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Significant spatiotemporal heterogeneity in drivers of water yield Service in Agro-pastoral Ecotone of Gansu, China

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Most previous studies on water yield service (WYs) analyzed the driving factors of WYs from a holistic perspective, but ignored the spatial heterogeneity and development of the driving factors. Using Invest model, Random Forest (RF) model and Geographically and Temporally Weighted (GTWR) model, we first examined the spatial distribution characteristics of WYs in agro-pastoral ecotone of Gansu China (AEGC) from 2000 to 2020. Secondly, the driving mechanism behind the spatiotemporal variation of WYs was discussed. The results show that: (1) In recent 20years, the average annual WYs of AEGC was 110.52mm, and the interannual variation showed an upward trend, with an increasing rate of 2.28 mm/a (p<0.05). WYs are high in the south, low in the north, and high in the northwest. Except for the southeast, WYs remained stable or increased in other regions. (2) The relative importance of the main influencing factors of WYs in AEGC successively were precipitation (1.57), evapotranspiration (1.29), temperature (1.12), population density (1.10), net primary productivity (NPP 1.06), and land use intensity (1.02). (3) Large-scale regional nature conditions are the primary force driving change in WYs, while in small-scale regions, human activities and land use are the primary drivers of WYs. Our research emphasizes that the effects of various influencing factors on WYs are significantly spatiotemporal heterogeneity, and WYs in different regions respond differently to the changes of influencing factors.

KEYWORDS

water yield service, climate change, human activities, driving factors, agro-pastoral ecotone of Gansu China

1. Introduction

Water connects various ecosystem processes and human activities, and water shortages are a common problem for people all around the world (Greve et al., 2018; Chung et al., 2021). Therefore, WYs are one of the most important ecosystem services (ES) and are a key indicator of ecosystem health (Costanza et al., 1997; Xia et al., 2021). Furthermore, WYs directly affect regional water supplies, impact many ecological functions such as food and raw material production, soil conservation, climate regulation, and biodiversity, and help maintain ecosystem stability and human social development (Graham et al., 2021; Jung et al., 2021; Queiroz et al., 2021).

WYs can be considered "ecosystem water production." The principle of dynamic water balance refers to the difference between water inflows and water expenditures over a certain period. Inflow is determined by precipitation, and expenditures are determined by actual evapotranspiration (Xia et al., 2017; Zhang and Zhang, 2018; Yang et al., 2019; Hu et al., 2021). Many studies have calculated the contribution of various factors to water balance. First, climate data and land use data from a

certain time period are modeled in different scenarios to represent, fluctuations in WYs caused by land use change. Additionally, they use fixed land use data and inputs different climatic conditions to simulate changes in WYs caused by climate change (Hu et al., 2020; He et al., 2021; Son et al., 2022). These methods consider precipitation to be an unpredictable factor and that human activities will change land use and alter evapotranspiration; however, this approach ignores the effects of climate change and the complexity of human activities (Rohatyn et al., 2018; Yin et al., 2020; Chen et al., 2022; Jia et al., 2022). Therefore, these simulations could overestimate the relative role of land-use change on WYs and underestimate the role of nature factors, resulting in inaccurate the conclusions.

Researchers have commonly explored the mechanism of WYs and related factors by assessing their relationship to the water cycle (Wang J. et al., 2021; Wang X. et al., 2021). Other studies have demonstrated that precipitation is significantly positively correlated with WYs and that actual evapotranspiration is significantly negatively correlated with WYs (Su and Fu, 2013; Zhang et al., 2020; Wu J. et al., 2021). Temperature, NPP, and normalized digital vegetation index (NDVI) have an insignificant impact on WYs, showing positive and negative bidirectionality. Some scholars (Yin et al., 2019) further believe that NPP positively affects WYs in regions with thriving ecosystems, and negatively affects regions with impaired ecosystems. However, our research analyzed these factors holistically and did not consider the spatial heterogeneity of WYs drivers; it also separated the continuity of development and change of WYs drivers over time. However, these results must be verified in other regions. Therefore, it is necessary to analyze the driving mechanism of WYs, using different methods at various regional scales.

