



Site Characteristics Mediate the Relationship Between Forest Productivity and Satellite Measured Solar Induced Fluorescence

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Solar-Induced Chlorophyll Fluorescence (SIF) can provide key information about the state of photosynthesis and offers the prospect of defining remote sensing-based estimation of Gross Primary Production (GPP). There is strong theoretical support for the link between SIF and GPP and this relationship has been empirically demonstrated using ground-based, airborne, and satellite-based SIF observations, as well as modeling. However, most evaluations have been based on monthly and annual scales, yet the GPP:SIF relations can be strongly influenced by both vegetation structure and physiology. At the monthly timescales, the structural response often dominates but short-term physiological variations can strongly impact the GPP:SIF relations. Here, we test how well SIF can predict the inter-daily variation of GPP during the growing season and under stress conditions, while taking into account the local effect of sites and abiotic conditions. We compare the accuracy of GPP predictions from SIF at different timescales (half-hourly, daily, and weekly), while evaluating effect of adding environmental variables to the relationship. We utilize observations for years 2018–2019 at 31 mid-latitudes, forested, eddy covariance (EC) flux sites in North America and Europe and use TROPOMI satellite data for SIF. Our results show that SIF is a good predictor of GPP, when accounting for inter-site variation, probably due

to differences in canopy structure. Seasonally averaged leaf area index, fraction of absorbed photosynthetically active radiation (fPAR) and canopy conductance provide a predictor to the site-level effect. We show that fPAR is the main factor driving errors in the linear model at high temporal resolution. Adding water stress indicators, namely canopy conductance, to a multi-linear SIF-based GPP model provides the best improvement in the model precision at the three considered timescales, showing the importance of accounting for water stress in GPP predictions, independent of the SIF signal. SIF is a promising predictor for GPP among other remote sensing variables, but more focus should be placed on including canopy structure, and water stress effects in the relationship, especially when considering intra-seasonal, and inter- and intra-daily resolutions.

Keywords: Gross Primary Production, Solar-Induced Chlorophyll Fluorescence, canopy conductance, canopy structure, photosynthesis

INTRODUCTION

Gross Primary Production (GPP), which is a measure of the flux of carbon taken up by vegetation through photosynthesis, is the largest components (along with ecosystem respiration) of CO₂ exchange between terrestrial ecosystems and the atmosphere. Solar-Induced Chlorophyll Fluorescence (SIF) has been gaining popularity as a tool to estimate GPP indirectly. SIF data can be obtained through either tower- or airborne-based measurements (Chang et al., 2020) and satellite remote sensing instruments (Frankenberg et al., 2011), where the latter presents a promising alternative for accurate global GPP modeling.

Solar-Induced Chlorophyll Fluorescence represents a small fraction of the Photosynthetically Active Radiation (PAR) that is absorbed by chlorophyll pigments and re-emitted as a faint glow mainly in the range of 650–800 nm (Papageorgiou, 1975; Baker, 2008). Since both light reaction of photosynthesis and SIF compete for the same excitation energy, SIF can be an indicator of the functioning of the photosynthetic mechanism (Porcar-Castell et al., 2014). Conceptually, both GPP and SIF are considered to be proportional to Absorbed PAR (APAR; Monteith, 1972; Guanter et al., 2014). Non-Photochemical Quenching (NPQ), a process for excess energy dissipation (Jonard et al., 2020), is a third pathway for light use, potentially playing a major role in the GPP:SIF relations (Wohlfahrt et al., 2018).

Based on the information provided by SIF on the actual electron transport from Photosystem II to Photosystem I, Gu et al. (2019) derived fundamental equations linking SIF to C₃ and C₄ photosynthesis at the canopy level using a big leaf approach, thus bridging leaf scale and canopy scales for GPP:SIF. However, many canopy-level quantities used in the derivation could not be estimated through remote sensing approaches or flux measurements, which limits the wide use of this formulation. Nevertheless, it is a benchmark for investigating GPP:SIF at larger spatial timescales.

Solar-Induced Chlorophyll Fluorescence is reported to be linearly related to GPP at the diurnal and seasonal scales in various studies across a variety of sites (Joiner et al., 2014; Yang et al., 2015; Zhang et al., 2016b; Yang H. et al., 2017; Du et al.,

2019; Magney et al., 2019; He et al., 2020b; Qiu et al., 2020). However, the slopes of these linear GPP:SIF relations differ across sites, biomes, and vegetation types (Smith et al., 2018; Sun et al., 2018; Zhang et al., 2018b). SIF is shown to be a relevant indicator of crop productivity (Guanter et al., 2014; Guan et al., 2016, 2017; Zhang et al., 2018a; He et al., 2020a) and seasonal phenology (Joiner et al., 2014; Jeong et al., 2017; Yang H. et al., 2017). However, at instantaneous to hourly temporal scales, the GPP:SIF correlation is not as strong as at longer timescales, i.e., from days to seasons and years (Zhang et al., 2018c; Marrs et al., 2020). Observations and models at short timescales are needed to characterize the environmental effects that cause rapid variations (i.e., intra-daily, and inter-daily within season) of GPP, such as light saturation and water stress.

At sub-diurnal or half-hourly timescales, GPP:SIF was reported to follow an asymptotic trend (Li et al., 2018a; Chen et al., 2020). Indeed, GPP saturates at high APAR, while SIF keeps increasing as APAR increases, leading to a hyperbolic relationship between SIF and GPP at the instantaneous timescale (Damm et al., 2015; Gu et al., 2019). This hyperbolic-shaped relationship is less apparent over longer timescales, when variability in SIF and GPP is dominated seasonal variations in canopy structure, as estimated with the Leaf Area Index (LAI), whose effects are present in both SIF and GPP signals (Lu et al., 2018; Balzarolo et al., 2019; Dechant et al., 2020). Thus, the GPP:SIF correlation becomes more linear and relatively less sensitive to faster variations in the environmental drivers and plant physiological stress. Beyond physiological effects that govern GPP:SIF at the leaf level, non-linear relationships of GPP:SIF at the flux-footprint scale can therefore also be attributed to canopy structure. In addition, SIF is function of the view geometry. SIF is emitted from leaves exposed to sunlight, thus all conditions involved in the interaction between leaves, incident radiation, and view angle are important drivers of SIF observed signal. Therefore, canopy structure, and specifically gap fraction, leaf angle, and clumping index, and their interactions with the incident radiation angle, and the satellite viewing angle are important factors influencing GPP:SIF (Dechant et al., 2020).

The ambiguity in the GPP:SIF relations across timescales opens the door for further investigations of the effects of

short-term environmental variables on this correlation, notably water stress. Beside variations in vegetation structure, GPP can be modulated by both non-stomatal and stomatal regulation. In periods of water shortage or stress, indicated by high vapor pressure deficit (VPD) and low soil moisture (Zhou et al., 2019), stomata tend to close in order to reduce water loss through transpiration, simultaneously resulting in a decrease in photosynthetic rate. However, the extent of stomatal response to water stress is species specific (Matheny et al., 2015; Konings and Gentine, 2017). Furthermore, reductions in stomatal conductance under high VPD do not necessarily translate into a reduction in photosynthesis of the same magnitude due to variations in intrinsic water use efficiency (Zhang et al., 2019; Green et al., 2020; Grossiord et al., 2020). Non-stomatal limitation of photosynthesis can be induced by water-stress through xylem cavitation, decrease in mesophyll conductance for CO₂ (Flexas et al., 2016), and reduction in metabolic efficiency of the enzyme Rubisco (Grassi and Magnani, 2005).

The partitioning of APAR between GPP, NPQ, and SIF is sensitive to environmental conditions, such as incoming PAR, canopy structure (as represented by the fraction of the absorbed PAR, which depends mainly on LAI, but also other structural characteristics of the canopy, such as gap fraction, leaf clustering, and leaf angles), and soil water availability. At low light, most of the absorbed PAR is utilized for photosynthesis, thus increasing its efficiency. However, at high light, energy-consuming biochemical reactions of CO₂ assimilation and electron transport chain saturate, leading to a re-allocation of excess energy into SIF and NPQ (Porcar-Castell et al., 2014), thus modifying the partitioning between SIF and GPP. Such light saturation or water stress scenarios should lead to deviation from a linear relationship between GPP and SIF.

