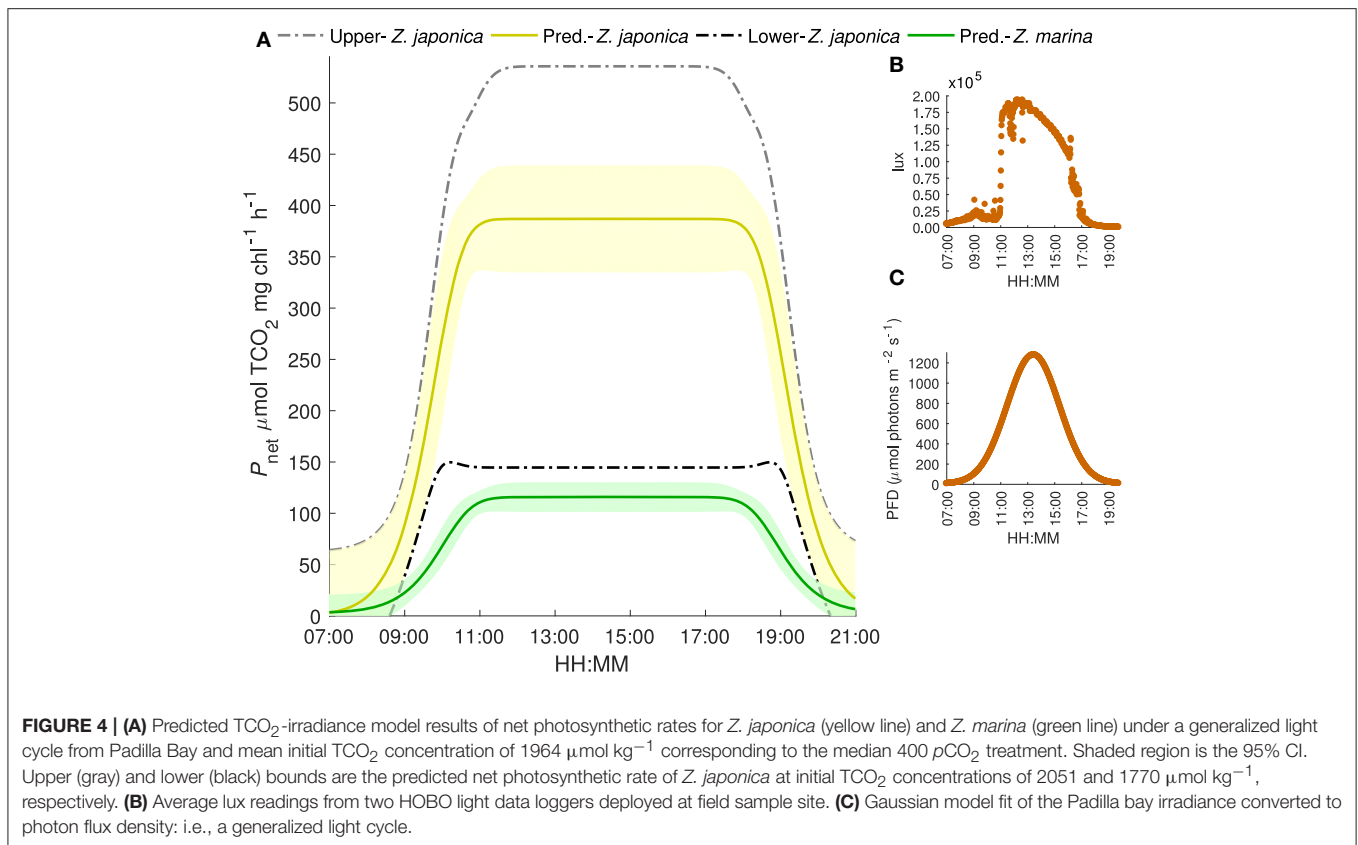


TABLE 3 | TCO_2 -irradiance model (Equation 3) predicted P_{max} ($\mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$) values for both *Z. japonica* and *Z. marina* at each initial TCO_2 concentration.

| Species | Predicted P_{max} | SE | Upper bound (95% CI) | Lower bound (95% CI) |
|--------------------|---------------------|------|----------------------|----------------------|
| <i>Z. marina</i> | 123.0 | 13.0 | 151.3 | 100.5 |
| | 117.8 | 9.1 | 138.6 | 102.9 |
| | 112.5 | 7.3 | 129.7 | 101.0 |
| | 106.4 | 9.6 | 128.2 | 90.5 |
| | 105.2 | 10.5 | 128.7 | 87.4 |
| <i>Z. japonica</i> | 223.6 | 40.4 | 303.0 | 144.6 |
| | 305.7 | 26.5 | 357.8 | 253.8 |
| | 380.1 | 25.9 | 431.1 | 329.5 |
| | 432.6 | 33.4 | 498.2 | 367.3 |
| | 459.7 | 38.7 | 535.7 | 384.1 |

potentially create OA refugia on short timescales. Our results indicate that, on a per chlorophyll basis, *Z. japonica* is able to take up more TCO_2 and more efficiently utilize available irradiance compared to *Z. marina* occurring in the same intertidal zone under similar TCO_2 and irradiance conditions (Figures 1, 2, Table 2). Because of this, *Z. japonica* may be more effective



at mitigating OA (increasing pH and Ω_{ar}) for more hours of the day than *Z. marina* on a per chlorophyll basis (Figure 4). The maximum potential drawdown of TCO_2 over the 6 h of saturating irradiance induces a positive change in pH (from 8.03 to 8.33 and 8.40) and Ω_{ar} (from 1.94 to 3.1 and 3.4) by *Z. marina* and *Z. japonica*, respectively, in our model. Even though these estimates of carbonate chemistry modification by *Z. marina* and *Z. japonica* are specific to our experimental conditions, we are able to detect the differential response of TCO_2 -dependent carbon uptake between the two species, and provide insight for potential OA mitigation by two PNW seagrasses.

Extrapolated Changes in Carbonate Chemistry

One focus of this study was to better understand the OA mitigation potential of seagrass clippings in isolation, and extend those results to various biomass to water volume ratios in an attempt to conceptualize how these attributes of seagrass beds drive carbonate chemistry. While these results are not representative of *in situ* conditions, by incorporating realistic biomass values, the box model results provide a point of comparison to field based studies and information about photosynthetic effects measured under controlled conditions to those quantifying potential seagrass contribution to a dynamic and complex system. The estimates presented here are considered the maximum potential possible by each species, as photosynthetic rates calculated from leaf segments are likely

an overestimate of whole plant photosynthetic rates (Herzka and Dunton, 1997); however, previous studies have shown that empirically derived photosynthetic rates can be similar between leaf segments and whole plants (see Table 2 in Lee et al., 2007; Table 6 in Shafer and Kaldy, 2014). It is important to note that the predicted changes in carbonate chemistry are derived from our measured mg chl:gDW^{-1} ratio (Equation 5), however, we do not use gDW normalized photosynthetic rates as the variance was extremely high within treatments and resulted in non-significant differences (see Miller, 2016).

