



Effects of Climate Change on Hydrological Indicators of Subsurface Drainage for a Representative French Drainage Site

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The aim of this study is to evaluate from a hydrological perspective and in the context of climate change the future of subsurface drainage of the *La Jaillière* site (western France), which is representative of the pedology of the majority of French subsurface drainage. We used a uniquely large and comprehensive range of 17 hydrological indicators (HIs), describing the temporal dynamics of drainage season, soil saturation, drained water balance and flood events. The HI values are calculated from simulated discharges provided by a subsurface drainage model, the SIDRA-RU model, fed by 12 climate projections from 1975 to 2100 (CMIP5 Euro-Cordex project), with three climate change scenarios: Representative Concentration Pathways (RCP) 2.6, RCP4.5 and RCP8.5. We first verified that the HIs simulated using climate projections in the SIDRA-RU model over the historical period were not critically biased compared to the HIs obtained from the reference climatic reanalysis (SAFRAN). Second, we analysed and compared the HI evolution over different periods and under different scenarios. Our results showed that the number of significant changes in HI values increased under climate change by 2100, depending on the RCP: 2 HIs out of the 17 changed under RCP2.6; 6 HIs under RCP4.5; 10 HIs under RCP8.5. The intensity of drainage peak flows linked to flood events and the annual maximal discharge changed significantly under all RCPs. The temporality of the drainage season was substantially affected according to how pessimistic the RCP was. The worst changes were observed under RCP8.5, which exacerbated extreme events: The wet period was shorter while the dry period was longer by about 67%; the drought index increased by 100%; the summer drained water balance decreased by 9%. On the contrary, in winter, the duration of the wet period decreased while maintaining the same drained water balance, thus inducing stronger flood events leading to an earlier saturation of the drainage networks. The sustainability of the drainage system design at *La Jaillière* is therefore threatened, with the risk of fulfilling its function less effectively by 2100, exposing current crops to more important runoff and affecting water quality by increasing the leaching of agrochemical inputs.

Keywords: subsurface drainage, hydrological indicators, climate change, modelling chain, uncertainties, adaptation

INTRODUCTION

In its successive reports, the IPCC stated that the climate will be subject to several changes by 2100, including an increasing number of rapid and intense events (Garner et al., 2015), such as flash floods (Kundzewicz et al., 2014; Yin et al., 2016; Xu et al., 2019; Zhang et al., 2019), and a global increase in air temperature (IPCC, 2008, 2014) resulting in more frequent drought events (Prudhomme et al., 2014). Numerous studies have been carried out to assess the impact of climate change on the hydrological cycle, both at the catchment and regional levels, mostly on hydrologically unmodified systems and at large spatial scales (Khaliq et al., 2009; Arnell and Gosling, 2016; Carvalho-Santos et al., 2016; Lemaitre-Basset et al., 2021).

The future of subsurface drainage hydrology under climate change is also assessed but sometimes studies reports contrasting results. Sojka et al. (2020) and Pease et al. (2017) found that the annual drained water balance decreases in central western Poland, and in the north-eastern United States (United States), respectively. Some studies, however, show the opposite pattern in eastern Canada and Quebec (Dayyani et al., 2012; Mehan et al., 2019; Jiang et al., 2020; Golmohammadi et al., 2021), even among sites from the same geographical area. Furthermore, making broad conclusions about the future of subsurface drainage under climate change, e.g. using a meta-analysis procedure, is particularly challenging due to the lack of available studies on this topic. A case-by-case study is then one of the remaining ways to assess the future of subsurface drainage system in view of possible and appropriate improvements. Indeed, the sustainability of the drainage network design, i.e. its capacity of fulfilling its function in the current design, is questioned (Deelstra, 2015; Abd-Elaty et al., 2019; Sojka et al., 2020), since the depth and spacing between each drain may no longer be suitable for future conditions. Stable hydric conditions in drained plots are no longer guaranteed, as they are exposed to more frequent flood events under climate change (IPCC, 2014) or drought events that increase crop water demand to which farmers may respond with increasing irrigation (Grusson et al., 2021).

Assessing subsurface drainage future is a key point to mitigate or adapt to the impacts of climate change on the concerned lands, especially regarding crop yields. Currently used on plots that show infiltration issues such as hydromorphic soils (Jamagne, 1968; Thompson et al., 1997; Baize and Jabiol, 2011; Lange et al., 2011), subsurface drainage is a soil management technique that consists in introducing a network of perforated pipes into the ground to facilitate water infiltration and reduce surface runoff (Henine et al., 2014; Tuohy et al., 2018). In Europe, the proportion of drained agricultural lands varies across countries depending on pedological properties, climate conditions and agricultural policies. Higher proportions are observed in northern Europe with more than 50% in the United Kingdom, the Netherlands and Finland, while in southern countries the proportions are lower, such as Belgium (from 10 to 20%), France and Spain (below 10%) (ICID, 2021). In France, subsurface drainage plays an important role in the environment and in the dynamics and quality of water resources (Tournebize et al., 2012; Tournebize et al., 2017; Lebrun et al., 2019). To adapt

subsurface drainage management sustainably, it is therefore essential to assess and understand its functioning under climate change.

Studies dealing with subsurface drainage differ in their view of the subsurface drainage and in their choice of hydrological indicators (HIs) with which to study the system. One limitation characterizing these studies comes from the low number of HIs commonly used to describe the subsurface drainage hydrology. Usually, HIs deal with the drained water balance (Dayyani et al., 2012; Mehan et al., 2019; Golmohammadi et al., 2021) and the daily discharge (Pease et al., 2017; Jiang et al., 2020). However, these characteristics do not offer a precise description of the future of subsurface drainage. On the contrary, indicators describing flood events, such as the flashiness index (Deelstra, 2015) that indicates the speed of change from one state to another, can be useful for describing drainage dynamics. Other HI classes helping farmers in the management of pollutants can also be particularly relevant, such as the beginning of the drainage season (Henine et al., 2022) or the duration of soil dryness or saturation. Introducing these new HI classes in drainage research can help gain knowledge about the future of subsurface drainage, providing useful information to the stakeholders concerned.

Running a hydrological model with climate change scenarios is a classic method for assessing the impact of climate change on hydrology (Arnell and Reynard, 1996; Prudhomme et al., 2010; Charlton and Arnell, 2014). A common approach is to use climate projections following different Representative Concentration Pathways (RCPs) induced by greenhouse gas emissions by 2100 (Moss et al., 2008; Kalantari, 2011; IPCC, 2014; Nam et al., 2015; Mukundan et al., 2020). Although useful, this method may induce biases in the calculation of the HIs. For instance, it is strongly recommended to use several climate projections (i.e. using different general circulation models, GCMs, Randall et al. (2001); Mechoso and Arakawa (2003)) per RCP to better account for uncertainties (Shrestha et al., 2018). Assessing and integrating them is therefore a necessary step that must be carefully addressed and is given specific attention in the present paper.

The aim of this study is thus to assess the future of subsurface drainage under climate change in one specific French drained plot in the *La Jaillièrè* site in western France, chosen for its representativeness of French subsurface drainage. To our knowledge, this is the first analysis of climate change effects on drainage in France. We use 30 Euro-Cordex future climate projections distributed over three RCPs: RCP2.6, RCP4.5 and RCP8.5. These climate projections feed the subsurface drainage model SIDRA-RU (Henine et al., 2022), a lumped and parsimonious model which has demonstrated good performance and temporal robustness in simulating subsurface drainage fluxes in France (Jeantet et al., 2021). The SIDRA-RU model provides subsurface drainage discharges, which then allow us to calculate 17 hydrological indicators specific to subsurface drainage, which constitute a uniquely large range of complementary indicators and give a comprehensive view of the subsurface drainage response. First, we assess the potential bias induced by the use of the climate projections in

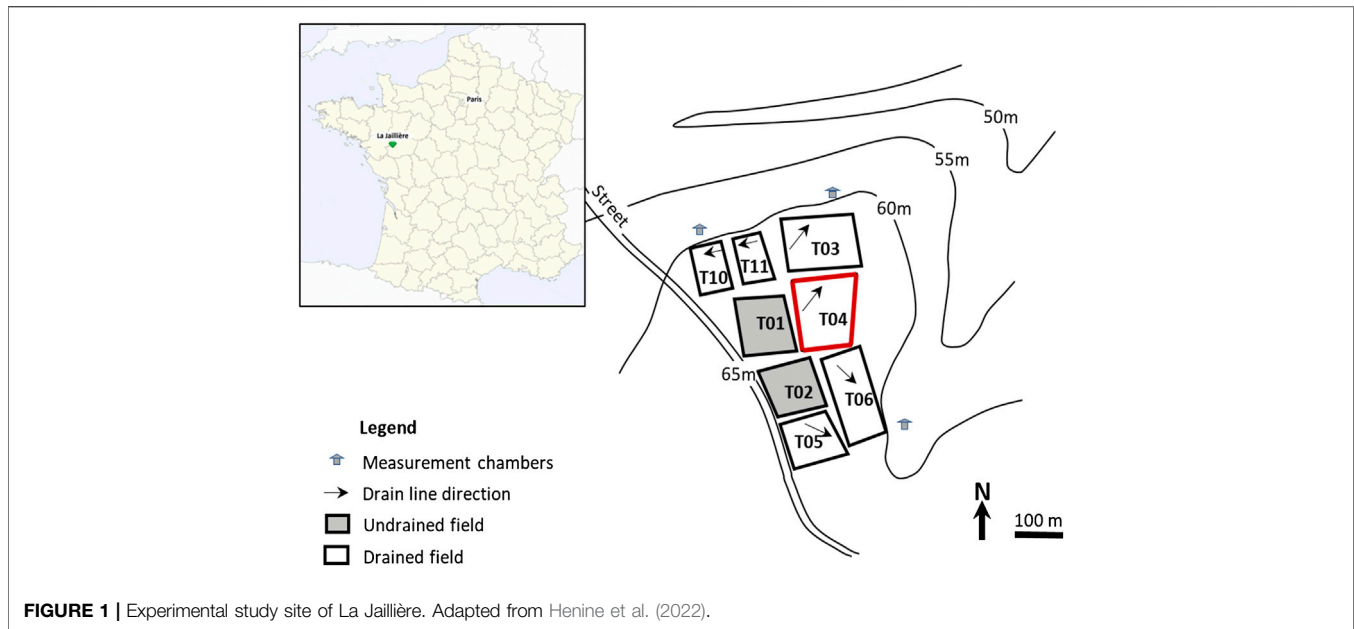


FIGURE 1 | Experimental study site of La Jaillière. Adapted from Henine et al. (2022).

the SIDRA-RU model on these indicators over a historical period (1975–2004). Second, we analyse the evolution of the indicators from the historical period to a future period (2006–2099) in order to assess the potential impact of climate change on subsurface drainage in *La Jaillière*.