Differences in SIF responses can also stem from SIF measurement methods, which can be either active, i.e., ground-based through pulse amplitude-modulated measurements (Goulas et al., 2017), or passive (remote sensing) methods. Active measurements can directly estimate the yield as they emit an active signal, but this can only be done at a local level (Moya et al., 2019). Passive measurements, such as from satellites, have a wider spatial coverage but they are available at a much lower temporal frequency (depending on satellite pass time) and cannot directly control the incoming PAR. Furthermore, differences in the time of signal acquisition by the satellite play an important role. For example, morning-time acquisitions (such as with the MetOp-A satellite using GOME-2) are conducted with limited APAR (Lin et al., 2019) and are less affected by water stress, because morning-time VPD is low, and the vegetation tends to recover from stress overnight. However, morning-time acquisitions represent a very low GPP, whereas noontime acquisitions (such as with TROPOMI) will have much larger APAR thus GPP would be stronger, but could be more easily affected by light saturation and water stress.

This study focuses on investigating the GPP:SIF relations at three short timescales: half-hourly, daily, and weekly, testing the predictability of inter-daily GPP variations within the growing season. We further study the effects of water stress, light saturation, and site/ecosystem characteristics on the GPP:SIF

relations. We use data from 31 eddy covariance (EC) sites in the northern hemisphere. SIF data are taken from the recent TROPOMI measurements, taking advantage of its high spatial and temporal resolution, and measuring SIF near noontime, which leads to a better assessment of water stress and light saturation. Data are restricted to the 2018–2019 growing seasons of each site, as we focus on evaluating the intra-seasonal effects of physiological connections between SIF and GPP, and aim to avoid the longer timescales, where GPP variation is driven primarily by seasonal phenology.

MATERIALS AND METHODS

The following section goes through the details of data acquisition, processing, and analysis. **Table 1** comprises the definitions for all variables used throughout the manuscript, including their names, acronyms, and units.

Study Sites

Eddy covariance data were obtained through the AmeriFlux database¹, and the European Fluxes Database Cluster (EFDC)². This study was focused on temperate forest ecosystems located between 35°N and 65°N and considers the following IGBP land cover classifications: Evergreen Needleleaf Forests (ENF), Evergreen Broadleaf Forests (EBF), Deciduous Needleleaf Forests (DNF), Deciduous Broadleaf Forests (DBF), Mixed Forests (MF), and Woody Savannas (WSA). The study period comprised of growing seasons 2018 and 2019, depending on data availability from each site. We selected all sites that reported carbon dioxide fluxes for at least 50% of the year in 2018 and/or 2019. After filtering, 31 sites with 47 site-year growing seasons' data were available for analysis: BE-Vie, CH-Lae, CZ-BK1, CZ-Lnz, CZ-RAJ, CZ-Stn, DE-Hai, DE-HoH, DE-Hzd, DE-Obe, DE-Tha, FR-Fon, IT-SR2, RU-Fyo, SE-Nor, US-Me2, US-Me6, US-MMS, US-NC3, US-NR1, US-PFa, US-Rpf, US-SRM, US-Syv, US-UMB, US-UMd, US-Vcm, US-WCr, YS-xBR, US-xDL, and US-xRM. Sites' details (location on map, number of data points, DOI, coordinates, IGBP) were listed in **Supplementary Material**. Two sites provided hourly data instead of half-hourly (US-MMS and US-PFa), and in these sites data were interpolated to half-hourly for consistency.

Solar-Induced Chlorophyll Fluorescence Data

Solar-Induced Chlorophyll Fluorescence data were calculated from spectral observations by the TROPospheric Monitoring Instrument (TROPOMI) satellite, launched on October 13th, 2017. TROPOMI spectral range met the 743–758 nm range for detecting SIF (Köhler et al., 2018), and was provided at a spatial resolution of 7 × 3.5 km² with global coverage. SIF emissions were detected daily at approximately same solar time (~13:30 pm) at the equator. Swath time correction was applied at all sites in order to determine the local solar time for each

¹<https://ameriflux.lbl.gov/>

²<http://www.europe-fluxdata.eu/>

TABLE 1 | Variable names, acronym definitions, and units.

Acronym	Full name	Unit
APAR	Absorbed Photosynthetically Active Radiation	[$\mu\text{molPhoton m}^{-2} \text{s}^{-1}$]
b	Fitted coefficient for Hyperbolic GPP:SIF Model	[$\text{mW m}^{-2} \text{sr}^{-1} \text{nm}^{-1}$]
C_{CO_2}	Carbon Dioxide Molar Fraction	[$\mu\text{mol mol}^{-1}$]
C_p	Specific Heat of Dry Air	[$\text{J g}^{-1} \text{K}^{-1}$]
EF	Evaporative Fraction	–
FC	Carbon Dioxide Flux	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
fPAR	Fraction of Absorbed Photosynthetically Active Radiation	–
g_c	Stomatal Conductance	[m s^{-1}]
GPP	Gross Primary Production	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
$\text{GPP}_{\text{gap-filled}}$	Gap-filled Gross Primary Production	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
GPP_{max}	Fitted coefficient for Hyperbolic GPP:SIF Model	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
H	Sensible Heat Flux	[W m^{-2}]
H_0	Step Function to zero negative $\cos(\theta)$	–
LAI	Leaf Area Index	–
LE	Latent Heat Flux	[W m^{-2}]
NEE	Net Ecosystem Exchange	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
NIR	Near-Infrared Reflectance	–
NPQ	Non-Photochemical Quenching	–
PAR	Photosynthetically Photosynthetic Active Radiation	[$\mu\text{molPhoton m}^{-2} \text{s}^{-1}$]
PRI	Photochemical Reflectance Index	–
R_{11}	Reflectance at band 11 (526–536 nm)	–
R_{12}	Reflectance at band 12 (546–556 nm)	–
r_{ah}	Aerodynamic Resistance	[s m^{-1}]
R_{eco}	Ecosystem Respiration	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
RH	Relative Humidity	–
R_n	Net Radiation	[W m^{-2}]
SC	Storage Term of Carbon Dioxide Flux	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
SIF	Solar Induced Fluorescence	[$\text{mW m}^{-2} \text{sr}^{-1} \text{nm}^{-1}$]
$\text{SIF}_{\text{Daily}}$	Upscaled Daily SIF Average	[$\text{mW m}^{-2} \text{sr}^{-1} \text{nm}^{-1}$]
SIF_{Inst}	SIF Value at Time of Measurement	[$\text{mW m}^{-2} \text{sr}^{-1} \text{nm}^{-1}$]
SM	Soil Moisture	–
SWC	Soil Water Content	–
TA	Air Temperature	[K]
t_m	SIF Time of Measurement	–
u^*	Friction Velocity	[m s^{-1}]
VPD	Vapor Pressure Deficit	[Pa]
WS	Wind Speed	[m s^{-1}]
α	Linear Mixed Effect (LME) Model Slope	[$\mu\text{molCO}_2 \text{s}^{-1} \text{mW}^{-1} \text{sr nm}$]
β_{Site}	Site Level Intercept of LME Model (Random Intercept)	[$\mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$]
γ	Psychrometric Constant	[Pa K ⁻¹]
γ_i	Response slope of additional variable in LME	Depends on the units of variable, in LME
Δ	Slope of the saturated vapor pressure curve	[Pa K ⁻¹]
θ	Solar Zenith Angle	Degree
ρ	Air density	[kg m^{-3}]
ρ_a	Molar Density of Air	[mol m^{-3}]

observation. Overall SIF satellite measurements fell between 10 am and 2 pm local time. Observations obstructed by cloud-cover conditions were filtered out. SIF data were reported in milliwatt per meter squared, per steradian, per nanometer [$\text{mW m}^{-2} \text{sr}^{-1} \text{nm}^{-1}$].

