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Effect of permeability anisotropy on the CO₂ saturation distribution and phase change during a leakage event in a saline aquifer

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Predicting impacts of potential carbon dioxide (CO₂) leakage into shallow aquifers that overlie geologic CO₂ storage formations is an important part of developing reliable carbon storage technology. To quantifying the effect of permeability anisotropy, a three-dimensional hypothetical reservoir model was formulated to analyze the migration behavior of CO₂ under diverse permeability anisotropy scenarios. Sensitivity analyses for parameters corresponding to the permeability anisotropy and the leakage rate are conducted, and the results suggest that permeability anisotropy significantly affect the CO₂ migration characteristics. Increasing the parameter of vertical/horizontal permeability ratio results in longer CO₂ migration distances, which enhances the aqueous phase ratio and safety through more interaction with the aquifer, but also raises the potential of the leakage reaching the ground surface due to higher gas ratio. A comprehensive understanding of these dynamics is crucial for implementing effective monitoring and management strategies.

KEYWORDS

CO₂ geological storage, permeability anisotropy, numerical simulation, leakage evaluation, phase change

1 Introduction

Porous media brine-filled aquifers provide a substantial potential storage capacity for the geological carbon dioxide (CO₂) sequestration technology (Eccles et al., 2012; Karvounis and Blunt, 2021; Li et al., 2023a), which (Allen et al., 2018; Rycroft et al., 2024) is world widely considered a key element in addressing the climate crisis and achieving carbon neutrality goals (Celia et al., 2005). Assessing the negative environmental impact is a crucial aspect to ensure the long-term effective operation of the storage system using saline aquifers (Bradshaw et al., 2007; Navarre-sitchler et al., 2013). During the decision-making stage, a rational monitoring plan should be formulated by comprehensively considering geological features, groundwater flow, environmental impact, and so on (Miocic et al., 2019). This involves identifying the most likely areas for leakage and potential leakage pathways by considering the geological structure, permeability distribution, and groundwater flow paths in the underground formations. Therefore, specific simulations based on a thorough geological survey are necessary to determine the optimal monitoring positions

(Chadwick et al., 2008). Besides the monitoring locations, the monitoring frequency is also a key factor in a leakage assessment (Gunter et al., 2000; Romanak and Dixon, 2022).

The permeability anisotropy of underground formation implies the permeability differences between different directions. The permeability anisotropy directly affects the CO₂ migration paths as it flows through the formations, and in turn its distribution in the leakage assessment of the stored CO₂ underground (Bear and Corapcioglu, 1986; Ghanbari et al., 2006; Pei et al., 2024). For example, anisotropy may lead to an uneven distribution of CO₂ in underground rock formations, potentially forming asymmetric leakage paths (Dai et al., 2022; Pavan and Govindarajan, 2023). In formations like shales, where geological CO₂ storage might occur, the horizontal permeability can be notably higher than the vertical permeability (Liang et al., 2021; Sun et al., 2023; Gu et al., 2024). The laminar structure of shale facilitates fluid movement along the bedding planes, making understanding this anisotropic nature crucial in assessing the potential pathways for CO₂ migration and the associated risk of leakage. Conversely, in formations like sandstones with a more uniform and interconnected porosity, the ratio of horizontal to vertical permeabilities may be closer to 1, reflecting a more isotropic nature (Li et al., 2023b). The grains in sandstones create a network of pathways that allows CO₂ to traverse vertically with a similar ease as flow horizontally. This has implications for the potential transport of CO₂ and the risk of fluid migration to shallower aquifers or, in worst-case scenarios, to the surface. It points out that one needs to characterize permeability very carefully to reduce uncertainty in permeability, making predictions of CO₂ transport and monitoring plan more accurate and allowing for a more reliable optimization of operational choices (e.g., properties and density of wells, etc.) (Oldenburg et al., 2023).

However, despite some qualitative descriptions of the impact of permeability anisotropy on CO₂ migration, there is a lack of quantitative research on its effects (Zhang et al., 2017; Gan et al., 2021; Liu et al., 2021). It is worth noting the relatively limited number of ongoing storage projects, resulting in a shortage of on-site data for researching leakage incidents. For better prediction, prevention, and response to leakage incidents, it is essential to conduct in-depth research and draw upon existing lessons and experiences (Saleem et al., 2021; Mahjour and Faroughi, 2023). Simple theoretical analysis and analogies fall short in providing sufficient information for engineering applications under complex geological conditions. Therefore, it becomes important to conduct in-depth numerical simulations to analyze the effect of permeability anisotropy on CO₂ migration patterns, providing a more accurate and reliable guidance for future geological storage projects. Numerical simulations also need to consider anisotropy to model the migration of CO₂ in underground rock formations more realistically (Carrera and Neuman, 1986). While the numerical simulation method plays a critical role in the leakage assessment, it faces a series of challenges. These challenges mainly stem from the limited sample quantities and the use of approximate values for underground rock properties (Hortle et al., 2014; Wang et al., 2021; Legentil et al., 2023). As the number of samples from the underground formations is often limited, it is extremely challenging to fully characterize the geological structure and properties of these three-dimensional and highly anisotropic formations. This leads to the parameters and properties used in

numerical simulations often being based on approximate values from the limited samples (Bianchi et al., 2016). The impact of this issue is primarily manifested in the uncertainty of simulation results. Numerical simulations cannot accurately reflect the real situation since the estimated range for anisotropy ratio spans from tens to hundreds, showcasing the substantial variability across different geological settings. Faced with this challenge, researchers need to better understand the reliability of simulation results by flexibly adjusting simulation parameters, employing sensitivity analysis, and pushing for innovation in geological exploration technologies to enhance the direct acquisition of underground rock properties (Nooraiepour, 2018).

Therefore, conducting an analysis of permeability anisotropy on leakage scenarios in geological CO₂ sequestration can assist researchers in better predicting potential leakage locations, leakage rates, and the non-uniformity of leakage paths. A three-dimensional hypothetical reservoir model was set up to analyze the migration behavior of CO₂ following its potential leakage into a saline aquifer under diverse permeability anisotropy scenarios. Through the numerical simulation examination of multiple permeability anisotropy configurations, this research could contribute insights essential to understanding and managing CO₂ sequestration and leakage risks in geological storage reservoirs, providing a good technical support for achieving carbon neutrality goals (Dai et al., 2022; Raats, 1973).