When extending our laboratory results of seagrass photosynthetic carbon uptake to changes in carbonate chemistry given specific biomass to water volume ratios, the expected potential change in pH, Ω_{ar} , and the substrate-to-inhibitor ratio— $[\text{HCO}_3^-]/[\text{H}^+]$ —would be positively affected by seagrass carbon uptake (Figure 5). Estimated increases in pH per hour from this model at a volume corresponding to 1 m depth over 6 h result in maximum pH changes from 0.09 to 0.73 under low and high TCO_2 conditions for *Z. japonica* (Figure 5). Since *Z. japonica* photosynthesis is carbon sensitive, the more realistic pH change is likely somewhere in the middle and more accurately predicted by the mean TCO_2 treatment, which would result in a pH change of 0.37 over 6 h. This would not be a concern for *Z. marina*, however, as we did not find sensitivity to TCO_2 over the range tested in this study. These rates are based on maximum photosynthetic rates, but also on conservative assumptions about short residence time in seagrass

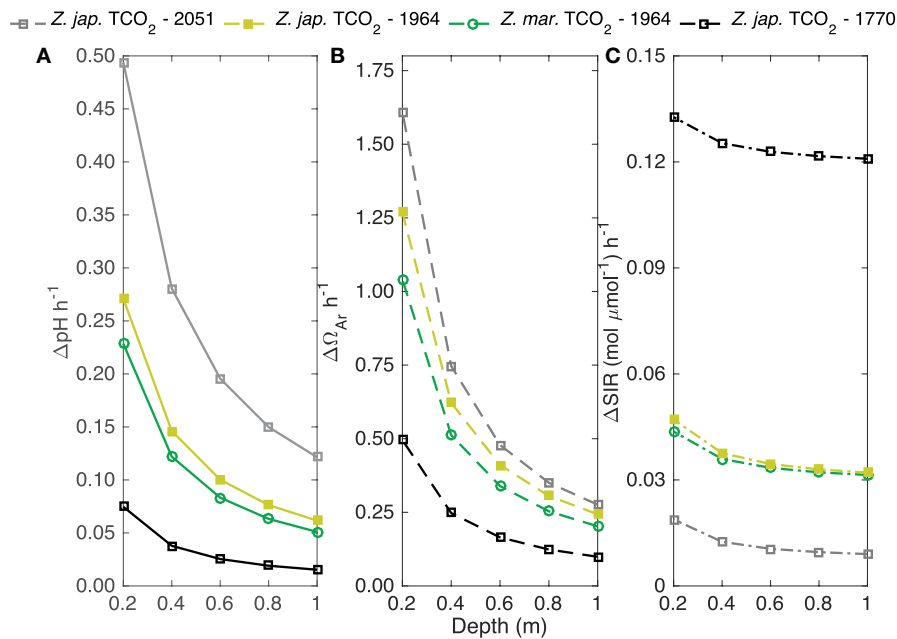


FIGURE 5 | Yellow and green lines are the predicted change in (A) pH, (B) aragonite saturation state (Ω_{ar}), and (C) substrate-to-inhibitor ($\text{mol } \mu\text{mol}^{-1}$) ratio based on the predicted P_{max} ($\mu\text{mol TCO}_2 \text{ mg chl}^{-1} \text{ h}^{-1}$) of *Z. japonica* (closed squares) and *Z. marina* (open circles) at an initial $[\text{TCO}_2]$ of $1964 \mu\text{mol kg}^{-1}$. Gray and black lines are the upper and lower bound of the predicted change in pH, aragonite saturation state, and substrate-to-inhibitor ratio of *Z. japonica* at an initial $[\text{TCO}_2]$ of 2,051 and 1,770 $\mu\text{mol kg}^{-1}$, respectively. Predicted change is over a 1 m^2 area and given depth assuming a biomass of 100 gDW m^{-2} (Equation 5). TCO_2 concentrations correspond to the 140, 400, and 900 pCO_2 treatments for *Z. japonica* and 400 pCO_2 treatment for *Z. marina* (TCO_2 corresponding to 400 pCO_2 treatments was averaged from both experiments). TA (i.e., average TA of both 400 pCO_2 treatments), temperature, and salinity were utilized from each respective treatment (Table 1) to constrain the carbonate system to calculate predictions.

beds. Where residence time is longer, chemical changes can be compounded and result in large swings, or the modified water can be exported to other locales. Observed diurnal change of pH in whole seagrass systems ranged from 0.2 to 0.7 units over several hours measured within, or directly above, shallow seagrass beds at a depth no greater than 12 m (Frankignoulle and Bouqueneau, 1990; Invers et al., 1997; Unsworth et al., 2012; Hendriks et al., 2014). We estimated diurnal increases in Ω_{ar} from 0.59 to 1.6 at a 1 m depth over 6 h under low and high TCO_2 conditions for *Z. japonica* (Figure 5). Model predictions from an extensive data set of Indo-Pacific seagrass community metabolism estimated maximum diurnal changes in Ω_{ar} to be 2.9 (Unsworth et al., 2012), whereas field based calculations of diurnal changes ranged from 1.38 to 1.67 (Hendriks et al., 2014). At shallower water depths, the higher Indo-Pacific change in Ω_{ar} of 2.9 is easily reached under our box model conditions (Figure 5). The congruence of these estimates illustrates how the laboratory studies presented here corroborate the idea that seagrass meadows can be drivers of locally important changes in carbonate chemistry. We note that the results of our box model should be viewed as context for field results and not as estimates of *in situ* processes. In order to make accurate predictions of carbonate chemistry variability in seagrass beds, higher resolution *in situ* studies and whole plant laboratory studies need to be performed.

Photosynthetic Response to TCO_2

Based on our results from the TCO_2 -irradiance model (Equation 3), *Z. japonica* increased its P_{max} proportionally across TCO_2 treatments—where the slope (β) of the linear TCO_2 response was highly significant ($p = <0.001$). Conversely, we were unable to detect any positive response of *Z. marina* P_{max} with increasing TCO_2 . That is according to the TCO_2 -irradiance model, the slope of the linear TCO_2 response was non-significant and slightly negative ($p = 0.327$, $\beta = -0.054$). Due to the lack of any significant trend and high variance within treatments, we suggest that *Z. marina* P_{max} was not positively affected by increasing TCO_2 under our study conditions. Our findings here are somewhat contradictory as most studies have reported *Z. marina* to have enhanced photosynthetic rates with increasing TCO_2 (Beer and Koch, 1996; Thom, 1996; Zimmerman et al., 1997), however, a recent study also found there to be no positive effect of TCO_2 on *Z. marina* photosynthesis up to pCO_2 levels of 2000 μatm (Pajusalu et al., 2016).

