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Recovery of soil microbial diversity and functions along a tropical montane forest disturbance gradient

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Logging and forest conversion are occurring at alarming rates in tropical forests. These disturbances alter soil microbial community structure and functions. While direct links between changes in soil properties, such as pH and microbial community structure are well established, the indirect effects of logging and forest conversion on soil microbial community structure and functions are poorly understood. We used a space-for-time substitution to investigate the changes in soil microbial diversity and functions across a forest recovery gradient in the tropical montane forests of northern Borneo. We used surface (top 5 cm) soil to assess soil physicochemical and microbial (nextgeneration DNA sequencing) properties, and standardized litterbags (Tea Bag Index) to assess litter decomposition and stabilization. Our results show that bacterial and fungal diversity increases with recovery time and reaches predisturbance levels between 60- and 80-years post-disturbance. Litter decomposition rate constants increased linearly with increasing bacterial and fungal diversity. Litter stabilization also increased linearly with fungal diversity, but was highest at intermediate levels of bacterial diversity. Our results provide insights on the effects of forest logging and conversion on soils and highlight the tight coupling between soil microbial diversity and soil functions in tropical montane forests.

KEYWORDS

16s RNA, ITS, diversity-function relationship, soil microbiome, litter decomposition, litter stabilization, tropical forests

1 Introduction

Understanding the complex links between disturbance regimes, biodiversity, and ecosystem functions is a central theme of ecological research. Disturbance regimes such as logging and forest conversion are specific concerns in tropical forests, which store the largest amount of living carbon (C) (Pan et al., 2013), account for over half of the global annual net primary production (Melillo et al., 1993), and are the most biodiverse ecosystems (Dirzo and Raven, 2003) in the world. An estimated 27.2 million ha were cleared from tropical forests between 2000 and 2005 (Hansen et al., 2008). Logging and forest conversion can have lasting effects on canopy structure (Okuda et al., 2003), microclimate (Meijide et al., 2018), soil pH (Waldrop et al., 2000), soil nutrient content (McGrath et al., 2001), C sequestration, productivity (Pinard and Cropper, 2000; DeFries et al., 2002), and multitrophic biodiversity (Barnes et al., 2014; Barlow et al., 2016). For example, logging and forest conversion can decrease soil C and nitrogen (N) content as well as soil pH (Waldrop et al., 2000; Templer P. et al., 2005; Alele et al., 2014). While logging and forest conversion can affect microbial community composition and microbial biomass (Kerfahi et al., 2014; Maharjan et al., 2017), little is known about its role in shaping diversity and functions of soil microbial communities.

Microbes such as bacteria and fungi make up a significant portion of the biodiversity of tropical forests (Whitman et al., 1998; Arnold and Lutzoni, 2007), and are major drivers of the belowground C cycle through their role in processes such as litter decomposition and stabilization (Schimel and Schaeffer, 2012). However, there is a lack of consensus on how soil microbial diversity changes as a function of land use gradients and ultimately, relates to shifts in ecosystem functions. For example, while some have found no difference in microbial diversity between logged and unlogged tropical forests (Lee-Cruz et al., 2013; Tripathi et al., 2016), others have found an increase in microbial diversity in post disturbance tropical and temperate forests (Wallander et al., 2010; Zhu et al., 2012) and along a tropical land-use intensification gradient (Mendes et al., 2015; Carvalho et al., 2016). Studies that have looked for relationships between microbial diversity and C processing have also found mixed results. For example, while one study showed little to no effect of microbial diversity on litter decomposition in a manipulative grassland experiment (Griffiths et al., 2000a), another found positive relationships between the two in a controlled laboratory experiment (Maron et al., 2018).

With an overall goal of understanding the post-disturbance trajectory of soil microbial community structure (fungal and bacterial diversity) and functions (litter decomposition and stabilization) in tropical ecosystems, we used a space-for-time replacement design to investigate forest regeneration in northern Borneo. Specifically, we asked: 1) *do bacterial and fungal diversity exhibit recovery trajectories following logging and forest conversion similar to soil physicochemical properties, and do they differ in their trajectories*? Changes in fungal diversity are often attributed to successional changes in the plant community (Zhu et al., 2012),

while bacteria are often more strongly linked to soil properties such as pH (Tripathi et al., 2012). We hypothesize that soil microbial diversity of tropical forests will increase with increasing recovery time, but that recovery time will be different for bacterial and fungal communities. 2) do litter decomposition and stabilization follow the same recovery trajectories as bacterial and fungal diversity following logging and forest conversion? Based on previous comparative studies (Ribolzi et al., 2017; Sarneel et al., 2020) we know that litter decomposition rates are greater in mature forests, but there is a lack of information about the post-disturbance trajectory of litter decomposition and stabilization. 3) are there links between microbial community structure (bacterial and fungal diversity) and functions (litter decomposition and stabilization)? Given the mixed responses, ranging from no relationship (Griffiths et al., 2000b) to a positive relationship (Maron et al., 2018) between microbial diversity and ecosystem functions in controlled experiments, we hypothesize that there will be high uncertainty in these relationships in our observational study.