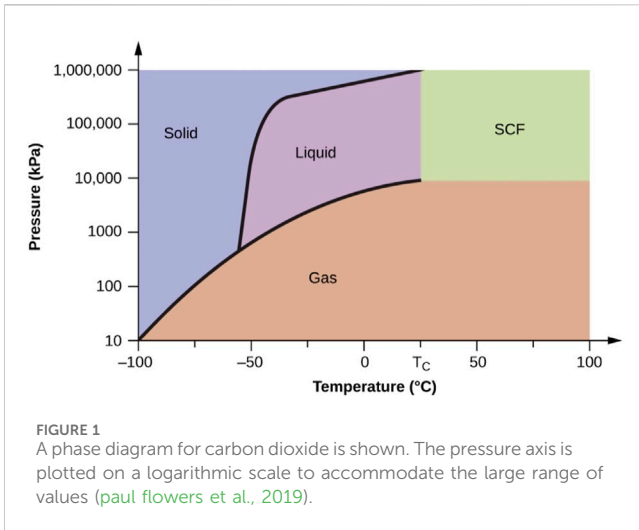
2 Materials and methods

2.1 Numerical methods

The CarbonFlow model was used as the simulation tool to do predictive modeling and sensitivity analyses of the fully coupled nonlinear equations describing mechanisms of CO₂ leakage. CarbonFlow is a module that has been derived from GPSFLOW (Cai et al., 2022), an general purpose subsurface flow simulator, which can model three-phase flow and transport over a wide range of pressures and temperatures encountered in reservoirs for water, CO₂ and NaCl. The governing equations and features related to permeability anisotropy and possible phase change of during CO₂ leakage to surface are listed following and other features can be found in the literature Cai et al. (2022), as shown in Table 1.

CarbonFlow uses the integral finite difference method of space discretization, which is the same in TOUGH3 (Jung et al., 2017). The permeability anisotropy could be characterized as $k_x \neq k_y \neq k_z$ of each element. To keep descriptions simple, we assume that the permeabilities in the two directions on the horizontal plane are the same, which means $k_x = k_y$. Hence, the permeability anisotropy is represented by the ratio ($\alpha = k_V/k_H$), where k_V is the vertical permeability and k_H is the horizontal permeability.

During a leakage event, a change of phase composition (appearance or disappearance of a phase) might occur when certain parameters such as pressure and temperature exceed certain threshold values, like the critical point (T_{crit} , P_{crit}) = (31.04°C, 7.38 MPa) (Pruess, 2011), as shown in Figure 1. For example, the injected CO₂ at 1 km depth is normally in supercritical phase and will change to gas phase when migrating upward to 500 m depth. In CarbonFlow, three different fluid



phases may be present: an aqueous phase that is mostly water but may contain some dissolved CO₂, a liquid CO₂-rich phase that may contain some dissolved water, and a gaseous CO₂-rich that may contain some water. The fundamental thermodynamic properties are calculated based on the NIST database (Linstrom and Mallard, 2024).

2.2 Basic model design

2.2.1 Conceptual model and leakage rate

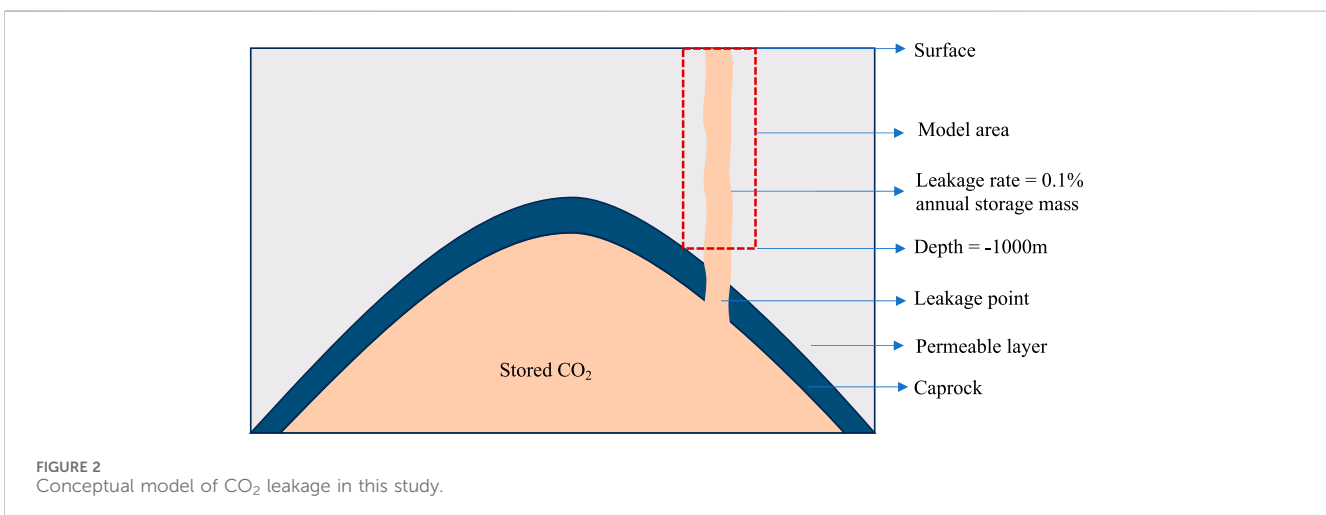
Potential pathways from saline formations have been detailed described in the IPCC’s special report (Allen et al., 2018): 1) through the pore system in low permeability caprocks if the capillary entry pressure is exceeded, 2) through opening in the caprock or fractures and faults, 3) through poorly completed and/or abandoned wells, and 4) through aquifers where dissolved CO₂ migrates laterally. It would be more than 100,000 years for CO₂ to leak through pore system in caprocks, and the flow in aquifers is estimated to be of the order of 1 mm/yr to 1 cm/yr.

As shown in Figure 2, this study focuses on leakage which can occur in just a few decades which operators can cope with and considers one simple scenario: leakage through faults or leakage through abandoned wells. Leakage through injection wells is excluded because these can be monitored relatively easily by operators and action can be taken when a leak occurs (Aoyagi et al., 2011). Leakage depth of 1,000 m is selected based on the depth description in Kalaydjian (Kalaydjian et al., 2011).

Since there is no sufficient on-site data research on leakage rate, and it depends on many factors including the geological conditions and operational conditions, 0.1% of annual storage mass is selected as the basic scenario based on the theoretical calculation considering natural carbon fluxes (Oldenburg et al., 2002). In the subsequent sensitivity analysis plan, the leakage rate is intended to be increased by 10 times to analyze this uncertainty factor. By increasing the leakage rate, the aim is to gain a deeper understanding of the system’s response and stability in extreme scenarios, thereby accurately assessing potential risks. In this study, we assume the stored CO₂ is 1 million tons per year, hence the leakage rate is 1,000 t/yr. The leakage continues 10 years, and we simulate to 20 years to study lateral migration characteristics.

