



# What Are the Principal Factors Affecting Ambient Ozone Concentrations in Czech Mountain Forests?

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The aim of our study was to identify the factors substantially affecting day-to-day variability in O<sub>3</sub> concentrations in Czech mountain forests and to describe their influence in detailed, quantitative way. We examined the effects of meteorology and ambient NO<sub>x</sub> recorded in regular long-term continuous monitoring at five mountain forest sites representing different regions, covering both polluted and relatively unpolluted areas over the time period of 1992–2018. To investigate the association between ambient O<sub>3</sub> concentrations on one hand, and precursor NO<sub>x</sub> concentrations, and meteorology on the other hand, we used a generalized additive model, GAM, with semiparametric (penalized-spline-based) components to capture properly the possible departures from linearity that is not captured by traditional linear regression approaches. Our results revealed that the O<sub>3</sub> concentrations showed significant associations with all selected explanatory variables, i.e., air temperature, global solar radiation (GLRD), relative humidity, and NO<sub>x</sub>. Apparently, both meteorology and air pollution are highly important for day-to-day O<sub>3</sub> concentrations, and this finding is consistent for all five rural sites, representing middle-elevated forested mountain areas in Central Europe. In addition to individual variables, we were able to detect interactions between three pairs of explanatory variables, namely temperature\*GLRD, temperature\*relative humidity, and GLRD\*relative humidity. Moreover, we confirmed non-linear O<sub>3</sub> behavior toward all individual explanatory variables.

**Keywords:** ambient ozone, generalized additive model, NO<sub>x</sub>, meteorology, non-linear effects

## INTRODUCTION

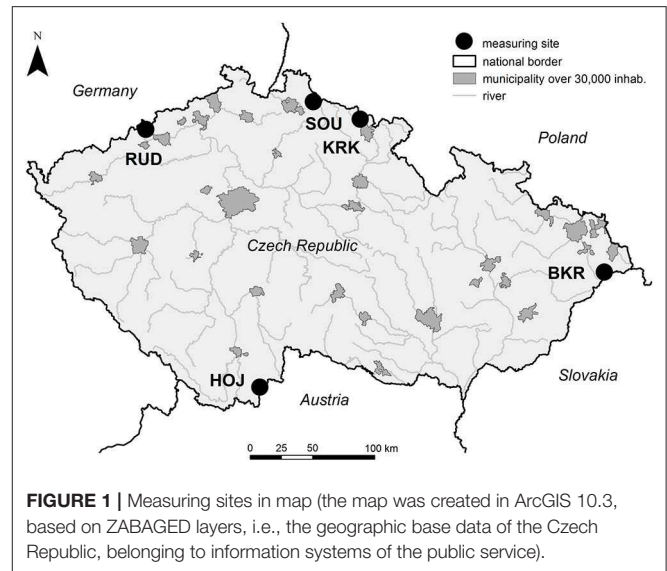
Ground-level ozone (O<sub>3</sub>), an important constituent of the atmosphere (Prinn, 2003; Singh and Fabian, 2003; Monks et al., 2015), belongs among the major factors exerting negative impacts on forests (Ferretti et al., 2015; EEA, 2016), and remains a challenging problem for current and future timber production and the conservation of natural plant communities, including species diversity (Krupa et al., 2001). A range of impacts due to elevated O<sub>3</sub> exposures have been reported by numerous authors, from changes in biochemical processes in living organisms to macroscopic injuries (e.g., Roschina and Roschina, 2003; Cape, 2008; Paoletti et al., 2010, etc.), although the field evidence for the impact of O<sub>3</sub> on forests remains less clear (Manning, 2005; De Vries et al., 2014; Cailleret et al., 2018).

The chemistry of O<sub>3</sub> is very complex (Finlayson-Pitts and Pitts, 1997; Seinfeldt and Pandis, 1998). Ozone is a product of the photochemical reactions of precursors, i.e., nitrogen oxides (NO<sub>x</sub>), volatile organic compounds (VOC), methane (CH<sub>4</sub>), and carbon monoxide (CO). Meteorology is strongly involved in O<sub>3</sub> formation, with atmospheric stability, high atmospheric pressure, high solar radiation, and air temperature as factors promoting O<sub>3</sub> buildup (Kovač-Andrić et al., 2009; Wang et al., 2017; Pyrgou et al., 2018). Many processes—physical, chemical, and biological—affect the formation, transportation, and destruction of O<sub>3</sub> and thus the final O<sub>3</sub> concentrations (Colbeck and Mackenzie, 1994; Seinfeldt and Pandis, 1998; Fowler et al., 2009). Ozone formation in the highly non-linear O<sub>3</sub>-VOC-NO<sub>x</sub> system is not yet fully understood (Sillman, 1999; Carillo-Torres et al., 2017).

Trends in O<sub>3</sub> are difficult to detect due to its large interannual variability (Jonson et al., 2006). The ambient O<sub>3</sub> in Europe has significantly decreased, though not as much as expected with respect to sharp reduction in precursor emissions since the 1990s (Sicard et al., 2013; Paoletti et al., 2014; Colette et al., 2017). It appears that the negative O<sub>3</sub> trend throughout Europe due to European emission controls has been counteracted by tendencies related to climate warming and the hemispheric transportation of pollutants from the source regions, such as Southeast Asia (Yan et al., 2018). Climate variability generally regulates the interannual variability in European O<sub>3</sub>, whereas the changes in anthropogenic precursor emissions predominantly influence O<sub>3</sub> trends (Yan et al., 2018).

In order to better understand O<sub>3</sub> behavior, it is of major concern to identify the factors that account for most of the day-to-day variability in O<sub>3</sub> concentrations. Numerous studies have been published tackling this issue from various perspectives, using different approaches, examining diverse variables and their measures (e.g., Duenas et al., 2002; Tarasova and Karpetchko, 2003; Abdul-Wahab et al., 2005; Özbay et al., 2011). Regression-based approaches—with the most often used multiple linear regression analysis—are commonly used for modeling O<sub>3</sub> concentrations as a response variable, and with meteorological characteristics and different ambient air pollutants as explaining variables (e.g., Abdul-Wahab et al., 2005; Sousa et al., 2006; Pavón-Domínguez et al., 2014).

Ambient ozone levels in Czech Republic are high (Hůnová and Bäumelt, 2018). Exposures over Czech forests exhibit high year-to-year variability (Hůnová et al., 2019), nevertheless they consistently exceed the critical level of 5 ppm h AOT40F (UN/ECE, 2004) since 1994, with peak values reaching 38–39 ppm h at some sites in different years. In some mountain forests, such as in the Jizerske hory Mts., the O<sub>3</sub> exposures are similar as in highly polluted sites in South Europe and higher altitudes (Hůnová et al., 2016). The critical level of 5 ppm h AOT40F is usually exceeded early in the growing season, generally in May (Hůnová and Schreiberová, 2012). The highest O<sub>3</sub> exposures, indicated by AOT40F, are permanently measured in south Czech Republic, the region much less affected by clouds, and thus receiving higher global radiation loads during growing seasons (Hůnová et al., 2019). Existing studies on O<sub>3</sub> biological effects on Czech forests are



equivocal (Šrámek et al., 2007, 2012; Hůnová et al., 2010, 2011; Zapletal et al., 2012; Vlasáková-Matoušková and Hůnová, 2015), however, and despite the high O<sub>3</sub> levels recorded, no serious damage attributable to O<sub>3</sub> has been reported so far.