2 Materials and methods

2.1 Study area

We selected five study sites along a forest recovery gradient (time since the last major disturbance ranged from 4 to 150 year) in the Tambunan District of central Sabah, Malaysia (5.71750°, 116.40055°) (Figure 1 and Table 1). Two of the sites, the Mahua Falls Forest (MaF) and the Malungung Forest (MuF), were located within the federally protected Crocker Range Park area which covers ~139,919 ha. MaF and MuF were heavily logged ~100 and ~70 years (respectively) prior to this study. The remaining three sites were in nearby privatelyowned lands within the Tambunan Valley. The oldest site, Angelo's Forest (AnF), had not been clear-cut/logged for over 150 years (150-500 years, personal communication with the landowner). The abandoned rubber plantation (ARP) was cleared, terraced, and planted 41 years ago; trees at this site were untapped at the time of the sampling. The Abandoned Agricultural Field (AAF) was previously cultivated for chili, cleared again in 2014, and left abandoned for 4 year prior to sampling. The study sites ranged in elevation from 870 to 1,150 m asl and experienced a similar regional climate with a mean annual temperature of 24.3°C (Supplementary Figure S1) and annual precipitation of ~1,968 mm. The underlying geology of this region consists mainly of Quaternary fluvial gravels and sands (Tjia, 2007), and the soils in this region are characterized as orthic acrisols, having low base saturation (<50%) and high clay content in the B Horizon (Chesworth, 2008).

2.2 Site land cover classification

To describe current land use surrounding our sites, we classified land cover within a 100-m radius (Figure 1B) of the



FIGURE 1

Description of the study sites and soil sampling scheme. Map of the study area (A) showing the location of the five study sites across a forest recovery gradient in Sabah, Malaysian Borneo (B). The average elevation of the sites was 934 m, mean annual air temperature (MAT) was 24.3°C, and mean annual precipitation (MAP) was 1968 mm. A 100-m buffer (yellow circle) was used around each site (yellow dot) to estimate percent forest cover using SENTINEL-2 imagery (B). Soil sampling scheme displaying a 45-m long transect with 10 sampling plots on an elevational gradient at each study site (C). Different sites including Angelo's Forest (AnF), Mahua Falls Forest (MAF), Malungung Forest (MuF), Abandoned Rubber Plantation (ARP), and Abandoned Agriculture Field (AAF) are displayed using different colors and symbols (B) and corresponding canopy photos (D–H) are displayed below. Canopy photos were taken using a Canon EOS 7D Mark II Digital SLR camera with an Altura Ultra Wide Angle Aspherical Fisheye Lens (8 mm f/3.0).

TABLE 1 Study site description including geographical location (latitude (Lat), longitude (Long) in decimal degree), altitude (m ± SE), slope (% ± SE),
dominant vegetative cover, recovery time (RT, year), and current/historical land use.

Site Name	Lat	Long	Altitude	Slope	Slope Aspect	Vegetation	RT	Land use current historical
Angelo's Forest (AnF)	5.713	116.335	900 ± 5	44 ± 19	WNW	Mixed dipterocarp forest	>150	Forested not logged or clear-cut for >150 year
Mahua Falls Forest (MaF)	5.797	116.406	1140 ± 2	78 ± 00	NNE	Mixed dipterocarp forest	100	Forested clear cut ~100 year ago, selectively logged until ~1983
Malungung Forest (MuF)	5.661	116.251	862 ± 2	26 ± 4	WNW	Mixed dipterocarp forest	70	Forested clear-cut ~ 70 year ago, selectively logged until 1993
Abandoned Rubber Plantation (ARP)	5.765	116.469	879 ± 2	26 ± 4	SSE	Rubber trees, roadside fern	41	Abandoned terraced and planted with rubber trees in 1977
Abandoned Agriculture Field (AAF)	5.766	116.470	891 ± 1	56 ± 11	SSW	Bamboo, mostly cleared	4	Abandoned cultivated for chili, cleared in 2014

transect center (Figure 1C) using a maximum likelihood supervised classification. Maximum likelihood supervised classification uses pixel color and pattern recognition aided by user-provided training samples to classify satellite imagery into land cover classes (Parece et al., 2014). We classified land cover as forest and non-forest in ArcMap (version 10.6.1, ESRI, Inc., © 1995–2018) using SENTINEL-2 satellite imagery. SENTINEL-2 imagery was downloaded from the Copernicus Open Access Hub (https://scihub.copernicus.eu/, last accessed 17 March 2019) and was selected due to its availability in our study area and high spatial resolution (10 m). Using all forested pixels within the 100-m radius buffer, we calculated percent forest cover for each site.

2.3 Field sampling

At each site, we collected the top 5 cm of soil (organic and mineral both) at 5 m intervals along a 45 m transect positioned along a slope gradient (Figure 1C; Table 1) during the peak growing season of June 2018. To sample maximum within-site variability, we collected two to five soil cores (corer diameter = 6.35 cm) within 1 m of each plot center to create one composite soil sample per plot (Figure 1C). Overall, 10 plots were sampled per site (sample size = 10 plots x 5 sites = 50) and each plot was a composite sample of 2-5 cores. Due to extremely rocky terrain at AnF, we were unable to use the 6.35 cm diameter soil auger and used a 2.54 cm diameter corer for soil sampling instead. We homogenized each composite sample and immediately preserved a 1 g subsample in RNALater stabilization solution (ThermoFisher Scientific). The preserved samples (n = 50) were shipped to the Molecular Research Laboratory (MR DNA, Shallowater, TX, United States) for microbial analysis within a week of sampling. In the remaining sample, we removed large roots, rocks, and undecayed litter. We then air-dried, ground, and sieved (2 mm) the soil to remove non-soil fractions; and shipped these samples to the Texas Tech University's soil laboratory for analysis of soil physicochemical properties. We measured surface (0-5 cm) soil temperature (digital soil thermometer probe, HANNA Instruments) in the field and collected additional surface soil samples in air-tight containers for assessing gravimetric soil water content (i.e., dried at 60°C until sample weight remained constant).