One reason for the different result in this study regarding *Z. marina* response to increasing TCO_2 may be due to the range of concentrations tested. The lowest enriched TCO_2 treatment in previous studies was $\sim 25\%$ higher than our highest TCO_2 treatment (Thom, 1996 does not give exact numbers, but minimum CO_2 increase was 25% higher than ambient), ranging up to 230% times greater than our highest treatment (Beer and

Koch, 1996). The first reason for the relatively low TCO_2 tested in this study is that ambient seawater from the Salish Sea has lower salinity and, therefore, a lower TCO_2 than open ocean values. In addition, the range of TCO_2 tested in this study was representative of conditions currently observed in Padilla Bay, not the much greater range in other studies, which were designed to highlight potential physiological responses, not necessarily *in situ* conditions. The result showing that *Z. marina* P_{max} does not respond to TCO_2 in this study may indicate that northern Salish Sea populations are not sensitive to TCO_2 at these lower concentrations.

Additionally, the high degree of variability within our treatments may have obscured a positive response. Zimmerman et al. (1997) showed that *Z. marina* increased its P_{max} 225% with an exposure to TCO_2 77% higher than their ambient treatment (from 2,074 to 3,673 $\mu\text{mol kg}^{-1}$). Over this range of TCO_2 , a linear relationship has been observed between P_{max} and TCO_2 (Beer and Koch, 1996). The linear equation derived from Zimmerman et al. (1997) can be extrapolated down to the range in this study (1,770–2,051 $\mu\text{mol kg}^{-1}$), which would therefore predict that the 18% increase in TCO_2 between our lowest and highest TCO_2 treatments would induce a 55% increase in *Z. marina* P_{max} . Given that the coefficient of variation for *Z. marina* P_{net} at, or above, E_k ranged from 30 to 60% over all initial TCO_2 treatments, any small positive responses to TCO_2 by *Z. marina* may likely be concealed by the variance—assuming an extrapolation of the Zimmerman et al. (1997) data is accurate given this is outside of the bounds tested. Physiological and methodological factors contributing to the large variance within treatments may be a result of, insufficient incubation period, variance in illumination, shoot to shoot variability of the photosynthetic apparatus, inefficiencies in chlorophyll extraction from leaf segments, or propagation of a diffusive boundary layer due to periodic rather than continuous stirring over the 90+ min incubation time, which would increase the variability of the relative TCO_2 conditions experienced by the leaf segment compared to the entire vial.

Interestingly, the increase in *Z. japonica* P_{max} was $\sim 50\%$ over an 18% increase in TCO_2 , which coincidentally is very close to the predicted increase from the linear equation derived from Zimmerman et al. (1997). This robust response by *Z. japonica* was observed despite a coefficient of variation in P_{net} at, or above, E_k identical to that of *Z. marina* (30–60%). This is the first time that a response in the photosynthetic rate of *Z. japonica* to TCO_2 has been documented, to our knowledge. Previous research has shown that *Z. japonica* photosynthesis is more sensitive to light and salinity (Shafer et al., 2011; Shafer and Kaldy, 2014) than is *Z. marina* photosynthesis (Hellblom and Björk, 1999; Shafer and Kaldy, 2014). Given that the two species appear to express different physiological response to environmental factors, it is likely that the utilization of TCO_2 would also be inherently different. The differential response of P_{max} to TCO_2 between species may be due to the mechanism by which TCO_2 is utilized for photosynthesis. Previous research suggests that *Z. marina* HCO_3^- utilization is saturated at pH ranging from 7.5 to 8.5 (Invers et al., 2001), which was the range of pH in our study (Table 1). Although we did not specifically measure independent HCO_3^- and CO_2 -uptake, an increase in HCO_3^- in addition to

CO_2 may be one mechanism that explains such a robust response by *Z. japonica* to TCO_2 ; however, further studies are needed to examine *Z. japonica*'s photosynthetic response to HCO_3^- and CO_2 independently.

Species-Specific Photosynthetic Rates

The relative difference between *Z. marina* and *Z. japonica* photosynthetic rates reported in our study are similar to previous findings, but differ in magnitude (Figure 1). The large degree of variability around the mean estimated model fit for *Z. japonica* photosynthesis can be partially explained by the increasing response of photosynthesis to TCO_2 (Figure 2). We reiterate, however, that the variability in our respiration measurements may somewhat lessen the predictive power of model estimates, particularly the photosynthetic efficiency, as the initial slope of the photosynthetic-irradiance curve is dependent on robust respiration values. For this reason, we have focused specifically on only our P_{max} results when interpreting photosynthetic carbon uptake induced changes on carbonate chemistry. A likely reason for the mix of positive and negative TCO_2 fluxes by our leaf segments in dark vials may be due to a short incubation time, volume of media to leaf segment ratio, or internal seagrass biorhythms of photosynthetic and respiration cycles. Similar to our findings, Shafer and Kaldy (2014) found that central Oregon populations of *Z. japonica* have a P_{max} that is $\sim 3\text{x}$ greater than *Z. marina* when exposed to the same light conditions; however, the chlorophyll standardized rates reported in that study were considerably lower than the rates found in this study. The differing magnitude of rates suggest that other factors such as organization of the photosynthetic package within seagrass tissue may have a significant impact on photosynthetic rates rather than chlorophyll pigment concentration alone. A multitude of environmental factors such as intertidal location, canopy density, age of shoot and leaf, time of year, and acclimatization to epiphytic growth can affect chlorophyll pigment concentration and the photoacclimation amongst shoots and along leaves, thus leading to drastically different mg chl:gdW^{-1} ratios (Dennison and Alberte, 1986; Durako and Kunzelman, 2002; Enriquez et al., 2002; Major and Dunton, 2002; Cummings and Zimmerman, 2003; Drake et al., 2003; Larkum et al., 2006 and references therein). Any of these factors may be responsible for the lower mg chl:gdW^{-1} ratio and dissimilar photosynthetic rates in our study. In addition, the *Z. marina* samples were collected at the upper limit of their distribution, potentially affecting the physiology and phenotypic expression of *Z. marina* at this location. Local acclimatization to *in situ* temperature and salinity can also impact photosynthetic rates via osmotic stress and changes in the photosynthesis-respiration ratio (Kenneth and Short, 2006 and references therein). In addition, our study examined photosynthetic rates by measuring the change in TCO_2 , which to our knowledge, has not been done before in the lab and rarely done in the field despite the robustness of the method (Silva et al., 2009). The photosynthetic quotient (O_2/CO_2), while assumed to be unity or close to in many studies, is based on community metabolism in seagrass beds or meadows (Oviatt et al., 1986; Leuschner and Rees, 1993; Mateo et al., 2001; Martin et al., 2005). Measuring TCO_2 rather than O_2 may result in photosynthetic