In order to get equivalent SIF values at the daily timescale, we used the scaling approach proposed by Frankenberg et al. (2011) to convert instantaneous SIF to daily average SIF. This method accounts for the variations in overpass time (including due to swath), length of day, and solar zenith angle:

$$\text{SIF}_{\text{Daily}} = \text{SIF}_{\text{Inst}} \times \frac{1}{\cos(\theta(t_m))} \int_{t=t_m-12\text{hours}}^{t=t_m+12\text{hours}} \cos(\theta(t)) \times H_0(\cos(\theta(t))) dt \quad (1)$$

where $\text{SIF}_{\text{Daily}}$ is the upscaled daily SIF average, SIF_{Inst} is the SIF value at time of measurement t_m , $\theta(t_m)$ is the corresponding solar zenith angle, and H_0 is a step function where $H_0(\cos(\theta(t)))$ equals zero when $\cos(\theta(t))$ is negative and equals one when $\cos(\theta(t))$ is positive (Köhler et al., 2018). The integral is calculated using a time interval, dt , of 10 min.

Gross Primary Production Data

All sites reported turbulent net carbon dioxide fluxes (column FC, in the EC data), but Net Ecosystem Exchange (NEE) and GPP were not provided by all sites. For consistency, a unified modeling approach for estimating GPP was followed across all sites. NEE was calculated as the sum of the turbulent Flux of CO_2 (FC) and Storage of CO_2 (SC):

$$\text{NEE} = \text{FC} + \text{SC} \quad (2)$$

Some of the sites did not provide SC, thus SC was approximated using the following equation:

$$\text{SC} = \rho_a \times \int_0^h \frac{dC_{\text{CO}_2}}{dt} dz \quad (3)$$

where ρ_a is the molar density of the air [mol m^{-3}], and C_{CO_2} is Carbon Dioxide molar fraction [μmolmol^{-1}]. Integration includes all available measurements for carbon dioxide concentration from the ground to the CO_2 flux measurement height, h .

Growing Season

The growing season was defined using the carbon uptake period (i.e., carbon flux phenology, Garrity et al., 2011). A 7-day moving average of NEE was calculated and then, the peak seasonal NEE, i.e., the most negative 7-day average NEE for CO_2 uptake. The start and end of each site-year's growing season were considered as the first and last day, respectively, where carbon uptake rates were above (more negative than) a threshold of 5% of the peak seasonal NEE. Growing season start and end dates of each site are provided in **Supplementary Table 1.3**. Following the identification of the time period of the growing season, a seasonal friction velocity (u^*) filter threshold value was defined following the approach by Reichstein et al. (2005) and filtered flux data during times of low turbulence, below the u^* threshold value.

Artificial Neural Network

An Artificial Neural Network (ANN) algorithm (Moffat et al., 2007), with parametric choices as described in Morin et al. (2014),

was used to model fluxes during each growing season at each site. For each EC flux variable, 50% of valid observations were used to train the network, 25% of the remaining valid data to evaluate the network goodness of fit, and 25% to validate the final model. 100 networks were run per site-season, and the final model was the ensemble average of the top-fitting 10% of these. Separate ANN models were trained for daytime and nighttime data for each site-season.

During the process of modeling GPP, we first used an ANN to gap fill sensible heat (H) and latent heat (LE) fluxes. Then, an ANN was used to model ecosystem respiration (R_{eco}). It was assumed that during the night there was no photosynthetic activity, therefore, nighttime NEE was equal to R_{eco} . Thus, the ANN was trained with nighttime data. The resulting model was used to gap fill nighttime R_{eco} and NEE, and to simulate daytime R_{eco} . We assumed no nighttime GPP. The daytime was determined based on the Greenwich Mean Time (GMT) offset of each site and its latitude (Tramontana et al., 2020). During daytime, GPP was:

$$GPP = NEE - R_{eco} \tag{4}$$

Where, during the daytime period of observations, GPP was a negative quantity and R_{eco} positive. This approach resulted in an observation gap whenever NEE observations were not available. We used these GPP data (corresponding with time when observed NEE values were available from EC and not gap-filled) for the half-hourly analysis. We used an ANN to gap fill NEE and used the complete time series of NEE and R_{eco} to calculate a gap-filled time series for GPP ($GPP_{gap-filled}$). Daily D and weekly GPP were calculated from $GPP_{gap-filled}$. The drivers of each of the ANN models for LE, H, R_{eco} , and $GPP_{gap-filled}$ are listed in **Table 2**. It should be noted that PAR or net radiation, and soil moisture are usually used as GPP drivers, however, many sites did not report them and thus, requiring them as ANN input will lead to losing many sites for missing the corresponding data.

Ecosystem State Variables

Several environmental variables were used in the analysis along with SIF and GPP: canopy conductance (g_c), evaporative fraction (EF), APAR, photochemical reflectance index (PRI), PAR, soil moisture (SM), VPD, RH, TA, LE, fPAR, LAI, and NIR. All these variable data were provided using EC and remote sensing measurements or derived from data from EC measurements.

Full descriptions for the source and calculation of each variable are listed below.

MODIS data product (MCD15A2H v006) was used for leaf area index (LAI) and fraction of absorbed PAR (fPAR), at 8-day time resolution and 500 m pixel size for each site using the Earthdata open-source repository³. APAR [$\mu\text{molPhoton m}^{-2} \text{s}^{-1}$] was calculated using the following equation:

$$APAR = fPAR \times PAR \tag{5}$$

where PAR [$\mu\text{molPhoton m}^{-2} \text{s}^{-1}$] observations were measured at the EC flux sites.

Surface conductance was determined by inverting the Penman-Monteith model (Monteith, 1972) with measured evapotranspiration and meteorological data. The Penman-Monteith model separates the effects of aerodynamic and surface conductances. In forest ecosystems, where soil evaporation is small relative to leaf transpiration, the surface conductance is a good proxy for canopy conductance, g_c [m s^{-1}] (Novick et al., 2016), which, when the leaf surfaces do not hold standing water (as is the case most of the time, except immediately after precipitation), represents a volume-weighted integral of the leaf-level stomatal conductance. Noting that observation times around rain events are excluded from the analysis because no SIF measurements are available for thick cloud conditions, thus:

$$g_c = \left[\frac{r_{ah} \times \Delta \times R_n + \rho \times C_p \times VPD}{LE \times \gamma} - \frac{r_{ah} \times (\gamma + \Delta)}{\gamma} \right]^{-1} \tag{6}$$

where, r_{ah} is the water vapor aerodynamic resistance [s m^{-1}], Δ is the slope of the saturated vapor pressure curve [Pa K^{-1}], R_n is the net radiation [W m^{-2}], ρ is the air density [kg m^{-3}], C_p is specific heat of dry air [$\text{J g}^{-1} \text{K}^{-1}$], VPD is the vapor pressure deficit [Pa], γ is the Psychrometric constant [Pa K^{-1}], and LE is the latent heat flux [W m^{-2}] measured at the site. r_{ah} was approximated using site observations of u^* and wind speed (\bar{u}) based on the empirical approach by Monteith and Unsworth (1990):

$$r_{ah} = \frac{\bar{u}}{u^{*2}} + \frac{6.2}{u^{*2/3}} \tag{7}$$

Evaporative Fraction was also used as an indicator for water stress conditions, where it is the ratio of latent heat flux to the

³<https://search.earthdata.nasa.gov/>

TABLE 2 | Variables used as drivers in the ANN models used to gap fill different fluxes.