The aim of our study is to explore selected factors substantially affecting the day-to-day variability of O<sub>3</sub> concentrations in Czech mountain forests. In the effort to decrease O<sub>3</sub> levels, it is truly of the utmost importance to know how meteorology and emission precursors influence ozone in a quantitative way.

## METHODS

### Sites and Period Under Review

We analyzed the observed data, recorded in regular long-term continuous monitoring at five mountain forest sites (**Figure 1**; **Table 1**) operated by the Czech Hydrometeorological Institute (CHMI). The sites under review are distributed unevenly across the territory of the Czech Republic (CR), and are situated in border mountains under different pollution loads: Krkonoše-Rýchory (KRK) and Souš (SOU) are situated in the northern CR; Rudolice v Horách (RUD) in the northwest CR, a region with cumulative major emission sources; Bílý Kříž (BKR) in the northeast CR in the highly polluted Silesia region; and Hojná Voda (HOJ) in the relatively unpolluted southern CR. All five measuring stations are situated in open areas. Nevertheless, in close vicinity of all these sites are extensive forested areas, covered predominantly by spruce forests (*Picea abies*). The KRK, BKR, and HOJ sites are roughly estimated to represent some tens to hundreds kilometers, whereas RUD and SOU sites represent somewhat less extensive regions, about several tens kilometers.

We examine the period of 1992–2018 (excluding days with missing ozone and/ or explanatory variables of our interest). This interval covers both the time with high emissions in the past (until 1998), and the “cleaner” period after profound socio-economic changes in Central Europe, including the CR, after the

**TABLE 1** | Measuring sites - basic characteristics.

Site	Code	Mountain area	Classification	Altitude [m a. s. l.]
Krkonoše-Rýchory	KRK	Krkonoše	B/R/N-REG	1,001
Bílý Kříž	BKR	Beskydy	B/R/N-REG	890
Rudolice v Horách	RUD	Krušné hory	B/R/N-REG	840
Hojná Voda	HOJ	Novohradské hory	B/R/N-REG	818
Souš	SOU	Jizerské hory	B/R/N-REG	771

introduction of novel, more effective legislation for ambient air protection, and after the adoption of diverse countermeasures. From this 27-years record numerous data are missing, in different periods for individual sites. The NO<sub>x</sub> ambient levels at four out of our five sites are available only until 2012. As a matter of fact, NO<sub>x</sub> was monitored at Czech mountain sites only until 2012 (due to very low concentrations recorded in these areas previously), with the exception of BKR where NO<sub>x</sub> monitoring continues. **Table S1** presents an overview of the numbers of daily data at disposal for our analysis.

## Input Data: Ambient Air Quality and Meteorology

The input data were retrieved from ISKO, the Czech nationwide ambient air quality database (CHMI, 2018b). All data we used were based on real-time, continuous measurements, from which 1 h averages were routinely calculated and stored—as the basic primary data—in ISKO database. With regard to quality of the dataset used: (1) the ambient O<sub>3</sub> and NO<sub>x</sub> concentrations were checked thoroughly for gross errors by a database procedure based on mathematical and statistical methods (CHMI, 2018a), whereas (2) meteorology data, considered as support data, were checked only based on logic. The input data for all five sites run by the CHMI were measured by well-established, standardized methods as follows.

### Ambient Ozone

We used the daily mean O<sub>3</sub> concentrations calculated from 1 h measurements, from the continuous, year-round, nationwide ambient air quality network. The measurement method was UV absorbance, the EC reference method (EC, 2008); the O<sub>3</sub> analyzers used were TEI-M49, manufactured by Thermo Environmental Instruments Inc., based in Franklin, Massachusetts, U.S. The sampling equipment was changed in 2015 to TAPI T-400, manufactured by Teledyne Advanced Pollution Instrumentation, Inc., based in San Diego, California, U.S. Standard QA/QC procedures in accordance with EU legislation (EC, 2008) were applied.

### Nitrogen Oxides

We used the daily mean NO<sub>x</sub> concentrations calculated from 1 h measurements, from the continuous, year-round, nationwide ambient air quality network. The measurement method was chemiluminescence, the EC reference method (EC, 2008); the NO<sub>x</sub> analyzers used were TEI-M42, manufactured by Thermo Environmental Instruments Inc., based in Franklin, Massachusetts, U.S. The sampling equipment was changed in

2015 to TAPI T-200, manufactured by Teledyne Advanced Pollution Instrumentation, Inc., based in San Diego, California, U.S. Standard QA/QC procedures in accordance with EU legislation (EC, 2008) were applied.

### Meteorology

Meteorology monitored continuously with a 1 h time resolution was used. Namely, we used the daily mean air temperature, daily mean relative humidity, and daily global solar radiation (GLRD). The temperature in this paper is given in Kelvins [K]—this unit is used in our nation-wide database to avoid negative numbers, as negative numbers are exclusively assigned to error codes. Ambient air temperature in 2 m above ground and relative humidity were measured by Thies Clima HTT Compact, manufactured by Adolf Thies GmbH & Co., based in Gottingen, Germany. GLRD was measured by CMP sensors, manufactured by Kipp & Zonen B.V., based in Delft, The Netherlands. The meteorological variables were measured at the same sites and at the same hourly intervals as the air pollution variables.

### Statistical Analysis

We have investigated the association between ambient O<sub>3</sub> concentrations on one hand, and precursor NO<sub>x</sub> concentrations and meteorology on the other hand, using a generalized additive model, GAM (Wood, 2006). We accounted for NO<sub>x</sub> as a proxy for ambient air pollution and a key player in O<sub>3</sub> chemistry. Unfortunately, VOCs as another key precursor group for O<sub>3</sub> formation could not be accounted for, as their concentrations are not recorded at the sites under review.

O<sub>3</sub> concentration was considered the dependent variable in the GAM model, whereas global solar radiation, air temperature, relative humidity, and NO<sub>x</sub> concentrations were considered as explanatory variables. We worked with daily mean values calculated from primary 1 h values stored in a nationwide air-quality database. As the effects of some of these variables are known not to be linear, we used flexible GAM model formulation with semiparametric (penalized-spline-based) components to capture departures from linearity instead of forcing the relationship into the traditional but unrealistic linear (or log-linear) framework (Crawley, 2005). On the other hand, the GAM-retrieved relationships essentially reduce to linearity when the data do not support a non-linear relationship.