2.4 Soil physicochemical analysis

We measured soil pH using the saturated paste method (Salinity Laboratory Staff, 1954), characterized soil elements (aluminum, sulphur, potassium, manganese, iron, nickel, copper, zinc, lead) using portable x-ray fluorescence (pXRF) (Goff et al., 2020), estimated soil organic matter (SOM) concentrations using the loss-on-ignition method (Schulte and Hoskins, 2011), total C and N using the Dumas method (Dumas, 1831), and soil texture (i.e., % sand, silt, clay) using a Hydrometer particle size analysis (152H Hydrometer) (Gee and Bauder, 2018). Due to an insufficient sample size at some plots, we performed soil texture analysis on 28 soil samples out of 50. Sites fell within the sandy clay (AnF, MaF), clay (ARP, AAF), and clay loam (MuF) soil texture classifications (Supplementary Figure S2).

2.5 Litter decomposition and stabilization

Inherent differences in litter quality can affect decomposition rates following disturbance (Keuskamp et al., 2013). The Tea Bag Index (TBI) method employs a standardized litter bag technique to compare plant litter decomposition and stabilization across sites without the influence of litter quality, and focuses on identifying environmental factors that influence decomposition rates across sampling locations (see global data in Keuskamp et al., 2013). We buried 10 pairs of pre-weighed litter bags (i.e., nonwoven tetrahedron-shaped polypropylene tea bags, Lipton, Unilever) per site, one green tea [89% green tea, fast decomposing (labile) litter] and one rooibos tea [93% rooibos tea, slow decomposing (recalcitrant) litter], at 8 cm below surface. Fine mesh size (0.25 mm) allowed microorganisms and mesofauna to enter the bags, but excluded macrofauna (Setälä et al., 1996). After 72-77 day, we retrieved the teabags and manually removed debris and roots following the "Tea Bag Index: Scientific Protocol, 2016" (http://www.teatime4science. org/method/stepwise-protocol/).

We used the initial and final oven-dried (70°C for 48 h) mass to calculate the stabilization factor (*S*, unitless) and the decomposition rate constant (k, d^{-1}).

During litter decomposition process, a portion of the labile compounds is stabilized and becomes recalcitrant (Prescott, 2010). This is calculated from the green tea bags (labile litter), as its labile component is expected to be decomposed within the incubation period (Eq. 1).

$$S = 1 - \frac{a_g}{H_g} \tag{1}$$

where a_g is the decomposable fraction from green tea, and H_g is the hydrolysable fraction of green tea. Assuming *S* is equal for both tea types, *S* can then be used with the hydrolysable fraction of rooibos tea (H_r) to calculated a_r , the decomposable fraction of rooibos tea (Eq. 2).

$$a_r = H_r \left(1 - S \right) \tag{2}$$

The assumption that the environmental stabilization of labile material is independent of the relative size and composition of the hydrolysable fraction has been tested using alternate analyses in previous work (Keuskamp et al., 2013).

The decomposition rate constant (k) can only be estimated from the early stages of decomposition. Since the green tea is collected after its labile material is decomposed, k is calculated from the rooibos tea (recalcitrant litter), using a single time point assuming an exponential decay function (Eq. 3).

$$W(t) = ae^{-kt} + (1 - a)$$
(3)

where W(t) is the weight of the substrate after incubation time *t*, *a* is the labile and (1 - a) is the recalcitrant fraction of the litter. Additional details can be found in Keuskamp et al. (2013).

2.6 Soil microbial diversity

We received final operational taxonomic units (OTUs) that were assigned based on 97% similarity and followed the proprietary MR DNA analysis pipeline (Dowd et al., 2008) from the Molecular Research Laboratory (MR DNA, Shallowater, TX, United States). Briefly, the 16S rRNA gene (bacteria) and ITS1 region (fungi) were amplified, respectively. PCR primers 515, ITS1 and ITS2, along with HotStarTaq Master Mix Kit (Qiagen, United States) were used in 30 PCR cycles with the following steps: 94°C for 3 min followed by 30 cycles of 94°C for 30 s; 53°C for 40 s and 72°C for 1 min; and a final elongation step at 72°C for 5 min. Following amplification, DNA fragments were sorted by molecular weight and concentration using 2% agarose gel. The sorted DNA samples were then purified using an 1X AmpureXP beads and checked on Agilent High Sensitivity (HS) chip on Bioanalyzer 2100 and quantified on a fluorimeter by Qubit dsDNA HS Assay kit (Brand from MR DNA). The library was loaded onto the Illumina Platform for clustering and sequencing. Paired-End sequencing allows the template fragments to be sequenced in both the forward and reverse directions. Sequencing was carried out on an Illumina MiSeq following the manufacturer's guidelines and data were analyzed using the proprietary MR DNA analysis pipeline (Dowd et al., 2008). Sequences with <150 bp or with ambiguous base calls were removed. Briefly, sequences were denoised, operational taxonomic units (OTUs) were assigned based on 97% similarity, and chimeras were removed. Singletons and OTUs appearing in only one sample were removed from OTU tables. Archaea and mitochondrial or chloroplasts OTUs were removed from the 16S data and non-fungi OTUs from ITS1 data. In total, 31,261 bacterial and 0 fungal chimeric were removed from further downstream analysis. Final OTUs were assigned using BLASTn; cloned sequences were searched against a database derived from RDP-II and NCBI (www.ncvi.nlm.nih.gov, http:// rdp.cme.msu.edu). A total of 3,439,724 (68,794 ± 3,167 on average) high quality 16S sequences and 2,920,915 (58,418 ± 1,185) high quality ITS sequences for the 50 soil samples were analyzed. The average of 1,274 \pm 180 bacterial and 2,105 \pm 59 fungal OTUs at 97% identity cutoff were obtained for 16S and ITS sequences.