ANN Model	Drivers	Goodness of fit (R^2)		RMSE	
		Day	Night	Day	Night
H [W m^{-2}]	VPD	0.61 ± 0.08	0.10 ± 0.10	50.4 ± 9.3	8.6 ± 3.9
LE [W m^{-2}]	VPD, U^*	0.63 ± 0.10	0.25 ± 0.14	42.2 ± 13.0	6.3 ± 3.4
R_{eco} [$\mu\text{molCO}_2 \text{ m}^{-2} \text{ s}^{-1}$]	LE, H, VPD, WS, U^* , TA	N/A	0.36 ± 0.2	N/A	1.0 ± 0.36
NEE [$\mu\text{molCO}_2 \text{ m}^{-2} \text{ s}^{-1}$]	LE, H, VPD, WS, U^* , RH, TA	0.62 ± 0.08	N/A	3.0 ± 0.96	1.1 ± 0.4
$GPP_{gap-filled}$ [$\mu\text{molCO}_2 \text{ m}^{-2} \text{ s}^{-1}$]	NEE- R_{eco}	N/A	N/A	N/A	N/A

The table includes goodness-of-fit statistics for ANN models used to gap fill fluxes. Values represent averages ± standard deviation of R^2 and root mean square error (RMSE) across all sites and years.

sum of latent heat flux and sensible heat flux, H [W m^{-2}]:

$$EF = \frac{LE}{H + LE} \quad (8)$$

Surface soil water content (SWC) observations were provided by ground-based measurements (half hourly) at the EC sites, and from the spaceborne SMAP dataset⁴ at 6 am and 6 pm local time at daily basis. Both data sets were used separately in the analysis.

Photochemical Reflectance Index (PRI) is sensitive to changes in carotenoid pigments and used as an inverse proxy for NPQ at short timescales. MODIS data product (MYDOCGA, v006) is used to get reflectance at bands 11 and 12 at daily resolution and calculate a proxy of PRI at each site as described in Wang et al. (2020):

$$PRI = \frac{R_{11} - R_{12}}{R_{11} + R_{12}} \quad (9)$$

where R_{11} and R_{12} are the reflectance at bands 11 (526–536 nm) and 12 (546–556 nm), respectively. MODIS data product (MOD13Q1 v006) was used for getting NIR. **Table 1** above includes all the acronyms used in this study.

Data Analysis

Data analysis involves three timescales: half-hourly, daily, and weekly. Instantaneous SIF measurements are used for half-hourly timescale, upscaled daily average SIF values (Eq. 1) are used for the daily timescale, while weekly averages of SIF measurements are used for the weekly timescale. As for GPP, the half-hourly window during which a SIF observation was available is used for the half-hourly timescale, while daily and weekly averaged GPP_{gap-filled} values are used for the daily and weekly timescales, respectively. For all statistical inferences, a 0.05 significance level is considered.

The GPP:SIF relations was evaluated at multiple spatial scales and at the three temporal scales: half-hourly, daily, and weekly for each spatial scale. First, it was tested at the site level, where both a linear and a hyperbolic fit models were tested at each site.

Second, data were pooled across all sites and fitted using a linear regression model. Fourth, data were fitted using a Linear Mixed Effect (LME) model, where sites is considered as a random effect, SIF as a fixed effect, and GPP as the response variable:

$$GPP \approx \alpha \times SIF + \beta_{Site} + \gamma_{Site} \times SIF \quad (10)$$

where α is the SIF-driven slope (fixed slope), β_{Site} is the site-level intercept (random intercept), and $\gamma_{Site} \times SIF$ is the interaction between SIF and site. Third, we used similarly structured LME models, but added additional environmental variables as fixed effects using a forward stepwise regression for the LME model:

$$GPP \approx \alpha \times SIF + \sum_i (\gamma_i \times Variable_i) + \beta_{Site} \quad (11)$$

where γ_i is the effect (response slope) of the i th environmental variable. Variables considered in the forward stepwise regression

are: g_c , TA, SM, VPD, LAI, PRI, APAR, EF, LE, fPAR, and NIR. Variables were added to the model based on their pairwise R^2 with the residuals of the current model, and only if their addition decreased the Akaike Information Criterion (AIC; Hosmer et al., 2013).

To develop a more reliable estimate of GPP using solely remote sensing data sources, a simple decision tree model was developed where SIF was considered as a predictor of GPP along with other remote-sensing variables. These predictors included top-surface soil moisture estimates, LAI, PRI, fPAR, NIR, LE, and APAR calculated using MODIS products, which was calculated following Eq. 5, but with PAR retrieved from MODIS instead of PAR observations at the EC-tower location. **Table 3** shows the source of each of the remote sensing product, in addition to their spatial resolution. For each variable, grids were aggregate consistent with a typical Eddy-Covariance footprint ($1.5 \times 1.5 \text{ km}^2$). For variables with larger spatial resolution, spatial interpolation was used. Site and IGBP were included as categorical predictors in this selection-tree model in order to account for site characteristics. The contribution portion of each variable explaining GPP was used to indicate the relative importance of each variable in predicting GPP. The decision tree includes 50 layers, and 3 splits per tree. Twenty percent of the data was used for validation. It should be noted that this model was different from the LME or linear regression, as the decision tree did not assume the effects of the drivers to be linearly continuous. An important feature of decision tree models was allowing predictors' classification by determining the contribution portion of each variable in predicting GPP.

Throughout the processing of the results, MATLAB (2018) was used for data processing, ANN modeling, plotting, and pairwise and multiple linear regressions. JMP Pro 14 was used for statistical inference using a general mixed-effect linear models (LME), and for the decision-tree model (JMP, 2018; MATLAB, 2018).

RESULTS

Gross Primary Production vs. Solar-Induced Chlorophyll Fluorescence at Each Site

At the half-hourly timescale, only 19% of the sites showed significant pairwise linear correlations between GPP and SIF. At

TABLE 3 | Remote sensing products used in the decision tree model and their corresponding spatial resolution.

Variable	Satellite product	Spatial resolution
SM	SMAP	9000 m × 9000 m
LAI	MCD15A2H v006	500 m × 500 m
PRI	MYDOCGA v006	1000 m × 1000 m
fPAR	MCD15A2H v006	500 m × 500 m
NIR	MOD09GQ v006	250 m × 250 m
LE	MOD16A2 v006	500 m × 500 m
PAR	MCD18A2 v006	5600 m × 5600 m

⁴<https://nsidc.org/data/smap>

the daily timescale, 52% of the sites showed significant correlation between GPP and SIF, and at the weekly timescale, 42% of the sites showed a significant correlation (**Supplementary Table 1.2**). The highest goodness of fit at all timescales was reached at the site SE-Nor, where weekly data had the highest R^2 of 0.66 (**Figure 1**). The number of data points for each fit ranged between 10 and 60, depending on the data availability at each site. The detailed results and statistics for each site are shown in **Supplementary Material**.

Alternatively, we tested the regression between GPP and SIF at half-hourly resolution assuming a hyperbolic relationship, following Damm et al. (2015):

$$GPP = GPP_{max} \times \frac{SIF}{SIF + b} \quad (12)$$

where GPP_{max} and b were fitting coefficients. This resulted in only two sites showing a significant but very weak relationship (**Figure 2**).

Gross Primary Production vs. Solar-Induced Chlorophyll Fluorescence Over All Sites

We tested the linear relationship between GPP and SIF, pooled across all sites of the same IGBP classification. No significant relationship was found for all sites, while a very weak correlation was found for ENF and DBF (**Figure 3**). As for MF a weak ($R^2 = 0.09$) but significant negative linear regression between GPP and SIF was found. However, this was due to the fact that data were clustered by site within the MF biome (**Figure 3D**), where some sites, such as CZ-Lnz and BE-Vie, had high GPP but low SIF relative to other sites, leading to an overall apparent negative GPP:SIF relations in this biome. Similar results were found for daily and weekly timescales.

When we included the variation between sites as a random effect (*Site*) in a LME, the resulting GPP:SIF correlation was highly significant, whereas the interaction term between SIF and site (i.e., assuming there was a different GPP:SIF slope at each site in addition to different intercepts) was found not to have a significant effect, and therefore was not included in the final model (**Figure 4**).