In order to maintain the comparability of results, the functional form of the GAM model is kept the same for all the sites we consider. Fitting was done separately for different sites (we stratified on site) so that we have site-specific parameters of the model with the same interpretation. For a given site, the GAM model is as follows:

$$Y_t = \beta_0 + s_T(\text{Temperature}_t) + s_G(\text{GLRD}_t) + s_H(\text{Humidity}_t) + s_{\text{NOX}}(\text{NOX}_t) + s_{\text{TG}}(\text{Temperature}_t, \text{GLRD}_t) + s_{\text{TH}}(\text{Temperature}_t, \text{Humidity}_t) + s_{\text{GH}}(\text{GLRD}_t, \text{Humidity}_t) + \varepsilon_t$$

where:

- $Y_t$  is the daily mean ozone concentration for day  $t$  (time is indexed by the number of days since the first day of the data),

- $\beta$  is an unknown constant (intercept) to be estimated from the data
- $Temperature_t$  is the mean temperature for day  $t$ ,
- $GLRD_t$  is the mean GLRD for day  $t$ ,
- $Humidity_t$  is the mean relative humidity for day  $t$ ,
- $NOX_t$  is the mean  $NO_x$  concentration for day  $t$ ,
- $\varepsilon_t$  is the random error term. We adopt a working assumption of  $\varepsilon_t \sim N(0, \sigma^2)$ , which is a normal, zero mean, homoscedastic distribution of errors,
- $S_T$  is an unknown univariate function (of temperature as an argument), of which the form is to be estimated from the data. We estimate it as a penalized spline (regularizing wiggleness via the penalization of the integral of the squared second derivative with respect to temperature). By that, we allow for potentially non-linear, smooth shapes of (marginal) dependence of  $O_3$  on temperature,
- $S_G$  is an unknown univariate function (of GLRD) to be estimated from the data as a penalized spline,
- $S_H$  is an unknown univariate function (of Humidity) to be estimated from the data as a penalized spline,
- $S_{NOX}$  is an unknown univariate function (of  $NO_x$ ) to be estimated from the data as a penalized spline,
- $S_{TG}$  is an unknown bivariate function (of Temperature and GLRD) to be estimated from the data as a tensor product penalized spline. This term allows us to investigate the possible interaction of Temperature and GLRD in influencing ozone concentration (i.e., departures from the simple additive effects of Temperature and GLRD). In other words, this effect enables us to investigate how the effect of Temperature is modified by GLRD (or, equivalently, how the effect of GLRD is modified by Temperature),
- $S_{TH}$  is an unknown bivariate function (of Temperature and Humidity) to be estimated from the data as a tensor product penalized spline. This term corresponds to the Temperature\*Humidity (parsimoniously formulated) interaction,
- $S_{GH}$  is an unknown bivariate function (of GLRD and Humidity) to be estimated from the data as a tensor product penalized spline. This term corresponds to the GLRD\*Humidity (parsimoniously formulated) interaction.

All unknown model components are estimated simultaneously (the model is identified) by optimizing the penalized likelihood. Unknown penalty coefficients are obtained via generalized cross-validation. All modeling computations were done in R core (2008). Results with  $p \leq 0.05$  are considered statistically significant. In graphs, the estimated spline functions are supplemented with 95% confidence intervals (constructed in pointwise fashion, so that they do not claim simultaneous coverage).

When estimating the model, we use all days with available data for a given station. From the model-fitting point of view, data are available when  $O_3$  and all explanatory variables (temperature, GLRD, humidity,  $NO_x$ ) are available. When at least one of these variables is missing, the day effectively appears as missing (i.e., it is not used in the fitting). Since different stations had missing days at different time-points, the model was estimated on a somehow

**TABLE 2 |** Descriptive statistics for included variables at all five sites (calculated from daily data).

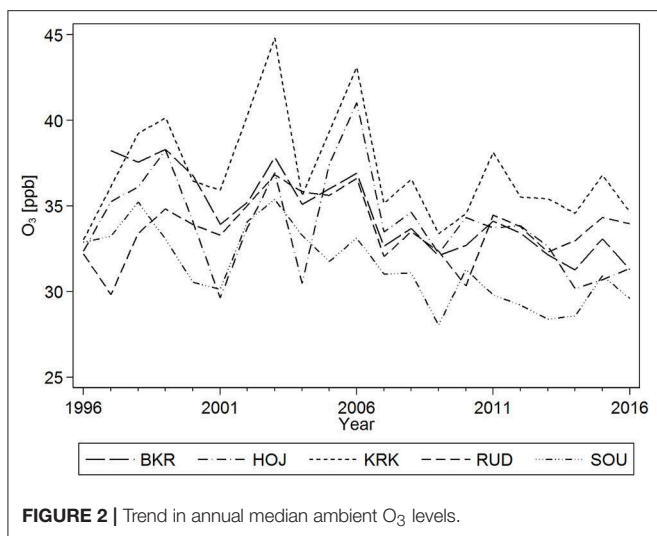
Site	Variable	N	25th perc	50th perc	75th perc
KRK	$O_3$ [ppb]	7,737	28.61	36.98	46.26
	$NO_x$ [ppb]	6,368	2.98	4.41	6.77
	H [%]	6,764	72.17	84.15	92.20
	Temp [K]	6,708	272.10	279.26	286.33
	GLRD [ $W\ m^{-2}$ ]	6,277	39.70	107.74	201.42
BKR	$O_3$ [ppb]	8,705	27.46	35.08	44.61
	$NO_x$ [ppb]	8,735	2.19	3.11	4.60
	H [%]	7,037	77.13	91.00	98.00
	Temp [K]	7,098	272.88	279.77	285.89
	GLRD [ $W\ m^{-2}$ ]	6,188	31.08	95.44	182.90
RUD	$O_3$ [ppb]	8,115	25.98	33.74	42.39
	$NO_x$ [ppb]	6,118	4.52	6.74	10.40
	H [%]	7,435	78.08	90.21	98.42
	Temp [K]	7,645	273.24	279.43	285.21
	GLRD [ $W\ m^{-2}$ ]	7,187	50.27	119.50	213.00
HOJ	$O_3$ [ppb]	8,409	26.32	33.45	41.96
	$NO_x$ [ppb]	6,405	2.43	3.34	4.94
	H [%]	7,847	71.13	82.96	91.29
	Temp [K]	8,079	273.69	280.20	286.21
	GLRD [ $W\ m^{-2}$ ]	7,752	48.80	106.35	193.10
SOU	$O_3$ [ppb]	8,945	24.65	31.94	40.03
	$NO_x$ [ppb]	7,030	3.29	4.79	7.14
	H [%]	7,454	82.17	90.88	95.96
	Temp [K]	7,673	273.64	280.13	286.89
	GLRD [ $W\ m^{-2}$ ]	7,987	31.69	92.69	183.54

different set of times, assuming that the ozone-to-explanatory-variables relationship is homogeneous in time.