2.2.2 Grids discretization

A three-dimensional hypothetical reservoir model was set up to analyze the migration behavior of CO₂ following its potential leakage into a saline aquifer under diverse permeability anisotropy scenarios. The mesh was created using mView software, which is a tool for mesh generation and management, as shown in Figure 3. The study area on the plane has been preliminarily chosen to be a 10 km by 10 km range based on prior research, which is larger than the affected area of the leak. Near the wellbore, at the center of the model, the mesh size is 2 m by 2 m. This finer grid extends to 50 m where the mesh size increases to 5 m by 5 m. Between 50 m and 5,000 m, the mesh size progressively increases with a multiplier of 2.5 starting from 5.0 m. There are a total of 1,521 cells in the horizontal plane. In the vertical direction, the mesh size near the leak point starts at 2 m and gradually increases to 20 m. Some grid refinements are conducted in the leakage and potential CO₂ phase change areas (from supercritical to



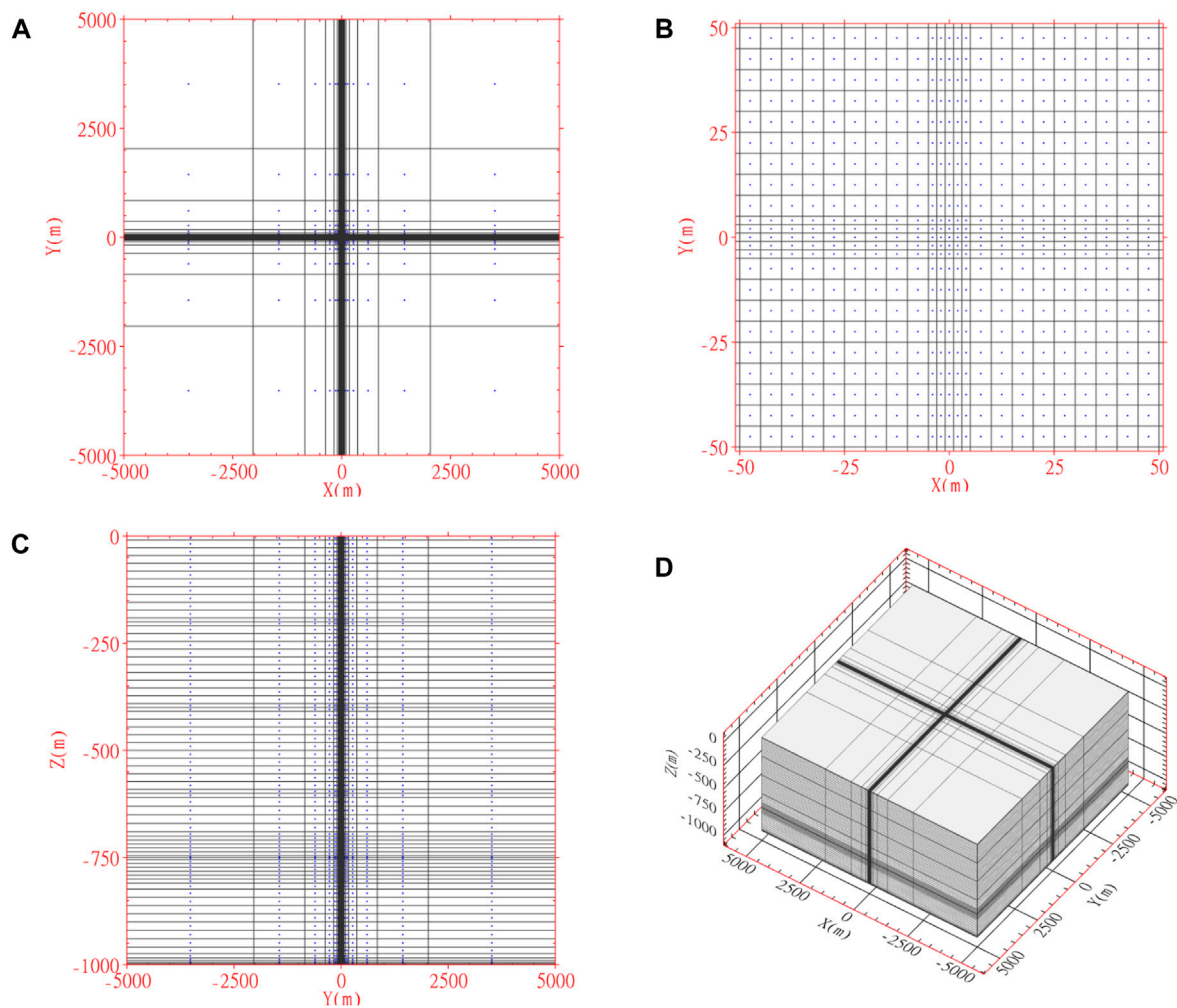


FIGURE 3 Domain discretization in plane perspective (A) with 1521 grids, grids refinement in the leakage area (B), domain discretization in a vertical profile (C) by 69 layers with refinement in leakage layer and phase change area (700–800 m depth), and the entire 3D model with 104,949 grids (D).

gas phases). There are 69 layers in total, comprising 104,949 cells overall.

2.2.3 Rock properties

We take the properties of Gudong Oilfield reservoir, which belongs to Shengli Oilfield, as a basic scenario in the hypothetical reservoir model. Data indicated porosity of 28%–35% in the Guantao Formation (Linze et al., 2004; Kalaydjian et al., 2011). Thus, an average porosity of 31% is determined to represent the sandstone in the basic model. The permeability of the sandstone is assumed to be 1,000 mD in the horizontal directions ($k_x = k_y$) as a basic scenario. This value is selected based on previous studies (evaluated permeability of 510–3118 mD) and is considered a reasonable approximation in a typical horizontal reservoir. However, it's important to note that the actual permeability of the sandstone can vary depending on factors such as the depth of the reservoir, the pressure and temperature conditions, and the presence of fractures or other features that can affect fluid flow. Therefore, the permeability

value used in the model should be considered as an approximation and would be adjusted in the next sensitivity studies. The other rock properties described in Kalaydjian (Kalaydjian et al., 2011) are used, as shown in Table 2.

The van Genuchten function (Van Genuchten, 1980) is used for the capillary pressure mode and the relative permeability mode, as in Pruess (Pruess et al., 1999). Hysteresis will certainly impact two-phase flow during the cycle of imbibition and drainage. But our focus is on the migration of CO_2 under constant leakage rate. Therefore, the hysteretic capillary pressure and relative permeability functions were not employed to account for a constant leakage rate since it was anticipated that the system would consistently remain on the drainage branch during the leakage phase, without undergoing repeated cycles of CO_2 leakage and water injection.