rates that are dissimilar if the photosynthetic quotient is not close to unity under these conditions. That is, given the higher photosynthetic rates found in this study, the change in CO₂ per O₂ would have to be much lower in seagrass communities than what was measured in our vials. Alternatively, differences in chlorophyll extraction methods, or inefficient extraction may have resulted in the higher photosynthetic rates found in our study. Additionally, recycling and movement of gases within the lacunal system are not well-understood (Mateo et al., 2001), further convoluting the comparison of instantaneous O₂ production and CO₂ uptake. It may be that deriving photosynthesis by measuring TCO₂ is not as robust as the O₂ normative, but due to the lack of studies measuring seagrass photosynthesis with TCO₂, this remains uncertain. Studies which capture O₂ production and CO₂ uptake would be best suited to determine the relative advantage of each in determining the interactions between carbonate chemistry and seagrass photosynthesis.

Variations of Carbonate Chemistry in Seagrass Beds

While this study provides a point of reference for how seagrass beds can modify carbonate chemistry, determining the effects of macrophytes on acidification is challenging due to the extreme spatial and temporal variability of carbonate chemistry in these zones (Hendriks et al., 2014, 2015; Krause-Jensen et al., 2015; Challenger et al., 2016). A multitude of factors such as seagrass epiphyte communities, heterotrophic respiration, tidal exchange, groundwater flux, and riverine input all contribute in modifying carbonate chemistry on various spatial and temporal scales. For example, differences in seagrass shoot density over a given area result in different mixing rates and flow regimes, which will further shift the carbonate chemistry of water parcels that have been modified by photosynthetic carbon uptake (Peterson et al., 2004; Koch et al., 2006; Marbá et al., 2006; Hendriks et al., 2014). In addition, variations in chlorophyll content along the leaf and amongst shoots results in differential photosynthetic rates and carbon uptake within the seagrass canopy, thereby creating disparate water parcels with respect to TCO₂ concentration on various spatial scales. These various interactions between water flow within a seagrass bed, irradiance, and TCO₂ are some of the determining factors of carbon assimilation and uptake (McPherson et al., 2015) and, therefore, directly affect the scale of OA mitigation by seagrasses via the formation of microzones with differential carbonate chemistry. Because organisms vulnerable to OA are more sensitive at particular life-stages and over hourly durations, the scale of carbonate chemistry variability (i.e., magnitude and duration of low saturation state and pH) is critical to understand when determining organismal resilience to OA (Kurihara, 2008; Barton et al., 2012; Hettinger et al., 2012; Onitsuka et al., 2014; Waldbusser et al., 2015; Miller and Waldbusser, 2016). Therefore, in order to elucidate a more nuanced amelioration of acidification by seagrasses, it will be necessary to conduct *in situ* studies that can effectively capture all the drivers of the carbonate system in spatiotemporal context.

CONCLUSION

A comparison of the photosynthetic potential between *Z. marina* and *Z. japonica* has implications for elucidating the contribution each species has on the carbonate system. *Z. japonica* exhibits a strong increase in photosynthetic rate in response to increasing TCO₂, while a similar response from *Z. marina* could not be identified. In the intertidal zone where species overlap occurs, our results indicate that *Z. japonica* also has a 3-fold greater photosynthetic potential than *Z. marina* when normalized to chlorophyll. By measuring photosynthetic potential as a change in TCO₂ and utilizing measured mg chl:gdW⁻¹ in addition to estimated m² biomass of each species, our results were extended to conceptualize how these additional attributes of biomass to depth ratios in seagrass beds affect hourly changes in carbonate chemistry. Based on our findings, *Z. japonica* appears to be better suited to mitigate OA on a per chlorophyll basis given its higher photosynthetic rate and efficiency and enhanced response to increases in TCO₂. Our study illuminates the potential of PNW populations of *Z. marina* and *Z. japonica* to modify the carbonate system, and provides a direct comparison of photosynthetic potential when exposed to varying levels of TCO₂. This is an initial step in attempting to determine the OA mitigation potential of seagrass systems in the PNW, where quantitative estimates can aid management practices and provide a better understanding for the local initiatives that aim to protect Washington state's aquaculture resources (Blue ribbon panel on OA 2012).

AUTHOR CONTRIBUTIONS

CM conceptualized and designed the experiment with modifications directed by BL and SY. CM led the writing of the paper with contributions by BL and SY. CM performed all data analysis, with statistical analysis supported by SY. BL provided substantial support for utilization of model projections. All authors reviewed and contributed to the writing of the final manuscript.

FUNDING

This work was supported by the Padilla Bay National Estuarine Research Reserve assistantship. Funding granted by the Padilla Bay foundation, Jude Apple and Sharon Riggs: Contract/Grant # Borman-01-2015.

ACKNOWLEDGMENTS

The authors would like to thank David Shull for thoughtful interpretation of our results. CM would like to thank Katherina L. Schoo for assistance with the ocean acidification experimental system. CM would also like to thank Rachel Blyth, Melissa Ciesielski, Taylor Clement, Rosie Gradoville, and Anne Harmann who were integral for carrying out the experiments.