2.7 Statistical analyses

We calculated Chao1 bacterial and fungal diversity at the OTU level for each soil sample in R (R Core Team, 2017) using the package "SpadeR" (Chao et al., 2016) and used SigmaPlot 14.0 (Systat Software Inc., San Jose California, United States) to display and fit the statistical models to the bivariate relationships between recovery time, soil microbial diversity, and litter decomposition and stabilization. All regression models were evaluated for strength (simpler models with greater R^2 were

selected) and significance (only significant ($p \le 0.05$) models were selected).

We used the "randomForest" package (Cutler and Wiener, 2018) in R to evaluate the relative importance of recovery time as a predictor of soil microbial diversity (bacterial and fungal) and soil functions (litter decomposition and stabilization) when compared with other predictors including soil physicochemical properties, aboveground vegetation cover, and environmental variables. Random forest is an ensemble machine learning technique that creates a series of uncorrelated regression trees from random subsamples of the data. These trees decide as a committee which variables are the most predictive, and rank the variables in order of their ability to increase the mean square error when removed from the model (Breiman, 2001). Before running the random forest analysis, we used the package "corrplot" in R to identify and remove highly correlated variables (|r| > 0.70) (Wei and Simko, 2017) (Supplementary Figure S3), retaining the most ecologically relevant and/or proximal predictor variables (Dormann et al., 2013). Through this process iron, C, manganese, sulfur, potassium, C:N, and zinc were removed from the random forest analysis. Although recovery time and pH were highly correlated ($|\mathbf{r}| = 0.82$), we chose to leave both variables in the analysis because of their unique ecological contributions to the model. To ensure that the importance given to each predictive variable would not be biased by the presence of the two correlated variables, we ran the random forest analyses again removing each of the two correlated variables one at a time and then simultaneously. We found reductions in the overall variance explained by the model, however the top predictive variables for each model were the same, excluding the variable(s) removed. Therefore, we concluded that the correlation of the variables had little effect on our model and chose to include them both in the final analyses. One extreme fungal diversity outlier was removed from the random forest and regression analyses but is displayed in figures and noted with an asterisk. We built 200 trees (ntrees = 200) for each run and sampled 3-4 variables (Mtry = 3-4, depending on the number of predictor variables) with replacement for each tree. Finally, we took the top four predictors and used regression analysis to further explore the trends.

3 Results

3.1 Surrounding forest cover

Surrounding forest cover varied among the study sites, with 100% forest cover measured at MaF, 99.4% at MuF, 88.4% at AnF, 60.3% at AAF, and 40% at ARP. While surrounding forest cover was not significantly related to recovery time (Supplementary Figure S4A), higher surrounding forest cover was found at the three naturally



FIGURE 2

Relationship between recovery time (year) and soil physicochemical properties (A-F) across a forest recovery gradient: Angelo's Forest (AnF), Mahua Falls Forest (MaF), Malungung Forest (MuF), Abandoned Rubber Plantation (ARP), and Abandoned Agriculture Field (AAF). Smaller symbols in the background of each plot represent raw data points (n = 10 per site), while larger symbols represent site means. Sample size for % sand varied from 1 to 10 depending on the site [AnF (n = 1), MaF (n = 6), MuF (n = 3), ARP (n = 8), AAF (n = 10] due to logistical difficulties of collecting enough soil. Bars represent the standard error of mean and gray lines represent 95% confidence intervals. All statistical models are significant at $p \le 0.05$.

forested sites than the agricultural (AAF) and converted (ARP) sites and surrounding forest cover had a cooling effect on daytime soil temperatures (Supplementary Figure S4B).

3.2 Relative influence of recovery time on soil properties

3.2.1 Physicochemical properties

There were a range of responses between soil physicochemical properties and recovery time (Figure 2). Soil pH (p < 0.001, Adj $R^2 = 0.67$, Figure 2A), sulfur (p = 0.05, Adj. $R^2 = 0.28$, Figure 2F), and % sand (p < 0.001, Adj. $R^2 = 0.60$ Figure 2B) increased linearly while aluminum decreased linearly (p < 0.001, Adj. $R^2 = 0.43$, Figure 2D) with increasing recovery time. SOM (p < 0.001, Adj. $R^2 = 0.36$) showed a peaked response at ~100 years following recovery (Figure 2E). Total N content increased linearly (p < 0.001, Adj. $R^2 = 30$) showed a saturated response with recovery time (Supplementary Figure S5); taken together this generated a peaked response of C:N at ~80 years following recovery (Figure 2C).