After considering the importance of ecological functions and the variance related to geographical locations, the AEGC was selected as our study area. The AEGC is an important ecological barrier and the largest climatic and ecological transition zone in China. It is extremely sensitive to nature conditions change and human interference. Therefore, it is a suitable region for studying environmental problems (Pei et al., 2021; Fang et al., 2022). As an ecologically fragile, this area is the site of national ecological restoration projects such as the Three-North Shelterbelt and returning farmland to a forest (Liu et al., 2021), making it a suitable location for exploring the driving mechanism of WYs and how it is affected by nature conditions change (Yang et al., 2021) and human activities.

In this study, the main objectives are to (1) explore the long-term change trend and spatial distribution characteristics of WYs in AEGC from 2000 to 2020. (2) establish a comprehensive database of natural conditions and human activities to quantify the relative importance of various factors affecting WYs. (3) identify the spatiotemporal differences of the main impact factors on WYs, and study the driving mechanism behind the spatiotemporal changes of WYs. This work can not only aid in comprehending the effects of natural conditions change and human activities on WYs, but also provides valuable information for the protection, utilization and management of water resources.

2. Materials and methods

2.1. Study area

The AEGC $(33^{\circ}40' \sim 37^{\circ}54'N, 101^{\circ}55' \sim 108^{\circ}34'E)$ includes 35 districts and counties (Figure 1), covering an area of 99,300 km². It is

located on the eastern edge of the Tibetan Plateau and is a transition zone for major geographical units. It is known as the "Asian water tower" and is connected to the Tibetan Plateau in the west, Qinba Mountain in the south, the Loess Plateau in the east, and the Hexi Corridor in the north (Zhang et al., 2016; Immerzeel et al., 2020). The study area spans the two major river basins of the Yellow River and the Yangtze River and is an important source of water for the Yellow River. Its main rivers are the Tao River, Wei River, Huangshui River, and Bailong River (Qiu et al., 2022). The ecological environment in the study area is complex and diverse, and its primary landforms are plateaus, mountains, hills, and river valleys (Nie et al., 2015). Its altitude is 1,102~5,339 m, and is generally higher in the west and lower in the east. Precipitation is higher in the south (about 800 mm per year) and lower in the north (about 200 mm per year), with an average annual temperature of 1~11°C. In the summer, temperatures are higher, and there is more precipitation, while in the winter, temperatures are lower with less precipitation. As a result, there are differences in topography, climate, hydrology, and land use among the regions of the AEGC (Ma et al., 2020).

2.2. Data sources and processing

Select population density, economic density and land use intensity as human activity factors (Sang et al., 2019; Liu et al., 2020; Jiang et al., 2023), precipitation, temperature, evapotranspiration, NDVI, NPP, (Digital Elevation Model) DEM and slope as natural condition factors (Yang et al., 2011; Ahmed et al., 2021; Hu et al., 2021; Wang J. et al., 2021; Wang X. et al., 2021), and establish a comprehensive factors database that may affect water production services, providing a basis for subsequent determination of impact factors (Figure 2). All data sources and preprocessing can be visible in Table 1.

2.3. Research method

2.3.1. Water yield

Many studies have confirmed that InVEST can effectively perform ES simulations in different regions.¹ The evaluation results of the model can be visualized at the landscape scale (Hamel and Guswa, 2015; Redhead et al., 2016). Therefore, this study uses the InVEST to evaluate WYs according to the following calculation equations:

$$WY_{xj} = \left(1 - \frac{AET_{xj}}{P_x}\right) \times P_x \tag{1}$$

$$\frac{AET_x}{P_x} = 1 + \frac{PET_X}{P_X} - \left[1 + \left(\frac{PET_X}{P_X}\right)^{\omega}\right]^{\frac{1}{\omega}}$$
(2)

$$PET_x = K_{c,j} \cdot ET_{o,X} \tag{3}$$

$$v_x = Z \frac{AWC_x}{P_x} + 1.25 \tag{4}$$

¹ https://naturalcapitalproject.stanford.edu/invest



$$AWC_x = \min(Soil \ depth, Root \ depth) \times PAWC$$
 (5)

$$PAWC = 54.509 - 0.132 \times SAND - 0.003 \times$$
$$SAND^{2} - 0.055 \times SILT - 0.006 \times$$
$$SILT^{2} - 0.738 \times CLAY + 0.007 \times$$
$$CLAY^{2} - 2.688 \times SOC + 0.501 \times SOC^{2}$$
(6)