2.2.4 Initial and boundary conditions

Initially, the pressure is in a state of hydrostatic equilibrium, and the temperature is set according to a geothermal gradient of $0.03^\circ\text{C}/$

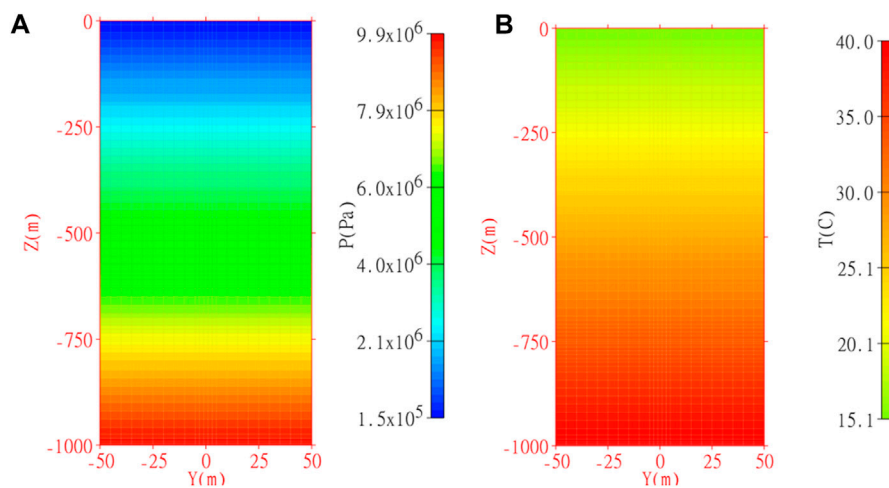


FIGURE 4 Initial pressure (A) and temperature (B) distribution.

TABLE 1 Governing equations used in CarbonFlow.

Parameter	Equations
Conservation of mass and energy	$\frac{d}{dt} \int_{V_n} M^i dV_n = \int_{\Gamma_n} F^i \cdot nd\Gamma_n + \int_{V_n} q^i dV_n$
Mass accumulation	$M^i = \Phi \sum_{\beta=1}^{NPH} S_{\beta} \rho_{\beta} X_{\beta}^i, i = 1, NK; \beta = 1, NPH$
Energy accumulation	$M^{NK+1} = \Phi \sum_{\beta=1}^{NPH} S_{\beta} \rho_{\beta} U_{\beta} + (1 - \Phi) \rho_R C_R T$
Mass flux	$F^i = \sum_{\beta=1}^{NPH} X_{\beta}^i \rho_{\beta} u_{\beta}$
Energy flux	$F^{NK+1} = -\lambda \nabla T + \varphi \sum_{\beta=1}^{NPH} h_{\beta} \rho_{\beta} u_{\beta}$

V_n , an arbitrary subdomain bounded by the closed surface Γ_n ; M^i , mass or energy accumulation term of component i per volume; F , mass or heat flux; n , normal vector on the surface element Γ_n pointing toward to V_n ; q , sink/source term of mass or energy. Φ , porosity; β , phase index (e.g., $\beta =$ gas, aqueous phase); S , saturation of phase β (the volume fractions of the pore space occupied by each phase); ρ_{β} , density of phase β ; X_{β}^i , mass fraction of component i in phase β ; NK, number of components; NPH, number of phases.

TABLE 2 Reservoir properties.

Property	Value	Unit	
Density	2,640 Aoyagi et al. (2011)	kg/m ³	
Porosity	0.31 Kalaydjian et al. (2011)	—	
Permeability of horizontal direction	1.0×10^{-12}	m ²	
Permeability of vertical direction	1.0×10^{-12}	m ²	
Capillary pressure (P_c) and relative permeability (kr)	Exponent in expressions, m	0.40	—
	Irreducible water saturation, S_{lr}	0.20	—
	Irreducible gas saturation, S_{lg}	0.05	—
	Strength coefficient, P_0	3,584	Pa
	Maximum possible value of P_{cap} , P_{max}	1.0×10^7	Pa
Saturated liquid phase, S_{ls}	0.999	—	