REFERENCES

- Backman, T. (1991). Genotypic and phenotypic variability of *Zostera marina* on the west coast. *Can. J. Bot. Rev. Can. Bot.* 69, 1361–1371. doi: 10.1139/b91-176
- Barton, A., Hales, B., Waldbusser, G. G., Langdon, C., and Feely, R. A. (2012). The Pacific oyster, *Crassostrea gigas*, shows negative correlation to naturally elevated carbon dioxide levels: implications for near-term ocean acidification effects. *Limnol. Oceanogr.* 57, 698–710. doi: 10.4319/lo.2012.57.3.0698
- Baumann, H., Wallace, R. B., Tagliaferri, T., and Gobler, C. J. (2015). Large natural pH, CO₂ and O₂ fluctuations in a temperate tidal salt marsh on diel, seasonal, and interannual time scales. *Estuar. Coasts* 38, 220–231. doi: 10.1007/s12237-014-9800-y
- Beer, S., and Koch, E. (1996). Photosynthesis of marine macroalgae and seagrasses in globally changing CO₂ environments. *Mar. Ecol. Prog. Ser.* 141, 199–204. doi: 10.3354/meps141199
- Bulthuis, D. A. (2013). *The Ecology of Padilla Bay, Washington: an Estuarine Profile of a National Estuarine Research Reserve*. Mount Vernon, WA: Washington State Department of Ecology, Padilla Bay National Estuarine Research Reserve. Available online at: www.padillabay.gov/researchrep/profile.asp
- Cai, W. J. (2011). “Estuarine and coastal ocean carbon paradox: CO₂ sinks or sites of terrestrial carbon incineration?,” in *Annual Review of Marine Science*, Vol. 3, eds C. A. Carlson and S. J. Giovannoni (Palo Alto, CA: Annual Reviews), 123–145.
- Campbell, J. E., and Fourqurean, J. W. (2013). Mechanisms of bicarbonate use influence the photosynthetic carbon dioxide sensitivity of tropical seagrasses. *Limnol. Oceanogr.* 58, 839–848. doi: 10.4319/lo.2013.58.3.0839
- Challener, R. C., Robbins, L. L., and McClintock, J. B. (2016). Variability of the carbonate chemistry in a shallow, seagrass-dominated ecosystem: implications for ocean acidification experiments. *Mar. Freshw. Res.* 67, 163–172. doi: 10.1071/MF14219
- Chung, I. K., Beardall, J., Mehta, S., Sahoo, D., and Stojkovic, S. (2011). Using marine macroalgae for carbon sequestration: a critical appraisal. *J. Appl. Phycol.* 23, 877–886. doi: 10.1007/s10811-010-9604-9
- Ciais, P., C., Sabine, G., Bala, L., Bopp, V., Brovkin, J., Canadell, A., et al. (2013). “Carbon and other biogeochemical cycles,” in *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, eds Stocker, T. F. D. Qin, G. K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex, and P. M. Midgley (Cambridge; New York, NY: Cambridge University Press), 465–570.
- Cox, T. E., Gazeau, F., Alliouane, S., Hendriks, I. E., Mahacek, P., Le Fur, A., et al. (2016). Effects of *in situ* CO₂ enrichment on structural characteristics, photosynthesis, and growth of the Mediterranean seagrass *Posidonia oceanica*. *Biogeosciences* 13, 2179–2194. doi: 10.5194/bg-13-2179-2016
- Cummings, M. E., and Zimmerman, R. C. (2003). Light harvesting and the package effect in the seagrasses *Thalassia testudinum* Banks ex König and *Zostera marina* L.: optical constraints on photoacclimation. *Aquat. Bot.* 75, 261–274. doi: 10.1016/S0304-3770(02)00180-8
- Dennison, W., and Alberte, R. (1986). Photoadaptation and growth of *Zostera marina* L. (eelgrass) transplants along a depth gradient. *J. Exp. Mar. Biol. Ecol.* 98, 265–282. doi: 10.1016/0022-0981(86)90217-0
- Dickson, A., and Millero, F. (1987). A comparison of the equilibrium constants for the dissociation of carbonic acid in seawater Media. *Deep Sea Res. Oceanogr. Res. Pap.* 34, 1733–1743. doi: 10.1016/0198-0149(87)90021-5
- Dickson, A. G., Sabine, C. L., and Christian, J. R. (2007). *Guide to Best Practices for Ocean CO₂ Measurements*, Vol. 3. Sidney, BC: PICES Special Publication.
- Doney, S. C., Fabry, V. J., Feely, R. A., and Kleypas, J. A. (2009). Ocean acidification: the other CO₂ problem. *Annu. Rev. Mar. Sci.* 1, 169–192. doi: 10.1146/annurev.marine.010908.163834
- Drake, L. A., Dobbs, F. C., and Zimmerman, R. C. (2003). Effects of epiphyte load on optical properties and photosynthetic potential of the seagrasses *Thalassia testudinum* Banks ex König and *Zostera marina* L. *Limnol. Oceanogr.* 48, 456–463. doi: 10.4319/lo.2003.48.1_part_2.0456
- Duarte, C. M., Hendriks, I. E., Moore, T. S., Olsen, Y. S., Steckbauer, A., Ramajo, L., et al. (2013). Is ocean acidification an open-ocean syndrome? Understanding anthropogenic impacts on seawater pH. *Estuar. Coasts* 36, 221–236. doi: 10.1007/s12237-013-9594-3
- Duarte, C. M., Marbà, N., Gacia, E., Fourqurean, J. W., Beggins, J., Barron, C., et al. (2010). Seagrass community metabolism: assessing the carbon sink capacity of seagrass meadows. *Glob. Biogeochem. Cycles* 24:GB4032. doi: 10.1029/2010gb003793
- Duarte, C. M., Middelburg, J. J., and Caraco, N. (2005). Major role of marine vegetation on the oceanic carbon cycle. *Biogeosciences* 2, 1–8. doi: 10.5194/bg-2-1-2005
- Dumbauld, B. R., and McCoy, L. M. (2015). Effect of oyster aquaculture on seagrass *Zostera marina* at the estuarine landscape scale in Willapa Bay, Washington (USA). *Aquac. Environ. Interact.* 7, 29–47. doi: 10.3354/aei00131
- Durako, M. J., and Kunzelman, J. I. (2002). Photosynthetic characteristics of *Thalassia testudinum* measured *in situ* by pulse-amplitude modulated (PAM) fluorometry: methodological and scale-based considerations. *Aquat. Bot.* 73, 173–185. doi: 10.1016/S0304-3770(02)00020-7
- Ekstrom, J. A., Suatoni, L., Cooley, S. R., Pendleton, L. H., Waldbusser, G. G., Cinner, J. E., et al. (2015). Vulnerability and adaptation of US shellfisheries to ocean acidification. *Nat. Clim. Change* 5, 207–214. doi: 10.1038/nclimate2508
- Enríquez, S., Merino, M., and Iglesias-Prieto, R. (2002). Variations in the photosynthetic performance along the leaves of the tropical seagrass *Thalassia testudinum*. *Mar. Biol.* 140, 891–900. doi: 10.1007/s00227-001-0760-y
- Fassbender, A. J., Sabine, C. L., and Feifel, K. M. (2016). Consideration of coastal carbonate chemistry in understanding biological calcification. *Geophys. Res. Lett.* 43:2016GL068860. doi: 10.1002/2016GL068860
- Feely, R. A., Alin, S. R., Newton, J., Sabine, C. L., Warner, M., Devol, A., et al. (2010). The combined effects of ocean acidification, mixing, and respiration on pH and carbonate saturation in an urbanized estuary. *Estuar. Coast. Shelf Sci.* 88, 442–449. doi: 10.1016/j.ecss.2010.05.004
- Feely, R. A., Sabine, C. L., Hernandez-Ayon, J. M., Ianson, D., and Hales, B. (2008). Evidence for upwelling of corrosive acidified water onto the continental shelf. *Science* 320, 1490–1492. doi: 10.1126/science.1155676
- Ferraro, S. P., and Cole, F. A. (2012). Ecological periodic tables for benthic macrofaunal usage of estuarine habitats: Insights from a case study in Tillamook Bay, Oregon, U.S.A. *Estuar. Coast. Shelf Sci.* 102, 70–83. doi: 10.1016/j.ecss.2012.03.009
- Fourqurean, J. W., Duarte, C. M., Kennedy, H., Marbà, N., Holmer, M., Angel Mateo, M., et al. (2012). Seagrass ecosystems as a globally significant carbon stock. *Nat. Geosci.* 5, 505–509. doi: 10.1038/ngeo1477
- Frankignoulle, M., and Bouquegneau, J. (1990). Daily and yearly variations of total inorganic carbon in a productive coastal Area. *Estuar. Coast. Shelf Sci.* 30, 79–89. doi: 10.1016/0272-7714(90)90078-6
- Gattuso, J. P., Frankignoulle, M., and Wollast, R. (1998). Carbon and carbonate metabolism in coastal aquatic ecosystems. *Annu. Rev. Ecol. Syst.* 29, 405–434. doi: 10.1146/annurev.ecolsys.29.1.405
- Grantham, B. A., Chan, F., Nielsen, K. J., Fox, D. S., Barth, J. A., Huyer, A., et al. (2004). Upwelling-driven nearshore hypoxia signals ecosystem and oceanographic changes in the northeast Pacific. *Nature* 429, 749–754. doi: 10.1038/nature02605
- Harris, K. E., DeGrandpre, M. D., and Hales, B. (2013). Aragonite saturation state dynamics in a coastal upwelling zone. *Geophys. Res. Lett.* 40, 2720–2725. doi: 10.1002/grl.50460
- Harrison, P., and Bigley, R. (1982). The recent introduction of the seagrass *Zostera japonica* Aschers. and Graebn. to the Pacific coast of North America. *Can. J. Fish. Aquat. Sci.* 39, 1642–1648. doi: 10.1139/f82-221
- Harrison, P. G. (1982). Spatial and temporal patterns in abundance of two intertidal seagrasses, *Zostera americana* den hartog and *Zostera marina* L. *Aquat. Bot.* 12, 305–320. doi: 10.1016/0304-3770(82)90024-9
- Hauri, C., Gruber, N., McDonnell, A. M. P., and Vogt, M. (2013). The intensity, duration, and severity of low aragonite saturation state events on the California continental shelf. *Geophys. Res. Lett.* 40, 3424–3428. doi: 10.1002/grl.50618
- Hellblom, F., and Björk, M. (1999). Photosynthetic responses in *Z. osteria marina* to decreasing salinity, inorganic carbon content and osmolality. *Aquat. Bot.* 65, 97–104. doi: 10.1016/S0304-3770(99)00034-0
- Hendriks, I. E., Duarte, C. M., Olsen, Y. S., Steckbauer, A., Ramajo, L., Moore, T. S., et al. (2015). Biological mechanisms supporting adaptation to ocean acidification in coastal ecosystems. *Estuar. Coast. Shelf Sci.* 152, A1–A8. doi: 10.1016/j.ecss.2014.07.019

