



# Relief of Phosphate Limitation Stimulates Methane Oxidation

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Aquatic ecosystems such as shallow lakes and wetlands are important emitters of the greenhouse gas methane (CH<sub>4</sub>). Increased phosphorus (P) loading is expected to increase CH<sub>4</sub> production in these ecosystems. This increased CH<sub>4</sub> production can potentially be mitigated by increased CH<sub>4</sub> oxidation, but how P availability affects methane-oxidizing bacterial (MOB) community composition and potential CH<sub>4</sub> oxidation remains to be tested. Here, we incubated MOB from sediments of four subtropical lakes of different trophic states for 7 days at different phosphate (PO<sub>4</sub><sup>3-</sup>) concentrations to determine the effects of P on MOB community composition and potential CH<sub>4</sub> oxidation. We measured CH<sub>4</sub> consumption daily and compared CH<sub>4</sub> oxidation during the exponential growth phase. Furthermore, we determined MOB community composition at the end of the incubations using qPCR of the *pmoA* gene. To test for differences in N and P uptake, we determined bacterial biomass N and P content. We found that increases in PO<sub>4</sub><sup>3-</sup> concentrations until 10 μM significantly increased CH<sub>4</sub> oxidation. PO<sub>4</sub><sup>3-</sup> also increased bacterial biomass P content, while N content was not affected. MOB community composition was not affected by PO<sub>4</sub><sup>3-</sup> but more strongly correlated to lake of origin, likely due to the short duration of the incubations. Our results show that PO<sub>4</sub><sup>3-</sup> can not only stimulate CH<sub>4</sub> oxidation indirectly through increased CH<sub>4</sub> production, but also directly by increasing MOB growth. Importantly, these effects only occur at low PO<sub>4</sub><sup>3-</sup> concentrations, indicating that at high nutrient loads the increased CH<sub>4</sub> oxidation will likely not mitigate the increased CH<sub>4</sub> production.

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## INTRODUCTION

Methane (CH<sub>4</sub>) is the second most important greenhouse gas, contributing to 16–25% of global warming (IPCC, 2014). Microbial processes in aquatic ecosystems, such as wetlands and shallow lakes, are the main source of this greenhouse gas, accounting for over 50% of global CH<sub>4</sub> emissions (Rosentreter et al., 2021). Currently, the nutrient status of many shallow lake ecosystems is changing because of eutrophication, which is expected to increase shallow lake CH<sub>4</sub> emissions (Aben et al., 2017; Davidson et al., 2018), mostly driven by phosphorus (P) and nitrogen (N) enrichment in water and sediment (Correll, 1998). While interactions between N cycling and CH<sub>4</sub> emissions have been well studied (Bodelier and Steenbergh, 2014), we still lack understanding about the effects of P on CH<sub>4</sub> emissions.

CH<sub>4</sub> cycling in shallow lakes is largely driven by the sediment microbiome. In the anoxic zone of the sediment, methanogens produce CH<sub>4</sub> (Conrad, 2007). Subsequently, a large part of the produced CH<sub>4</sub> is mitigated by methane-oxidizing bacteria (MOB), which consume CH<sub>4</sub> mostly aerobically, converting CH<sub>4</sub> into CO<sub>2</sub>. In freshwater ecosystems with high concentrations of alternative electron acceptors, such as nitrate and sulfate, CH<sub>4</sub> can also be converted anaerobically (Deutzmann, 2020). This study focuses on aerobic MOB, which belong to the Gammaproteobacteria (type I), Alphaproteobacteria (type II), and Verrucomicrobia, of which type I and type II MOB are dominant in most freshwater lakes. These MOB differ in their phylogeny, physiology, biochemistry, and morphology (Semrau et al., 2010). While both perform the same process—CH<sub>4</sub> oxidation—they prefer different environments. Type I MOB are most abundant in stable environments with high substrate availability while type II are most abundant in environments with disturbances such as heat waves and desiccation and low substrate availability (Ho et al., 2013). CH<sub>4</sub> oxidation by MOB is the sole biological CH<sub>4</sub> sink (Adamsen and King, 1993), and it is therefore important to know how MOB react to disturbances such as eutrophication.

The accumulation of phosphate (PO<sub>4</sub><sup>3-</sup>) in aquatic ecosystems can affect CH<sub>4</sub> oxidation. In a field study, higher potential CH<sub>4</sub> oxidation rates were found in sediments that had higher PO<sub>4</sub><sup>3-</sup> concentrations (Veraart et al., 2015). Increased [PO<sub>4</sub><sup>3-</sup>] is also hypothesized to indirectly affect CH<sub>4</sub> oxidation rates by changing the MOB community. It has been suggested that type I MOB thrive in eutrophic environments with high [PO<sub>4</sub><sup>3-</sup>] while type II MOB compete better in more oligotrophic environments with low [PO<sub>4</sub><sup>3-</sup>] (Ho et al., 2013). One of the reasons for this is that type II MOB generally harbor more enzymes which liberate P from organic molecules and have membranes with lower P content (Veraart et al., 2015). Changes in the dominant type of MOB can potentially also lead to changes in CH<sub>4</sub> oxidation rates (Reis et al., 2020; Nijman et al., 2021), thereby altering the mitigation capacity of the MOB community.