The effect of adding other environmental variables, in addition to SIF, as predictors for GPP was significant, but with different results among timescales (**Table 3**). It was found that water-status

related variables, g_c , LE, and EF had the strongest impact on the goodness of fit (evaluated in terms of improvement in correlation R^2) of the GPP at all three timescales. Daily and weekly timescales were characterized by having a set of environmental variables that improved the relationship between GPP and SIF, namely variables related to phenology and canopy structure (g_c , LAI, fPAR, NIR, and APAR), which was not the case for the half-hourly model.

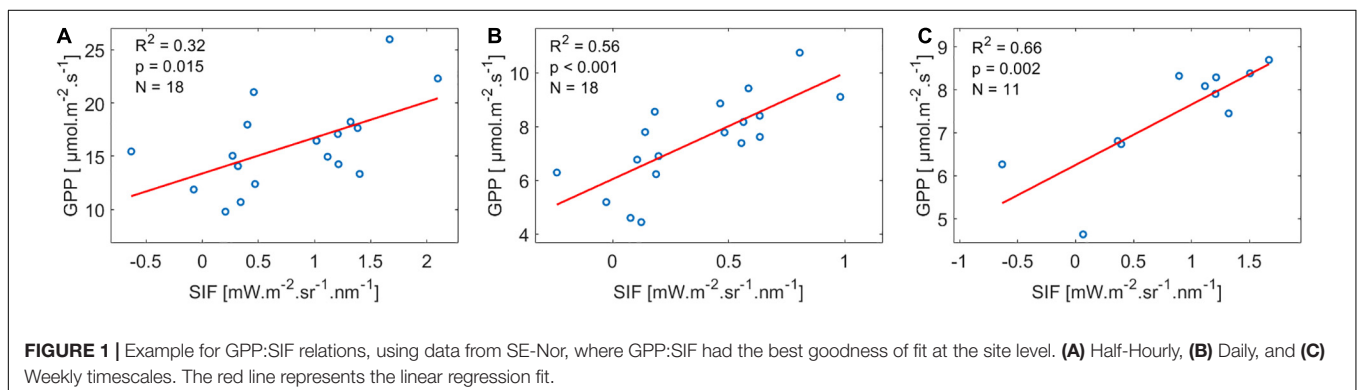
It should be noted that many of the different environmental variables involved in the analysis are inter-correlated, particularly those related to leaf color that vary with a similar seasonal phenological pattern, i.e., LAI, fPAR, and NIR. Nevertheless, each of these might have an independent component in its information content regarding GPP:SIF. In the forward stepwise regression, we sorted the variables in order of their pairwise regression with the residuals of GPP:SIF. We then added each element in order to the multivariate regression model. A variable that was perfectly correlated to another variable already included in the model would not have any information content remaining to improve the model. Our model showed that LAI, fPAR, and NIR all added a significant information components to the model at both daily and weekly timescales. These three variables were therefore, further considered in the more complex decision tree model.

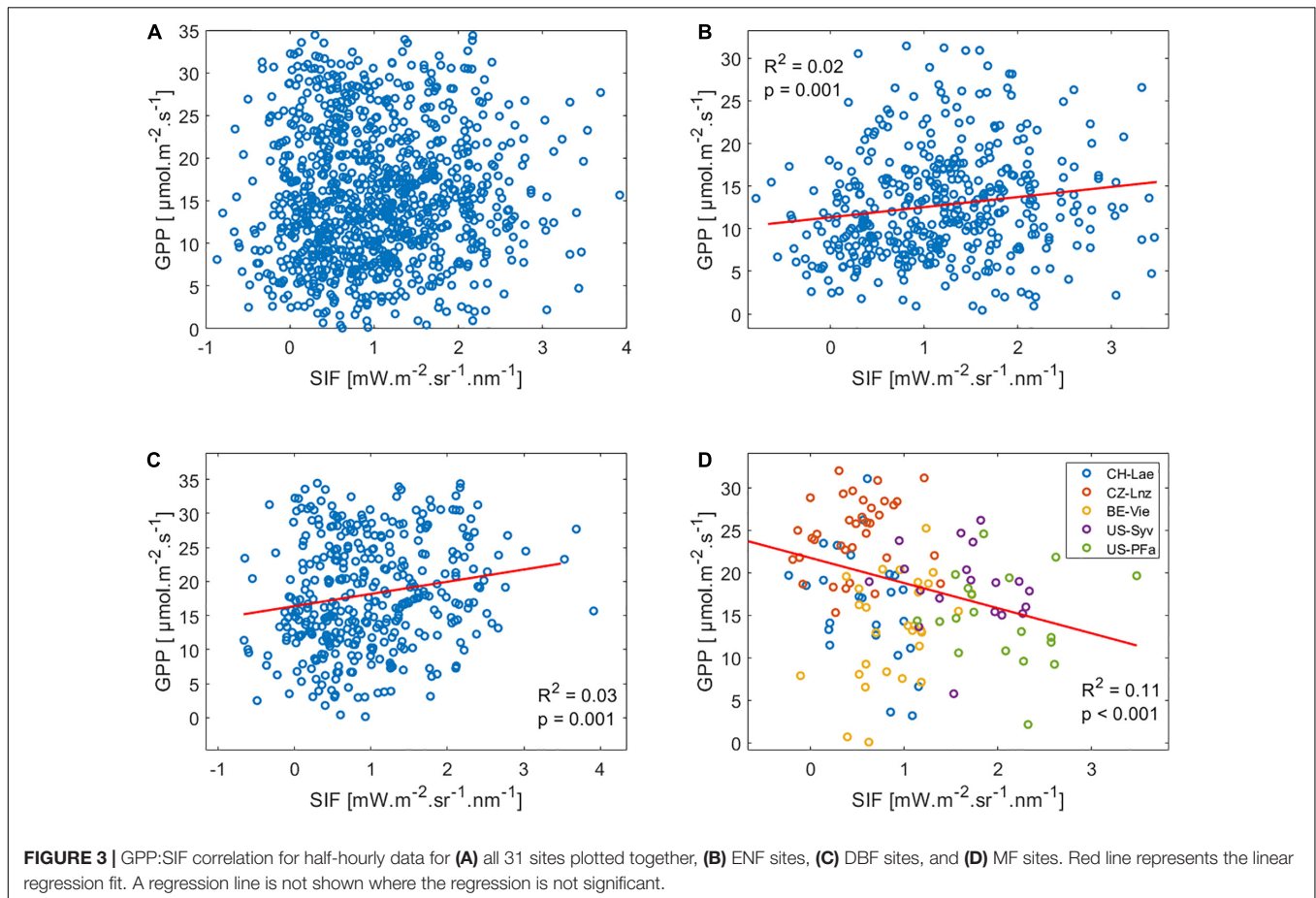
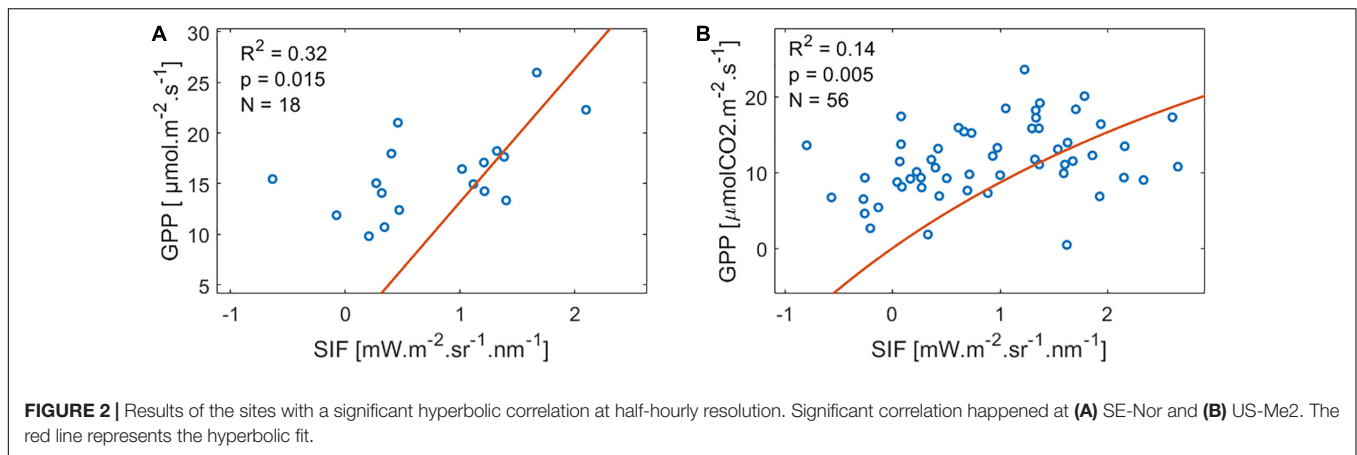
In order to further test the influence of g_c and canopy on the empirical relationship between GPP and SIF, the distribution of the LME model errors of GPP vs. SIF (expressed as model residuals = observations – model predictions) is studied under low (the lowest 33% at each site), intermediate (33th to 67th percentile per site), and high (highest 33% at each site) g_c and fPAR separately (**Figure 5**).