## RESULTS

**Table 2** presents an overview of descriptive statistics calculated from the mean daily values for all considered variables at all five measuring sites. **Figures 2, 3** present the dynamics in the response variable (given as annual median values of respective variable), i.e., ambient  $O_3$  concentrations, and explanatory variables: air temperature, GLRD, relative humidity, and ambient  $NO_x$  concentrations at all five sites under review. Though we used data over the 1992–2018 time period in our model, the dynamics in the variables (**Figures 2, 3**) are shown only for 1996–2016 due to the fact that at the beginning of measuring period, many data are missing, which would misrepresent the annual median in graphs. The GAM model, however, can cope with the missing data automatically (assuming homogeneity of the  $O_3$  relationship to the explanatory variables in time and missing at random, or the MAR mechanism (Little and Rubin, 2002). The year-to-year variability in  $O_3$  is high; the median concentrations at individual sites range between 28 and 45 ppb with clear peaks at some sites in 2003 and 2006. In spite of a clear decrease in  $NO_x$  concentrations, the ambient  $O_3$  levels apparently do not decrease accordingly. Annual median temperatures ranged between 277.5





and 282 K, GLRD between 74 and 200 W m<sup>-2</sup>, and relative humidity between 74 and 98%.

The day-to-day variability in ambient O<sub>3</sub> concentrations is well described by a model including the additive effects of air temperature, GLRD, relative humidity, and ambient NO<sub>x</sub> concentration, and the interactions of air temperature\*GLRD, air temperature\*relative humidity, and GLRD\*relative humidity (Figures 4–10). Though Figures 4–10 show the smooth terms with 95% confidence intervals from GAM model only for one site, namely BKR, the models for other four sites are looking quite similar. The relationship between O<sub>3</sub> expected value and temperature (Figure 4) shows higher O<sub>3</sub> values not only with higher temperatures, but, surprisingly, also with lower temperatures below 273.15 K (i.e., 0°C). Likely explanation for this finding rests in the fact that, mountain stations the data from which we analyze, are frequently above the inversion cloud layer in winter time, and receive high solar radiation capable of splitting the NO<sub>2</sub>, a precursor molecule to O<sub>3</sub>. Consequently, O<sub>3</sub> is recorded also in conditions of simultaneous low temperatures and high solar radiation, though O<sub>3</sub> levels are not that high as in summer when much higher temperatures promote O<sub>3</sub> formation reactions. The expected ozone value increases with increasing GLRD, though at some point the curve shows a plateau or even a slight decline (the decline is not significant, however, since it is outweighed by the increased variability) for higher GLRD values (Figure 5). Ozone concentration increases with decreasing relative humidity (Figure 6), right in line with physically-based intuition. The expected ozone value reaches its peak much more quickly than it decreases and shows a local maximum at about ambient NO<sub>x</sub> concentration 3 ppb (Figure 7). Apart from the significant relationship between the response variable and individual explanatory variables, we also detected significant interactions between the three pairs of explanatory variables. Figure 8 shows the complex effect of interactions of air temperature\*GLRD upon the expected ambient O<sub>3</sub> concentration. Ozone evidently increases not only with increasing temperature and GLRD, which is a well-known

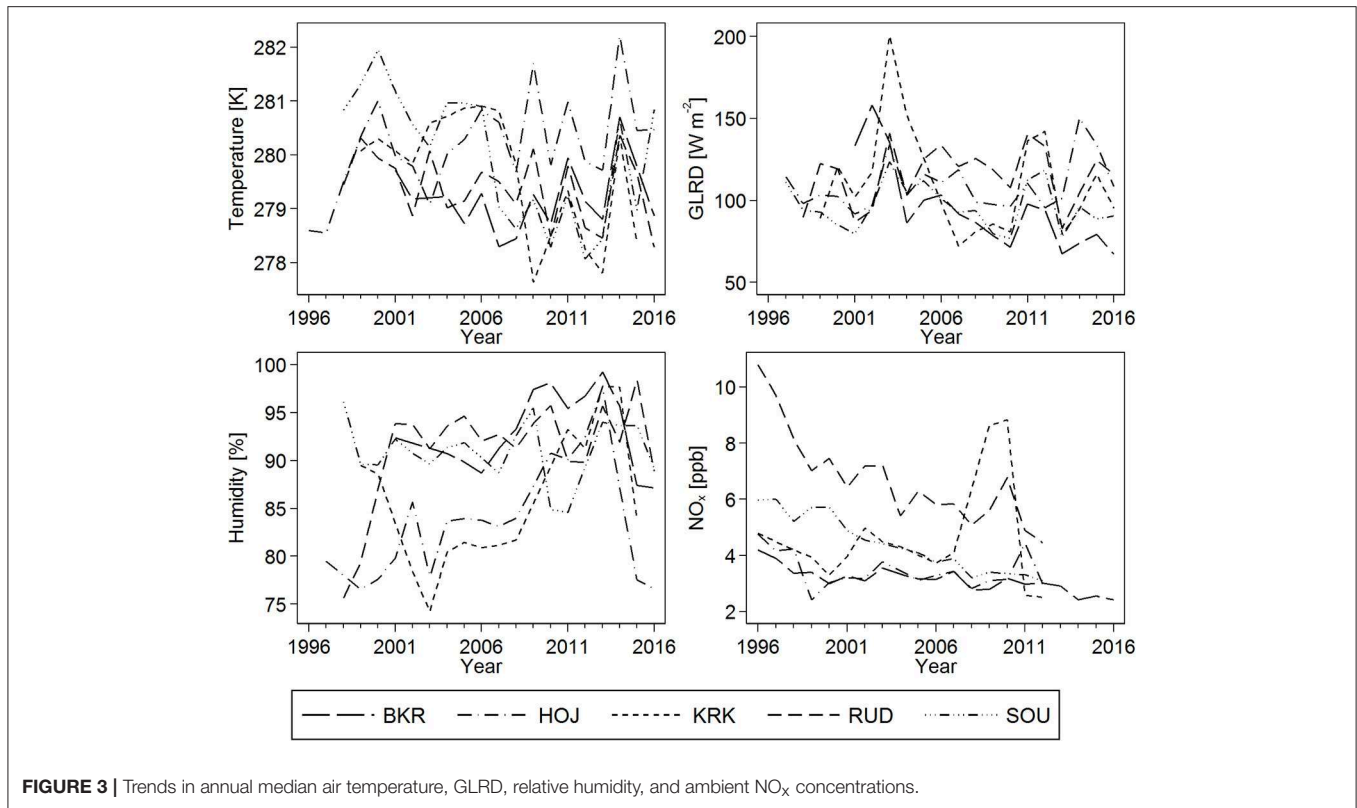
fact, but also, surprisingly, with increasing GLRD at low temperatures. The dividing line is apparently around the ambient air temperature 280 K (i.e., 7°C). The ambient O<sub>3</sub> concentration generally increases with increasing temperature and decreasing humidity (Figure 9), whereas effect of interactions of relative humidity\*GLRD upon expected ambient O<sub>3</sub> concentration appears less straightforward (Figure 10). We can see a sharp increase of O<sub>3</sub> concentrations with decreasing relative air humidity at a GLRD around 250 W m<sup>-2</sup>, whereas above that GLRD value, O<sub>3</sub> decreases.