MATERIALS AND METHODS

Study Site

The *La Jaillière* site (Figure 1) is an experimental station located in the Loire-Atlantique region in western France, managed by Cemagref and ITCF institutes between 1987 and 1999, and by Arvalis-“Institut du végétal” (Vegetal institute in France) from 1999 to the present. The soil of *La Jaillière* is hydromorphic and brown belonging to the Luvisol category (IUSS Working Group WRB, 2015; Dairon et al., 2017) with a loamy soil texture. This combination is found in approximately 80% of the total French drained areas (Lagacherie and Favrot, 1987) and is thus a good representative of French subsurface drained areas. *La Jaillière* is also a reference site as part of the EU agricultural experimental sites for the assessment of the dynamics of active pesticide substances in drained soils (FOCUS, 2012). Our study focuses on plot T4 (Figure 1) that covers 0.9 ha and is characterized by a mainly silty-sandy soil texture at the soil horizon (20–25% clayey, 40–45% silty, 30–40% sandy) and silty-clayey texture in the lower horizon (40–50% clayey, 35% silty, 15–25% sandy).

The climate in the study site is oceanic with an annual cumulative rainfall and a mean potential evapotranspiration, respectively, of 709 and 738 mm, and with a mean annual temperature of approximately 11°C. A subsurface drainage system composed of perforated polyvinyl chloride (PVC) pipes drains the plot, with an inter-drain spacing of 10 m at an average depth of 0.9 m. The drainage network was settled in

the 80s during campaigns managed by INRAE (formerly CEMAGREF) to evaluate drainage modalities according to pedological and climatic conditions. The water-holding capacity was estimated to be 104 mm (Henine et al., 2022). The crop system is a two-year rotation of winter wheat and maize. winter wheat is cultivated from the second half of October to the second half of the next July, while maize is cultivated from late April/early May to September. Cover crops are used once every two years in the period included from the harvest of winter wheat to the sowing of maize. The surface runoff is deemed negligible on this site (Kuzmanovski et al., 2015; Henine et al., 2022) and thus it was not studied here.

Input Data

The meteorological data used as a reference for the historical period were provided by the SAFRAN database (Vidal et al., 2010), supplying a meteorological reanalysis of precipitation (P) and potential evapotranspiration (PE, based on the FAO-56 Penman–Monteith PE formulation (Córdova et al., 2015)) over France. Data are available from 1959 to 2019 at a daily time step and on a regular grid of 8 × 8 km.

In this study, climate projection data were provided by six GCMs coupled with nine regional circulation models (RCMs) in order to spatially disaggregate data from GCMs. Due to the computation time limitations in generating the climate projections (CPs), not all GCM–RCM couples are available for every RCP (Table 1, Lemaitre-Basset et al., 2022). In our study, we considered 12 CPs (Soubeyrou et al., 2021) from the Euro-Cordex project (Jacob et al., 2014) selected by Météo-France to ensure the representativeness of future conditions in France. Post-treatment through the ADAMONT method (Verfaillie et al., 2017) was carried out on the CPs, first to correct bias versus SAFRAN data using the quantile–quantile mapping method (Maurer et al.,

TABLE 1 | Availability of climate projections. The numbers (2.6, 4.5 and 8.5) refer to the RCPs used by the GCM (rows)/RCM (columns) pairs. “-” indicates the absence of data. Adapted from Lemaître-Basset et al. (2022).

GCM/RCM	Aladin63	CCLM4-8	HIRHAM5	Racmo22E	RCA4	RegCM4	REMO2009	REMO2015	WRF381P
CNRM-CM5	2.6–4.5–8.5	—	—	2.6–4.5–8.5	—	—	—	—	—
EC-EARTH	—	—	—	2.6–4.5–8.5	2.6–4.5–8.5	—	—	—	—
IPSL-CM5A	—	—	—	—	4.5–8.5	—	—	—	4.5–8.5
HadGEM2-ES	—	4.5–8.5	—	—	—	2.6–8.5	—	—	—
MPI-ESM-LR	—	2.6–4.5–8.5	—	—	—	—	2.6–4.5–8.5	—	—
NorESM1-M	—	—	4.5–8.5	—	—	—	—	2.6–8.5	—

2010; Navarro-Racines et al., 2020; Potter et al., 2020), and second to downscale the CPs to the same spatial resolution as SAFRAN. These CPs provide simulations following three RCPs: the very stringent pathway RCP2.6 (van Vuuren et al., 2007), the intermediate pathway RCP4.5 (Thomson et al., 2011) and the worst-case pathway RCP8.5 (Riahi et al., 2011). Projections are available at a daily time step from 1975 to 2100 and are split into two parts:

- “Historical” from 1975 to 2004: All climate models are forced by the observed greenhouse gases concentrations. According to the hydrological year concept, we considered data from August 1, 1975 to July 31, 2004;
- “Projected” from 2006 to 2100: The projections are led by the RCPs, defining the evolution of greenhouse gas concentrations. As the year 2005 is the transitional year between the two periods, we excluded it from the analyses. Thus, we considered data from August 1, 2006 to July 31, 2099.

The SIDRA-RU Model

The hydrological model used in this study was the SIDRA-RU model (Henine et al., 2022), a four-parameter lumped and semi-conceptual model, designed to simulate subsurface drainage discharges from artificially drained plots. The model uses the principle of rainfall–discharge conversion being fed by daily precipitation P(t) (without distinguishing liquid and solid parts) and PE(t) to predict drainage discharge Q(t) at the drainage network outlet. First, an evapotranspiration module converts PET(t) into an approximate value of actual evapotranspiration, called corrected evapotranspiration CET(t), from the available water level S(t) in a conceptual storage to satisfy the evapotranspiration (see Eq. 1):

$$CET(t) = PET(t) * e^{-\frac{S_{RFU}-S(t)}{S(t)}} \tag{1}$$

With S_{RFU} a threshold assigning the minimal water level to fully satisfy PET(t). Let us note that current crop is not used to simulate CET, only depending on the water availability of the storage. Second, the net infiltration $P_{net}(t)$ is calculated by subtracting CET(t) from P(t) in order to simulate the water table recharge term R(t) (mm) using the RU module (“Réserve Utile” in French designated the water holding capacity of a soil). R(t) is simulated from the meteorological input and the available water level S(t) in the reservoir (see Eq. 2) distinguishing three situations:

$$\begin{cases} \text{if } S(t) < S_{inter}: S(t) = S(t-1) + P_{net}(t) \mid R(t) = 0; \\ \text{if } S(t) \in [S_{inter}; S_{max}]: S(t) = (1 - \alpha) P_{net}(t) + S(t-1) \mid R(t) = \alpha * P_{net}(t); \\ \text{if } S(t) \geq S_{max}: S(t) = S_{max} \mid R(t) = P_{net}(t). \end{cases} \tag{2}$$

With:

- α (-), proportion of $P_{net}(t)$ experimentally set at 1/3 (Henine et al., 2022);
- S_{inter} (mm), intermediate threshold of the soil reservoir defining the water quantity needed to generate flow in the reservoir before saturation of the storage ($S_{RFU} = 0.4 * S_{inter}$, Jeantet et al. (2021));
- S_{max} (mm), parameter representing the maximal capacity of the soil reservoir from which the net infiltration is fully converted into R(t).

Let us note that S_{inter} and S_{max} are the two main parameters of the RU module, used to simulate some hydrological indicators (see *Hydrological Indicators of Subsurface Drainage Section*). Third, the physically-based SIDRA module (“Simulation du DRAINage” for drainage simulation in French) converts R(t) into Q(t) (see Eqs 3, 4), following the Boussinesq equation (Boussinesq, 1904):

$$\frac{dh(t)}{dt} = \frac{R(t) - K \frac{h(t)^2}{L^2}}{A_2 \mu} ; h(t+1) = h(t) + \frac{dh(t)}{dt} \tag{3}$$

$$Q(t) = AK \frac{h(t)^2}{L^2} + (1 - A)R(t) \tag{4}$$

With:

- L: space between drain and inter-drain (m);
- A_2 : second water table shape factor (Lesaffre, 1989), $A_2 \approx 0.89$ (-);
- A: third water table shape factor (Lesaffre, 1989), $A = 0.869$ (-). Here, we assume that the shape of the water table between the drain and the mid-drain is an ellipse. Therefore, A is obtained by integrating $\frac{1}{4}$ of this reference ellipse.

The SIDRA module is mainly controlled by two parameters: the horizontal hydraulic conductivity K ($m.d^{-1}$), and drainage porosity μ (-). The four calibrated parameters are K, μ , S_{inter} and S_{max} . In this study, their values were extracted from the analysis of Jeantet et al. (2021). Unlike most current subsurface drainage models, the SIDRA-RU model use strong assumptions, such

TABLE 2 | Hydrological indicators (HIs) and corresponding variables analysed.

Indicator Classes	HIs	Analysed Variables from HIs	Units	Key Index for Illustrations
1. Temporality of the drainage season	Beginning of the drainage season	Annual date of the drainage season	date	1. BegDS 1. LenDS
	Length of the drainage season	Number of days in a year	d	
2. Water content in the storage	Wet index	Number of days with a saturated storage $S(t) = S_{max}$	d	2. WetInd 2. DrouInd
	Drought index	Number of days with a level in the storage $S(t) \leq 25$ mm	d	
3. Drained water balance	Annual drained water balance (ADWB)	Quantity of drained water over a year	mm	3. ADWB 3. SDWB
	Summer drained water balance (SDWB)	Quantity of drained water from April to September	mm	
4. Temporal flow dynamics	Length of the dry period	Number of days in a year	d	4. LenDryP 4. LenRechP 4. LenWetP 4. LenReceP
	Length of the recharge period	Number of days in a year	d	
	Length of the wet period	Number of days in a year	d	
	Length of the recession period	Number of days in a year	d	
5. Annual flood events	Number of annual flood events	Number of annual flood events	—	5. NumFE 5. CumLenFE 5. MaxDisch 5. PkFlowAve
	Annual cumulative length in flood	Number of days in a year	d	
	Annual maximal discharge	Annual maximal discharge	mm.d ⁻¹	
	Annual average of peak flows	Annual mean of peak flows from floods	mm.d ⁻¹	
6. Sustainability of current networks	Return period (RP) of the regime change	RP at the regime change in the $Q_{spe} = f(RP)$ equation	y	—
	Associated specific discharge	Specific discharge Q_{spe} corresponding to the RP at the regime change	mm.d ⁻¹	—
	QRP05, QRP10, QRP20 and QRP50	Specific discharge respectively for RP = 5, 10, 20 and 50 years	mm.d ⁻¹	—

neglecting current crop to simulate CET. This assumption is based on the fact that, despite being managed by the two-year crop rotation winter wheat/maize and the absence/presence of cover crops, the water balance of the studied site is not affected by current crop (Henine et al., 2022). A more detailed description of the SIDRA-RU model is provided by Henine et al. (2022) and Jeantet et al. (2021).