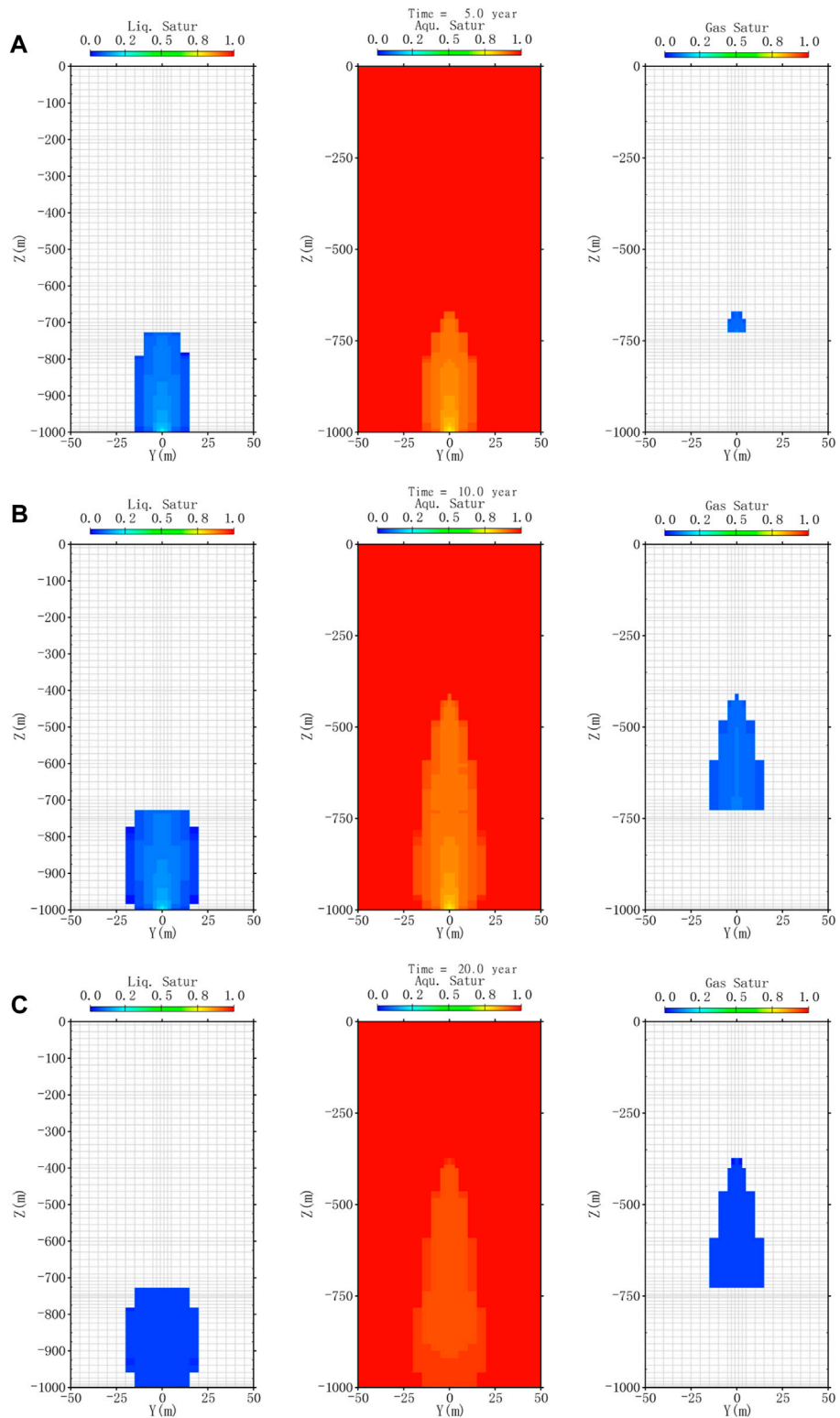
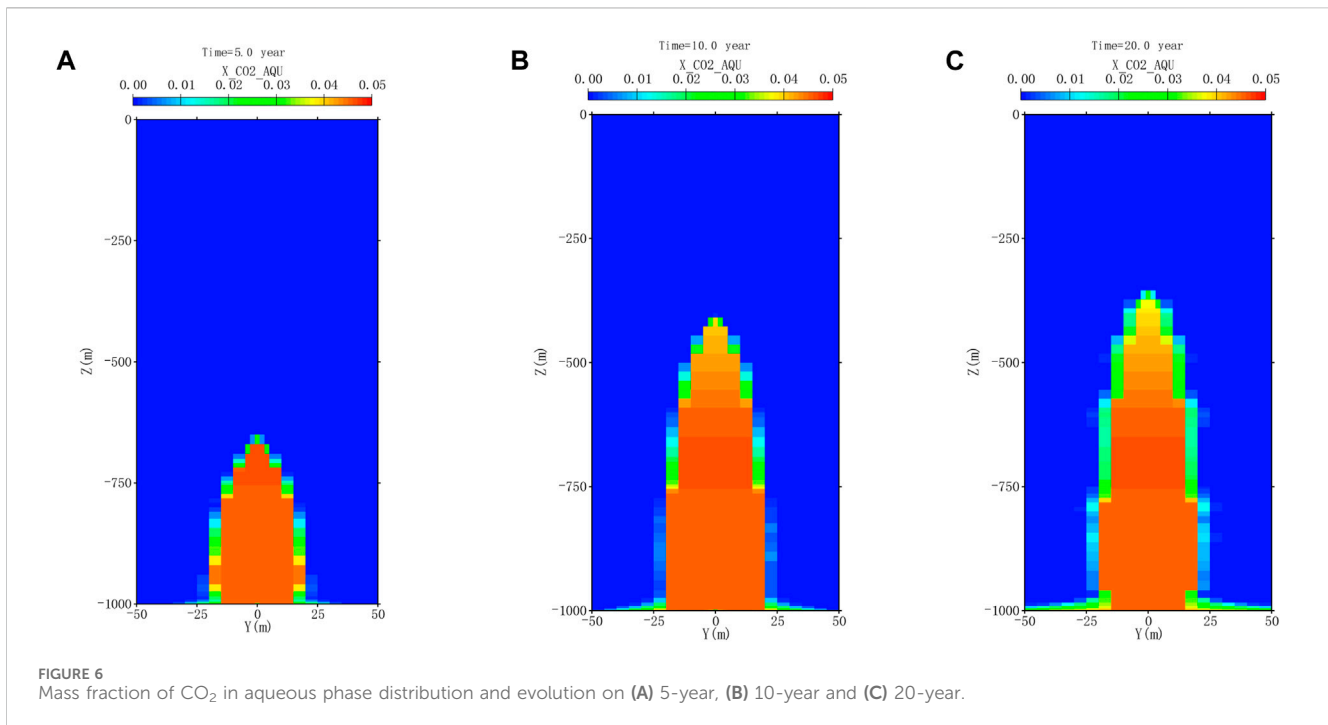


FIGURE 5 Phase saturation distribution and evolution on (A) 5-year, (B) 10-year and (C) 20-year.

m. The initial salinity of brine is 3% with no gas saturation, which has been used in the [Aoyagi et al. \(2011\)](#) and [Wang et al. \(2013\)](#). For visualization, we show initial conditions only for a portion of the

domain ($x = 0$ m, $y = -50$ m– 50 m, $z = -1,000$ m– 0 m), because conditions are the same in the rest of the domain, as shown in [Figure 4](#).



3 Results and discussion

3.1 CO₂ migration characterization

Figure 5 shows the phase saturation distribution and evolution. During the initial leakage period (0–5 years), as shown in Figure 5A, carbon dioxide primarily exists in the supercritical phase. Specifically, in our software, we categorize different phases using terms such as liquid phase, aqueous phase, and gas phase. Therefore, in the figure, we use the term “liquid phase” to represent the supercritical phase of carbon dioxide. From Figure 6A, in approximately the fifth year of the model, as it migrates upward to a depth of about 730 m, corresponding to the triple point conditions (pressure $p = 7.38$ MPa, temperature $T = 31.04^{\circ}\text{C}$), CO₂ undergoes a significant transition from the liquid phase to the gas phase. This transformation plays a crucial role in the distribution and evolution of phase saturation within underground reservoirs. Changes in underground conditions lead to this shift in CO₂ state, thereby influencing the distribution of phase saturation.

As shown in Figure 5B, at the end of the leakage period (10th year), CO₂ migrates upward to a position at -400 m, covering an upward leakage distance of 600 m. Ten years after the leakage ceases, Figure 5C, a notable transformation occurs in the behavior of CO₂. During this period, the majority of CO₂ undergoes a phase transition, converting into the aqueous phase, as shown in Figure 6B. Despite the passage of time, the distance of its upward migration remains relatively unchanged at -360 m (Figure 6C). The relatively stable migration distance suggests that the dissolution process efficiently absorbs the leaked CO₂, preventing significant vertical movement. This phenomenon might be attributed to the limited quantity of leaked CO₂. The scarcity in leakage volume results in a predominant reliance on dissolution processes. Unlike scenarios with larger leaks, where gaseous CO₂ might travel greater

distances due to buoyancy and pressure differentials, the modest volume of leaked CO₂ in this case allows dissolution to play a dominant role.