Only one study has addressed the direct relationship between [PO<sub>4</sub><sup>3-</sup>] and MOB functioning (Sawakuchi et al., 2021). In this study, surface water MOB from ice-covered lakes oxidized more CH<sub>4</sub> at 500 μg/L (~5 μM) PO<sub>4</sub><sup>3-</sup> than without PO<sub>4</sub><sup>3-</sup> in long-term incubations at 4 °C. However, the specific conditions of this experiment make it difficult to apply to MOB from other ecosystems. Most important, the dominant MOB in the ice-covered lake experiment are very rare outside of arctic regions (50.3% *Methylobacter tundripaludum* and 44.6% *Methylovulum psychrotolerans*). Second, the CH<sub>4</sub> oxidation rates were very low (2.5–29.4 μmol CH<sub>4</sub> L<sup>-1</sup> d<sup>-1</sup>), compared to many other MOB communities, such as those in lake sediments (e.g. Shelley et al., 2015). Finally, the question arises whether higher PO<sub>4</sub><sup>3-</sup> concentrations would further increase potential CH<sub>4</sub> oxidation rates, and up to which concentration PO<sub>4</sub><sup>3-</sup> could still be the limiting growth factor for MOB.

Therefore, we still lack a good understanding of the direct relationship between [PO<sub>4</sub><sup>3-</sup>] and CH<sub>4</sub> oxidation. Other studies of MOB, which showed effects of PO<sub>4</sub><sup>3-</sup> on CH<sub>4</sub> oxidation rates (Veraart et al., 2015) and on the type of MOB (Ho et al., 2013),

were based on correlative field observations, and thus only showed an indirect relationship between [PO<sub>4</sub><sup>3-</sup>] and MOB functioning. Since many biogeochemical variables are correlated, the relationship between [PO<sub>4</sub><sup>3-</sup>] and CH<sub>4</sub> oxidation found in these studies could also have been due to other variables, such as organic matter content (Czepiel et al., 1995) or N availability (Bodelier and Laanbroek, 2004). Furthermore, since increased [PO<sub>4</sub><sup>3-</sup>] stimulates lake CH<sub>4</sub> emissions, likely through enhanced CH<sub>4</sub> production (Aben et al., 2017; Davidson et al., 2018), oxidation may also be indirectly stimulated by PO<sub>4</sub><sup>3-</sup> due to higher CH<sub>4</sub> availability. However, considering that bacteria need P for growth, and especially type I MOB since they contain phospholipid-rich stacked membranes (Bodelier et al., 2009), it is likely that PO<sub>4</sub><sup>3-</sup> also directly affects MOB, especially at low PO<sub>4</sub><sup>3-</sup> concentrations.

Here, we investigated the effect of PO<sub>4</sub><sup>3-</sup> concentration on potential CH<sub>4</sub> oxidation and MOB community composition. We expected increases in PO<sub>4</sub><sup>3-</sup> concentration would increase potential CH<sub>4</sub> oxidation. Furthermore, we expected that increases in PO<sub>4</sub><sup>3-</sup> concentration would favor type I MOB over type II MOB.