The marginal effects of all selected explanatory variables are significant, but they differ in the strength of their effects upon ozone. All variables and their selected interactions are highly significant ( $p < 0.001$ ) for all five sites, with a minor exception for the interaction of GLRD\*relative humidity for the KRK site, which is not significant, and for the BKR site with the  $p$ -value slightly less than the significance level of 0.05. The approximate significance of smooth terms is summarized in Table 3. At the KRK site, situated in the unpolluted top of the Krkonoše Mountains, influenced only by the long-range transport of air pollutants, all explanatory variables show similar strength, which also applies for interactions, which show a similar though lesser strength, as compared to the individual explanatory variables. A very similar pattern, though with a somewhat higher strength of relative humidity and lesser strength of air temperature, is obvious for the HOJ site, situated in the Novohradské hory Mountains, in the unpolluted (with regard to primary emissions and anthropogenic precursors of ambient ozone) south of the Czech Republic at the Austrian border. Ambient NO<sub>x</sub> concentrations have a greater effect at sites representing more polluted regions. This holds true for the RUD site, situated at the Krusne hory mountain plateau and influenced by nearby large emission sources down in the valley (Bridges et al., 2002); for the SOU site, influenced by car exhaust from the nearby local road; and for the BKR site, influenced by emission sources in the polluted Silesian region (both on the Czech and the Polish sides of the border). The deviance explained for the individual sites is high, and ranges between 74.3% for the BKR and 63.4% for the SOU sites.

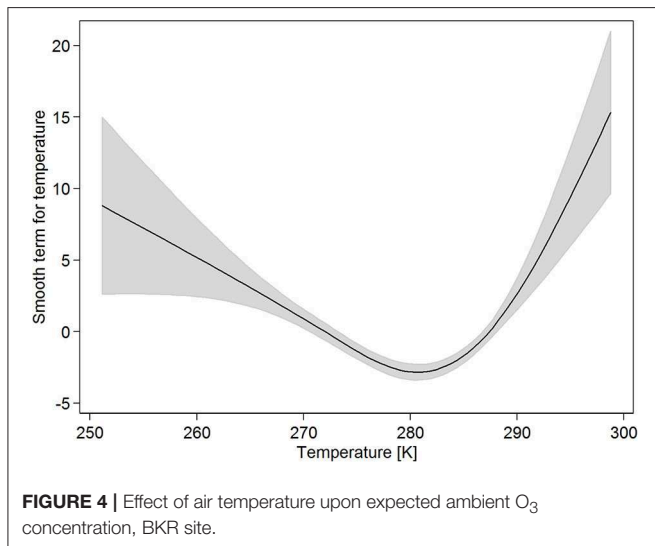
## DISCUSSION

### Complex Atmospheric Chemistry of O<sub>3</sub>

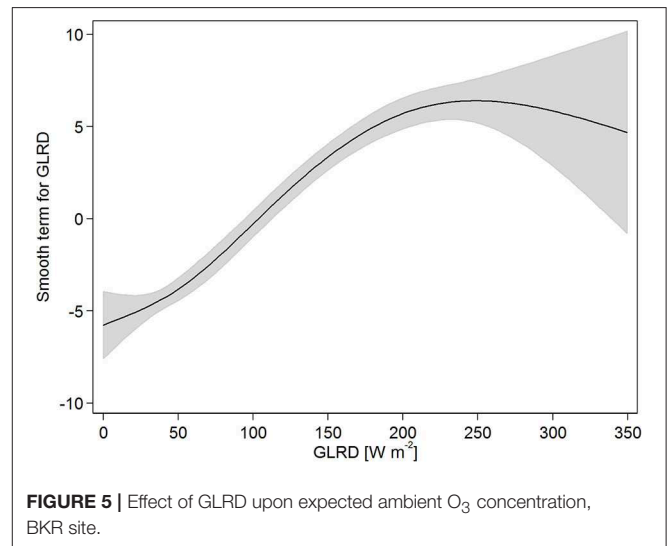
Numerous studies examining O<sub>3</sub> chemistry, meteorology, and precursor emissions in Europe and North America, in different field research programs conducted under diverse geographical and climatic conditions, as well as studies of the combination of ground-based ozone data and meteorological observations, have resulted in enhancement of our knowledge of photochemical processes under various tropospheric conditions (Solomon et al., 2000; Monks et al., 2015). Nevertheless, it remains a challenge to interpret ambient O<sub>3</sub> behavior, levels, and trends. Atmospheric chemistry resulting in O<sub>3</sub> formation and destruction is, as a matter of fact, extremely complex due to numerous factors influencing these processes, and hence, O<sub>3</sub> concentrations (Cape, 2008).



**FIGURE 3 |** Trends in annual median air temperature, GLRD, relative humidity, and ambient  $NO_x$  concentrations.



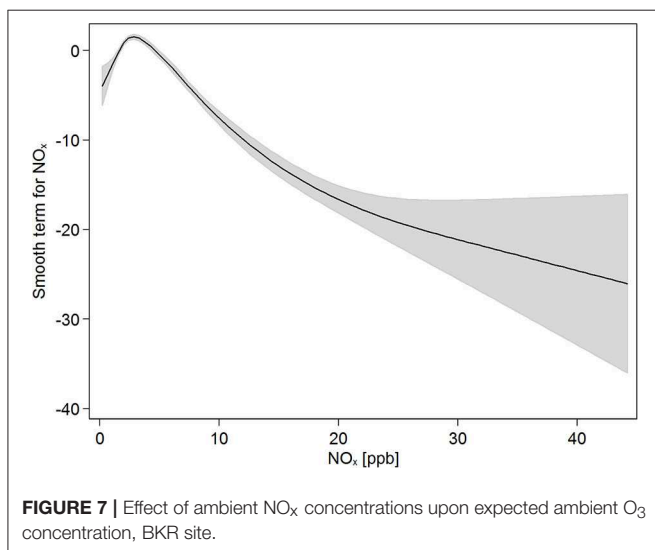
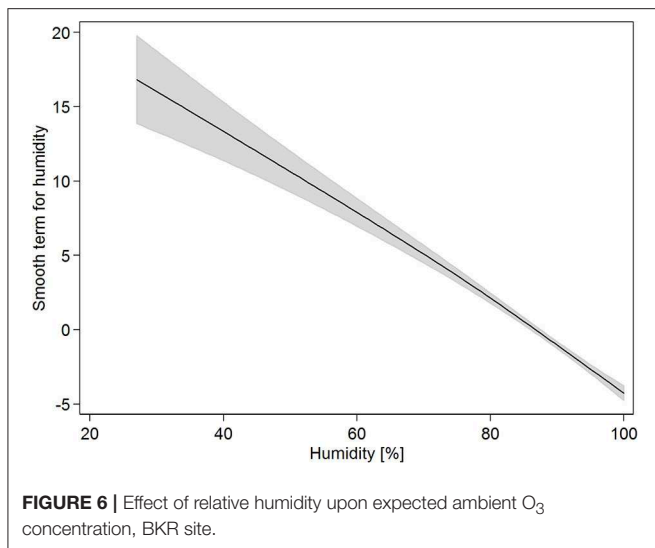
**FIGURE 4 |** Effect of air temperature upon expected ambient  $O_3$  concentration, BKR site.