3.2 Permeability anisotropy

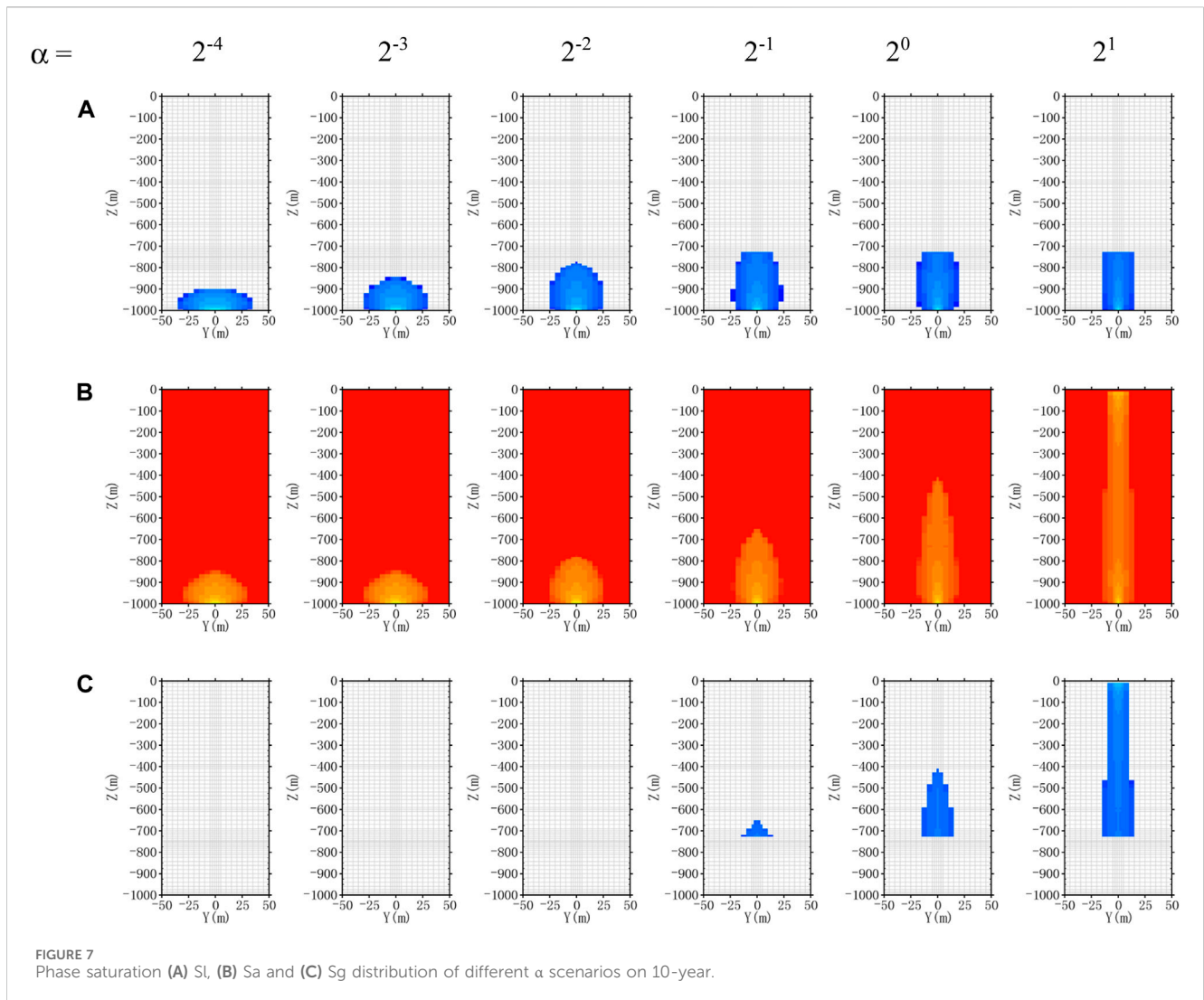
Various scenarios are designed as detailed in Table 3, where α ranges from 2^{-4} to 2 while the horizontal permeability remains constant. Ennis-king et al. (2005) conducted a study on how anisotropy affects the onset of instability, employing both linear and global stability analyses (Erfani et al., 2022). Their findings revealed that reducing α , while maintaining constant horizontal permeability and decreasing vertical permeability, stabilizes the process and impedes the initiation of instability. Due to the layered structure of natural formations, α is usually less than 1. In our design, we chose a maximum value of 2 to represent the broadest possible scenario. This design allows for realistic variations within the constraints of geological realities, providing a comprehensive exploration of permeability anisotropy.

3.2.1 CO₂ phase saturation distribution

Figures 7, 8 illustrate the distribution of phase saturation for various scenarios, offering a detailed glimpse into migration patterns over both 10 and 20 years. The analysis of figures shows a noticeable gathering of CO₂ around the leakage point at smaller scales. This is linked to a decrease in vertical permeability, constraining the paths for CO₂ migration and resulting in a significant reduction in movement within subsurface geological layers.

3.2.2 Trapping performance

In a storage unit, CO₂ in the aqueous phase is considered more stable than it in the gaseous or liquid phase. Therefore, we analyze



CO₂ trapped in the aqueous phase for a more in-depth examination, aiming to understand the impact of permeability anisotropy, as shown in Figure 9.

At first glance, with an increase in the parameter α , CO₂ migration seems to move away from the leakage point, covering longer distances and encountering the aquifer more extensively. Consequently, the aqueous phase ratio becomes larger, suggesting a heightened level of safety in this context. This positive aspect is attributed to increased interaction with the aquifer, contributing to a more stable and secure scenario.

However, it is essential to note that the larger values of α also signify a higher gas ratio. This implies a greater proportion of the gas phase, characterized by increased mobility, which, in turn, raises the potential for leakage to the ground. Despite the advantageous aspect of a larger aqueous phase ratio, the elevated gas ratio introduces a concern related to the mobility of the CO₂, posing a risk of ground leakage.