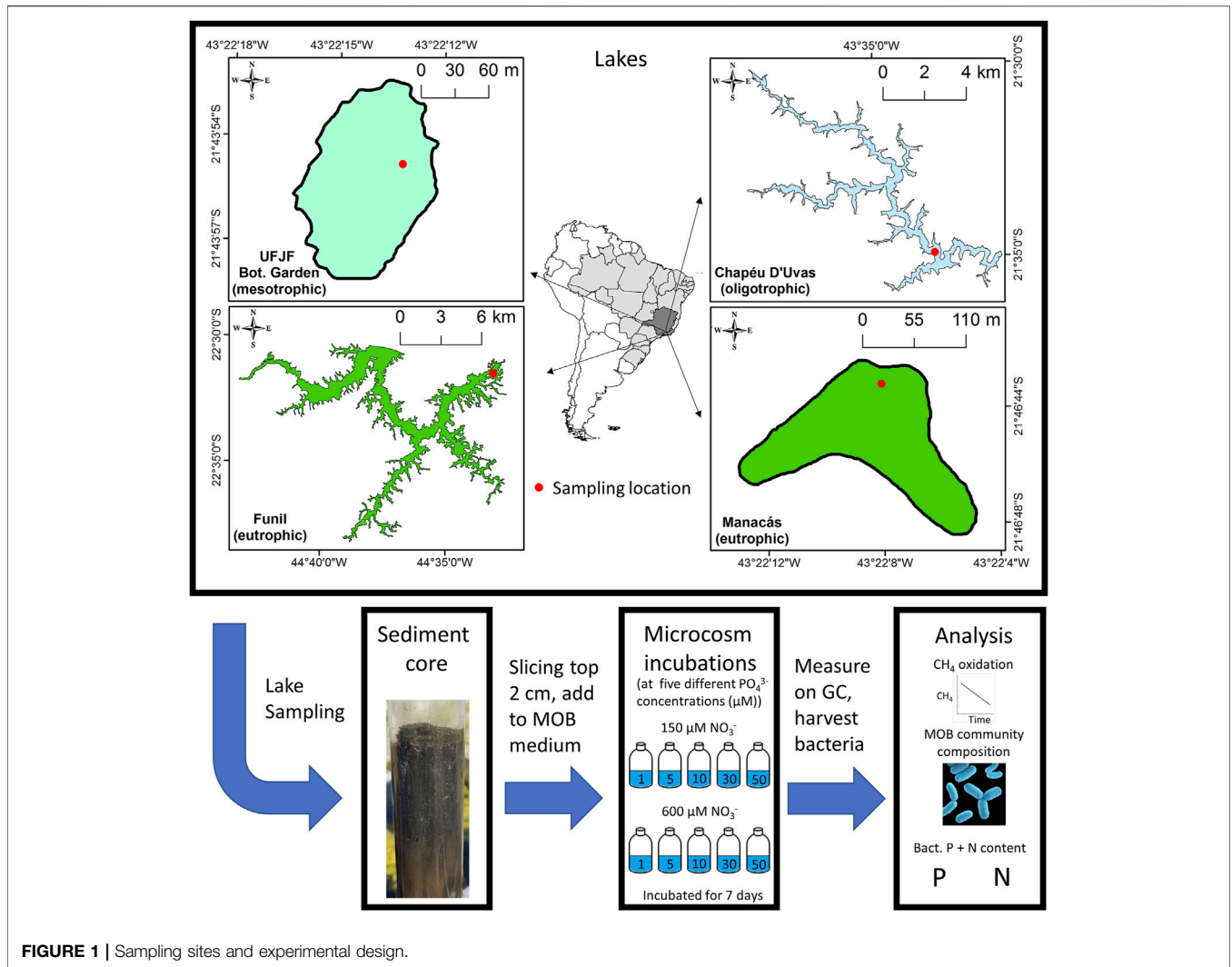
## METHODS

To test the effects of [PO<sub>4</sub><sup>3-</sup>] on potential CH<sub>4</sub> oxidation and MOB community composition, we incubated MOB from four different subtropical lake sediments in 48 microcosms at different [PO<sub>4</sub><sup>3-</sup>], analyzing CH<sub>4</sub> oxidation rates during exponential growth and final MOB community composition (Figure 1).

### Sediment Sampling and Inoculum Preparation

Four subtropical lake sediments in south-eastern Brazil were sampled in april 2019 to obtain different MOB communities for the microcosms (Supplementary Table S1). We sampled Funil lake (latitude, longitude: 2231'31''S, 44°33'1''W), Manacás lake (2146'43''S, 4322'8''W), Botanical Garden lake of Universidade Federal de Juiz de Fora (UFJF; 2143'55''S, 4322'14''W), and Chapéu D'Uvas lake (CDU; 2134'55''S, 4333'16''W). The lakes were chosen because of their different trophic state. Funil lake (Pacheco et al., 2015) and Manacás lake (Lawall et al., 2005) are eutrophic, UFJF botanical garden lake is mesotrophic (Nunes et al., 2021) and CDU lake is oligotrophic (Almeida et al., 2019). We took one sediment core from each lake with a Uwitech corer (ø 60 mm), of which we sliced off the top 2 cm (~56.5 ml), and used those for the microcosms. While one core is not enough for a representative picture of the community of a single lake sediment, the use of single cores from four different lakes allowed us to investigate the relationship between [PO<sub>4</sub><sup>3-</sup>] and CH<sub>4</sub> oxidation for MOB communities with different community composition.

The sediments were used to prepare inocula for the microcosms. The sliced sediment was first diluted with lake water from the corresponding lake to obtain a total volume of



150 ml, and homogenized using a blender. While blending could have lysed some cells, the low speed of the blender was likely not enough to significantly affect the microbial community. After mixing, the suspended sediment was pipetted into 15 ml falcon tubes and centrifuged at 4,000 rpm for 5 min. Then, the supernatant, containing a large number of bacteria, was taken. The supernatant was filtered through 1.2  $\mu\text{m}$  glass fiber filters to remove most protozoa and remaining sediment particles, and the filtrate was used as inoculum for the microcosms.

## Microcosms

We used microcosms to test how  $[\text{PO}_4^{3-}]$  affected MOB growth and community composition. We prepared Nitrate Mineral Salts (NMS) medium specific for MOB (Whittenbury et al., 1970, ATCC 1306) with 10  $\mu\text{M}$  copper and five different  $\text{PO}_4^{3-}$  concentrations (1, 5, 10, 30, and 50  $\mu\text{M}$ , added as  $\text{KH}_2\text{PO}_4$  and  $\text{NaHPO}_4$  in a 1:2.5 ratio). We also used two different  $\text{NO}_3^-$  concentrations (150 and 600  $\mu\text{M}$ , added as  $\text{KNO}_3$ ) to control for effects of N and N:P ratio, and included a negative

control (0  $\mu\text{M}$   $\text{PO}_4^{3-}$  and 0  $\mu\text{M}$   $\text{NO}_3^-$ , no inoculum added) and positive control (1 mM  $\text{PO}_4^{3-}$  and 1 mM  $\text{NO}_3^-$ ), resulting in twelve microcosms per lake. The  $\text{PO}_4^{3-}$  and  $\text{NO}_3^-$  concentrations were chosen based on comparable concentrations in mesotrophic to eutrophic freshwater sediments (Geurts et al., 2008; Veraart et al., 2015). We ran one series of microcosm experiments per lake, yielding four replicates per combination of  $[\text{PO}_4^{3-}]$  and  $[\text{NO}_3^-]$ . Microcosms were prepared by adding 57 ml NMS medium and 3 ml inoculum of the respective lake to 120 ml serum bottles. Bottles were closed with butyl stoppers and capped, after which 10% of the total headspace was replaced with  $\text{CH}_4$  (purity grade 4.5). This left an oxygen ( $\text{O}_2$ ) concentration of 18.9%. We incubated the microcosms in the dark for 7 days at 20°C, while stirring continuously. Because MOB use  $\text{O}_2$  in a 2:1 ratio to oxidize  $\text{CH}_4$ , there was not enough  $\text{O}_2$  in the microcosms to completely consume all  $\text{CH}_4$ . However, we found that in all microcosms, the  $[\text{CH}_4]$  still decreased after the time of exponential growth, indicating that  $\text{O}_2$  was not yet depleted. It was also unlikely  $\text{CH}_4$  decreased because of anaerobic oxidation of methane (AOM), as described below.