The decision-tree model shows that sites has the highest contribution to predicting GPP among other remote sensing predictors, while SIF has a minimal contribution, along with IGBP, PRI, SM, and fPAR. APAR, LAI, LE, and NIR showed a relatively high contribution with divergence in contribution portions across timescale (**Figure 6**). Model precision (estimated through R^2) increased with decreasing temporal resolution.

DISCUSSION

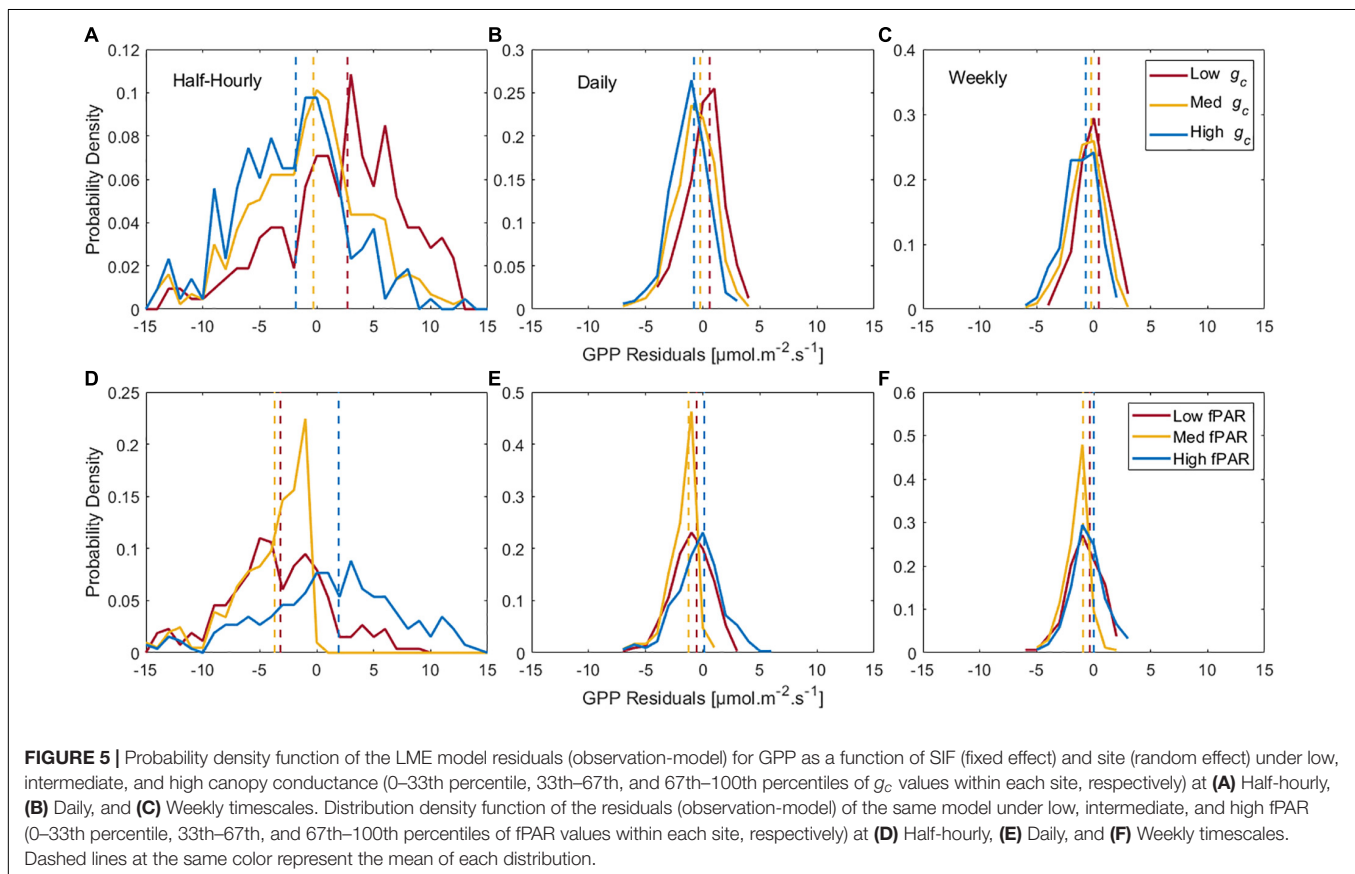
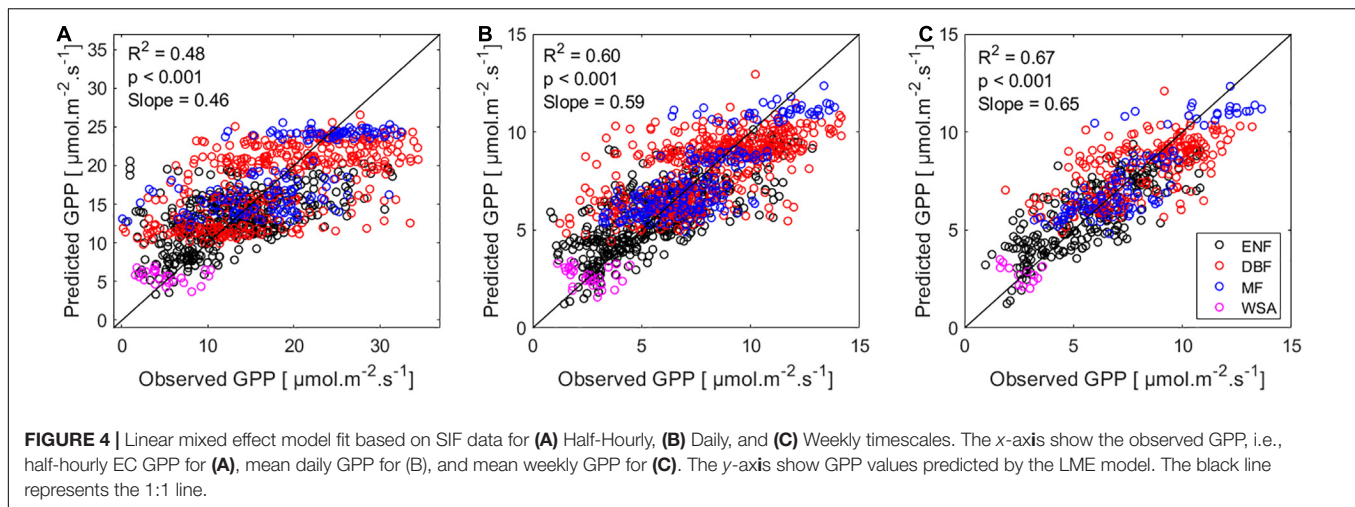
Other studies of GPP:SIF using EC method and remote sensing (e.g., Li et al., 2018b) showed stronger and more consistent GPP:SIF correlations than the ones we found. We hypothesize





that the main reason for this difference in the goodness of fit of the SIF models is that our study focused on the inter-daily and inter-weekly variation within the growing season. Therefore, diurnal and seasonal variations of GPP and SIF, which are stronger and more predictable than the inter-daily variations, are not emphasized in this study. Our findings that weekly and daily average GPP are better correlated to SIF than half-hourly GPP are consistent with earlier studies, which have shown a stronger linear relationship at lower temporal resolution. This trend across timescales indicates that SIF is a good predictor for seasonal