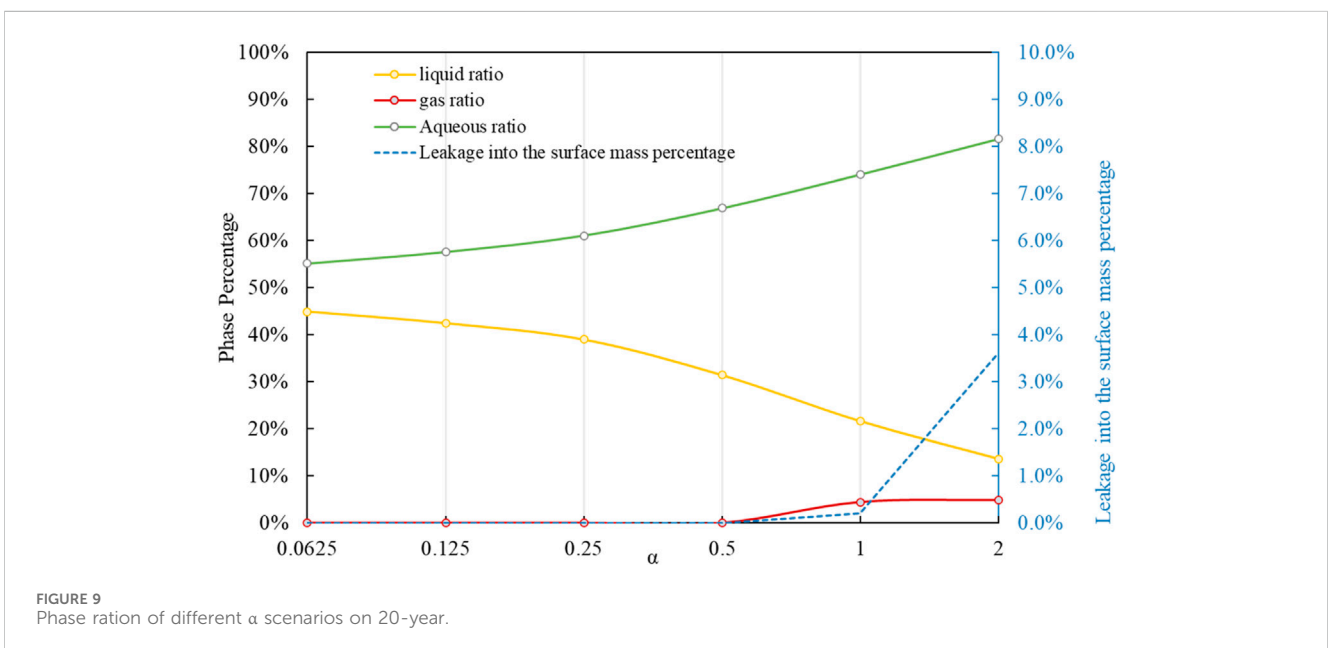
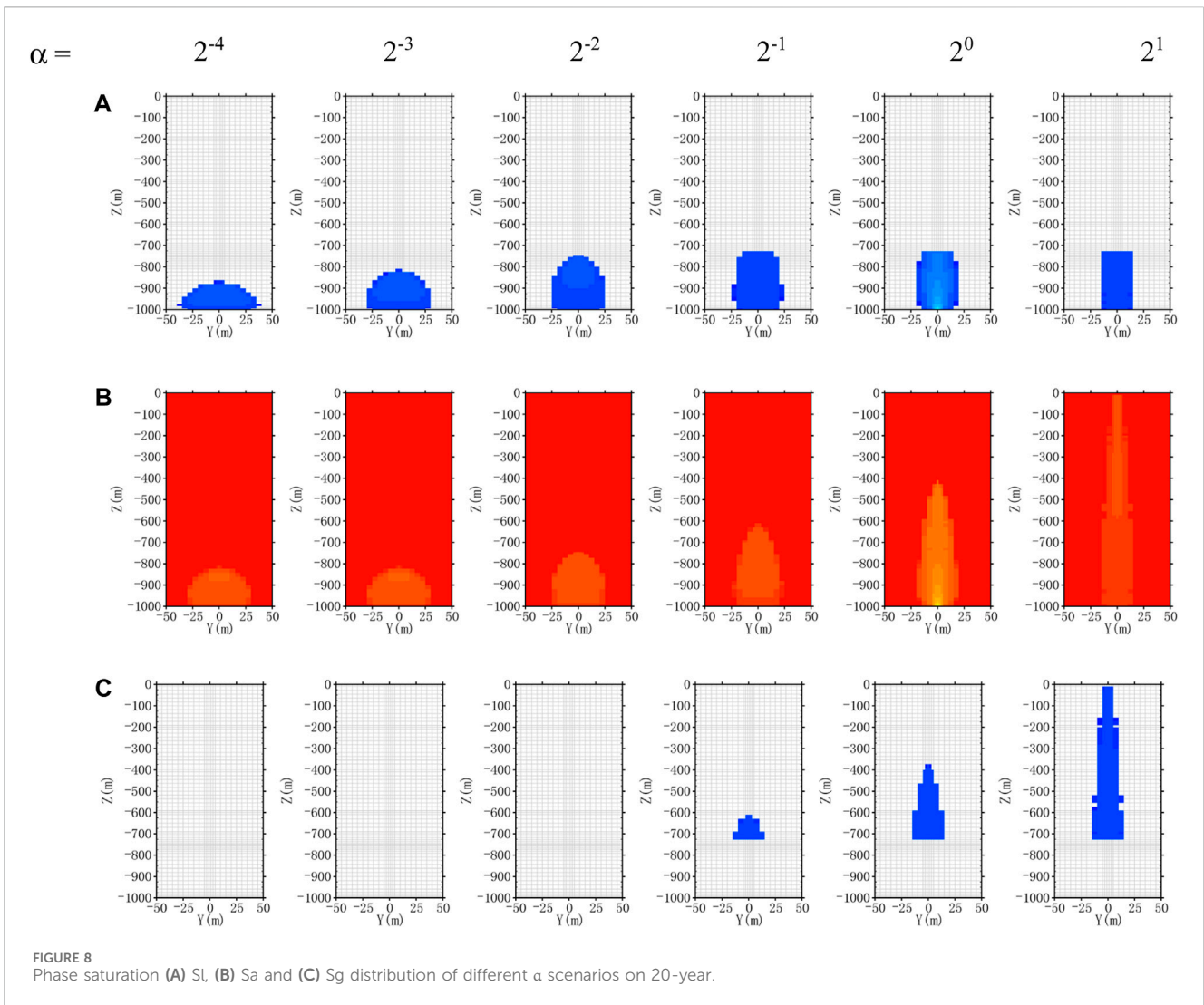
Notably, when α equals 1, a critical threshold is reached. At this point, some CO₂ may migrate out of the ground, potentially causing damage to the surrounding environment. This emphasizes the delicate balance between safety and risk associated with varying values of α in

the context of CO₂ migration. A comprehensive understanding of these dynamics is crucial for implementing effective monitoring and management strategies to ensure the safe storage and containment of CO₂, especially in scenarios where the parameter α reaches a critical value.

3.3 Leakage rate

The phase transitions and dissolution processes of CO₂ can vary with different leakage amounts. Two additional leakage rate scenarios (5 and 10 times the leakage rate in the basic model) were introduced to explore the interaction between leakage volume and permeability anisotropy.

The phase ratio for different leakage scenarios over 20 years, with a α equal to 1.0, is illustrated in Figure 10. Despite the leakage rate increasing to 10 times that of the basic model, the phase ratio undergoes a relatively moderate change, reaching 40% of the leakage mass into the surface. Reducing α representing smaller vertical permeability, proves effective in preventing leakage to the surface, as indicated by the decreased leakage ratio in both Figures 11, 12.