- Hendriks, I. E., Olsen, Y. S., Ramajo, L., Basso, L., Steckbauer, A., Moore, T. S., et al. (2014). Photosynthetic activity buffers ocean acidification in seagrass meadows. *Biogeosciences* 11, 333–346. doi: 10.5194/bg-11-333-2014
- Herzka, S. Z., and Dunton, K. H. (1997). Seasonal photosynthetic patterns of the seagrass *Thalassia testudinum* in the western Gulf of Mexico. *Mar. Ecol. Prog. Ser.* 152, 103–117. doi: 10.3354/meps152103
- Hettinger, A., Sanford, E., Hill, T. M., Russell, A. D., Sato, K. N. S., Hoey, J., et al. (2012). Persistent carry-over effects of planktonic exposure to ocean acidification in the *Olympia* oyster. *Ecology* 93, 2758–2768. doi: 10.1890/12-0567.1
- Hönisch, B., Ridgwell, A., Schmidt, D. N., Thomas, E., Gibbs, S. J., Sluijs, A., et al. (2012). The Geological record of ocean acidification. *Science* 335, 1058–1063. doi: 10.1126/science.1208277
- Invers, O., Romero, J., and Pérez, M. (1997). Effects of pH on seagrass photosynthesis: a laboratory and field assessment. *Aquat. Bot.* 59, 185–194.
- Invers, O., Zimmerman, R. C., Alberte, R. S., Perez, M., and Romero, J. (2001). Inorganic carbon sources for seagrass photosynthesis: an experimental evaluation of bicarbonate use in species inhabiting temperate waters. *J. Exp. Mar. Biol. Ecol.* 265, 203–217. doi: 10.1016/S0022-0981(01)00332-X
- Jassby, A. D., and Platt, T. (1976). Mathematical formulation of the relationship between photosynthesis and light for phytoplankton. *Limnol. Oceanogr.* 21, 540–547. doi: 10.4319/lo.1976.21.4.0540
- Kaldy, J. E. (2006). Production ecology of the non-indigenous seagrass, dwarf eelgrass (*Zostera japonica* Ascher. & Graeb.), in a Pacific Northwest estuary, USA. *Hydrobiologia* 553, 201–217. doi: 10.1007/s10750-005-5764-z
- Kaldy, J. E., Shafer, D. J., and Magoun, A. D. (2015). Duration of temperature exposure controls growth of *Zostera japonica*: implications for zonation and colonization. *J. Exp. Mar. Biol. Ecol.* 464, 68–74. doi: 10.1016/j.jembe.2014.12.015
- Kenneth, A., and Short, F. T. (2006). “*Zostera*: biology, ecology, and management,” in *Seagrasses: Biology, Ecology and Conservation*, eds A. W. D. Larkum, R. J. Orth, and C. M. Duarte (Dordrecht: Springer), 361–386.
- Koch, E. W., Ackerman, J., Verduin, J., and van Keulen, M. (2006). “Fluid dynamics in seagrass ecology—from molecules to ecosystems,” in *Seagrasses: Biology, Ecology and Conservation*, eds A. W. D. Larkum, R. J. Orth, and C. M. Duarte (Dordrecht: Springer), 193–225.
- Koch, M., Bowes, G., Ross, C., and Zhang, X.-H. (2013). Climate change and ocean acidification effects on seagrasses and marine macroalgae. *Glob. Change Biol.* 19, 103–132. doi: 10.1111/j.1365-2486.2012.02791.x
- Krause-Jensen, D., Duarte, C. M., Hendriks, I. E., Meire, L., Blicher, M. E., Marbà, N., et al. (2015). Macroalgae contribute to nested mosaics of pH variability in a sub-Arctic fjord. *Biogeosciences Discuss* 12, 4907–4945. doi: 10.5194/bg-12-4907-2015
- Kroeker, K. J., Kordas, R. L., Crim, R., Hendriks, I. E., Ramajo, L., Singh, G. S., et al. (2013). Impacts of ocean acidification on marine organisms: quantifying sensitivities and interaction with warming. *Glob. Change Biol.* 19, 1884–1896. doi: 10.1111/gcb.12179
- Kroeker, K. J., Kordas, R. L., Crim, R. N., and Singh, G. G. (2010). Meta-analysis reveals negative yet variable effects of ocean acidification on marine organisms. *Ecol. Lett.* 13, 1419–1434. doi: 10.1111/j.1461-0248.2010.01518.x
- Kurihara, H. (2008). Effects of CO₂ driven ocean acidification on the early developmental stages of invertebrates. *Mar. Ecol. Prog. Ser.* 373, 275–284. doi: 10.3354/meps07802
- Larkum, A. W. D., Drew, E. A., and Ralph, P. J. (2006). “Photosynthesis and metabolism,” in *Seagrasses: Biology, Ecology and Conservation*, eds A. W. D. Larkum, R. J. Orth, and C. M. Duarte (Dordrecht: Springer), 323–345.
- Lee, K. S., Park, S. R., and Kim, Y. K. (2007). Effects of irradiance, temperature, and nutrients on growth dynamics of seagrasses: a review. *J. Exp. Mar. Biol. Ecol.* 350, 144–175. doi: 10.1016/j.jembe.2007.06.016
- Leuschner, C., and Rees, U. (1993). CO₂ gas exchange of two intertidal seagrass species, *Zostera marina* L. and *Zostera noltii* Hornem., during emersion. *Aquat. Bot.* 45, 53–62.
- Lorenzen, C. J. (1966). A method for the continuous measurement of *in vivo* chlorophyll concentration. *Deep Sea Res. Oceanogr. Abstr.* 13, 223–227.
- Love, B. A., O’Brien, C., and Bohlmann, H. (2016). “Photosynthetically driven cycles produce extreme pCO₂ variability in a large eelgrass meadow and readily measured proxies can be used to estimate these changes,” in *EC54A–1311* 2016 Ocean Sciences Meeting. AGU/ASLO/TOS (New Orleans, LA).