In Eqs (1)–(6), WY_{xj} is WYs of land-use type *j* in grid *x* (mm); AET_{xj} is the actual evapotranspiration of land-use type *j* in grid *x* (mm); PET_{xj} is the annual precipitation in grid *x* (mm); $K_{c,j}$ is the evapotranspiration coefficient of land-use *j*; $ET_{0,x}$ is the reference crop evapotranspiration in grid *x* (mm); w_x is the empirical parameters in grid *x*; *Z* is seasonal content with a range of 1 to 30, which was assigned a value of 30 according to the estimates of actual evapotranspiration; AWC_x is the soil available water capacity in grid *x* (mm); *Soil_depth* is the root burial depth (mm) from HWSD; $Root_depth$ is the plant root depth (mm); *PAWC* is the plant available water content [0,1]. where the SAND, SILT, CLA Y, and SOC are the proportion of sand, silt, clay, and organic matters in the soil, respectively.

2.3.2. Trend analysis and coefficient of variation

A linear regression analysis was used to analyze the variation trend of WYs by grid unit, which is widely used and easy to interpret (Belay et al., 2021; Kilama Luwa et al., 2021). The coefficient of variation is the ratio of the standard deviation to the mean value, which can reflect the spatial fluctuation characteristics of the WYs of the grid unit. The following calculation equations were used:

$$k = \frac{n \times \sum_{i=1}^{n} i \times WY_i - \sum_{i=1}^{n} i \sum_{i=1}^{n} WY_i}{n \times \sum_{i=1}^{n} i^2 - \left(\sum_{i=1}^{n} i\right)^2}$$
(7)

$$C = \frac{S}{\left|\overline{x}\right|} \tag{8}$$

In Eqs (7)–(8), WY_i is the WYs in years of *i* (mm); where *i* is the year; *n* is the number of years; *k* is the slope of the trend of water production; the significance test indicates 95% confidence, *C* is the coefficient of variation; *S* is the standard deviation; and \bar{x} is the average value.

2.3.3. Random forest

Using RF to evaluate the importance of independent variables is to quantify the contribution degree of each independent variable to the classification performance of the constructed K decision trees (Arabameri et al., 2019; Peng et al., 2020). Contribution degree is usually evaluated by out of bag (OOB) error rate. Contribution degree is represented by feature importance measures (FIM).



FIGURE 2

The impact factors database is briefly displayed. Due to the large amount of data, except DEM and slope data, which are basically fixed, other data are displayed spatially based on the average value from 2000 to 2020. The population density (A) and economic density (B) to some extent indicate the intensity of human activities in the region, and the intensity of land use directly (C) reflects the transformation of the natural environment by human beings. (D–F) belongs to climate data, which are average precipitation, average temperature and average evaporation respectively. (G, H) are geographic data, where are DEM and topographic slope respectively. (I, J) are ecological data, which are average NDVI and average NPP respectively.

TABLE 1 Data descriptions and sources used in this study.

Data type	Data name	Data source	Description	Resolution
Human activity data	Population density	Open Spatial Demographic Data and Research (https:// www.worldpop.org)	Population density raster data. 2000, 2005, 2010, and 2020 are public data. Data of other years are calculated by Arcgis10.5 according to the population growth rate.	100 m × 100 m
	Economic density	Resource and Environment Science and Data Center (http://www.resdc.cn)	Economic density raster data. 2000, 2005, 2010, and 2020 are public data. Data of other years are calculated by Arcgis10.5 according to the population growth rate.	$500\mathrm{m} imes500\mathrm{m}$
	Land use intensity		LUCC raster data was interpreted by Landsat TM remote sensing satellite images. The land use intensity of unused land was 1, that of grassland, woodland, wetland and water area was 2, that of cultivated land was 3, and that of construction land was 4	100 m × 100 m
Climatic data	Precipitation Temperature	Chinese Meteorological Sharing	Interpolate the precipitation and temperature grid data of 37	500 m × 500 m 500 m × 500 m
		Service System (http://data.Data.cma.cn/data)	meteorological stations in and around the study area (use the meteorological interpolation software ANUSPLIN to carry the BVTPS and TVPTPS models to interpolate the precipitation and temperature related data, and use Arcgis10.5 to synthesize the annual precipitation and average temperature data).	
	Evapotranspiration	Earth Data Open Access for Open Science (https://www.earthdata.nasa. gov)	MOD16A2 raster data, with a time resolution of 8 days, were synthesized into annual evapotranspiration data by Arcgis10.5	$500\mathrm{m} imes500\mathrm{m}$
Geographic data	DEM Slope	Resource and Environment Science and Data Center (http://www.resdc.cn)	DEM raster data Slope raster data are obtained by slope calculation of DEM raster data by Arcgis10.5 software	30 m × 30 m 30 m × 30 m
	Soil	Harmonized World Soil Database (https://www.fao.org/soils- portal/data-hub)	Soil raster data	1,000 m × 1,000 m
Ecological data	NDVI	Earth Data Open Access for Open Science (https://www.earthdata.nasa. gov)	MOD13A1 raster data, with a time resolution of 16 days, were synthesized into annual NDVI data by Arcgis10.5	500 m × 500 m
	NPP		MOD17A2 raster data, with a time resolution of 16 days, were synthesized into annual NPP data by Arcgis10.5	$500\mathrm{m} imes500\mathrm{m}$