Hydrological Indicators of Subsurface Drainage

To describe the future reaction of subsurface drainage systems as precisely as possible so as to help farmers and decision-makers, 17 HIs gathered into six classes were defined (see Table 2). The drained water balance and the water content in the storage categories are mainly used to evaluate the arable nature of a soil. The categories of annual flood events and the sustainability of the current networks are analysed to assess whether the drainage network is still able to reduce flooding or whether it requires resizing. Finally, the temporality of the drainage season and the temporal flow dynamic categories are useful for assessing the effects of hydrological changes on water quality due to pollutant leaching. These classes and indicators are further detailed below:

1) Temporality of the drainage season. Considering a hydrological year defined from August 1 to next July 31, we considered two HIs: the beginning of the drainage season and its length. The beginning of the drainage season is set to

day d using the cumulative discharge at d $Cum_Q(d)$ since the last August 1 (see Eq. 5):

$$\left\{ \begin{array}{l} \text{on the day } d : Cum_Q(d) \geq Cum_{Drainage} \\ [d + 1 : d + 5] : Cum_Q(d) \geq Cum_{Drainage} + Cum_{reliable} \end{array} \right\} \quad (5)$$

$Cum_{Drainage}$ and $Cum_{reliable}$ correspond to thresholds of cumulative discharges from the previous August 1, considered as minimum water quantities required to assess that the drainage season started at day d (Henine et al., 2022). They were experimentally set to 1.4 and 2.7 mm, respectively. The duration of the drainage season is defined as the number of days between the start and the end of the drainage season. The end of the drainage season is set on the first day d when the difference between the annual drained water balance on July 31, i.e. the end of the hydrological year, and $Cum_Q(d)$ is lower than $Cum_{Drainage}$:

- 2) Water content in the soil profile: the number of days the storage is saturated ($S(t) = S_{max}$) and the number of days the storage is close to being empty ($S(t) \leq 25$ mm). Below 25 mm, the remaining stock is assimilated to non-available water to crops or PE demand. These two HIs are, respectively, defined as the wet index and the drought index;
- 3) Drained water balance: annual drained water balance (ADWB) and summer drained water balance (SDWB), for each hydrological year. The latter is set from April to September, i.e. until the end of the winter runoff to the end of the period when the soil storage is empty in most cases;

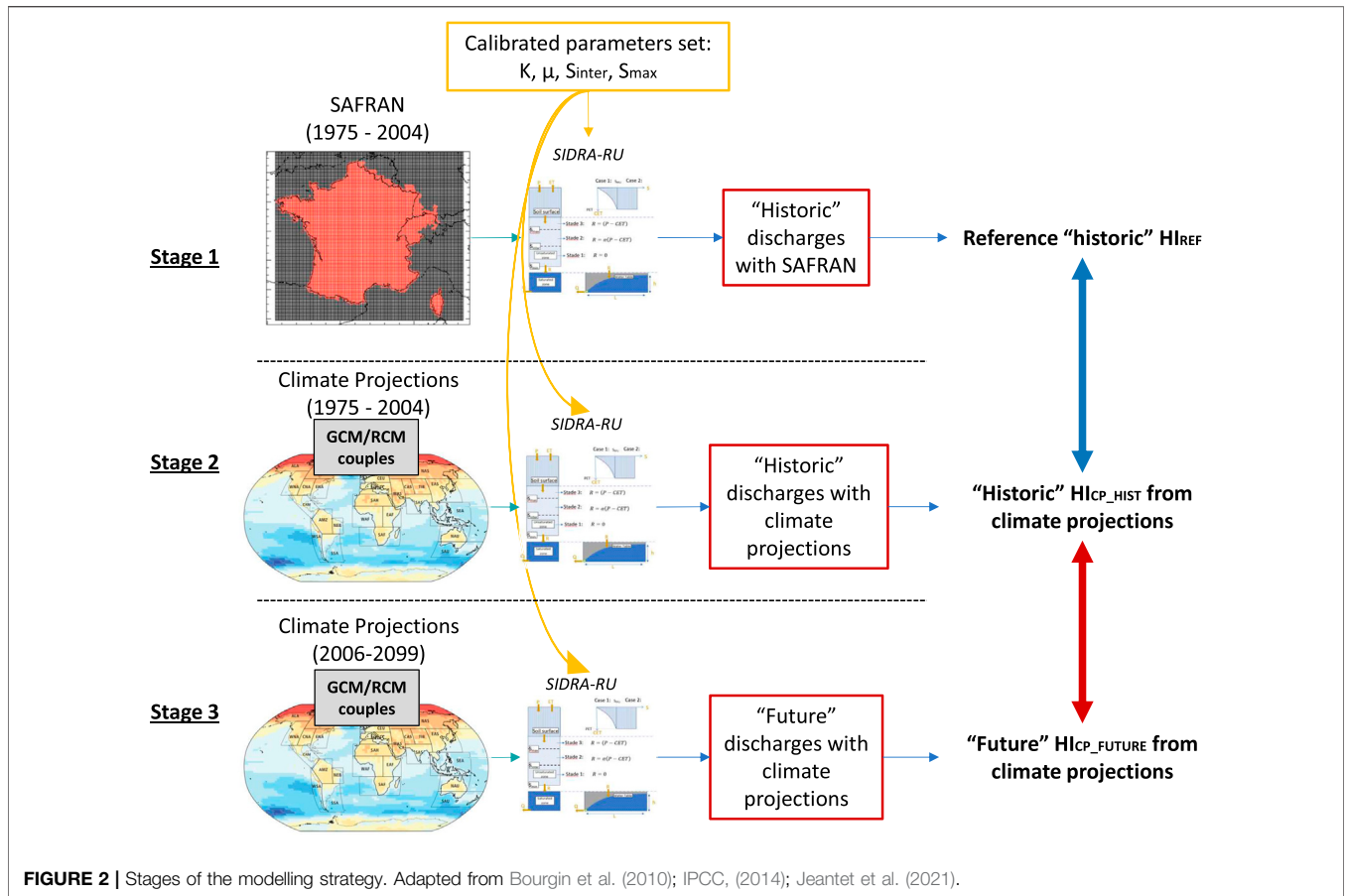


FIGURE 2 | Stages of the modelling strategy. Adapted from Bourgin et al. (2010); IPCC, (2014); Jeantet et al. (2021).

- 4) Temporal flow dynamics: the annual flow hydrograph is split into four phases per hydrological year (Humbert et al., 2015; Strohmenger et al., 2020): dry period, recharge period, wet period and recession period. The duration of each phase was analysed;
- 5) Annual flood events: a discharge Q_d on day d is considered as a flood event following Eq. 6:

$$\left\{ \begin{array}{l} \text{on the day } d : Q_d \geq T_{flood} \\ \text{on the day } d + 1 : Q_{d+1} \geq Q_d + T_{flood} \end{array} \right\} \quad (6)$$

The threshold T_{flood} is set to 1 mm d^{-1} (Nedelec, 2005). A flood ends on day d following Eq. 7:

$$\left\{ \begin{array}{l} Q_d < T_{flood} \\ Q_d < 0.8 * Q_{flood_max} \end{array} \right\} \quad (7)$$

with Q_{flood_max} the maximal discharge of the flood event (peak flow). The number of annual flood events, the annual cumulative duration in flood events, the annual maximal discharge and the annual mean of peak flows were considered;

- 6) Sustainability of the current networks: a regression analysis is carried out on the evolution of a specific discharge Q_{spe} according to the associated return period (RP). The regime from the relation $Q_{spe} = f(RP)$ is characterized by two

different linear regimes split by a breaking point close to 2 years, due to design criteria in order to intercept a rainfall of 15 mm d^{-1} (Augeard et al., 2008; Henine et al., 2012). Above this value, the pipe is pressurized (Henine et al., 2010) and the specific discharge may be strongly reduced (Nedelec, 2005; Henine et al., 2012), according to the size of the networks and depending on the intensity of the event. If the RP of the breaking point decreases, the size of the networks might be considered as no longer suitable to fulfil efficiently their flood protection function. In this study, Hazen’s relation and the Gumbel distribution function (Gumbel, 1954) were used to establish the piecewise function of the regime and determine the breakpoint (see Eqs 8, 9) so as to assess its evolution under climate change:

$$Q_{spe} = \begin{cases} a_1 * u + b_0 & \text{if } u \leq u_{RC}; \\ a_2 * u + a_1 * u_{RC} + b_0 & \text{if } u > u_{RC}; \end{cases} \quad (8)$$

$$u = \ln \left[- \ln \left(1 - \frac{1}{RP} \right) \right] \quad (9)$$

with:

- > u : the change of variable in the Gumbel relation;
- > u_{RC} : the change of variable corresponding to RP_{RC} , the RP at the regime change;

TABLE 3 | Temporal split of the climate projections.

Sub-period years	1975–2004	2006–2040	2041–2070	2071–2099
Sub-period name	Historical	PP1	PP2	PP3

> a_1 , a_2 and b_0 : parameters from the non-linear regression.

The evolution of Q_{spe} at RP = 5, 10, 20 and 50 years (QRP05, QRP10, QRP20 and QRP50) was also assessed.