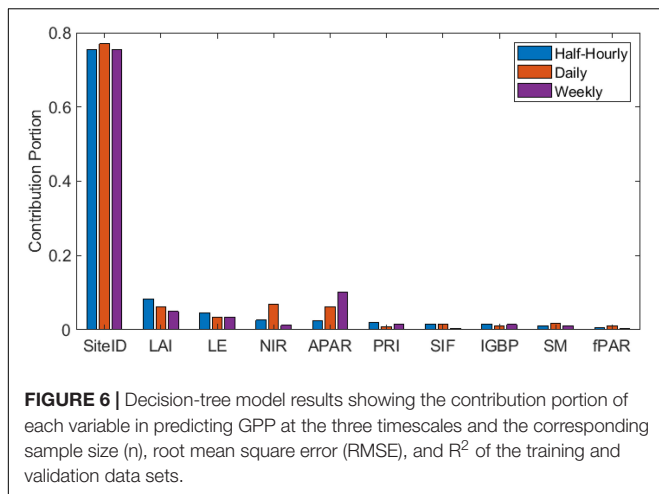
variations of GPP (Magney et al., 2019), but not necessarily for the instantaneous photochemical activity (Magney et al., 2020). Site-level analysis shows limited relationship between SIF and GPP at midday, half-hourly resolution using both linear and asymptotic fits, while it is expected to have a hyperbolic correlation at this time of day. This can be due to light saturation that dominates midday photosynthesis under most conditions (Zhang et al., 2016a). Thus, TROPOMI-sampled points at solar noon would lie mostly on the asymptotic end of the GPP vs. SIF hypothetical non-linear curve (Damm et al., 2015). In order



to get the hypothetical non-linear correlation between GPP and SIF, data points with minimal light saturation are needed to complement the hyperbolic shape of the curve, i.e., morning SIF data should be available, which is not the case with TROPOMI, or midday data with less stress conditions.

It is possible that variability in species specific parameters that governs GPP:SIF creates high levels of within-site variation in species-rich and heterogeneous sites. If this hypothesis is true, the sites with the strong GPP:SIF correlation are expected to be the ones that have the lowest species diversity. Indeed, such is

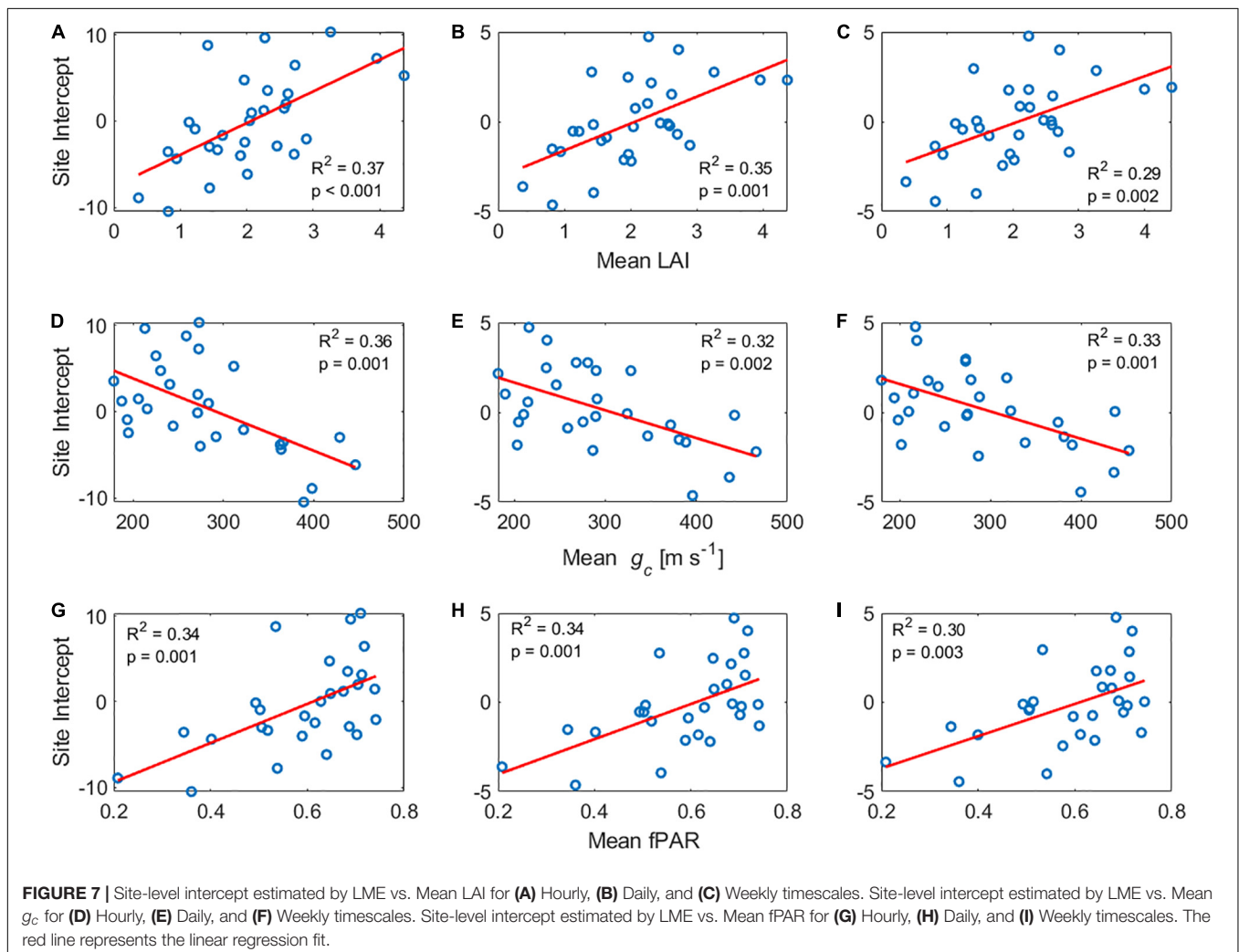
the case in some sites, for example SE-Nor is composed of only pine and spruce (Lindroth et al., 1998), CZ-BK1 and CZ-RAJ are composed of monoculture Norway spruce (Sedláč et al., 2010; McGloin et al., 2018), in BE-Vie three species represent more than 80% of vegetation (Aubinet et al., 2001), and US-NC3 is composed almost exclusively of loblolly pine (Yang Y. et al., 2017). However, other sites, where GPP and SIF are significantly and strongly correlated, are among the most diverse. For example, US-UMB has seven different species with relatively equal dominance among three of them (Matheny et al., 2014),



species diversity index for each site, and we could not test the significance of the assumed negative relationship between species diversity and GPP:SIF goodness of fit. Nonetheless, we investigated the community-composition complexity using the spatial heterogeneity of LAI (similar to the approach by Chu et al., 2021). This quantity is an indirect estimation of the variation in vegetation structure, and could be used as an indirect proxy for the diversity of the dominant species. The coefficient of variability for LAI (using spatial standard deviation) was provided by the MODIS MCD15A2H product at each site. LAI values were calculated at a 1.5 × 1.5 km² spatial resolution, which is close to a typical EC footprint. We found no significant relationship between the R² of site-level GPP:SIF correlations and the LAI coefficient of variability (plots and detailed results are shown in **Supplementary Material**). It should be noted that the LAI coefficient of variability is an indicator of overstorey dominant species, thus neglecting the effect of understorey species, which would significantly contribute to the total fluorescence signal when light conditions allow.