3.2.2 Bacterial diversity

The random forest analysis explained 67.08% of the total variance in Chao1 bacterial OTU diversity (hereafter, bacterial diversity) with recovery time, soil pH, percent forest cover, and SOM as the top four predictors (Figure 3A). Bacterial diversity peaked at a soil pH of 5 (p < 0.001, Adj. $R^2 = 0.63$, Figure 3B), and linearly increased with both recovery time ($p \le 0.01$, Adj. $R^2 = 0.56$, Figure 3C) and surrounding forest cover (p < 0.001, Adj. $R^2 = 0.43$, Figure 3D). Bacterial diversity levels saturated around 6% SOM (p < 0.001, Adj. $R^2 = 0.29$, Figure 3E).

3.2.3 Fungal diversity

The random forest analysis explained 18.69% of the total variance in Chao1 fungal OTU diversity (hereafter, fungal diversity) with recovery time, percent forest cover, soil pH, and soil aluminum as the top four predictors (Figure 4A). Fungal diversity increased with recovery time, saturating at ~100 years (p < 0.001, Adj. $R^2 = 0.34$, Figure 4B), but showed weaker relationship with recovery time as compared to bacterial diversity ($R^2 = 0.63$, Figure 3B). Surrounding forest cover was the second most important predictor of fungal diversity, with a positive linear relationship (p < 0.001, Adj. $R^2 = 0.21$,



time (RT, year), surrounding forest cover (SFC, %), soil organic matter (SOM, %), daytime soil temperature (ST, °C), N (%), gravimetric water content (GWC, g g⁻¹), Ni (mg g⁻¹), Al (mg g⁻¹), Cu (mg g⁻¹), and Pb (mg g⁻¹). Sites including Angelo's Forest (AnF), Mahua Falls Forest (MaF), Mahungung Forest (MuF), Abandoned Rubber Plantation (ARP), and Abandoned Agriculture Field (AAF) are displayed using different colors and symbols. Smaller symbols in the background of each bivariate plot represent raw data points (n = 10 per site), while larger symbols represent site means. Bars represent the standard error of mean and gray lines represent 95% confidence intervals. All statistical models are significant at $p \le 0.05$.

Figure 4C). Fungal diversity peaked at a soil pH of ~5 ($p \le 0.01$, Adj. $R^2 = 0.17$, Figure 4D) and had a weak inverse relationship as a function of soil aluminum concentration ($p \le 0.01$, Adj. $R^2 = 0.09$, Figure 4E).

3.3 Relative influence of recovery time on ecosystem functions

3.3.1 Litter decomposition rate constant (k)

The random forest analysis explained 15.08% of the overall variance in *k* with N, pH, fungal diversity, bacterial diversity, and recovery time as the top five predictors (Figure 5A). *k* increased linearly with increasing soil N (p < 0.05, Adj. $R^2 = 0.22$, Figure 5B), fungal ($p \le 0.05$, Adj. $R^2 = 0.22$, Figure 5D), and bacterial diversity (p < 0.001, Adj. $R^2 = 0.19$, Figure 5E). *k* peaked at a soil pH of ~5 (p < 0.05, Adj. $R^2 = 0.13$, Figure 5C) and

increased with recovery time, saturating at ~100 years (p < 0.05, Adj. $R^2 = 0.52$, Supplementary Figure S7 in appendix).

3.3.2 Litter stabilization factor (S)

The random forest analysis explained 47.09% of the overall variance in *S* with recovery time as one of the top five predictors along with surrounding forest cover, bacterial diversity, and daytime soil temperature (Figures 6A–E). *S* increased with surrounding forest cover, saturating at ~80% (p < 0.005, Adj. $R^2 = 0.63$, Figure 6B), increased linearly with recovery time (p < 0.01, Adj. $R^2 = 0.14$, Figure 6D), and showed a peaked response to bacterial diversity and increasing daytime soil temperature. *S* peaked at intermediate bacterial diversity (p < 0.001, Adj. $R^2 = 0.32$, Figure 6C) and between 20°C–21°C (p < 0.001, Adj. $R^2 = 0.23$, Figure 6E). All three naturally regenerating forests had higher mean *S* estimates (AnF = 0.239 ± 0.014,



MaF = 0.245 \pm 0.011, and MuF = 0.242 \pm 0.017) than the abandoned rubber plantation (S = 0.070 \pm 0.025) and agriculture (S = 0.201 \pm 0.017) sites.

3.4 Influence of microbial diversity on ecosystem functions

Looking at each site individually, we didn't find significant relationships among bacterial diversity (k in Supplementary Figure S8, S in Supplementary Figure S9, in appendix) or fungal diversity (k in Supplementary Figure S10, S in Supplementary Figure S11, in appendix) and ecosystem functions. However, significant patterns emerged at the biome scale when all sites across the disturbance gradient were evaluated together. For example, k increased linearly with increasing bacterial diversity (Figure 5E, also shown as Supplementary Figure S8F, in appendix) and fungal diversity (Figure 5D, also shown as Supplementary Figure S10F, in appendix). S showed a peak response with bacterial diversity (Figure 6C, also shown as Supplementary Figure S9F, in appendix), while it had a weak positive linear response to fungal diversity (Supplementary Figure S11F, in appendix).