**FIGURE 5 |** Effect of GLRD upon expected ambient  $O_3$  concentration, BKR site.

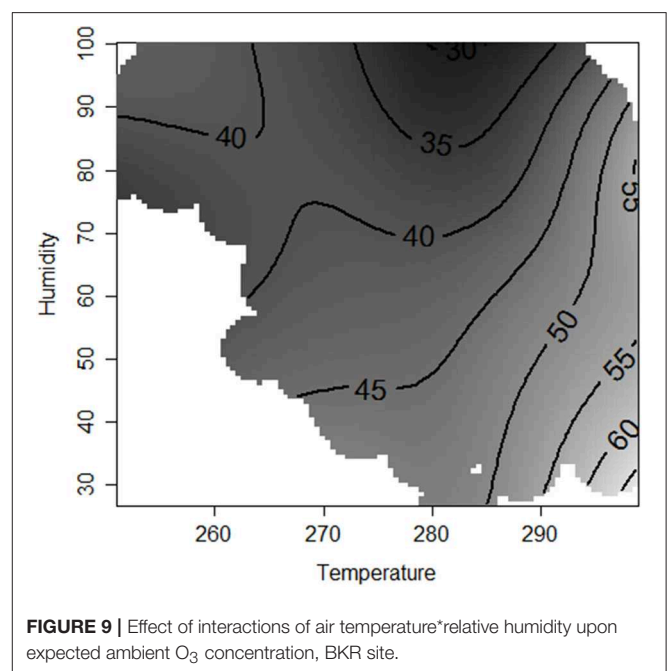
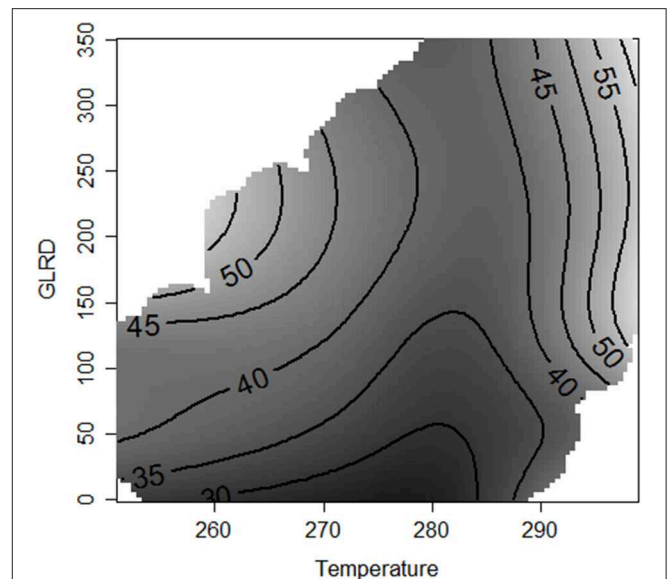
Meteorology is extremely important for  $O_3$  formation. It affects  $O_3$  concentrations not only directly (via horizontal advection, vertical diffusion, and photolysis rates), but also indirectly, by influencing the concentrations of its precursors and chemistry of its formation and destruction (Oikonomakis et al., 2018).  $O_3$  has a temperature-dependent chemistry (Pusede et al., 2015), and temperature was used in some studies as a surrogate to indicate  $O_3$  formation via the  $O_3$ -temperature association, both at individual measuring sites

and on a greater regional scale (Oikonomakis et al., 2018). The  $O_3$ -temperature relationship originates in: (1) temperature-dependent biogenic VOC emissions, (2) thermal decomposition of PAN to  $HO_x$  and  $NO_x$ , (3) increased likelihood of favorable meteorological conditions for ozone formation (Abeleira and Farmer, 2017). However, there are major uncertainties in the mechanisms underlying the temperature-dependent changes in  $O_3$  concentrations, their interactions, and relative contributions in rural and remote regions (Romer et al., 2018). Ozone chemistry



regimes are shifting as precursor emissions are changing (Abeleira and Farmer, 2017). Moreover, meteorology-dependent O<sub>3</sub> chemistry implicates the impact of ongoing climate change on O<sub>3</sub>. Increasing temperature is expected to increase O<sub>3</sub> concentrations (Abeleira and Farmer, 2017). Additionally, the projected rise in global precursor emissions over the twenty-first century is also assumed to have a strong effect on O<sub>3</sub> throughout the world (Vingarzan, 2004). In contrast to temperature, field studies on GLRD influence on ground-level O<sub>3</sub> concentrations are scarce (Duenas et al., 2002). Though it is a well-known fact, that ambient O<sub>3</sub> is strongly dependent on GLRD (Finlayson-Pitts and Pitts, 1997; Seinfeldt and Pandis, 1998), some authors showed that solar radiation had a lower effect than expected upon O<sub>3</sub> concentrations as compared to more important effect of temperature (Abdul-Wahab and Al-Alawi, 2002).

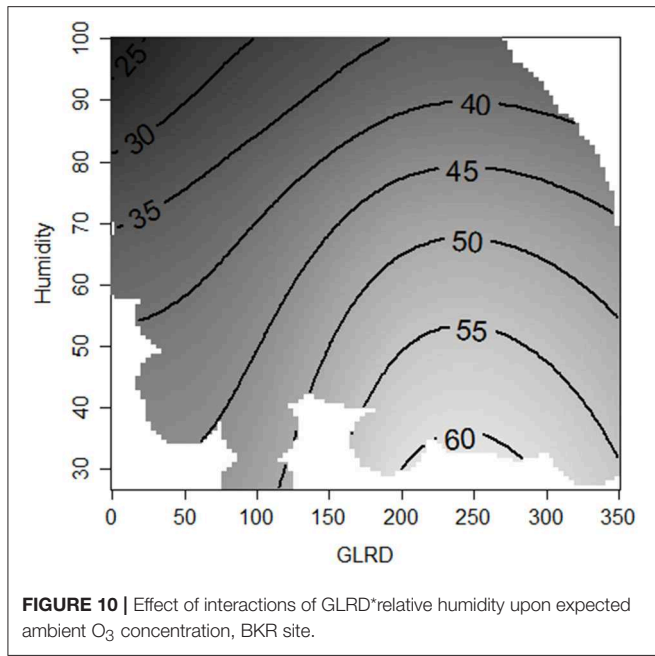
The association between meteorological factors, precursor emissions, and O<sub>3</sub> daily variability might differ in urban and rural areas. Whereas the precursor emissions used to be much



higher in urban areas, in rural regions the precursors are lower, in particular with respect to NO<sub>x</sub> (Romer et al., 2018). Most papers on this issue examine the O<sub>3</sub> regime in urban areas (e.g., Duenas et al., 2002; Tan et al., 2018), whereas the papers on rural regions are less frequent (e.g., Pudasainee et al., 2006).