- Love, B. A., Olson, M. B., and Wuori, T. (2017). Technical Note: a minimally-invasive experimental system for pCO₂ manipulation in plankton cultures using passive gas exchange (Atmospheric Carbon Control Simulator). *Biogeosciences Discuss.* 2016, 1–19. doi: 10.5194/bg-2016-502
- Mach, M. E., Wyllie-Echeverria, S., and Chan, K. M. A. (2014). Ecological effect of a nonnative seagrass spreading in the Northeast Pacific: a review of *Zostera japonica*. *Ocean Coast. Manag.* 102, 375–382. doi: 10.1016/j.ocecoaman.2014.10.002
- Major, K. M., and Dunton, K. H. (2002). Variations in light-harvesting characteristics of the seagrass, *Thalassia testudinum*: evidence for photoacclimation. *J. Exp. Mar. Biol. Ecol.* 275, 173–189. doi: 10.1016/S0022-0981(02)00212-5
- Manzello, D. P., Enochs, I. C., Melo, N., Gledhill, D. K., and Johns, E. M. (2012). Ocean acidification refugia of the Florida reef tract. *PLoS ONE* 7:e41715. doi: 10.1371/journal.pone.0041715
- Marbà, N., Arias-Ortiz, A., Masque, P., Kendrick, G. A., Mazarrasa, I., Bastyan, G. R., et al. (2015). Impact of seagrass loss and subsequent revegetation on carbon sequestration and stocks. *J. Ecol.* 103, 296–302. doi: 10.1371/journal.pone.0041715
- Marbà, N., Holmer, M., Gacia, E., and Barrón, C. (2006). “Seagrass beds and coastal biogeochemistry,” in *Seagrasses: Biology, Ecology and Conservation*, eds A. W. D. Larkum, R. J. Orth, and C. M. Duarte (Dordrecht: Springer), 135–157.
- Martin, S., Clavier, J., Guarini, J. M., Chauvaud, L., Hily, C., Grall, J., et al. (2005). Comparison of *Zostera marina* and maerl community metabolism. *Aquat. Bot.* 83, 161–174. doi: 10.1016/j.aquabot.2005.06.00
- Mateo, M. A., Renom, P., Hemminga, M. A., and Peene, J. (2001). Measurement of seagrass production using the ¹³C stable isotope compared with classical O₂ and ¹⁴C methods. *Mar. Ecol. Prog. Ser.* 223, 157–165. doi: 10.3354/meps223157
- McLeod, E., Chmura, G. L., Bouillon, S., Salm, R., Bjork, M., Duarte, C. M., et al. (2011). A blueprint for blue carbon: toward an improved understanding of the role of vegetated coastal habitats in sequestering CO₂. *Front. Ecol. Environ.* 9, 552–560. doi: 10.1890/110004
- McPherson, M. L., Zimmerman, R. C., and Hill, V. J. (2015). Predicting carbon isotope discrimination in eelgrass (*Zostera marina* L.) from the environmental parameters—light, flow, and [DIC]. *Limnol. Oceanogr.* 60, 1875–1889. doi: 10.1002/lno.10142
- Mehrbach, C., Culberso, C. H., Hawley, J., and Pytkowic, R. M. (1973). Measurement of apparent dissociation constants of carbonic acid in seawater at atmospheric pressure. *Limnol. Oceanogr.* 18, 897–907.
- Miller, C. A. (2016). *Seagrasses (Zostera marina) and (Zostera japonica) Display a Differential Photosynthetic Response to TCO₂: Implications for Acidification Mitigation*. Master’s thesis, Western Washington University.
- Miller, C. A., and Waldbusser, G. G. (2016). A post-larval stage-based model of hard clam *Mercenaria mercenaria* development in response to multiple stressors: temperature and acidification severity. *Mar. Ecol. Prog. Ser.* 558, 35–49. doi: 10.3354/meps11882
- Onitsuka, T., Kimura, R., Ono, T., Takami, H., and Nojiri, Y. (2014). Effects of ocean acidification on the early developmental stages of the horned turban, *Turbo cornutus*. *Mar. Biol.* 161, 1127–1138. doi: 10.1007/s00227-014-2405-y
- Oviatt, C., Rudnick, D., Keller, A., Sampou, P., and Almquist, G. (1986). A comparison of system (O₂ and CO₂) and ¹⁴C measurements of metabolism. *Mar. Ecol. Prog. Ser.* 28, 57–67.
- Ow, Y. X., Uthicke, S., and Collier, C. J. (2016). Light levels affect carbon utilisation in tropical seagrass under ocean acidification. *PLoS ONE* 11:e0150352. doi: 10.1371/journal.pone.0150352
- Pajusalu, L., Martin, G., Pöllumäe, A., and Paalme, T. (2016). The influence of CO₂ enrichment on net photosynthesis of seagrass *Zostera marina* in a brackish water environment. *Front. Mar. Sci.* 3:239. doi: 10.3389/fmars.2016.00239
- Peterson, C. H., Luettich, R. A., Micheli, F., and Skilleter, G. A. (2004). Attenuation of water flow inside seagrass canopies of differing structure. *Mar. Ecol. Prog. Ser.* 268, 81–92. doi: 10.3354/meps268081
- Pierrot, D. E., Lewis, E., and Wallace, D. W. R. (2006). *MS Excel Program Developed for CO₂ System Calculations*. Oak Ridge, TN: Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory (ORNL).
- Poppe, K. (2016). *An Ecogeomorphic Model to Assess the Response of Padilla Bay’s Eelgrass Habitat to Sea Level Rise*. Master’s thesis, Western Washington University.