4 Discussion

Logging and forest conversion continue to craft the tropical montane landscape, yet we still have a limited understanding of its consequences to the global C cycle. In this study, we show strong links between forest regeneration and soil physicochemical properties, microbial diversity, and C cycling



in a tropical montane region of Sabah, Borneo. We found that many of the physical and biological variables covary across this forest regeneration gradient and take 80–100 years to reach predisturbance levels; this highlights the vulnerability of these systems to forest conversion and logging.

4.1 Relative influence of recovery time on soil properties

4.1.1 Physicochemical properties

We compared soil properties at five sites across a forest recovery gradient (space-for-time substitution), with sites ranging from four to over 150 years of post-disturbance regeneration. As expected, recovery time significantly affected the soil physicochemical properties (Figure 2). Soil pH had the most significant change along the gradient, increasing from 3.7 at the youngest site (AAF) to 5.4 at the oldest site (AnF) (Figure 2A). This pH range is similar to the range reported by others across a forest disturbance gradient in Sabah, Borneo (Tripathi et al., 2016), and is consistent with findings in other tropical systems that reported lower pH after forest logging and conversion (Templer P. H. et al., 2005; Alele et al., 2014; Tripathi et al., 2016). Lower pH levels measured in recently recovered sites could be attributed to recent harvesting. For example, plants are capable of altering the pH of the rhizosphere and bulk soil through ion uptake. The removal of soil cations (e.g., calcium and magnesium) can lead to soil acidification when plant material is removed from a site during harvest (Rengel, 2003). Previous logging and cultivation of the sites may have resulted in a decrease in soil pH, with soil pH gradually increasing over time in older forests with organic matter decomposition (Allen, 1985).

We found that sulfur (Figure 2F) increased linearly, while SOM (Figure 2E), C content (Supplementary Figure S5B, in

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appendix), and C:N (Figure 2C) showed peaked responses with recovery time. Our results are consistent with prior work that reported higher concentrations of total soil C, N (Dinesh et al., 2003), and SOM (Templer P. H. et al., 2005) in forested sites compared to both agricultural and plantation sites, mainly attributed to a decrease in quantity and quality of plant residues returned to the soil. Though there is no information on how sulfur changes along a forest regeneration gradient, soil sulfur concentrations can be tightly coupled with total soil C content and SOM levels (Kirkby et al., 2011). Our results reflect this nutrient coupling, with the soil sulfur concentration increasing proportionally with soil C across the forest recovery gradient. Previous studies have observed significant losses of soil nutrients due to leaching (Brouwer and Riezebos, 1998), biomass removal (Martinelli et al., 2000; Olander et al., 2005), and altered nutrient allocation (Swinfield et al., 2020) associated with logging. The increased soil C pools and nutrients along this forest regeneration gradient might be due to the soil nutrient stocks rebuilding to predisturbance levels.

Given that rates of organic matter accumulation typically slow over time (Silver et al., 2000), the observed peaked relationships were not surprising. Percent SOM has been found to return to pre-disturbance levels within 40–50 years of succession in tropical forests (Brown and Lugo, 1990), however, others have observed this return in less than a decade (Templer P. H. et al., 2005). The recovery time required to accumulate organic matter in the soil is variable and can be affected by vegetation (i.e., quantity and quality of litter inputs), abiotic and biotic soil properties (i.e., soil texture and microbial activity), and land use history (Post and Kwon, 2000). Additionally, inherent differences among our forested sites in terms of their soil textures and terrains (the oldest site, AnF, had rocky terrain) could have resulted in the high variability and lower amounts of SOM at AnF as compared to MaF. Similar to soil C and SOM, C:N peaked at ~100 years and subsequently declined at the oldest site (Figure 2C), due to the combined effects of total soil C decreasing and total soil N increasing (Supplementary Figure S5, in appendix). Nitrogen fixing organisms, particularly microbial plant symbionts, are considered an important source of new N in tropical soils (Gehring et al., 2005). Some late-successional plant species have higher leaf N content, with correspondingly high soil N content, as compared with early-successional forests (Yan et al., 2006). Greater N content in litter could result in enhanced decomposition rates, driving soil C and SOM levels down.

Soil texture also changed significantly along the recovery gradient; % sand increased until around 100-years postrecovery (Figure 2B). Fine particles can be lost from soil following disturbance events such as tillage due to erosion and runoff, leaving a higher percentage of larger particles such as sand (Brady and Weil, 2017). Our results contradict these findings and could be a result of several factors including variation in soil parent material and natural erosion patterns. While we identified no effects of recovery time on soil microclimate, we found a decrease in daytime soil temperature with increasing forest cover (Supplementary Figure S4, in appendix). Removal of the forest canopy during logging events increases daytime soil temperature (Yashiro et al., 2008), and reforestation can have significant surface cooling effects (Novick and Katul, 2020).

4.1.2 Bacterial diversity

Our results show that bacterial diversity increased linearly with recovery time (Figure 3C), which is consistent with prior studies that found an increase in microbial diversity with stand age (Wallander et al., 2010; Zhu et al., 2010). While previous work has reported a peak in bacterial diversity at ~40-years post-disturbance in a subtropical secondary forest (Zhu et al., 2010), our study showed no decline in bacterial diversity with recovery time. We also found that the surrounding forest cover is a significant predictor of bacterial diversity (Figure 3D).