### Strengths and Weaknesses of the GAM Model Used

Various statistical approaches to model O<sub>3</sub> dependence on meteorology and ambient air pollutants have been employed recently and have been reviewed by Thompson et al. (2001) and



Schlink et al. (2003). The main objectives are O<sub>3</sub> forecasting, estimating O<sub>3</sub> time trends, and investigating the underlying processes based on observation data. The input data may vary widely both in terms of the variables considered, and in terms of observation scales in space and time (Thompson et al., 2001). The multiple linear regression method is frequently used (e.g., Abdul-Wahab et al., 2005; Sousa et al., 2006; Pavón-Domínguez et al., 2014), though it encounters serious difficulties when the independent variables are correlated with each other, i.e., when they exhibit multicollinearity (Al-Alawi et al., 2008). This is exactly the case of O<sub>3</sub>, when e.g., the explanatory variables of temperature and GLRD are correlated. One method used to remove this multicollinearity is principal component analysis (PCA), selecting variables with high loadings to be used as predictors in a regression equation (e.g., Rajab et al., 2013). In our data (and moderate number of explanatory variables), collinearity is not a problem. The variance inflation factor, VIF (Rawlings et al., 1998), is smaller than 2.5 at all five sites. Even more importantly, O<sub>3</sub> formation and destruction are complex non-linear processes, and principal component regression cannot adequately model the non-linear relationship (Al-Alawi et al., 2008), and hence it cannot be used to retrieve true functional relationships, only linear (often hard-to-interpret) approximation. Neural networks and generalized additive models are recommended to address this issue, as they can handle non-linear associations and can be adapted easily to site-specific conditions (Schlink et al., 2003), but since they are used in a black-box style, they are much more useful for prediction than for analysis purposes.

Since we used a rather flexible class of statistical models (Generalized Additive Model with penalized spline components), we allow for a smooth relationship between O<sub>3</sub> and explanatory variables (temperature, GLRD, humidity, and NO<sub>x</sub>). This class

**TABLE 3 |** Approximate significance of smooth terms.

	KRK			BKR			RUD			HOJ			SOU		
	F	EDF	p-value	F	EDF	p-value	F	EDF	p-value	F	EDF	p-value	F	EDF	p-value
T	20.72	3.661	<0.001	30.07	3.638	<0.001	21.23	3.463	<0.001	15.73	3.999	<0.001	35.32	3.859	<0.001
GLRD	23.98	2.497	<0.001	84.85	3.691	<0.001	65.10	3.512	<0.001	53.24	3.796	<0.001	76.40	4.000	<0.001
I(T, GLRD)	5.47	12.451	<0.001	7.13	11.374	<0.001	2.56	7.811	<0.001	4.56	12.193	<0.001	10.286	8.269	<0.001
NO <sub>x</sub>	32.44	4.89	<0.001	112.92	6.116	<0.001	195.53	5.190	<0.001	43.67	6.187	<0.001	110.13	7.029	<0.001
H	39.40	3.75	<0.001	150.63	1.738	<0.001	50.19	3.967	<0.001	64.09	1.879	<0.001	7.09	3.807	<0.001
I(T, H)	5.42	14.711	<0.001	8.14	8.748	<0.001	8.32	9.852	<0.001	7.54	7.786	<0.001	4.41	11.092	<0.001
I(GLRD, H)	2.21	3.859	0.054	2.53	4.150	<0.05	4.00	8.904	<0.001	14.86	10.904	<0.001	8.03	12.779	<0.001
Deviance explained			66.5%			74.3%			71.2%			66.3%			63.4%

F, test statistics; EDF, equivalent degrees of freedom; I, interactions; T, temperature; GLRD, global solar radiation; H, relative humidity.



would not be ideal for relationships showing very sharp bends, sudden jumps at particular values of covariates, etc. From the physical and chemical backgrounds of O<sub>3</sub> behavior, however, we do not expect this sort of behavior and take GAM as a tool allowing for non-linearity in explanatory variables (unlike standard multiple regression, which insists on linearity, i.e., on the gradient with respect to a given explanatory variable being constant throughout its range). Since we use tensor product splines, we can also assess possible interactions in influences of certain pairs of explanatory variables (i.e., to assess whether their influence is simply additive or more complicated). Since we concentrate only on two-variable interactions, we can assess not only their overall significance (via the *p*-values of formalized hypothesis tests), but also display the interactions in graphical form (as contour plots showing how the mean ozone concentration depends on a pair of explanatory variables).

### Volatile Organic Compounds Are Lacking in Our Model

The weak point of our model is the fact that the VOCs, an important O<sub>3</sub> precursor group, are not included in our model. This is due simply to the fact that the measured data are not available. There are no significant anthropogenic sources of VOCs near the five mountain sites under review; nevertheless, the surrounding forests are undoubtedly a major source of natural biogenic VOCs (further, BVOCs). BVOCs are a large group of organic chemicals including isoprene, terpenes, fatty acid derivatives, benzenoids, phenylpropanoids, and amino acid derived metabolites (Esposito et al., 2016). They are emitted into the atmosphere by plants, and significantly affect its chemical composition and physical properties (Laothawornkitkul et al., 2009; Pinto et al., 2010), contributing substantially, among other effects, to the oxidizing capacity of the atmosphere through the recycling of hydroxyl radicals, OH (Lelieveld et al., 2008). Interestingly, on a global scale, BVOC emissions exceed by far those emitted due to anthropogenic activities (Seinfeldt and Pandis, 1998), and are more reactive (Holzinger et al., 2005). BVOCs belong among the key players in ambient O<sub>3</sub> chemistry in rural areas (Atkinson and Arey, 2003; Di Carlo et al., 2004).

We assume that natural BVOCs at the investigated sites are very important, as all five sites under review are situated in the close vicinity of vast forested areas, with a clear predominance of Norway spruce (*Picea abies*). It has been widely accepted that the vegetative emissions display large inter-species and inter-individual variability (Aydin et al., 2014). As compared to some other woody species, such as poplar or beech, spruce is a lower BVOC emitter (Bortsoukidis et al., 2014); however, it is still known to emit considerable amounts of reactive trace gases, and is considered particularly as an emitter of monoterpenes, such as  $\alpha$ -pinene,  $\beta$ -pinene, limonene, and myrcene, whereas it is only a low isoprene emitter (Esposito et al., 2016; Jurán et al., 2017). Though some indicative information exists on BVOC emission for the CR (Zemankova and Brechler, 2010; Jurán et al., 2017), observation-based data in high time-resolution suitable for use in our model are lacking at present. Nevertheless, it would be

very interesting in the future to include the data on BVOC concentration to observe how they influence our model.