$$FIM_{km}^{(OOB)} = \frac{\sum_{P=1}^{n_o^k} I(Y_P, Y_P^k)}{n_o^k} - \frac{\sum_{P=1}^{n_o^P} I(Y_P, Y_{P,\pi_m}^k)}{n_o^k} \qquad (9)$$

$$FIM_{m}^{(OOB)} = \frac{\sum_{k=1}^{K} FIM_{km}^{(OOB)}}{K\sigma}$$
(10)

In Eqs (9)–(10), n_o^k is the number of observed samples of the *kth* tree. Y_P is the *Pth* sample pair true classification labels should be applied. Y_P^k is the prediction classification result of the *p* observation of OOB data by the *kth* tree before random replacement F_m . Y_{P,\neq_m}^k is the classification result of the *pth* sample by the *k*th decision tree after random replacement F_m . When the feature F_m does not appear in the *k*th tree, $FIM_{km}^{(OOB)} = 0$. *K* shows the number of decision trees in the RF. σ is the standard deviation of $FIM_{km}^{(OOB)}$. The importance score $FIM_m^{(OOB)}$ of feature F_m represents the contribution of feature F_m to the classification accuracy. FIM was determined by mean and standard deviation of OOB error rate.

2.3.4. Geographically and temporally weighted

Based on the classical geographically-weighted regression model, the GTWR model adds time to the two-dimensional spatial coordinates from the geographically-weighted regression model, producing threedimensional coordinates (He and Huang, 2018; Liu et al., 2018; Li and Managi, 2021; Wu S. et al., 2021). The spatiotemporal non-stationarity relationship was modeled by constructing a local model of spatial and temporal dependence, which can better reveal the local effects of various driving factors on WYs. It was calculated as follows:

$$y_i = \beta_0 \left(u_i, v_i, t_i \right) + \sum_{k=1}^p \beta_k \left(u_i, v_i, t_i \right) x_{ik} + \varepsilon_i$$
(11)

In Eq. (11), y_i is the fitting value of the WYs change of sample i; (u_i, v_i, t_i) is the space-time coordinate of sample i; $\beta_0(u_i, v_i, t_i)$ is the intercept term; p is the number of explanatory variables; $\beta_k(u_i, v_i, t_i)$ is the estimated coefficient of the kth explanatory variable; and ε_i is the model residual.

2.3.5. Dynamic analysis of the driving area

The dynamic degree is a measure of changes in the area of WYs driven by dominant driving factors (Li et al., 2022; Yao et al., 2022) and is calculated as follows:

$$D = \frac{U_b - U_a}{U_a} \times \frac{1}{T} \times 100\%$$
(12)

In Eq. (12), D is the dynamic degree of the dominant driving WYs area of a factor; U_a , U_b are the areas dominated by the element at the beginning and end of the study; T is the length of the study period.