Modelling Strategy

The modelling strategy is divided into three parts (Figure 2). First, we established reference hydrological indicators in the historical period (HI_{REF}) from 1975 to 2004 by simulating historical discharge using the SIDRA-RU model fed with the SAFRAN meteorological data. Second, the SIDRA-RU was fed with each historical CP to calculate all historical hydrological indicators (HI_{CP_HIST}) for the same historical period. Third, the SIDRA-RU was fed with the CPs in the future projected period, from 2006 to 2099. The future discharge series obtained were split into three continuous sub-periods, each one of approximately 30 years (Table 3) to represent the dispersion of climatic patterns (Soubeyroux et al., 2021). Then, we calculated future hydrological indicators (HI_{CP_FUTURE}) in the whole future period for HI classes 1–5 and in each future sub-period for HI class 6. Two sequences of analysis were performed: 1) the HI_{REF} and the HI_{CP_HIST} were compared to assess the bias induced by the use of climate projections in the SIDRA-RU model on the HIs; 2) the HI_{CP_FUTURE} were compared to the HI_{CP_HIST} to assess the impact of climate change on subsurface drainage hydrology.

Evaluation of Bias From Projections in the SIDRA-RU Model Over the Historical Period

To assess the potential bias induced by the use of CPs in the SIDRA-RU model, the distribution functions of the HI_{CP_HIST} were compared with those from the HI_{REF} . We calculated the distribution functions from the HI values extracted on an annual basis and we then ranked them in ascending order. For each HI from classes 1 to 5 (Table 2), comparisons were made between the HI_{CP_HIST} from 12 different CPs (Table 1), i.e. from 12 series of HI_{CP_HIST} obtained for 30 hydrological years from 1975 to 2004. To avoid the comparison of each of these 12 series while preserving two samples of equivalent sizes, we averaged the 12 distribution functions of the HI_{CP_HIST} and compared the mean of the distribution function obtained to the corresponding HI_{REF} . The mean was calculated with a confidence interval at 95%, following a normal distribution (see Eq. 10, Saporta, (2006)):

$$\left[\mu - t_{0.025, n-1} * \frac{\sigma}{\sqrt{n}} ; \mu + t_{0.025, n-1} * \frac{\sigma}{\sqrt{n}} \right] \quad (10)$$

with:

> μ : the mean quantile from the distribution function from the CPs, per HI;

> σ : the associated standard deviation;

> n : the size of the sample, corresponding to the number of CPs tested (here, $n = 12$);

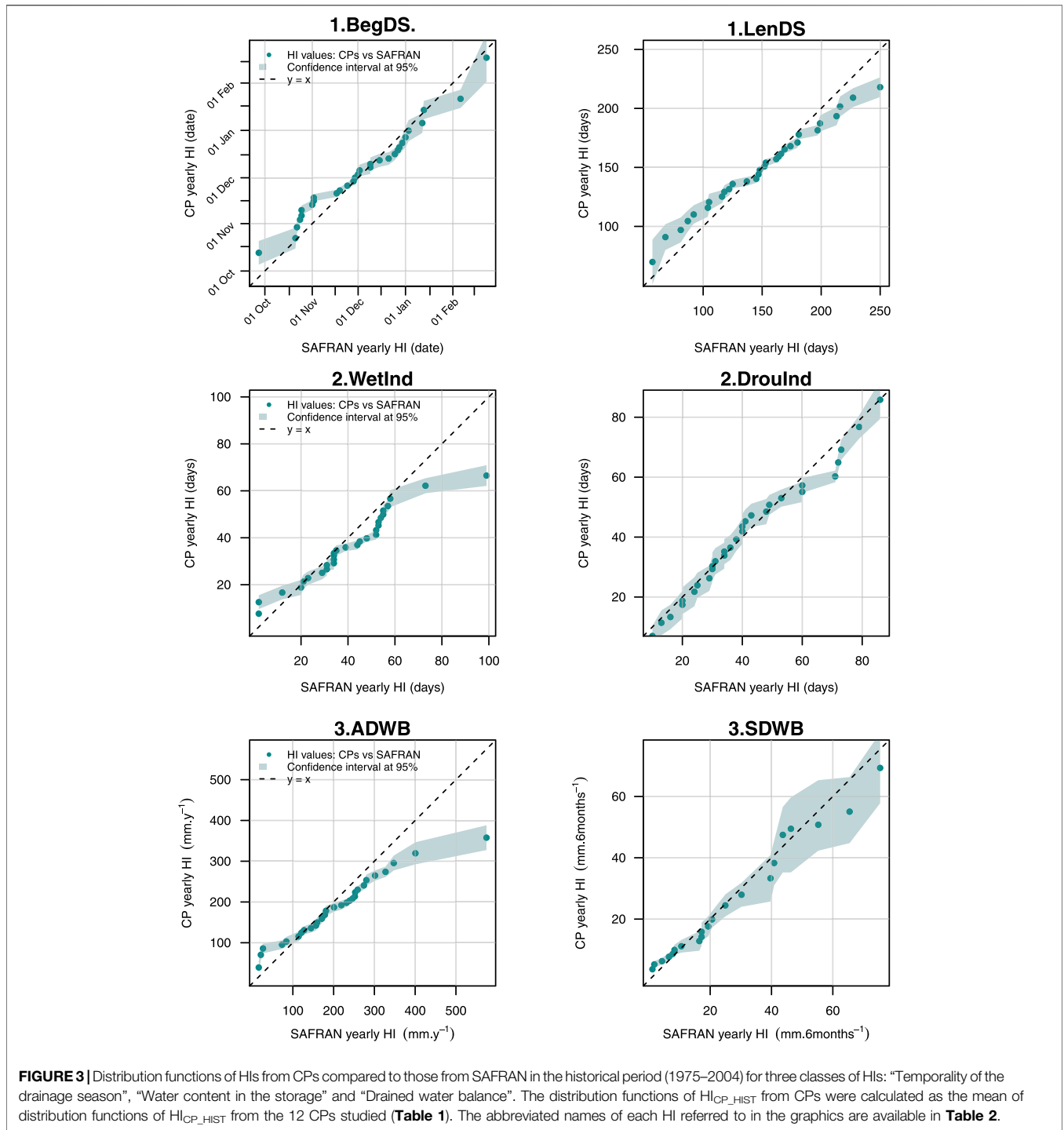
> $t_{0.025}$: a coefficient used to establish a theoretical confidence interval at 95% according to n (Barrett et al., 2015), assuming a normal distribution.

In practice, an initial graphic evaluation of quantile-quantile plots (Saporta, 2006) made it possible to compare the mean of the distribution functions of HI_{CP_HIST} with the distribution function of HI_{REF} . Second, we statistically tested the similarity between the HI_{CP_HIST} and the HI_{REF} by two non-parametric similarity tests: the Mann–Whitney–Wilcoxon test (Mann and Whitney, 1947; Hollander and Wolfe, 1973) and the Kolmogorov–Smirnov test (Birnbaum and Tingey, 1951; Conover, 1971). Third, two correlation tests were performed to assess the correlation between these two distribution functions: the Pearson test (Pearson and Galton, 1895) and the Spearman test (Spearman, 1904). The R package “stats” (R Core Team, 2021) was used for all these tests. The threshold for statistical significance was set at $p \leq 0.05$.

For HI class 6 (see Table 2), the analyses consisted in a graphic comparison of specific discharges Q_{spe} associated with RPs in the historical period (Table 1), namely between Q_{spe} from the historical reference and those from CPs. Here, Q_{spe} was associated with the annual maximal discharge Q_{max} , defined in HI class 5 (Table 2). Q_{max} values from CPs were extracted from the mean of the distribution functions of Q_{max} from the 12 CPs, calculated with a confidence interval at 95%. Then, piecewise functions (see Eq. 8) from both situations were fitted by a non-linear regression using smoothed Q_{max} series and the R package “stats” to extract the parameters a_1 , a_2 and b_0 . Finally, RP values for the regime change were compared, along with the corresponding Q_{max} . Moreover, the QRP05, QRP10, QRP20 and QRP50 values from the historical reference were compared with those from CPs.

Evolution of the Hydrological Indicators in the Projected Period

For HI classes 1–5 (Table 2) and for each of the three RCPs, we calculated a 5-years rolling mean for the period 1975–2099 of each HI_{CP_HIST} and the associated HI_{CP_FUTURE} to assess the temporal trend of HIs in the future, with a confidence interval at 95%. Because of the CP data availability from 1975 to 2099, the rolling period was not centred on the first two years, i.e. 1975 and 1976, or the last two years, i.e. 2098 and 2099. As mentioned in Input data Section, the year 2005 was excluded from the analyses because it marks the transition from the projected past to the projected future. Therefore, the rolling period was not centred on 2003, 2004, 2006 and 2007 either. The significance of the trend of each HI_{CP_FUTURE} was assessed by performing a Mann–Kendall test (Mann, 1945; Hipel and McLeod, 1994) on both means and lower and upper limits of the confidence interval (R package “Kendall”, McLeod, (2011)). The threshold for statistical significance was set at $p \leq 0.05$. In those cases, the trends were calculated as a percentage change from the HI_{CP_HIST} from the historical period 1975–2004 and the corresponding HI_{CP_FUTURE}



from the future period PP3 (Table 3). The trends of HI_{CP_HIST} were calculated using the 12 CPs whereas the trends of HI_{CP_FUTURE} were calculated using the available CPs according to the corresponding RCP (Table 1).

For HI class 6, the analyses consisted in a graphic comparison between specific discharges Q_{spe} associated with RPs from the historical period and those from the three future sub-periods (Table 1), all calculated from CPs. Similarly to the bias analysis

for the historical period (*Evaluation of Bias From Projections in the SIDRA-RU Model Over the Historical Period Section*), the Q_{spe} values from each future sub-period under each RCP (Table 3) were assimilated to the mean of the distribution functions of Q_{max} from the available corresponding CPs (Table 1), supported by a confidence interval at 95%. Then, piecewise functions (see Eq. 8) were fitted by non-linear regression from the smoothed Q_{max} of each case using the R package “stats”, and RP values for the

TABLE 4 | Results of the significant tests and the correlation tests comparing the HI_{CP_HIST} and HI_{REF} distribution functions in the historical period (1975–2004).