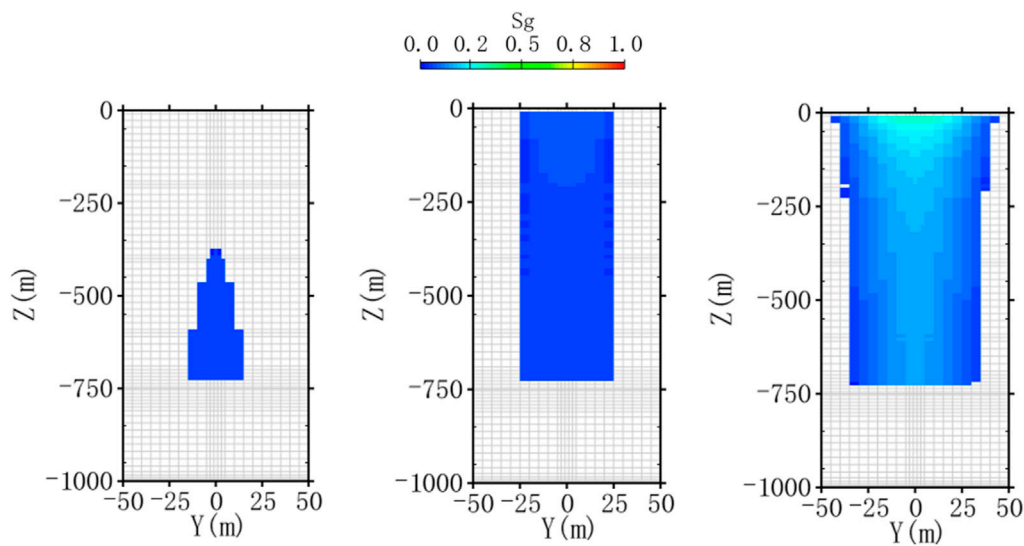


FIGURE 10 Gas phase saturation distribution of different leakage rate scenarios (from left to right refer to 0.1%, 0.5%, 1.0%) at 20 years ($\alpha = 1.0$).

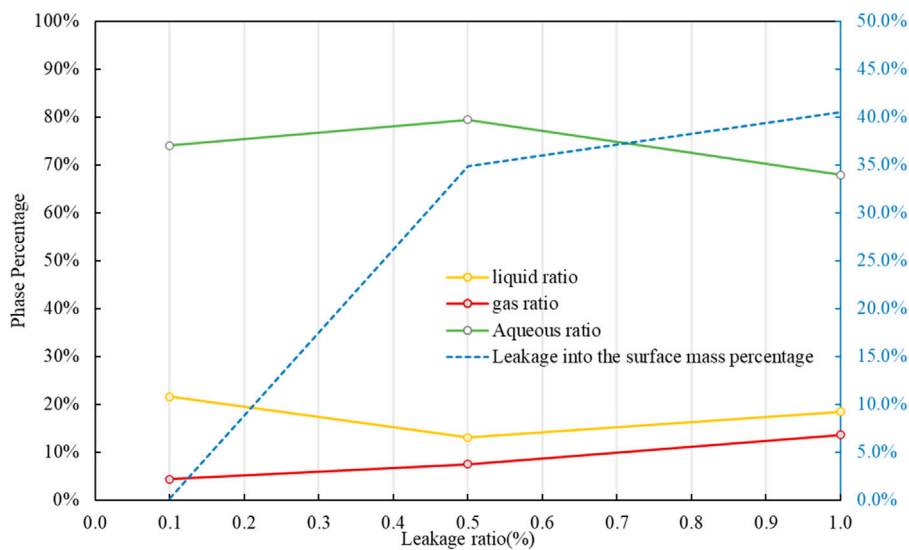


FIGURE 11 Phase partitions of different leakage ratio scenarios at 20 years ($\alpha = 1.0$).

TABLE 3 Scenarios design of permeability anisotropy.

Parameter	Value	Note (m^2)
α	$2^{-4}, 2^{-3}, 2^{-2}, 2^{-1}, 2^0, 2^1$	$kH = 1.0 \times 10^{-12}$

4 Conclusion

A comprehensive three-dimensional reservoir model was set up to analyze the migration behavior of CO₂ under diverse permeability anisotropy scenarios. The investigation

demonstrated that the complex interaction of permeability anisotropy, manifesting as varied permeability values in horizontal and vertical directions within the rock structure, can substantially modify the spread of CO₂ both horizontally and vertically. Sensitivity analyses indicated that the response of the reservoir to changes in permeability anisotropy is not uniform, suggesting that a one-size-fits-all approach to site evaluation may be inadequate. Instead, the simulation outcomes underscore the necessity for tailored acceptance criteria that reflect the unique properties and operational parameters of individual storage sites. Consequently, these

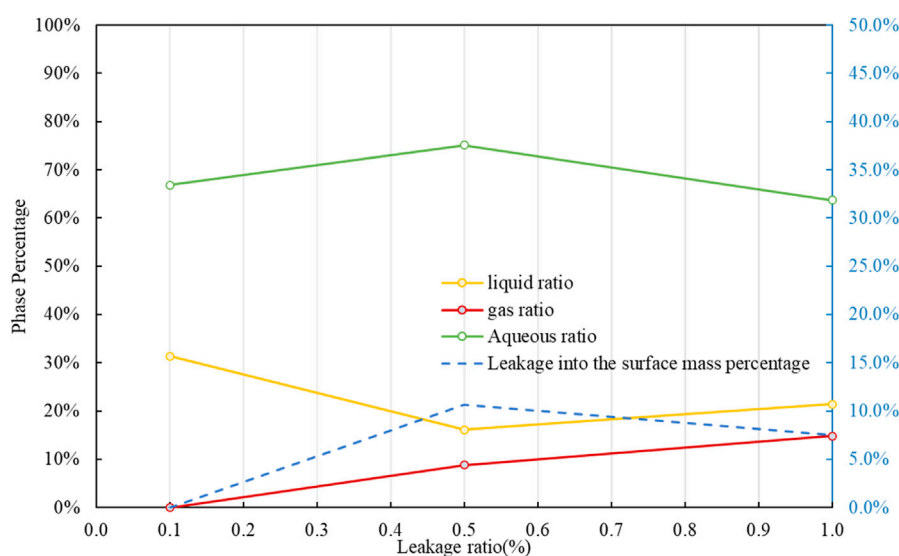


FIGURE 12
Phase partitions of different leakage ratio scenarios on 20-year ($\alpha = 0.5$).

results provide a robust framework for the assessment and design of potential storage sites, enabling a more precise estimation of the risks associated with CO₂ leakage.

In practice, the variability in these ratios across different geological settings highlights the need for detailed site-specific investigations to enhance our understanding of subsurface processes related to CO₂ storage. Hence, a thorough consideration of the impact of permeability anisotropy on the assessment of CO₂ geological sequestration leakage is crucial. Advances in technology and comprehensive data collection methods continue to contribute to refining our knowledge of rock permeability, ultimately supporting sustainable and informed decision-making in geological carbon dioxide storage.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary material, further inquiries can be directed to the corresponding author.

Author contributions

CG: Writing—original draft, Writing—review and editing. XW: Conceptualization, Methodology, Software, Writing—original draft.

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Conflict of interest

Author XW was employed by Petroleum Exploration and Development Research Institute, SINOPEC.

CG is a member of SINOPEC Key Laboratory of Carbon Capture, Utilization and Storage and Petroleum Exploration and Development Research Institute, SINOPEC.

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