3. Results

3.1. Spatiotemporal changes in WYs

3.1.1. Interannual variation of WYs

From 2000 to 2020, the average annual WYs in AEGC was 110.52 mm, and the total average annual WYs was 1.1×10^{10} m³(Figure 3). The WYs in AEGC generally showed an upward trend, with an average increase rate of 2.28 mm/a. The annual WYs were the smallest in 2015 (35.48 mm), and the largest in 2019 (172.63 mm). Subsection statistics better reflect the changing trend of WYs. From 2000 to 2010, the annual WYs steadily increased, with an average WYs of 99.59 mm and an increased rate of 0.58 mm/a. From



2011 to 2020, the fluctuation range of annual WYs became larger; the average WYs was 122.56 mm, and the rate of increase was 5.21 mm/a (p < 0.05).

3.1.2. Spatial characteristics and changing trends of WYs

Based on the spatial distribution characteristics of the multiyear average WYs in AEGC from 2000 to 2020 (Figure 4A), the WYs have obvious spatial differentiation. The overall WYs are high in the south, low in the north, and high in the northwest. High-value areas with WYs exceeding 300 mm are primarily concentrated in the south and northwest of the AEGC, and the low-value areas with WYs below 100 mm are mainly distributed in the north and northeast of the AEGC. In terms of subregions, there were significant differences in WYs among districts and counties (Figure 4B). The average annual WYs of Jingtai was the lowest (1.54 mm), and Hezuo was the highest (295.41 mm). The average WYs of all units were divided into 5 categories using the ArcGIS10.5 natural breaking point method. The spatial distribution of the units with low WYs formed a regional cluster, with Jingyuan at the center. The spatial distribution of high WYs units in the south is centered around Zhuoni.

According to the results of the change trend analysis, the change trends of WYs in AEGC are divided into 4 categories (Figure 4C).

Areas with insignificant increases in WYs account for the largest proportion (57.18%) and are widely distributed in the middle and north of the AEGC. The significantly increased area was concentrated in the southwest and northeast of the AEGC (24.29%). Only in the southeast, WYs showed a downward trend (5.22%). Using the ArcGIS10.5 natural breaking point method, the WYs fluctuations in AEGC were divided into five categories (Figure 4D). WYs fluctuations are higher in the north, lower in the south, and higher along the Yellow River. Areas with high WYs fluctuations are concentrated in the Lanzhou New District and the Yellow River Valley and account for 3.82% of the AEGC. The relatively low fluctuation and low fluctuation areas are widely distributed in the southern, central-eastern, and northwestern parts of the AEGC, accounting for 49.17 and 18.53% of the AEGC, respectively.

3.2. Driving mechanism of WYs

3.2.1. WYs independent variable importance measure

The driving factors of WYs have always been the focus of relevant research. In this study, based on the establishment of a comprehensive database including human activity factors and natural condition factors, the RF was applied to evaluate the relative importance of these factors. According to the results (Figure 5), Ranking of importance of independent variables: precipitation (1.57) > evapotranspiration (1.29) > temperature (1.12) > population density (1.10) > NPP (1.06) > land use intensity (1.02) > NDVI (0.83) > DEM (0.71) > economic density (0.68) > slope (0.51). Indicators with relative importance greater than 1 were selected for GTWR analysis.

3.2.2. Spatiotemporal heterogeneity of the main influencing factors on WYs

The GTWR model reflects the intensity and difference of these effects through the different regression coefficients of the driving factors in each region. Table 2 shows the fitting results of the GTWR model on the driving intensity of WYs by various factors from 2000 to 2020. Precipitation has a significant positive driving effect on WYs in AEGC (Figure 5), and the driving strength increases from the northeast to the south and northwest. The actual evapotranspiration negatively affects WYs (Figure 5). The driving intensity increases from the northeast to the south, and the actual evapotranspiration in the southeast most clearly inhibits the southeast WYs. The impact of temperature on WYs in the central and eastern regions showed a negative effect, and in the western region, it showed a positive effect, especially in the northwest region (Figure 5). The inhibition of population density on WYs is strong in the middle AEGC (Figure 5). NPP inhibits WYs in the south and northwest and promotes WYs in the northeast (Figure 5). Land-use intensity has a significant negative driving effect on WYs (Figure 5), especially in the south and northwest.