Since the microcosms had a much higher  $O_2$  concentration than is present in sediments (Wang et al., 2014), we did not aim to show how sediment MOB would respond to increased  $[PO_4^{3-}]$ . Rather, the aim was to show how model MOB communities would respond to different  $PO_4^{3-}$  concentrations. Using an initially high  $[O_2]$  had two advantages. First, it allowed the MOB to reach exponential growth because they could consume the majority of  $CH_4$  before  $O_2$  would become limiting. Second, it prevented AOM and methanogenesis. For AOM, 2%  $O_2$  decreased  $CH_4$  oxidation of ‘*Candidatus Methylobacterium oxyfera*’ by 60% within 22 h without recovery (Kampman et al., 2018), while 5%  $O_2$  caused ‘*Candidatus Methanoperedens nitroreducens*’ to be outcompeted by *Methylobacter*, an aerobic MOB (Guerrero-Cruz et al., 2018). Also, while some methanogens can survive  $O_2$  exposure (Peters and Conrad, 1995), these only showed activity after 15 days without  $O_2$ . Therefore, it is unlikely there was AOM or methanogenesis in the microcosms.

To determine  $CH_4$  consumption rates, we measured  $CH_4$  concentration in the headspace directly after  $CH_4$  addition and once per day afterwards. The  $CH_4$  concentration was measured by taking 1 ml of the headspace and injecting into a Bruker 450-GC (Bruker Technologies, Australia) equipped with a flame ionization detector. The concentration was calculated by comparing to a standard curve. Furthermore, 1 ml of medium was taken daily for MOB analysis, although these samples were not used because the number of cells was too low for DNA extractions.

## CH<sub>4</sub> Oxidation During Exponential Growth

To estimate  $CH_4$  oxidation rates, we calculated the rate at which  $CH_4$  decreased per day per ml of medium, accounting for removal of water and air from the microcosms (supplementary methods). Surprisingly, there was no  $CH_4$  oxidation in the microcosms of mesotrophic UFJF botanical garden lake. Little data on this lake are available thus far, but potential MOB inhibitors such as low pH, high  $NH_4^+$  and humic substance concentration have not been observed. However, one possibility is the lakes’ high oxygen concentration (11 mg/L, Nunes et al., 2021) may have prevented most  $CH_4$  production, exemplified by very low  $CH_4$  emissions (3.39 mg  $CH_4$   $m^{-2}$   $day^{-1}$  September 2021, personal communication Nasário and Barros), and thereby the development of an active MOB community, which remains to be experimentally confirmed. Because of the lack of  $CH_4$  oxidation, we removed those microcosms from our analysis. For the remaining microcosms, we used the time of exponential growth to compare  $CH_4$  oxidation at different  $PO_4^{3-}$  concentrations, based on the day with the largest decrease. All incubations had  $>0.75\%$   $CH_4$  (12  $\mu$ M  $CH_4$  in the water) left after this day, which is not yet limiting (Shelley et al., 2015). Exponential growth occurred 1 day earlier for the eutrophic sediments than for the oligotrophic sediment (Supplementary Figure S4).

## MOB Community Composition

To evaluate MOB community composition, we harvested the bacteria from the microcosms after 7 days. Thirty ml of the final

medium was filtered through 0.2  $\mu$ m filters, using half for DNA extractions. Filters of the positive control were used to test the DNA extraction method, while the negative control was not filtered as there was no bacterial growth. Four filters were accidentally lost during transportation and the remaining 26 filters were used for qPCR. Bacteria were scraped off the filters using sterile cell scrapers (Greiner Bio-One 541070) and DNA was extracted with the DNeasy PowerSoil kit (Qiagen 12888-100) according to manufacturer’s instructions with one modification: bead-beating was performed using a Powerlyzer for 45 s at 2,500 rpm. We then determined copy number of type Ia (*Methylobacter*, *Methylomonas*), type Ib (*Methylocaldum*, *Methylococcus*) and type II (*Methylosinus*, *Methylocella*, *Methylocystis*) MOB by qPCR using different primers for the *pmoA* gene, according to Kolb et al. (2003) with some small modifications (Supplementary Table S2), using iQ SYBR® Green Supermix as mastermix. All samples and standards were measured in triplicate on a BioRad iQ5 Multicolour Real-Time PCR detection system (Vers. 2.0, BioRad, Gothenburg Sweden), on which we checked the melt curves.