Hydrological Indicators		Kolmogorov–Smirnov's D (p)	Mann–Whitney's W (p)	Pearson's Correlation (p)	Spearman's Correlation (p)
1. Temporality of the drainage season	Beginning of the drainage season	0.200 (0.586)	437 (0.853)	0.998 (<0.001)	1.000 (<0.001)
	Length of the drainage season	0.133 (0.952)	449 (0.994)	0.995 (<0.001)	1.000 (<0.001)
2. Water content in the storage	Wet index	0.241 (0.367)	481 (0.351)	0.970 (<0.001)	0.998 (<0.001)
	Drought index	0.103 (0.998)	424 (0.957)	0.988 (<0.001)	0.999 (<0.001)
3. Drained water balance	ADWB	0.200 (0.586)	495 (0.513)	0.978 (<0.001)	1.000 (<0.001)
	SDWB	0.143 (0.983)	223 (0.960)	0.988 (<0.001)	0.999 (<0.001)
4. Temporal flow dynamics	Length of the dry period	0.138 (0.945)	387 (0.613)	0.996 (<0.001)	1.000 (<0.001)
	Length of the recharge period	0.172 (0.782)	386 (0.597)	0.981 (<0.001)	0.996 (<0.001)
	Length of the wet period	0.138 (0.945)	387 (0.663)	0.993 (<0.001)	1.000 (<0.001)
	Length of the recession period	0.138 (0.945)	396 (0.709)	0.971 (<0.001)	0.995 (<0.001)
5. Annual flood events	Number of annual flood events	0.138 (0.945)	430 (0.882)	0.966 (<0.001)	0.994 (<0.001)
	Annual cumulative duration in flood	0.172 (0.782)	456 (0.586)	0.960 (<0.001)	1.000 (<0.001)
	Annual maximal discharge	0.207 (0.564)	473 (0.367)	0.964 (<0.001)	1.000 (<0.001)
	Annual average of peak flows	0.241 (0.367)	480 (0.362)	0.946 (<0.001)	1.000 (<0.001)

The p -values of the statistical tests.

regime change were compared, along with the corresponding Q_{max} . The deviations between the historical and future RPs at the regime changes RP_{RC} were calculated only if they were significant, i.e. if the confidence intervals from scatter plots were not superimposed close to the regime change, the same as for the corresponding discharges Q_{RC} . Moreover, the QRP05, QRP10, QRP20 and QRP50 values from the three future sub-periods were extracted from the previously fitted piecewise functions (see Eq. 8) and were compared with the values from the historical period.

RESULTS

Evaluation of Potential Biases in HI Values From Climate Projections Over the Historical Reference

For the large majority of HIs, quantiles of HI_{CP_HIST} distributions were close to those of HI_{REF} following the 1:1 line, i.e. the line corresponding to the equation $y = x$ (Figure 3 for HI classes 1–3 and Supplementary Appendix Figure SA1, for HI classes 4 and 5). Only a few extreme values (1–3 out of 30) from HI_{CP_HIST} showed stronger deviations compared to those from HI_{REF} in particular for the wet index (2.WetInd in Figure 3), the annual drained water balance (3.ADWB in Figure 3), showing the strongest deviations although rarely exceeding 10%, and the length of the wet period (4.LenWetP on Supplementary Appendix Figure SA1) that were deemed negligible.

Results from the Mann–Whitney test and the Kolmogorov–Smirnov test showed no significant difference

between the distribution functions of HI_{CP_HIST} and the distribution functions of HI_{REF} from classes 1 to 5, with all p values being higher than 0.05 (Table 4). In addition, the results from Spearman and Pearson correlation tests showed that for HI classes 1–5 the HI_{REF} and HI_{CP_HIST} trends were significantly correlated (p values ≤ 0.05), with correlation coefficients currently over 0.95. Overall, these results (graphical exploration and statistical testing) supported the hypothesis that for all the HIs from classes 1 to 5, the HI_{CP_HIST} and HI_{REF} distributions did not substantially differ, which allowed us to consider that no critical bias was exposed in the use of SIDRA-RU.

The discrepancies between Q_{max} associated with the RPs from the historical reference and those obtained from the CPs in the historical period were relatively small (≤ 2 mm d^{-1}) and deemed negligible (Figure 4). Let us note that Q_{max} from the CPs appeared slightly underestimated compared to those from the historical reference, except for RPs shorter than 1.3 years. The two fitted piecewise functions, on Q_{max} from the CPs and on Q_{max} from the historical reference, were well fitted with respective small residual sum-of-squares of 0.79 and 1.98. The corresponding lines in Figure 4 show the regime change from CPs was 2 months late compared to the change from the historical reference, for an associated Q_{max} , respectively, equal to 12.3 mm d^{-1} and 12.8 mm d^{-1} . The coordinates from both breaking points appeared close.

Table 5 gathers the discharges calculated from specific RPs (QRPs) from the historical reference and those from CPs. Both sets of QRPs were calculated with the corresponding fitted piecewise functions from the analyses in Figure 4.

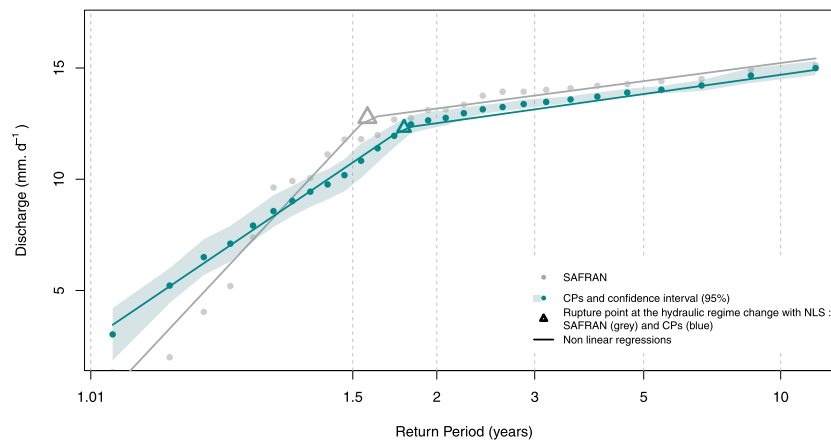


FIGURE 4 | Comparison of the specific discharges Q_{spe} (here the annual maximal discharge) associated with return periods RPs in the historical period 1975–2004 between those obtained from the CPs and those from the historical reference. The equation $Q_{spe} = f(RP)$ was fitted as a piecewise function (see **Eq. 8**) according to a non-linear regression, for both situations. The equation from the CPs was established from the mean of the distribution functions of the annual maximal discharge from the 12 CPs, supported by a confidence interval at 95%.

TABLE 5 | Comparison between the discharges from a specific return period (QRP) from the historical reference SAFRAN and those from CPs. The latter were obtained from the equation $Q_{spe} = f(RP)$ established from the mean of the distribution functions of Q_{max} from the 12 CPs.

Return Period (years)	QRP from SAFRAN (mm.d ⁻¹)	QRP from CPs (mm.d ⁻¹)
5	14.4	14.0
10	15.2	14.9
20	16.0	15.7
50	17.0	16.9

Results showed that QRPs from CPs and those from the historical reference were quite similar across the four RPs studied, except for RP = 5 years that showed the largest deviation of 0.4 mm d⁻¹, although remaining low and therefore negligible.

Evolution of the HIs From CPs in the Future

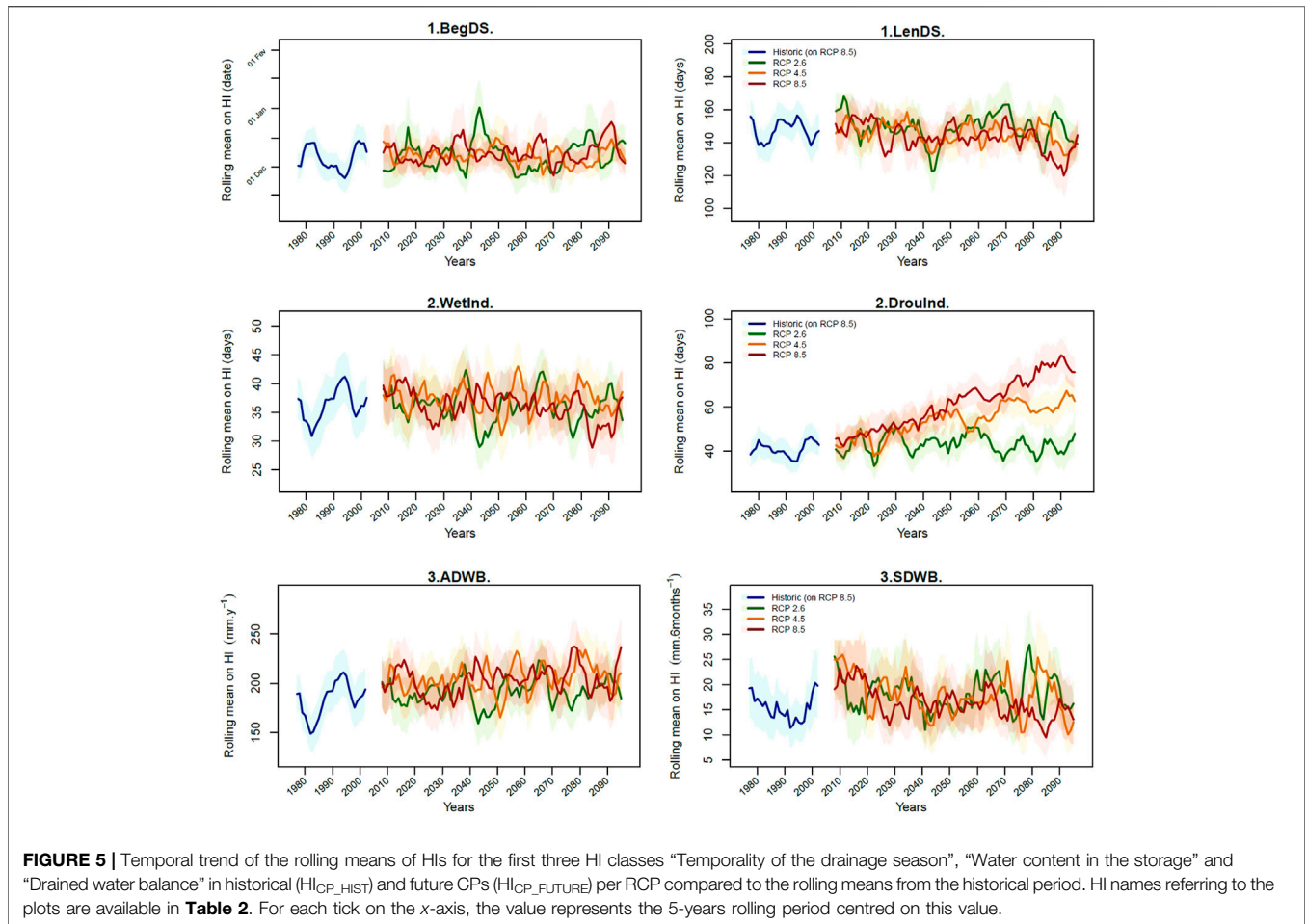
Figure 5 shows the evolution of the rolling mean in HI classes 1–3 for both the historical and the future CPs under the three RCPs (see **Supplementary Figure Appendix SB1** for the graphical results for HI classes 4–5). The results from the Mann–Kendall tests assessing the significance of trends are presented in **Table 6**. First, all of the HI_{CP_FUTURE} showed strong inter-annual variations but their trends remained relatively linear in most cases from 2006 to 2099. Second, the results showed that the number of changing HIs depended on the RCPs. The annual maximal discharge (5.MaxDisch on **Supplementary Figure Appendix SB1**) and the annual average of peak flow (5.PkFlowAve) were the only two HI_{CP_FUTURE} which changed significantly compared to the HI_{CP_HIST} under all the RCPs: They increased by 2 and 4%, respectively, under RCP2.6, by 9 and 12% under RCP4.5 and by 10 and 16% under RCP8.5. For the remaining HIs from classes 1 to 5, the more pessimistic the RCP, the higher the number of changed HIs.