The time axis of the GTWR model was adjusted to fit the driving intensity of various factors on WYs in stages from 2000 to 2010 and 2011 to 2020. The dominant driving factors in different regions in each stage were then identified (Figure 6). The area of WYs driven by



The relative importance of the impact factors of WYs. The importance value has no substantive meaning and can be used for comparison.

precipitation is the largest (54.7%). Evapotranspiration inhibits WYs in the southeast. The low WYs in the northern part of AEGC are negatively driven by temperature. WYs in Lanzhou was negatively driven by population density. WYs in the northeast region are positively driven by NPP. WYs was inhibited by land use intensity in northwest AEGC. Table 3 shows that the WYs area driven by population density and temperature increases significantly, while the WYs area driven by precipitation and actual evapotranspiration decreases significantly. From the dynamic point of view, the population density driven WYs area growth rate is the highest, and the land use intensity and temperature driven WYs area growth rate is more than 10%. These results indicate that human activities have an increasingly significant influence on WYs in local areas (see Figure 7).

4. Discussion

4.1. Natural factors affecting WYs

Based on the interannual variation of WYs in AEGC, WYs was lowest in 2015 and highest in 2019 Since 2010, the overall WYs of AEGC has been on the rise. Meteorological research shows (Lai, 2012; Wang et al., 2016; Lu et al., 2021; Yang et al., 2021) that 2015 was a strong El Niño year with little precipitation in eastern Gansu. In 2019, precipitation significantly increased in the AEGC, and precipitation in the eastern part of Northwest China showed an upward trend. Consistent changes of precipitation and WYs (Dai et al., 2020). This study demonstrates that the area where precipitation dominated WYs was the largest, accounting for 52.59% of the AEGC. Overall, after analyzing regions with different WYs levels, precipitation was the most important factor affecting WYs.

Evapotranspiration, an important parameter to quantify WYs, combines water and energy balances and links the climate system to terrestrial ecosystems (Li et al., 2016; McEvoy et al., 2016; Zhang et al., 2020). In the southeastern AEGC with higher WYs, evapotranspiration is the dominant driver. This is because in areas with high WYs, WYs and evapotranspiration levels are in the same order of magnitude, meaning that WYs are sensitive to changes in evapotranspiration (Dai et al., 2020; Dai and Wang, 2020; Zhang Y. et al., 2021; Zhang J. et al., 2021; Jiang et al., 2022). However, in areas with low WYs, evapotranspiration is several times higher than WYs, and interannual changes in evapotranspiration could be greater than the average annual WYs (Zeng et al., 2022). Therefore, WYs is not sensitive to evapotranspiration response.

The effect of temperature on WYs shows positive and negative bidirectionality. In the northern AEGC of low WYs, temperature inhibits the rapid growth of WYs area. This because warming and drying of the drylands in recent decades, most notably in the current decade (Daramola and Xu, 2022; Yang et al., 2022). The increase of surface temperature leads to increase of precipitation intensity in the Tibetan plateau, especially for convective precipitation (Zhao et al., 2022). Therefore, temperatures are positive effecting WYs positively in the western AEGC.

TABLE 2 Result of GTWR model.

Bandwidth	AICc	Sigma	Spatiotemporal distance ratio	R ²	R ² Adjusted
0.1150	7,144	0.4719	0.4588	0.8397	0.8392



Driving factors	2000–2010		2011–2020		Variation /km ²	Dynamic / %
	Area /km ²	Proportion /%	Area /km ²	Proportion /%		
Precipitation	57,253	57.64	51,408	51.76	-5,845	-1.02
Evapotranspiration	22,338	22.49	14,292	14.39	-8,046	-3.60
Temperature	3,654	3.68	8,523	8.58	4,869	13.33
Population density	2,957	2.98	7,173	7.22	4,216	14.26
NPP	11,304	11.38	13,563	13.66	2,259	2.00
Land use intensity	1818	1.83	4,365	4.39	2,547	14.01

TABLE 3 Summary of control area of the main impact factors of WYs.

The AEGC is located in the monsoon region, and the interannual precipitation and evapotranspiration have a large fluctuation. In order to reduce the impact of volatility and better explore the role of impact factors on WYs, we take 10 years as a stage, and use the average value of each index for GTWR calculation to find the long-term change trend of the impact of driving factors on WYs.