## Bacterial Biomass N and P Content

To investigate whether the  $PO_4^{3-}$  concentrations changed the N and P content of the MOB, we filtered 10 ml of the final medium through two 0.7  $\mu$ m glass fiber filters per microcosm. Filters were weighted before and after filtering, and bacterial N and P concentration was determined from the bacterial biomass retained on the filters after potassium persulfate digestion (Carmouze, 1994). N concentration in bacterial biomass was measured on a total organic carbon analyzer, equipped with a TNM-1 total nitrogen module (TOC-V CPN, Shimadzu), whereas P in bacterial biomass was measured spectrophotometrically (Lambda 365 UV/VIS, Perkin Elmer) by the ascorbic acid method (Mackereth et al., 1979). Filters were pre-combusted (550°C for 4 h) before filtration and oven dried before analysis (60°C for >24 h). The filtration method for microbial biomass nutrient contents using 0.7  $\mu$ m pore size filters may not account for all the microbial community - smaller cells may pass through filters - but a previous study showed an average 80% retention of bacterial community using this pore size (They et al., 2017). Because the main available carbon substrate for microbial growth was  $CH_4$ , and because we used NMS medium specific for MOB, we assumed that MOB were the predominant bacteria, and hence bacterial biomass N and P content was closely related to MOB N and P content. Also, because most  $CH_4$  was consumed in all incubations, similar total growth of MOB had taken place, indicating that N and P content per cell were related to total N and P content of the bacterial community. However, these results should be interpreted with caution, because P could also have formed suspended particles, increasing P content of the filters.

## Statistical Analysis

Statistical analyses were performed in R 3.6.3 (R Core Team, 2019). To analyze how  $[PO_4^{3-}]$  affected  $CH_4$  oxidation during exponential growth, we calculated potential  $CH_4$  oxidation rates at different  $[PO_4^{3-}]$  (supplementary methods). Next, we

determined which model fit the data best to compare effects of  $[\text{PO}_4^{3-}]$  on  $\text{CH}_4$  consumption for the different lakes. We tested a linear regression, a nonlinear least-squares regression using the *nls* function, and a segmented regression using the *segmented* function (Muggeo, 2003). Based on a comparison of AIC values (Supplementary Table S3), the best-fitting model was the nonlinear least-squares regression, which we further analyzed. We also compared the nonlinear least-squares regression for 150 and 600  $\mu\text{M}$   $\text{NO}_3^-$  separately (Supplementary Figure S1), but since the differences between these models were small, we finally used one model combining all data.

To determine at which  $\text{PO}_4^{3-}$  concentrations  $\text{CH}_4$  oxidation rates during exponential growth significantly differed from each other, we determined the confidence intervals (CIs) of  $\text{CH}_4$  oxidation at different  $\text{PO}_4^{3-}$  concentrations and calculated the *p*-value at which CIs of different concentrations overlapped.

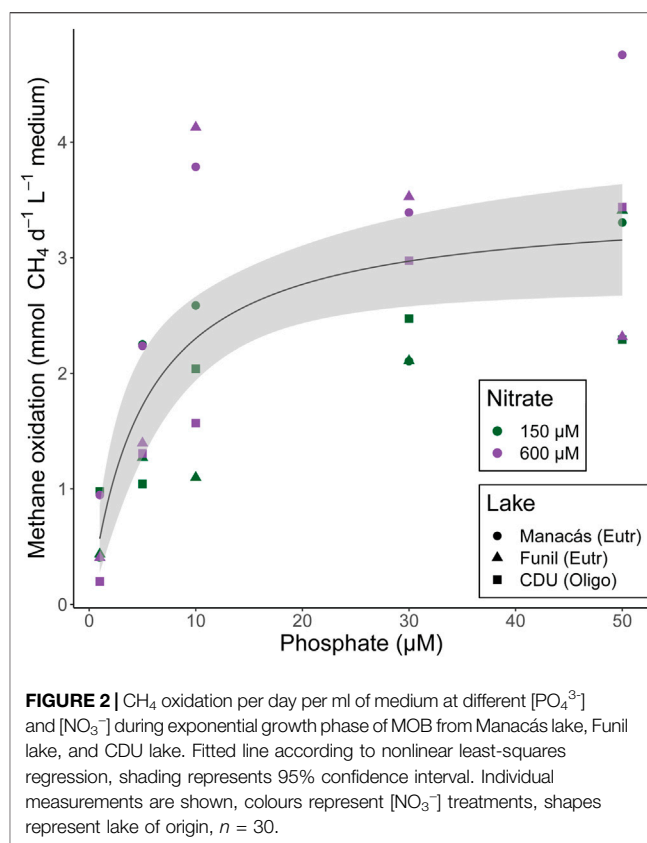
Relative abundance of different MOB types was calculated by dividing the copy number per MOB type by the copy number of all MOB. To achieve normality, the proportions of MOB type Ia and II were logit transformed while the type Ib proportion was log transformed. Linear models were used to test the effects of  $[\text{PO}_4^{3-}]$ ,  $[\text{NO}_3^-]$ , and lake of origin on relative abundance of different MOB types, and the *emmeans* function (Lenth et al., 2020) was used for post-hoc between-lake comparisons. We also inspected the relationship between N:P ratio and proportions of MOB (Supplementary Figure S2), but found no clear patterns.

Bacterial P content was log-transformed to achieve normality and linear models were used to test the effects of  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  on bacterial N content and log-transformed bacterial P content. A similar model comparison to  $\text{CH}_4$  oxidation showed the best fitting model was a linear regression rather than a nonlinear least-squares regression or segmented regression (Supplementary Table S4). Similar to  $\text{CH}_4$  oxidation, we compared the models for 150 and 600  $\mu\text{M}$   $\text{NO}_3^-$  separately (Supplementary Figure S3), but finally decided to combine all data since the differences between the models were very small.