- Pörtner, H. O. (2008). Ecosystem effects of ocean acidification in times of ocean warming: a physiologist's view. *Mar. Ecol. Prog. Ser.* 373, 203–217. doi: 10.3354/meps07768
- Ruesink, J. L., Hong, J. S., Wischart, L., Hacker, S. D., Dumbauld, B. R., Hession-Lewis, M., et al. (2010). Congener comparison of native (*Zostera marina*) and introduced (*Z. japonica*) eelgrass at multiple scales within a Pacific Northwest estuary. *Biol. Invasions* 12, 1773–1789. doi: 10.1007/s10530-009-9588-z
- Shafer, D. J., and Kaldy, J. E. (2014). Comparison of photosynthetic characteristics of the seagrass congeners *Zostera marina* L. and *Zostera japonica* Ascher. & Graeb. *Aquat. Bot.* 112, 91–97. doi: 10.1016/j.aquabot.2013.09.002
- Shafer, D. J., Kaldy, J. E., Sherman, T. D., and Marko, K. M. (2011). Effects of salinity on photosynthesis and respiration of the seagrass *Zostera japonica*: a comparison of two established populations in North America. *Aquat. Bot.* 95, 214–220. doi: 10.1016/j.aquabot.2011.06.003
- Silva, J., Sharon, Y., Santos, R., and Beer, S. (2009). Measuring seagrass photosynthesis: methods and applications. *Aquat. Biol.* 7, 127–141. doi: 10.3354/ab00173
- Sunda, W. G., and Cai, W. J. (2012). Eutrophication induced CO₂-acidification of subsurface coastal waters: interactive effects of temperature, salinity, and atmospheric Pco₂. *Environ. Sci. Technol.* 46, 10651–10659. doi: 10.1021/es300626f
- Talmage, S. C., and Gobler, C. J. (2009). The effects of elevated carbon dioxide concentrations on the metamorphosis, size, and survival of larval hard clams (*Mercenaria mercenaria*), bay scallops (*Argopecten irradians*), and Eastern oysters (*Crassostrea virginica*). *Limnol. Oceanogr.* 54, 2072–2080. doi: 10.4319/lo.2009.54.6.2072
- Thimijan, R., and Heins, R. (1983). Photometric, radiometric, and quantum light units of measure - a review. *HortScience* 18, 818–822.
- Thom, R. (1990). Spatial and temporal patterns in plant standing stock and primary production in a temperate seagrass system. *Bot. Mar.* 33, 497–510.
- Thom, R. M. (1996). CO₂-enrichment effects on eelgrass (*Zostera marina* L.) and bull kelp (*Nereocystis luetkeana* (mert.) P & R.). *Water Air Soil Pollut.* 88, 383–391.
- Thomsen, J., Haynert, K., Wegner, K. M., and Melzner, F. (2015). Impact of seawater carbonate chemistry on the calcification of marine bivalves. *Biogeosciences* 12, 4209–4220. doi: 10.5194/bg-12-4209-2015
- Unsworth, R. K. F., Collier, C. J., Henderson, G. M., and McKenzie, L. J. (2012). Tropical seagrass meadows modify seawater carbon chemistry: implications for coral reefs impacted by ocean acidification. *Environ. Res. Lett.* 7:024026. doi: 10.1088/1748-9326/7/2/024026
- Waldbusser, G. G., Hales, B., Langdon, C. J., Haley, B. A., Schrader, P., Brunner, E. L., et al. (2015). Saturation-state sensitivity of marine bivalve larvae to ocean acidification. *Nat. Clim. Change* 5, 273–280. doi: 10.1088/1748-9326/7/2/024026
- Waldbusser, G. G., and Salisbury, J. E. (2014). "Ocean acidification in the coastal zone from an organism's perspective: multiple system parameters, frequency domains, and habitats," in *Annual Review of Marine Science*, Vol. 6, eds C. A. Carlson and S. J. Giovannoni (Palo Alto, CA: Annual Reviews), 221–247.
- Wallace, R. B., Baumann, H., Grear, J. S., Aller, R. C., and Gobler, C. J. (2014). Coastal ocean acidification: the other eutrophication problem. *Estuar. Coast. Shelf Sci.* 148, 1–13. doi: 10.1016/j.ecss.2014.05.027
- Washington State Blue Ribbon Panel on Ocean Acidification and Ocean Acidification. (2012). *From Knowledge to Action, Washington State's Strategic Response*. Washington State Department of Ecology. Available online at: <https://fortress.wa.gov/ecy/publications/publications/1201015.pdf>
- Zimmerman, R. C., Kohrs, D. G., Steller, D. L., and Alberte, R. S. (1997). Impacts of CO₂ enrichment on productivity and light requirements of eelgrass. *Plant Physiol.* 115, 599–607.

Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Copyright © 2017 Miller, Yang and Love. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) or licensor are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.