The lowest bacterial diversity was found in the landscapes dominated by human activities (ARP and AAF), evident from the existence of social trails, grazing (AAF), greater amount of exposed soil, and the proximity to inhabited areas. The three naturally regenerated forested sites (MuF, MaF, and AnF) were relatively undisturbed and surrounded by a large tropical forest buffer. Previous research suggests that anthropogenic disturbances such as conversion to pasture or plantation can induce biotic homogenization (Devictor et al., 2008; Lôbo et al., 2011) of the soil bacterial community, leading to the loss of some more endemic taxa and to the broader dispersal of the existing taxa within a site (Rodrigues et al., 2013; Alele et al., 2014). This pattern is also supported by our findings that the naturally forested sites showed the highest variation along their transects for several soil properties, including pH, SOM, and total C content. Associated losses of endemic taxa are likely to be better quantified by the Chao1 diversity index, used in this study, as it is weighted by rare taxa.

Soil pH was the top predictor of bacterial diversity (Figure 3B), consistent with previous work in the tropics (Tripathi et al., 2012; Tripathi et al., 2016), temperate forests, grasslands (Fierer and Jackson, 2006; Bahram et al., 2018; Malik et al., 2018), and alpine meadows (Shen et al., 2019). Bacterial diversity peaked at a pH of ~5; this is a consistent pattern in other forested systems (Tripathi et al., 2012; Kaiser et al., 2016), but peak diversity occurred at a pH of ~7 in one of these studies (Tripathi et al., 2012). Bacteria are known to experience physiological constraints at extreme pH levels and optimum soil pH varies among bacterial taxa (Tripathi et al., 2012). However, the intracellular pH of bacteria is often close to neutral (Slonczewski et al., 1981; Krulwich et al., 2011), and a more extreme external pH may place stress on bacterial cells (Lauber et al., 2009). Bacterial diversity may also be indirectly affected by pH through its effect on the availability of various soil nutrients (Lauber et al., 2009; Tripathi et al., 2012).

SOM is the fourth most important predictor of bacterial diversity, showing a saturating response of bacterial diversity at ~6% SOM (Figure 3E). Soil bacterial communities rely on the decomposition of organic matter for energy (Trivedi et al., 2013), and SOM has previously been identified as one of the main drivers of soil bacterial diversity on a global scale (Tian et al., 2018).

4.1.3 Fungal diversity

We found that fungal diversity increased significantly with recovery time up until ~100 years (Figure 4B), which is consistent with prior studies that found an increase in microbial diversity with stand age (Wallander et al., 2010; Zhu et al., 2010). This increase in microbial diversity over time may reflect spatial heterogeneity in accumulation of SOM (Klein et al., 1995) or increased variability in litter quality and quantity as a result of changes in plant diversity (Zhu et al., 2010) as forests mature. Soil disturbance and homogenization from tillage and logging can negatively impact fungal communities (Brinkmann et al., 2019). As mentioned earlier, disturbed sites, AAF and ARP, showed lower variation along the transect for several soil parameters including pH, SOM, and total C content. These findings suggest that the sites incurred resource homogenization during forest clearing, cultivation, and continued anthropogenic disturbance. Many fungal species (e.g., namely ectomycorrhizal fungi), are plant-symbionts and often exhibit strong host tree preference (Tedersoo et al., 2008). Thus, the alteration of tree species composition and abundance can also have direct effects upon fungal community structure and diversity (Gao et al., 2013). While we didn't document forest composition, we did find that the current

surrounding forest cover (i.e., cover by a mature tree canopy) is a significant predictor of fungal diversity, with AAF and ARP exhibiting the lowest levels of fungal diversity, and mature forests (AnF and MaF) exhibiting higher fungal diversity (Figure 4C).

Future studies in this region, relating plant community cover, composition, and root structure to the relative abundance of fungal taxa, would aid in establishing these linkages. The random forest analysis explained only 18.69% of variation in fungal diversity. The lack of strong direct ties between fungal diversity and soil physicochemical properties, in contrast to bacterial diversity that showed strong links to soil properties and recovery time, could be tied to the fungal scale of operation; fungal hyphae can access a larger volume of soil, making fungi more able to survive in a variety of soil conditions, while bacteria are more heavily impacted by local substrate and nutrient availability (Zheng et al., 2019). Of the physicochemical properties evaluated we found that soil pH (Figure 4D) and soil aluminum concentrations (Figure 4E), which were negatively correlated with each other $(|\mathbf{r}| = 0.66)$, were the strongest predictors of fungal diversity. Although similar results have been found in previous studies (Zheng et al., 2019) these relationships are often weaker (Lauber et al., 2008; Zheng et al., 2019) than those between soil pH and bacterial community composition and diversity (Fierer and Jackson, 2006; Lauber et al., 2009; Tripathi et al., 2012; Tripathi et al., 2016).