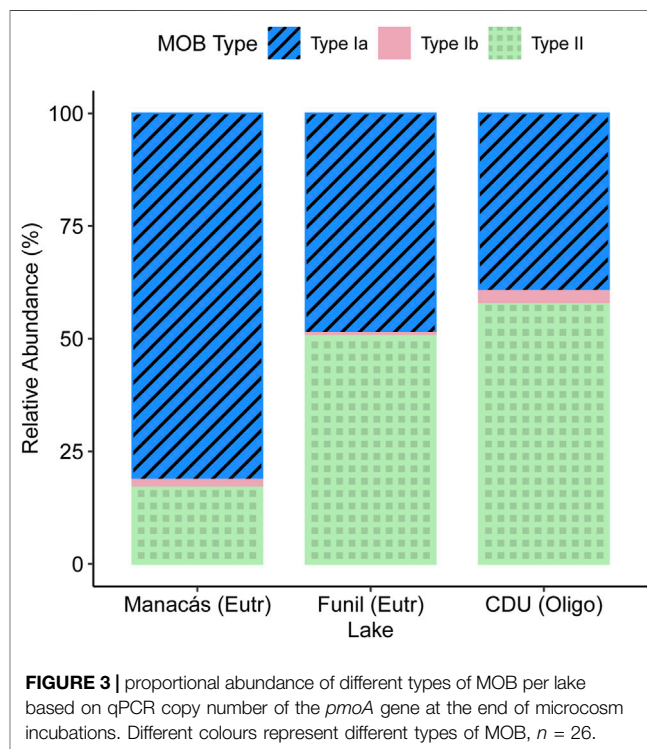
## RESULTS

To investigate the effects of  $[\text{PO}_4^{3-}]$  on model MOB communities, 30 microcosms from three different lakes (Manacás, Funil, CDU) were incubated at five different  $\text{PO}_4^{3-}$  concentrations (1–50  $\mu\text{M}$ ), while a fourth lake (UFJF botanical garden) yielded no MOB growth and was therefore removed from the analysis.

$\text{CH}_4$  oxidation increased with  $[\text{PO}_4^{3-}]$  (Figure 2). Interestingly,  $\text{PO}_4^{3-}$  stimulated  $\text{CH}_4$  oxidation during exponential growth mostly at low  $[\text{PO}_4^{3-}]$  up to 10  $\mu\text{M}$ , as shown by the good fit of the nonlinear least-squares regression to the  $\text{CH}_4$  oxidation data (Supplementary Table S3). Analysis of confidence intervals showed that  $\text{CH}_4$  oxidation at 1  $\mu\text{M}$   $\text{PO}_4^{3-}$  was significantly lower than at all higher concentrations (5  $\mu\text{M}$ :  $p = 0.0036$ ; 10, 30, and 50  $\mu\text{M}$ :  $p < 0.0001$ ) while  $\text{CH}_4$  oxidation at 5  $\mu\text{M}$   $\text{PO}_4^{3-}$  was significantly lower than at 30  $\mu\text{M}$  ( $p = 0.0051$ ) and 50  $\mu\text{M}$   $\text{PO}_4^{3-}$  ( $p = 0.0039$ ), and slightly lower at 10  $\mu\text{M}$   $\text{PO}_4^{3-}$  than at 50  $\mu\text{M}$   $\text{PO}_4^{3-}$  ( $p = 0.048$ ). There was no effect of  $[\text{NO}_3^-]$  or lake of origin on  $\text{CH}_4$  oxidation.