### Relevance of Our Results to Forests

On a long-term basis, ambient O<sub>3</sub> has been considered as a stress factor contributing to impairment of forest health status and its influence has been studied from different aspects. In spite of clear evidence of O<sub>3</sub> harmful effects observed in laboratory experiments, fumigation chambers, or FO<sub>3</sub>X (Free air O<sub>3</sub> eXposure) experiments (Sandermann et al., 1997; Paoletti et al., 2017; Franz et al., 2018; Hoshika et al., 2018), the field evidence for impacts of O<sub>3</sub> exposure on tree growth is not that clear (De Vries et al., 2014; Cailleret et al., 2018). Moreover, observations in real stand conditions from numerous regions show that measured high O<sub>3</sub> exposures or modeled high O<sub>3</sub> stomatal flux do not correspond with unclear impacts on forest ecosystems (e.g., Ferretti et al., 2007; Matyssek et al., 2007; Waldner et al., 2007; Baumgarten et al., 2009).

European-wide assessment shows that, ambient O<sub>3</sub> exposures recorded at Czech background rural sites are medium in European context, being lower than in Southern Europe including Mediterranean, but higher than in Scandinavia, similar as in Germany and Switzerland (e.g., Waldner et al., 2007; Baumgarten et al., 2009; EEA, 2016). In our earlier work we assessed long-term time trends and spatial variability in ambient O<sub>3</sub> concentrations throughout the Czech Republic (Hůnová and Bäumelt, 2018), O<sub>3</sub> exposure over Czech forested areas (Hůnová and Schreiberová, 2012), compared different GIS methods to create a reliable O<sub>3</sub> and AOT40F map based on observed data (Hůnová et al., 2012), studied O<sub>3</sub> exposure, stomatal flux, and ozone injury on native vegetation in mountain forests (Hůnová et al., 2011; Vlasáková-Matoušková and Hůnová, 2015; Hůnová, 2017), and indicated the Czech forest regions which are stressed by permanent high O<sub>3</sub> exposures (Hůnová et al., 2019).

In present study we have investigated exclusively the O<sub>3</sub> behavior with respect to its day-to-day response to selected meteorology characteristics and NO<sub>x</sub> precursors. Our results confirmed non-linear behavior of ambient O<sub>3</sub> not only to NO<sub>x</sub> precursors, which is a well-known fact (Finlayson-Pitts and Pitts, 1997; Seinfeldt and Pandis, 1998), but also to meteorology, including all variables considered. Our results thus point out to certain deficiency of models investigating the underlying processes of O<sub>3</sub> formation based on measured data not considering this non-linear behavior toward ambient air temperature, GLRD, and relative humidity. This might be beneficial, for example, in modeling future scenarios for forested regions, accounting for changes in ambient O<sub>3</sub> toward local ambient air temperature, GLRD and relative humidity. Furthermore, our results indicated the necessity of considering not only individual explaining variables but also their interactions, such as the interactions of air temperature\*GLRD, air temperature\*relative humidity, and GLRD\*relative humidity. This might be of high relevance in particular in context of global climate change (Bytnerowicz et al., 2007; Sicard et al., 2017). In this respect Czech Republic heads toward hotter and drier conditions (Hlásny et al., 2011; Trnka et al., 2015; Štěpánek et al., 2016), which is likely to affect future ambient O<sub>3</sub> concentrations.

The associations indicated by our study might enhance the models estimating potential future impacts of environmental/climate changes on forests.

Ambient O<sub>3</sub> is, alongside with nitrogen deposition, considered on a long-term basis as an important factor affecting negatively European forests (EEA, 2016). Nevertheless, in recent years, this factor has been in Central Europe and elsewhere overshadowed by two other major threats for forests, namely unprecedented drought and unprecedented bark beetle outbreaks (Allen et al., 2010). Forests with compromised ecological stability, such as traditional forest monocultures in particular, are especially prone to large scale damage (Hlásný and Turčáni, 2013). This holds for most of the Czech forests, which can be classified as managed low-diversity ecosystems (Vacek et al., 2013), with decreased stand stability in respect to damages caused by natural abiotic agents (wind, snow, glaze, drought), biotic factors (insects including bark-beetle, pathogens), and also by anthropogenic effects, such as ambient air pollution. This adverse development is manifested by recent unprecedented salvage felling in Czech forests (Pajtk et al., 2018; Zahradník and Zahradníková, 2019). That present Czech forest decline is a result of a complex interplay of multiple factors was stressed by several recent observation-based studies (Kolár et al., 2015; Cienciala et al., 2016; Altman et al., 2017).

Czech forests in history were substantially changing in terms of area, species representation, and tree age. Presently, out of 26,664 km<sup>2</sup> of forested area representing some 34% of the Czech territory, Norway spruce (*Picea abies*), traditionally an important timber tree, is still a dominant species covering ca. 51% of the total forested area, following pine trees (*Pinus* spp.) cover 17%, beech (*Fagus sylvatica*) 8%, and oak (*Quercus* spp.) 7% (Ministry of Agriculture, 2015). Spruce is a somewhat ambiguous species with respect to O<sub>3</sub> sensitivity—it is assumed not to be especially sensitive in relation to visible injury but regarded as O<sub>3</sub>-sensitive in relation to growth responses (UN/ECE, 2004). Nevertheless, we have to keep on mind that ambient air pollution is a factor affecting forests. We have already witnessed earlier that, predisposition of Norway spruce by air pollution (namely by high ambient SO<sub>2</sub> concentrations in the 1970s and 1980s) to stress from climatic events, such as drought and temperature changes, resulted in forest decline in Czech Republic (Vacek et al., 2015).

## CONCLUSIONS

The generalized additive model confirmed that selected explanatory variables, i.e., air temperature, GLRD, relative humidity, and NO<sub>x</sub>, significantly affect daily O<sub>3</sub> concentrations in Czech forests. Apparently, both meteorology and air pollution are highly influential on day-to-day O<sub>3</sub> concentrations, and this finding is consistent for all five rural sites representing

middle-elevated forested mountain areas in Central Europe. We believe it would be of the utmost importance to include BVOCs in our analysis. Nevertheless, at present this is impossible, as the relevant data for BVOCs for these sites do not exist. Additionally, in contrast to a standard approach based on multiple linear regression used for similar studies, we were able to access non-linear relationships of various covariates to ozone, and also to characterize the interactions of the three pairs of explanatory variables, namely temperature\*GLRD, temperature\*relative humidity, and GLRD\*relative humidity. We extracted functional relationship of various explanatory variables, demonstrating deficiencies of standardly used linear approximations. Local extrema visible in some of the bivariate relationships (O<sub>3</sub> to temperature or to NO<sub>x</sub>) can, in addition to saturation (upper asymptote) visible in the relationship of O<sub>3</sub> to GLRD, easily distort quantification of covariate effects, e.g., to forests injury and/ or to smear out various climatic scenario comparisons based on standard linear approach. Non-linear relationships (qualitatively clear from the O<sub>3</sub> atmospheric chemistry) should be taken seriously also in real, observed data analysis. GAM methodology offers a powerful tool for this.

## AUTHOR CONTRIBUTIONS

IH coordinated the study. IH, MB, and MM contributed to the conception and design of the study. MM organized the database and prepared the final version of **Figures 2–10**. MB performed the statistical analysis. IH wrote the first draft of the manuscript. All authors contributed to manuscript revision, read, and approved the submitted version.

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## SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/ffgc.2019.00031/full#supplementary-material>

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