Five HI_{CP_FUTURE} started to change from RCP4.5: the summer drained water balance (3.SDWB) increased by 12% and decreased by 9%, respectively, under RCP4.5 and RCP8.5; the drought index (2.DrouInd) increased by 50 and 85%, respectively, under RCP4.5 and RCP8.5; the length of the drainage season (1.LenDS) decreased by 3 and 7%, respectively, under RCP4.5 and RCP8.5; the length of the dry period increased by 2 and 6%, respectively, under RCP4.5 and RCP8.5. According to the Mann–Kendall test, the length of the recharge period (4.LenRechP) was the only HI_{CP_FUTURE} to change under RCP4.5, decreasing by 16%. Finally, two HI_{CP_FUTURE} changed significantly only under RCP8.5: the wet index (2.WetInd) and the length of the wet period (4.LenWetP) both decreased, respectively, by 5% and by 7%.

Some HI_{CP_FUTURE} showed no significant changes under any RCP: the beginning of the drainage season (1.BegDS), the annual drained water balance (3.ADWB), the length of the recession period (4.LenReceP), the annual number of flood events (5.NumFE) and the annual cumulative length in flood (5.CumLenFE). This was congruent with the graphic analysis showing that HI_{CP_FUTURE} had mean values close to the corresponding HI_{CP_HIST} .

An analysis performed on spring flood events, from April to June, showed that the number of spring flood events significantly changed from the historical period to PP3, depending on the RCP (**Supplementary Table Appendix SB1**). It increased by 43 and 10% under RCP2.6 and RCP4.5, respectively, and decreased by 5% under RCP8.5. The cumulative duration of spring floods started to change under RCP4.5, increasing by 23% and decreasing by 5% under RCP4.5 and RCP8.5, respectively. The average of spring peak flows increased by 12 and 3%, respectively, under RCP4.5 and RCP8.5. Finally, the spring maximal discharge increased only under RCP8.5 by 5%.

The comparison of Q_{max} associated with RPs from CPs between the historical period and PP3 (**Table 3**), for the three RCPs, revealed that no significant evolution in the future was



observed under RCP2.6, with the confidence intervals merging (**Figure 6**). However, Q_{max} values from the future period were slightly higher than values from the historical period under RCP4.5 and RCP8.5 for PP2 and PP3, with non-superimposing intervals (**Figure 6** and **Supplementary Figure Appendix SB2**). This was consistent with the previous result from HI class 5 showing that the more pessimistic the RCP, the higher the increase of Q_{max} . The same applies to the deviations, reaching 4 mm d^{-1} from the same RP under RCP8.5. The lines in **Figure 6** represent the fitted piecewise functions (see **Eq. 8**) on Q_{max} series from the available CPs per RCP (**Table 1**).

Under RCP4.5, those from the future period were higher than the ones from the historical period and the regime change in RC appeared earlier by about 2.7 and 1.8 months, with the corresponding discharge Q_{RC} increasing by 0.3 and 0.4 mm d^{-1} , respectively, in PP2 and PP3 (**Table 7**). Regarding RCP8.5, the regime change appeared earlier by 2.5 and 2.9 months, with the corresponding Q_{RC} values both increasing by 0.5 mm d^{-1} , respectively, in PP2 and PP3. The discharge at the break point from the historical period was reached 4 months earlier by the end of the 21st century, i.e. twice every 3 years instead of once every 2 years.

Figure 7 shows the evolution of the QRP05, QRP10, QRP20 and QRP50 values per RCP and sub-period in the projected time.

The associated errors were established using the confidence intervals in the mean distribution function of Q_{max} from CPs. The results showed that no QRP in RCP2.6 varied from the historical period to the future, with the confidence intervals intersecting each other. In PP1, RCP8.5 induced no significant change due to intersecting confidence intervals, whereas RCP4.5 led to an increase in QRP20 and QRP50, respectively, by 0.3 and 0.5 mm d^{-1} . In PP2, neither RCP4.5 nor RCP8.5 induced significant changes in any QRP. In PP3, results showed that under RCP4.5 and RCP8.5, the four QRPs each increased by about 1 mm d^{-1} . There was no significant difference between RCP4.5 and RCP8.5 with regard to intersecting confidence intervals.

DISCUSSION

Impact of the Use of Climate Projections in SIDRA-RU on Hydrological Indicators

In this study, the future of subsurface drainage was assessed feeding the subsurface drainage model SIDRA-RU with climatic projections, resulting in climatic chronicles for a historical period (1975–2004) and a projected period (2006–2100). For HI classes 1–5, we showed that HI_{CP_HIST} and HI_{REF} trends were not

TABLE 6 | Results of Mann–Kendall tests regarding the trends from the first five HI classes—“Temporality of the drainage season”, “Water content in the storage”, “Drained water balance”, “Temporal flow dynamics” and “Annual flood events” in the projected future per RCP. Per HI, the p value corresponding to the rolling mean was supported by the p values from the tests on the confidence interval (lower limit; upper limit). The trends were defined as the relative deviation between a mean historical value from 1975–2004 and a mean future value from 2071–2099. Only trends with significant p values were calculated.

Hydrological Indicators		RCP2.6		RCP4.5		RCP8.5	
Families	Variables	Trend	Rolling mean p value (lower limit; upper limit)	Trend	Rolling mean p value (lower limit; upper limit)	Trend	Rolling mean p value (lower limit; upper limit)
1. Temporality of the drainage season	Beginning of the drainage season	=	0.113 (0.021; 0.385)	=	0.232 (0.355; 0.267)	⇒	0.26 (0.036; 0.554)
	Durability of the drainage season	=	0.374 (0.463; 0.196)	-3%	0.002 (0.027; <0.001)	-7%	<0.001 (<0.001; <0.001)
2. Water content in the storage	Wet index	=	0.552 (0.974; 0.208)	=	0.368 (0.487; 0.250)	-5%	<0.001 (<0.001; <0.001)
	Drought index	=	0.469 (0.385; 0.781)	+50%	<0.001 (<0.001; <0.001)	+85%	<0.001 (<0.001; <0.001)
3. Drained water balance	Annual drained water balance	=	0.060 (0.035; 0.222)	=	0.012 (0.064; 0.006)	=	0.024 (0.211; 0.001)
	Summer drained water balance	=	0.905 (0.174; 0.373)	+12%	0.001 (0.001; 0.001)	-9%	<0.001 (<0.001; <0.001)
4. Temporal flow dynamics	Dry period	=	0.894 (0.951; 0.934)	+2%	0.001 (<0.001; 0.013)	+6%	<0.001 (<0.001; <0.001)
	Recharge period	=	0.223 (0.358; 0.211)	-16%	0.005 (0.006; 0.004)	=	0.017 (0.003; 0.067)
	Wet period	=	0.832 (0.934; 0.737)	=	0.052 (0.284; 0.015)	-7%	<0.001 (<0.001; <0.001)
	Recession period	=	0.067 (0.177; 0.090)	=	0.057 (0.567; 0.011)	=	0.024 (0.119; 0.005)
5. Annual flood events	Number of annual flood events	=	0.542 (0.443; 0.860)	=	0.743 (0.447; 0.945)	=	0.197 (0.986; 0.038)
	Annual cumulative duration in flood	=	0.621 (0.294; 0.871)	=	0.145 (0.086; 0.262)	=	0.857 (0.601; 0.336)
	Annual maximal discharge	+2%	0.002 (0.001; 0.004)	+9%	0.002 (0.005; <0.001)	+10%	<0.001 (<0.001; <0.001)
	Annual average of peak flows	+4%	<0.001 (<0.001; <0.001)	+12%	0.001 (0.002; 0.001)	+16%	<0.001 (<0.001; <0.001)

The trends if the p-values are significant (p-values < 0.005).

significantly different and were strongly correlated (95%) in most cases. Although the analysis of HI class 6 revealed a possible bias in Eq. 8, it remained negligible as the regime change appeared almost at the same time and with the same specific discharge. Moreover, the QRP05, QRP10, QRP20 and QRP50 values from CPs were similar to those obtained using the observed SAFRAN reanalysis. Consequently, we assumed that bias correction of the indicators was not required to interpret their response under climate change; they have been used for long-term perspectives.

However, some of the deviations observed in extreme values between the HI_{CP_HIST} and the HI_{REF} , especially for the annual drained water balance, can be questioned. First, we can assume that the deviations arise from the method used to calculate the distribution function of each HI_{CP_HIST} . The latter was extracted from the mean of the distribution functions of the concerned HI from the 12 CPs studied in the historical period under RCP8.5 (Table 1). Consequently, the associated standard deviation was reduced compared to the reference, thus affecting extreme values. Second, the deviations arise from the modelling chain (Figure 2). There are different sources of uncertainty from the estimation of the greenhouse gas emission scenarios to the calculation of HIs, and each stage has its own contribution to the total uncertainty (Vidal et al., 2016; Lemaitre-Basset et al., 2021). The determination of these contributions is a key point in climate change impact studies (Hattermann et al., 2017), which requires further analyses. Third, the deviations arise from the

use of the SIDRA-RU model to simulate future discharges. One limitation of the study is that the specific impact of the SIDRA-RU model on the studied HIs was not analysed. One way to analyse it is to extract hydrological indicators from observation (HI_{OBS}) and to compare them to the HI_{REF} on the historical period. However, due to significant gaps in observed hydrological data to generate HI_{OBS} , we were unable to perform such a study. Another way joins the above-mentioned uncertainty analysis from climate components of the modelling chain, consisting of using several hydrological models to relatively establish the contribution of the used hydrological model to the total uncertainty. This will be done in a further study.