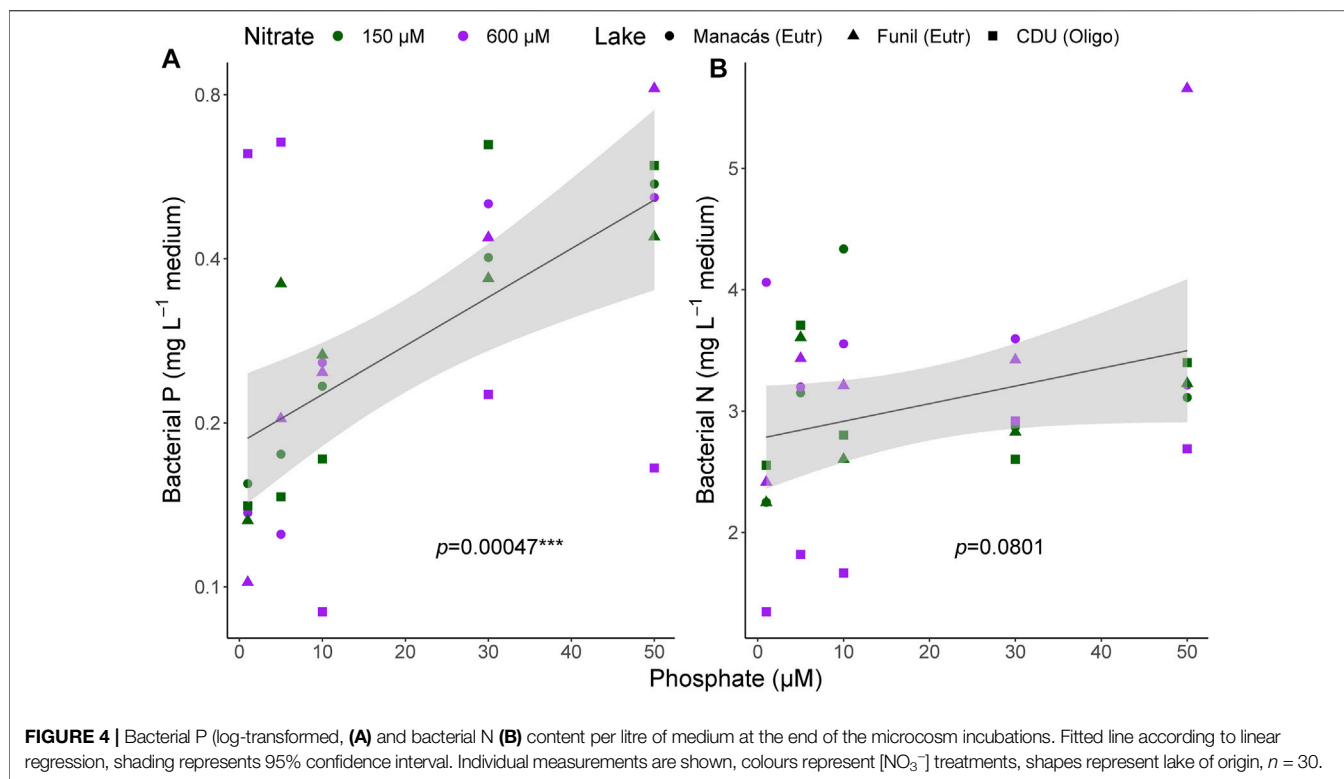


**FIGURE 2** |  $\text{CH}_4$  oxidation per day per ml of medium at different  $[\text{PO}_4^{3-}]$  and  $[\text{NO}_3^-]$  during exponential growth phase of MOB from Manacás lake, Funil lake, and CDU lake. Fitted line according to nonlinear least-squares regression, shading represents 95% confidence interval. Individual measurements are shown, colours represent  $[\text{NO}_3^-]$  treatments, shapes represent lake of origin,  $n = 30$ .



**FIGURE 3** | proportional abundance of different types of MOB per lake based on qPCR copy number of the *pmoA* gene at the end of microcosm incubations. Different colours represent different types of MOB,  $n = 26$ .

The change in  $\text{CH}_4$  oxidation did not coincide with changes in relative abundance of different MOB types, which were mostly influenced by the lake of origin (Supplementary Table S5;



**Figure 3).** Most MOB were either type Ia (58.8%) or type II (39.6%), while only a small percentage was type Ib (1.7%). Type Ia MOB were more abundant in the Manacás microcosms than in the CDU microcosms ( $p < 0.05$ ) but not influenced by  $[\text{NO}_3^-]$  or  $[\text{PO}_4^{3-}]$  concentrations. Type II MOB showed the opposite trend and were more abundant in the CDU lake microcosms than in the Manacás microcosms ( $p < 0.05$ ), while also not being affected by  $[\text{NO}_3^-]$  or  $[\text{PO}_4^{3-}]$ . Type Ib were more abundant in the Manacás and CDU microcosms than in the Funil microcosms ( $p < 0.05$  for both, **Supplementary Table S5**). Interestingly, the number of type Ib MOB significantly decreased at higher  $[\text{NO}_3^-]$  ( $p = 0.05$ , **Supplementary Table S5**) but since the abundance of type Ib was relatively low, this had little impact on overall community composition.

Bacterial P content showed a strong relationship with  $[\text{PO}_4^{3-}]$  ( $p = 0.00047$ , **Figure 4A**), while bacterial N content (**Figure 4B**) was not influenced by either  $[\text{NO}_3^-]$  or  $[\text{PO}_4^{3-}]$ . In contrast to  $\text{CH}_4$  oxidation rates, the increase in bacterial P content was also observed at high  $[\text{PO}_4^{3-}]$  and not only at low concentrations, as there was a linear relationship between the log transformed bacterial P content and  $[\text{PO}_4^{3-}]$ .

## DISCUSSION

In our microcosm experiment, we found that increased  $[\text{PO}_4^{3-}]$  led to an increase in potential  $\text{CH}_4$  oxidation rates, in agreement with our hypothesis. This was likely because MOB incorporated more P to increase their growth rates, as shown by the increase in bacterial P biomass content at high  $[\text{PO}_4^{3-}]$ . However, the  $[\text{PO}_4^{3-}]$  did not affect the relative abundance of different types of MOB.

$\text{PO}_4^{3-}$  increased  $\text{CH}_4$  oxidation during exponential growth in a nonlinear way, indicating that  $\text{PO}_4^{3-}$  was likely the limiting growth factor for the MOB communities at low concentrations. Increasing  $\text{PO}_4^{3-}$  to concentrations up to  $10 \mu\text{M}$  relieved MOB from  $\text{PO}_4^{3-}$  limitation and thereby increased their growth. This is in line with previous research showing higher rates of  $\text{CH}_4$  oxidation in systems with higher  $\text{PO}_4^{3-}$  availability (Veraart et al., 2015; Nijman et al., 2021). Increased P content suggested that bacteria used more P at higher  $\text{PO}_4^{3-}$ , possibly to invest in rRNA to increase their growth rate - as suggested in the growth rate hypothesis (Elser et al., 2000) - thereby increasing  $\text{CH}_4$  oxidation rates. At  $\text{PO}_4^{3-}$  concentrations above  $10 \mu\text{M}$ , P on the filters still increased but this did not coincide with increased  $\text{CH}_4$  oxidation, possibly because bacteria stored the excess  $\text{PO}_4^{3-}$  in organic and inorganic molecules inside the cytoplasm (Kulakovskaya, 2014) rather than investing in new cells. Likely, another limiting factor was preventing faster growth of MOB, such as copper (Semrau et al., 2010) or maximum growth rate (Ho et al., 2013). Future studies could investigate if  $\text{PO}_4^{3-}$  concentrations lower than  $1 \mu\text{M}$ , which are common in surface waters (Geurts et al., 2008), limit  $\text{CH}_4$  oxidation even more.