RU-Fyo has 49% of spruce, 12% of pine forests, 28% birch, 14% aspen, and 1% alder (Sogachev et al., 2002). The site descriptions in AmeriFlux and EFDC do not enable calculating a formal

Another hypothetical cause of the low correlation between GPP and SIF may be the lack of representativeness of the SIF pixel



relative to the EC-flux footprint (Chu et al., 2021). While these land cover overlap, but potentially different areas, it is possible that species specific response within the EC-footprint captures a somewhat different evapotranspiration dynamic than the SIF pixel. However, the finding that the spatial variability of LAI is not correlated with the goodness of fit of GPP:SIF is indicative that low footprint representativeness (typical to sites with high LAI variation) is not the lead cause of the low GPP:SIF fit. Furthermore, TROPOMI offers a higher resolution ($7 \times 3.5 \text{ km}^2$) compared than other satellites that provide SIF data, such as GOME that has a resolution of $80 \times 40 \text{ km}^2$, and a better spatial coverage than OCO-2 which have a higher spatial resolution of $2 \times 1.3 \text{ km}^2$ (Köhler et al., 2018).

Given the lack of significant effect of within-site vegetation heterogeneity, we hypothesize that site-specific environmental and structural conditions (water status, canopy structure, degree of isohydricity, soil texture) are probably more important in driving GPP:SIF than the ecosystem characteristics that classify its IGBP type or its species richness and heterogeneity (Zhang et al., 2018c; Dechant et al., 2020). This hypothesis is further supported by the LME models results, which show a strong variability in GPP:SIF relations among sites, and the decision tree model, which shows that Site is the main contributor to GPP prediction. Li et al. (2020) shows that the linear GPP:SIF relations is affected by the growth stages of maize during the growing season. Migliavacca et al. (2017) shows that the relationship is a function of nutrient addition (Nitrogen, and Phosphorous), which induces changes in canopy structure and functional traits. We find that the size of the site-specific random effect in the LME model can be predicted by the site-level mean LAI, fPAR and canopy conductance at all timescales (Figure 7). A multiple-linear regression of sites' intercept with LAI and g_c together yielded an R^2 of 0.56 and the interaction between g_c and LAI was not significant. This result further substantiates the hypothesis that both canopy structure and plant function control GPP:SIF, to a large degree, and offers hope for effectively predicting the site-level intercept that is needed for the global applicability of GPP:SIF (Dechant et al., 2020; Kim et al., 2021).

Table 4 shows that water stress indicators, provide significant improvement to the GPP:SIF correlation among other environmental variables at all timescales. Water stress variables, specifically stomatal conductance significantly improved the model at all timescales. Phenology indicators (LAI, NIR, and fPAR) and APAR improve GPP:SIF at daily and weekly resolutions, which is in agreement with earlier studies, notably Magney et al. (2020), but not at half hourly resolution. Contrary to our expectations and earlier findings (Helm et al., 2020; Marrs et al., 2020), we found that the GPP:SIF LME model is over-estimating GPP under high and intermediate g_c (Figures 5A–C) and underestimating GPP under low g_c . Overestimation of GPP happens at low and intermediate fPAR at the half-hourly timescale. SIF is driven by canopy properties like chlorophyll content, LAI, and angle distribution of leaves more than by canopy biochemistry (Frankenberg and Berry, 2018). At high chlorophyll content, light absorption per unit of chlorophyll decreases, thus resulting in a non-linear

TABLE 4 | Effects of environmental variables on the goodness-of-fit of LME of GPP vs. SIF using forward stepwise regression.

Half-Hourly		Daily		Weekly	
Variable	R^2	Variable	R^2	Variable	R^2
SIF + Site	0.476	SIF + Site	0.602	SIF + Site	0.672
g_c	0.532	LE	0.693	g_c	0.717
EF	0.567	g_c	0.711	APAR	0.786
–	–	LAI	0.722	fPAR	0.802
–	–	VPD	0.724	EF	0.815
–	–	TA	0.732	NIR	0.818
–	–	fPAR	0.733	–	–
–	–	APAR	0.746	–	–
–	–	SWC	0.740	–	–

Values that are not reported correspond to non-significant contributions to the model (as tested by the AIC). Reported R^2 are the cumulative, whole-model R^2 (including all variables up to the current).

TABLE 5 | Decision tree model statistics.

Timescale	Model	R^2	RMSE [$\mu\text{mol}/\text{CO}_2 \text{ m}^{-2} \text{ s}^{-1}$]	n
Half-Hourly	Training	0.582	5.021	781
	Validation	0.355	6.100	196
Daily	Training	0.689	1.569	957
	Validation	0.649	1.735	211
Weekly	Training	0.793	1.180	454
	Validation	0.704	1.417	128

relationship between chlorophyll content and light absorption (Porcar-Castell et al., 2014). In such case, a high SIF signal is expected to overestimate GPP. Since our study focuses on the growing season, high chlorophyll content governs the state of vegetation, resulting in GPP overestimation under low and intermediate fPAR. This effect can also be responsible for the overestimation of GPP under high stomata conductance. At low fPAR, low GPP is driven by low APAR instead of stomatal conductance.

The decision-tree model considered only remote sensing variables as GPP predictors in addition to site. The model results identified the strong effect of inter-site variability in predicting GPP. It also found low contribution of SIF to GPP predictability, compared to other variables. However, surprisingly, the decision-tree model resulted in an overall similar performance as the GPP:SIF LME model when comparing the R^2 of the two models at each timescale (Table 5 and Figure 4). We hypothesize that the strong cross-correlation of many environmental variables have limited the information content of the overall ensemble of variables and explains the similar goodness of fit of these two very different models. Decision-tree model results emphasized the role of variables indicative of canopy structure (NIR, APAR, and LAI) and water status (LE) in supplementing SIF-based prediction of GPP, noting that fPAR showed a lower contribution portion probably due to its high covariance with LAI.

CONCLUSION

Our study evaluates spaceborne SIF from TROPOMI as a predicting variable for inter-daily variation of GPP during the growing season in 31 EC sites. A strong inter-site variability in the intercept of GPP:SIF regression is found. The need for site-specific intercepts as coefficients in a model for accurately predicting GPP from SIF limits the applicability of SIF as a globally observable surrogate of GPP. However, our results suggest that this intercept is driven by canopy structure and site-level vegetation function and is predictable using site-level, season-long, mean fPAR, LAI, and canopy conductance. These latter showed as well to significantly improve the LME model at all timescales. Thus, canopy structure and water status variables at site level are important factors to account for when using SIF as a predictor of GPP intra-daily variations.

DATA AVAILABILITY STATEMENT

Observations of carbon fluxes and other site-level meteorological variables are available through AmeriFlux (<https://ameriflux.lbl.gov/data/aboutdata/>) and European Fluxes Database Cluster (EFDC) (<http://www.europe-fluxdata.eu/home/data/data-policy>). All Site IDs and DOIs for all sites used are listed in **Table S1.1** in **Supplementary Material 1**. MODIS datasets MCD15A2H v006, MCD18A2 v006, MOD13Q1 v006, MOD16A2 v006, and MYDOCGA v006 provide LAI and fPAR, PAR, PRI, NIR, LE, and PRI data and are available through NASA's EarthData (<https://search.earthdata.nasa.gov/search>). SMAP soil moisture data are available through the National Snow and Ice Data Center (NSIDC) (<https://nsidc.org/data/smap>). SIF data from the TROPOMI satellite are available through (<http://www.tropomi.eu/data-products/data-access>). Merged datasets, with the modeled, site-level, GPP and SIF values, and other environmental variables that were used in the analysis at half-hourly, daily, and weekly timescales are provided in **Supplementary Material 3–5**.

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AUTHOR CONTRIBUTIONS

TY performed the data analyses and prepared the figures. TY and GB led the writing of the manuscript. GB and PG conceived the idea of the manuscript. LY processed TROPOMI data. NA, CB, PB, AD, DD, AK, NK, SM, MM, AN, KN, RS, LŠ, KS, MU, and AV contributed to EC data. All authors discussed the results and participated in writing and editing of the manuscript.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/ffgc.2021.695269/full#supplementary-material>

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