Evolution of Subsurface Drainage Under Climate Change

Our results showed that 2–10 out of the 17 HIs studied here were modified under climate change depending on the RCP considered. This change gradually increased as the RCP became more severe and the HIs that changed in a less severe RCP also changed in a more severe RCP: two under RCP2.6 (≈12% of HIs), six under RCP4.5 (≈35% of HIs), 10 under RCP8.5 (≈59% of HIs). This suggests that a very optimistic policy (RCP2.6) does not enable the prevention of all the changes but strongly reduces the number and the

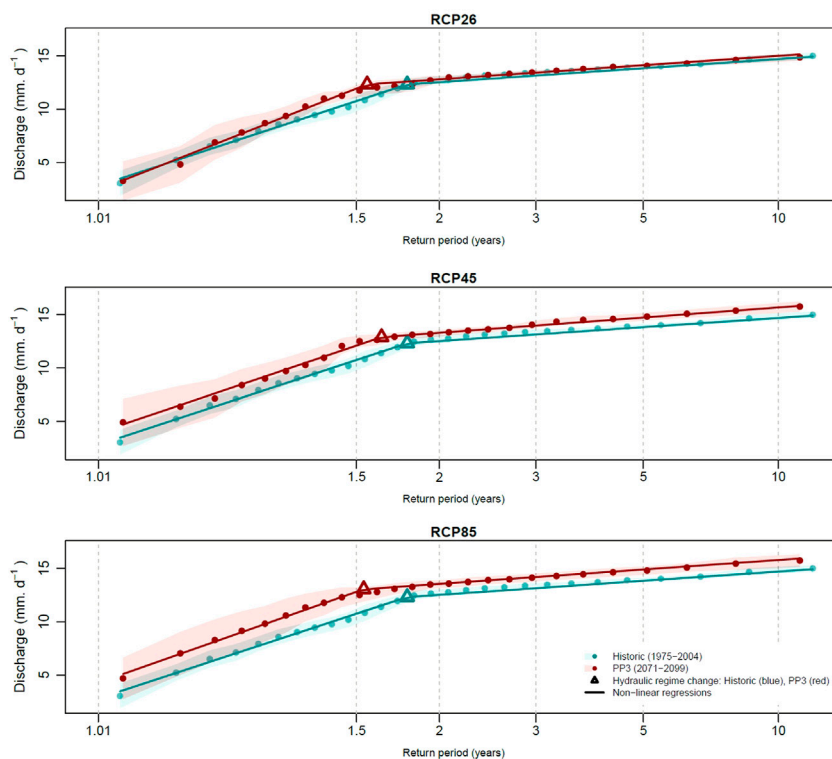


FIGURE 6 | Comparison of the specific discharges Q_{spe} (here Q_{max}) associated with return periods RPs between the projected future period 2071–2099 (red points) and the projected historical period 1975–2004 (blue points), for the three RCPs. Each scatter plot was calculated from the mean of the distribution functions of Q_{max} from the available CPs according to the corresponding RCP (Table 1), supported by a confidence interval at 95%. The equation $Q_{spe} = f(RP)$ was fitted as a piecewise function (see Eq. 8) by non-linear regression, for all cases.

TABLE 7 | Evolution of the return period in the regime change RP_{RC} and the corresponding discharges Q_{RC} in the projected future for the three RCPs. The deviations (Dev.) for RP_{RC} and Q_{RC} were only calculated when the deviations from the historical period to the future were significant, i.e. when the confidence intervals were not superimposed close to the regime change in the graphics.

RCP	Projected Time		RP_{RC} (y)	Dev. In RP_{RC} (months)	Q_{RC} (mm.d ⁻¹)	Dev. In Q_{RC} (mm.d ⁻¹)
	Historic					
			1.8	—	12.5	—
RCP 2.6	PP1		1.6	—	12.5	—
	PP2		1.7	—	12.6	—
	PP3		1.5	—	12.3	—
RCP 4.5	PP1		1.6	—	12.4	—
	PP2		1.6	-2.7	12.8	0.3
	PP3		1.6	-1.8	12.9	0.4
RCP 8.5	PP1		1.7	—	12.8	—
	PP2		1.6	-2.5	13	0.5
	PP3		1.5	-2.9	13	0.5

magnitude of changes. On the contrary, a policy that does not tackle climate change (RCP8.5, “Business as usual”) causes a larger number of changes in the present-day subsurface drainage in *La Jaillière*. First, changes appear in flood discharge. Second, the saturation in the storage is impacted, as is the hydraulic saturation of the drainage network. Finally, the temporal flow dynamics and the complete temporality of the drainage season are modified.

The HI showing the strongest change was the drought index, remaining stable under RCP2.6 but increasing significantly from 50% (+20 days) to 85% (+34 days), respectively, under RCPs 4.5 and 8.5 by 2100. This behaviour can be related to the evolution of the other HIs, such as the wet index and the length of the drainage season, both decreasing by approximately 5% under RCP8.5. Similarly, the temporal flow dynamics in the plot was modified, the dry period was longer while the wet period was shorter, by

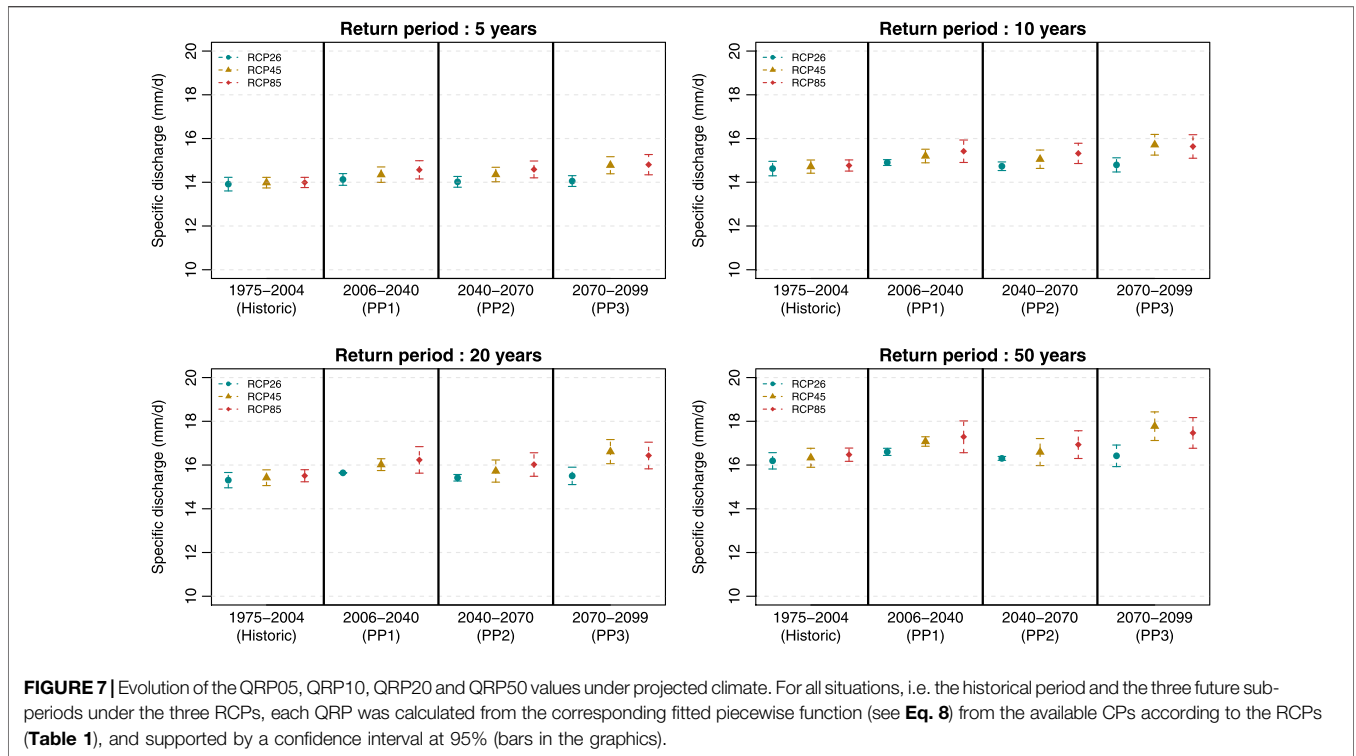


FIGURE 7 | Evolution of the QRP05, QRP10, QRP20 and QRP50 values under projected climate. For all situations, i.e. the historical period and the three future sub-periods under the three RCPs, each QRP was calculated from the corresponding fitted piecewise function (see **Eq. 8**) from the available CPs according to the RCPs (**Table 1**), and supported by a confidence interval at 95% (bars in the graphics).

approximately 6% under RCP8.5. These decreases represented an annual 15–30-days longer period of drought. Finally, the summer drained water balance under RCP8.5 by 2100 was reduced by about 10%. All these changes show that drought events are stronger according to how severe the RCP is.

On the other hand, our results showed an intensification of flood events. Indeed, the annual maximal discharge and the annual mean of peak flows were the only HIs to significantly change under all RCPs, with the strongest changes under a pessimistic RCP. Under RCP2.6, this increase was weak, approximately 2–4%. However, it was greater than 10% under RCP8.5. The intensity of flood events influences the regime of hydraulic saturation in the drainage network. Under RCP4.5 and RCP8.5, the regime change appeared 2–3 months earlier, and the same discharge appeared 4 months earlier compared to the historical period. Consequently, under these new conditions, the sustainability and effectiveness of the drainage network could be partially compromised. Moreover, the faster saturation of the drainage network may generate an earlier runoff, which is harmful for crops if the drainage system is less effective in reducing peak flow impacts.