It is not clear whether the response of the model MOB communities in our experiment would be comparable to the response of sediment MOB communities, as those generally experience much lower  $\text{O}_2$  concentrations (Wang et al., 2014). Interestingly, MOB communities in the water column have the highest growth rates at low  $[\text{O}_2]$ , and the question remains whether the effects of  $\text{PO}_4^{3-}$  that we found in this experiment would also be found at low  $[\text{O}_2]$ . A future experiment at lower initial  $[\text{O}_2]$ , in which  $[\text{O}_2]$  is closely monitored and added when necessary, could give further insight into these dynamics.

Since the increased  $\text{CH}_4$  oxidation in the microcosms was only apparent at low  $\text{PO}_4^{3-}$  concentrations, the direct effect of  $\text{PO}_4^{3-}$  on  $\text{CH}_4$  oxidation might contribute little to the  $\text{CH}_4$  cycle in highly eutrophic lakes. Rather, these lakes might be more strongly influenced by an increase in methanogenesis because of higher algal carbon decomposition (West et al., 2012) and decreased oxygen (Bastviken et al., 2004). If MOB in sediments are stimulated by  $\text{PO}_4^{3-}$  in a comparable way to those in the microcosms, that may be one of the reasons why  $\text{CH}_4$  emissions are higher in eutrophic lakes (Aben et al., 2017; Davidson et al., 2018), while MOB are a more effective  $\text{CH}_4$  filter in mesotrophic lakes where they still benefit from increased  $\text{PO}_4^{3-}$  concentrations (Veraart et al., 2015).

Our results are in line with the increased  $\text{CH}_4$  oxidation rates found in P-amended incubations of surface water from ice-covered lakes (Sawakuchi et al., 2021). In both studies, adding  $5 \mu\text{M}$   $\text{PO}_4^{3-}$  increased  $\text{CH}_4$  oxidation rates. In contrast, the rates in our experiment were on average about 100-fold higher ( $0.19\text{--}4.75 \text{ mmol CH}_4 \text{ d}^{-1} \text{ L}^{-1}$  vs.  $0.0025\text{--}0.029 \text{ mmol CH}_4 \text{ d}^{-1} \text{ L}^{-1}$ ), and type II MOB were far more abundant (40% vs. <5%). That is likely because our model communities originated from freshwater sediments, which often harbour highly active MOB communities (Oremland and Culbertson, 1992).

Increased  $[\text{PO}_4^{3-}]$  had little effect on MOB community composition, which was more related to the lake of origin. The eutrophic Manacás lake had a higher proportion of type Ia MOB, while the oligotrophic CDU lake had more type II MOB. This is in line with predictions that type I MOB are better competitors and thrive in high nutrient environments, while type II are better stress-tolerators that can deal with low nutrient conditions (Ho et al., 2013). However, the fact that MOB community composition did not change in response to  $[\text{PO}_4^{3-}]$  was in contrast with these predictions. Most likely, the short duration of 1 week for our experiment explains this result, which was too short to change MOB community composition. Future research could simulate MOB growth at different  $[\text{PO}_4^{3-}]$  for a longer period to see possible shifts in community composition.

The microcosms did not show relationships between N:P ratio and  $\text{CH}_4$  oxidation, bacterial N and P content, and MOB community composition (**Supplementary Figures S1–3**). Likely, this was because  $[\text{PO}_4^{3-}]$  was limiting and therefore driving most changes in the microcosms, while changes in MOB community composition would require more time. A future experiment using concentrations of  $\text{NO}_3^-$  and  $\text{PO}_4^{3-}$  that are not limiting, e.g.,  $100\text{--}600 \mu\text{M}$   $\text{NO}_3^-$  and  $10\text{--}100 \mu\text{M}$   $\text{PO}_4^{3-}$ , could provide further insight into the effects of N:P ratio on  $\text{CH}_4$  oxidation.

To conclude, we show that MOB originating from lake sediments can be  $\text{PO}_4^{3-}$  limited and show a nonlinear

response to increased  $\text{PO}_4^{3-}$  concentrations, indicating that they can potentially mitigate increased methanogenesis in mesotrophic lakes that become more eutrophic.

## DATA AVAILABILITY STATEMENT

The datasets presented in this study can be found in online repositories. The names of the repository/repositories and accession number(s) can be found below: (10.17026/dans-x59-warf) (DANS EASY).

## AUTHOR CONTRIBUTIONS

TN, AA, and AV designed the experiment. TN and AA carried out the practical work, with support from PB for the qPCR analyses. All authors discussed the results and contributed to the final manuscript.

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## SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fenvs.2022.804512/full#supplementary-material>

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