4.2 Relative influence of recovery time on ecosystem functions

4.2.1 Litter decomposition rate constant (k)

Our results suggest that soil N (Figure 5B) and pH (Figure 5C) were significant predictors of k. Because microbially mediated litter decomposition is generally thought to be nutrient limited (Crews et al., 1995), it is likely that increases in nutrient levels in these ecosystems increased k. These results are consistent with a previous meta-analysis (Knorr et al., 2005) that found that N addition often increased decomposition of high-quality litter. k was also related to recovery time, saturating quickly at ~40 years (Supplementary Figure S7, in appendix). During litter decomposition, some labile plant tissues are microbially and biochemically transformed into recalcitrant, stable compounds. The diversion of labile plant tissues, which would otherwise decompose, into stabilized organic matter builds upon humus pools, thereby sequestering more C in the soil (Prescott, 2010).

4.2.2 Litter stabilization factor (S)

S increased linearly with recovery time (Figure 6D) and showed a saturating response to surrounding forest cover (Figure 6B), reaching a maximum at ~ 90% cover in the older

forests (AnF and MaF). Previous studies have found that the presence/absence of woody vegetation (Sarneel et al., 2020) can affect C turnover. It is likely that increased forest cover and live root mass in older sites protected the litter bags from decomposition and led to higher stabilization. The three forested sites (AnF, MaF, and MuF) had the highest S followed by AAF, and ARP. The youngest site (AAF) had relatively low tree cover, but had a dense herbaceous understory that could have provided some physical protection to the soil, an effect that has been previously documented in fallow land in tropical montane regions (Ribolzi et al., 2017). The ARP site on the other hand, had a considerable amount of exposed soil both beneath the tree cover and in the surrounding area. Possible mechanisms for litter stabilization include the formation of a specialized microbial community with distinct stabilizing capabilities (Keiser et al., 2011), the addition of a more preferential microbial energy source through root exudates (Sparling et al., 1982), the increased input of recalcitrant root tissue, the physical protection of organic matter by the formation of soil aggregates (Prescott, 2010), and reduced erosion and leaching (Ribolzi et al., 2017).

4.2.3 Comparison with the global data

Our results match the global pattern (Keuskamp et al., 2013) of the relationship between the litter decomposition rate constant (k, d^{-1}) and the litter stabilization factor (S, unitless) (Supplementary Figure S6 in appendix). The litter decomposed faster in the two old growth forests (AnF = 0.032 and MaF = 0.036 d⁻¹), and k values were comparable to the other lowland tropical forests from global data. All three naturally regenerating forests showed relatively high S (AnF = 0.239, MaF = 0.245, and MuF = 0.242) similar to mixed and birch forests from global dataset. Interestingly, the younger natural forest (MuF) closely matched the agriculture (AAF) site with intermediate k (MuF = 0.020, AAF = 0.016 d⁻¹) and relatively high S (MuF = 0.242, AAF = 0.201). Contrastingly, the rubber plantation (ARP) had a low S (0.070) similar to the lowland tropical forest, coupled with an intermediate k (0.024 d⁻¹), placing it near the geothermal wet grassland ecosystem within the global dataset.

4.3 Influence of microbial diversity on ecosystem functions

Soil microbial communities are one of the main drivers of soil C cycling through processes such as litter decomposition and stabilization (Schimel and Schaeffer, 2012). Previous studies connecting microbial diversity to C cycling have found mixed results. For example, some have found little to no effect of microbial diversity on litter decomposition (Griffiths et al., 2000b), while others have found positive relationships

between the two (Maron et al., 2018). When evaluating sites individually we found no significant relationships between microbial diversity metrics and ecosystem functions (Supplementary Figures S8-S11, in appendix). However, patterns emerged at the biome scale. For example, we found that k increased linearly with increasing fungal diversity (Figure 5D, also shown as Supplementary Figure S10F, in appendix) and bacterial diversity (Figure 5E, also shown as Supplementary Figure S8F, in appendix). These results suggest a structural metric, such as microbial diversity, can be used as an indicator of functional shifts in the tropical montane systems. Our findings contribute insights into the current debate about diversity versus functional redundancy (Giller et al., 1996); they suggest that reduced microbial diversity results in a reduction of functional capabilities in the soil. Changes in microbial growth efficiency, or the ability of the microbial community to incorporate substrates into biomass and byproducts, can alter litter transformation rates (Six et al., 2006). Our data matched the global patterns of litter decomposition and stabilization, but both processes were decoupled unlike the global data which showed an inverse relationship between k and S (Keuskamp et al., 2013) (Supplementary Figure S6, in appendix).

Conclusion

Considering the rapid rate of deforestation in the tropics and microbial role in nutrient cycling, it is critical to investigate how logging and forest conversion affect soil microbial structure and functions in these systems. Our results suggest that logging and forest conversion significantly affect soil microbial diversity and can have prolonged effects on C cycling in tropical montane forests. Considering the potential global impacts of land-use on the terrestrial C cycle in one of the most productive biomes on Earth, these results have wide implications for land management and biodiversity conservation.

Data availability statement

The datasets presented in this study can be found in online repositories. The names of the repository/repositories and accession number(s) can be found below: https://datadryad. org/stash, https://doi.org/10.5061/dryad.0vt4b8h1t.

Author contributions

KN and JM designed the study; MF collaborated with private landowners for land access and selected sites across a gradient of

land use disturbance; RS, JM, AR, JG, and KN conducted field work; DW conducted soil physicochemical analysis; RS analyzed the data with assistance from KN; RS and KN wrote the manuscript with assistance from JM, and JS; all authors approved the paper.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/fenvs.2022. 853686/full#supplementary-material

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