Furthermore, the flood-specific discharge, i.e. the QRP05, QRP10, QRP20 and QRP50, potentially increasing by 2100 by approximately 0.5–1 mm d⁻¹ under RCP4.5 and RCP8.5, respectively, may enhance this effect. Considering that the risk of nitrate leaching during a flood event is highly correlated with the value of the peak flood flow, the increase in the peak flood flow may increase nitrate leaching, especially in winter (Billy et al., 2011; Tournebize et al., 2017). Pesticides, for their part, are less affected by the increase of peak flow in winter, and mostly depend

on the interaction between the beginning of the drainage season, which did not change, and the pesticide application (Branger et al., 2009; Delcour et al., 2015). This pertains to most pesticides except for the most mobile ones, e.g. glyphosate or isoproturon (Le Cor et al., 2021). Moreover, despite RCP8.5 being the most harmful RCP for water quality in winter due to nitrate leaching, our results showed that the frequency of spring flood events increased according to how optimistic the RCP was, changing from three events every 5 years to four to five events every 5 years and being more damaging for pesticide leaching during the main periods of use. The risk is weaker as the scenario becomes pessimistic, most likely correlated with the aforementioned reduction in summer discharge. A pessimistic scenario has a positive effect on water quality in terms of pesticide leaching in spring.

Some HIs did not vary under climate change, such as the beginning of the drainage season. Despite observing stronger flood events, the annual number of flood events and the annual cumulative duration of floods did not change under any RCP, which reinforces the fact that each flood event is stronger under climate change without being longer. The transitory periods from the temporal flow dynamics, i.e. the recharge and recessions periods, did not change either. Although the statistical analyses showed that the recharge period increased under RCP4.5, the increase was approximately 16% (+2 days) which is marginal given the level of uncertainty involved in this analysis cascade. This is reinforced by the fact that the drought index increased while the beginning of the drainage season did not change by 2100, meaning that the soil profile is empty during a longer period but this does not influence its recharge.

Finally, the annual drained water balance does not change under any RCP. This result contradicts the findings from the cited literature that shows either an increase (Pease et al., 2017; Sojka et al., 2020) or a decrease (Dayyani et al., 2012; Mehan et al., 2019; Jiang et al., 2020; Golmohammadi et al., 2021) by the end of the 21st century. Yet, these studies are based on methodologies similar to the one used here: A hydrological model is fed with climate projections simulated from greenhouse gas emission scenarios, mostly RCP4.5 and RCP8.5, to generate future trends. Similarly, four out of these six studies used the DRAINMOD model, including studies with divergent results, excluding the hydrological model effect on the deviation. The deviation therefore more likely arises from the differences in site locations and local climatic or pedologic conditions. This reinforces the point supported in the Introduction about how difficult it is to generalize the impacts of climate change on subsurface drainage. The results of our study are thus limited to the *La Jaillière* site. To sum up, the results can be used to make general, qualitative projections for the future of subsurface drainage in French sites with similar pedoclimatic contexts, especially the silty sites. However, it is necessary to carry out more precise studies of other drained sites in France to more accurately quantify the impact of climate change on French subsurface drainage.

To summarize, our results show an increase in drought and flood events, which is congruent with the literature predicting an intensification of extreme events such as droughts over France (IPCC, 2014; Marx et al., 2018), and depending on the socio-environmental path our societies follow, some adaptive strategies will be urgently needed. However, the reader has to keep in mind that this study is based on significant assumptions. Among them, current crop is neglected assuming that it does not significantly influence the water balance (see *The SIDRA-RU model Section*). For example, if the purpose of a future study is to assess the impact of climate change on subsurface drainage from a more agronomic point of view, the SIDRA-RU model does not appear as the wisest choice. This fact limits the emerging interpretations from this study.

Relevant Hydrological Subsurface Drainage Indicators

The major contribution of this study is to introduce 17 HIs to characterize subsurface drainage hydrology. Some HIs already existing in the literature were kept such as the annual drained water balance (Dayyani et al., 2012; Pease et al., 2017; Awad et al., 2020; Wang et al., 2021). To improve the accuracy of the analysis of subsurface drainage, new HIs were set—the length and beginning of the drainage season (Henine et al., 2022); the temporal flow dynamics class distinguishing each annual flow type (Molenat et al., 2008; Humbert et al., 2015; Strohmenger et al., 2020); HIs regarding the water content in the soil storage; HIs dealing with flood events and the discharge specific to an RP (QRP)—which is an approximate version of the indicators currently used to study the impact of climate change on conventional hydrology (Neves et al., 2020; Henriksen et al., 2021). The whole set

of HIs allows us to provide farmers and decision-makers with potentially useful tools and information (Kobierska et al., 2020) on crop management, on the risk of pesticide and nitrate leaching and on the sustainability of drainage networks designs.

Each HI informs on a different aspect of subsurface drainage. By assessing the future of 17 HIs (partially independent) instead of one-by-one, our study provides a more integrative understanding of the agrosystem response to climate change. Our study highlights three major possible changes, obviously depending on the scenario that is followed: 1) the increasing intensity of flood events enhances the exposure of crops; 2) this reinforcement also degrades agricultural water quality by favouring the leaching of pollutants; 3) drought events are longer in summer. Considered separately, each of these changes might lead decision-makers to opt for an adaptive strategy that could be inadequate for the other changes. Therefore, understanding the overall and combined impacts of these changes provides a general view of the effects of climate change on subsurface drainage and offers guidance towards the most suitable strategies to adapt to various changes.

Possible Adaptations of Subsurface Drainage to Climate Change

In the “Business as usual” RCP 8.5, the most important changes concern the increase in annual maximal discharge and the peak flows from flood events. Considering the possibly underdimensioned network of *La Jaillière*, these changes may result in the faster saturation of pipes and lead to surface runoff potentially damaging current crops. Farmers need to be informed about the risks and potential adaptation options.

One solution to prevent damage to crops from floods is to resize the network by reducing the spacing between drains and/or increasing the pipe diameter. This would increase the rate of water infiltration into the soil, so that the pipes would be filled less quickly thereby reducing the risk of crop-damaging runoff (Broadhead and Skaggs, 1982; Mulqueen, 1998; Nijland et al., 2005). Moreover, this would also reduce the risk of pollutant leaching, which is experienced when high discharges occur shortly after the application period. However, this solution is very expensive and seems difficult to achieve. Drainage system installation on new drained sites may be designed according to these requirements but complementary studies are required to show their efficiency in context of climate change.

Setting retention ponds is another way to prevent damage to crops from flood events (Ferk et al., 2020), to soften the impacts of stronger rainfall events (IPCC, 2014). Set downstream of drained plots, these ponds may also help reduce the leaching of agrochemical inputs (Blankenberg et al., 2007; Tournebize et al., 2012; Tournebize et al., 2017; Vymazal and Březinová, 2015). To be efficient, e.g. cutting 50% of nitrate leaching, the ratio between the pond size and the area of the corresponding drained field must be greater than 2% (Tournebize et al., 2017). Finally, the retention ponds can be useful for saving the excess water in winter from rainfall or direct discharge from the surface drainage network, to be reused for the possible increasing

irrigation needs in view of the expected reduction in summer drained water balance under RCP8.5 (Watanabe, 2016; Awad et al., 2020). This strategy was tested in a French context similar to *La Jaillière* and the ratio of water from winter storage to summer represented 14% of the annual irrigation supplies (Tournebize et al., 2015). However, even if the French government provides financial support to promote and preserve wetlands (Tuffnell and Bignon, 2019), the construction of ponds requires either the allocation of part of the field area, involving a net loss in crop productivity for farmers, or dedicating an equivalent surface connected to the hydraulic network from the drained plot. Both solutions might be troublesome in terms of land management, and therefore such a strategy is inherently challenging to implement in practice.

Instead of adapting the hydraulic dimension of subsurface drainage to climate change, changing agricultural practices can be useful to adapt to its effects. Indeed, the rising frequency of stronger events in a shorter period may increase the number of days of tillage by farmers and then compel them to adapt their agricultural calendar, e.g. to bring forward spring crop sowing. Another solution could be to use crop varieties that grow better in a warmer climate that lead to a decrease in water-holding capacity in summer, e.g. spring crops with a shorter development period, or that are more resilient to flood events (Teixeira et al., 2018) instead of the usual winter wheat and corn. Similarly, shifting the current cultural strategy to more resilient strategies such as crop diversification (Nazir and Das Lohano, 2022) might help to adapt agriculture to the increase in flood events. Changes in tillage practices may also play a role in the adaptation of practices to climate change (Dairon et al., 2017). Further studies are required to validate such proposals. Other less invasive strategies might also be adopted to prevent subsurface drainage from climate change effects, such as numerical tools to help farmers make decisions according to a reasoned use of agrochemicals, e.g. targeting specific times of pesticide supply depending on the season (Lewan et al., 2009; Kobińska et al., 2020).

CONCLUSION

The aim of this study was to assess the possible future of subsurface drainage under climate change using a large and comprehensive range of hydrological indicators. We first verified that we could use the projected historical hydrological indicators as reference in our subsequent analysis without including critical bias. We then showed that the more pessimistic the scenario of climate change, the more significant the changes in the indicators—a trend that is amplified with time until 2100. Under RCP8.5, the “Business as usual” scenario, the most harmful expected changes are that: 1) the increasing intensity of flood events enhances crop exposure; 2) this reinforcement also deteriorates agricultural water quality by favouring pollutant leaching; 3) drought events are longer in summer. Subsurface drainage will still be usable by 2100, but following the adopted

environmental policy, some adaptation strategies are required to best deal with these changes and preserve the environment. However, although the *La Jaillière* pedology might be a relevant reference for the French drained areas, its climate is location-specific. Analysing the projection of the indicators presented here in other French drained sites subject to different climate conditions and defined by a greater variety of soil types is therefore necessary, thereby allowing for a broader assessment of the future of French subsurface drainage under climate change.

DATA AVAILABILITY STATEMENT

The datasets presented in this article are not readily available because climatic data for this work were provided by Météo-France and are not publicly available since they are part of a commercial product. Météo-France can freely provide them on request for research and non-profit purposes. Requests to access the datasets should be directed to Météo-France, contactmail@meteo.fr.

AUTHOR CONTRIBUTIONS

AeJ, GT, and JT conceptualized the work. AiJ provided the methodological guidelines on statistical tests. AeJ drafted the paper. GT, AiJ, PM, and JT all revised the paper and contributed to its analyses and discussions.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fenvs.2022.899226/full#supplementary-material>

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