This study found that NPP could positively drive WYs in fragile ecological areas with an average NPP of less than 200 gc/m², which is in addition to the previous conclusions (Yin et al., 2020). This might be directly related to the nature reserve construction projects and the Conversion of Cropland to Forest Project, projects had the climate and water conservation effects (Liu et al., 2020; Qu et al., 2020). However, in areas experiencing water shortages, such as the Loess Plateau, it is necessary to simulate and evaluate the water consumption of vegetation restoration projects to ensure that the ecological water demand matches local water resources, thereby avoiding construction that further stresses local water resources (Xin and Jérôme, 2019).

4.2. Human factors driving force on WYs

In the densely populated Lanzhou-Baiyin urban agglomeration, human density significantly inhibited WYs, and the area with strong inhibitory effect increased by 142.6%. China has witnessed unprecedented urbanization during the last three decades, with 901.99 million people (i.e., 63.89% of its population) living in urban areas in 2020 (NBSC, 2020). Undoubtedly, human activity and urban expansion has emerged as a highly significant driver of regional climate change (Griggs et al., 2013; Nagendra et al., 2018). The urban, characterized by



high heat-storage capacity, impediments to atmospheric motion, and increased solar radiation absorption, greatly alter the surface energy budget and hydrological cycle, resulting in profound changes of WYs (Grimm et al., 2008; Georgescu, 2015; Jiang et al., 2022).

Land use is one of the most important and lasting ways that humans affect the natural environment and affects the hydrological processes of an ecosystem (Zhu and Woodcock, 2014; Pan et al., 2021; Fang et al., 2022). Land-use intensity significantly negatively affects areas with high WYs because there are significant differences in soil water content, evapotranspiration capacity, litter water holding capacity, and canopy interception among different land-use types (Rutter et al., 1971; Gash, 1979). The higher the land-use intensity, the lower the WYs. The south and northwest of the AEGC have high WYs and are important water conservation areas in the upper reaches of the Yellow River. At the same time, there is low land-use intensity in these areas, which are primarily grassland and forest land and are sensitive to land-use changes (Yang et al., 2018; Xia et al., 2022). Therefore, economic development must be balanced with ecosystem conservation, particularly in water conservation areas with high WYs. Land use should be strictly controlled, full use should be made of existing construction land, and land-use efficiency should be improved, to avoid the decrease in WYs caused by the transformation of ecological land and agricultural land to construction land.

By studying the spatiotemporal heterogeneity and evolution of the factors driving WYs, we found that large-scale regional nature conditions are the primary force driving changes in WYs (Su and Fu, 2013; Wang J. et al., 2021; Wang X. et al., 2021; Jia et al., 2022), while in small-scale regions, human activities and land use are primary forces altering WYs (Rohatyn et al., 2018; Hu et al., 2020; He et al., 2021).

4.3. Limitations and future work

There are certain uncertainties when evaluating WYs. The spatial distribution of meteorological data is affected by factors such as the meteorological stations and interpolation methods; in mountainous areas with vertical effects, the interpolation method will increase uncertainty (Guo et al., 2020; MacDonald et al., 2020). Future research can produce more accurate results by increasing the amount of data and comparing various interpolation methods. While climate and anthropogenic factors are included, the internal relationship between these factors has not been systematically explored (Wang et al., 2016, 2022; Dai et al., 2020, 2022; Cui et al., 2021), which is an area worthy of further research. Research on multi-scale spatial differentiation is the focus of current geographic research, while the spatial differentiation attributes of WYs should be studied at different scales to better reveal the factors influencing WYs spatial differentiation. Climate changing (Jiao K. et al., 2021; Jiao W. et al., 2021) will increase precipitation in Northwest China, also will increase the probability of sudden drought events (Yang et al., 2021; Gong et al., 2022). Therefore, fluctuations in precipitation and temperature increase will likely continue to negatively impact WYs and ecosystems, which is of significant concern.

5. Conclusion

This study explored the spatiotemporal distribution and driving mechanism of WYs in AEGC from 2000 to 2020. To achieve the goal, a comprehensive database of natural conditions and human activities affecting WYs has been built, and algorithms such as Invest model, RF model and GTWR model have been used. The main conclusions are as follows: (1) In recent 20 years, the average annual WYs of AEGC was 110.52 mm, and the interannual variation showed an upward

trend, with an increasing rate of 2.28 mm/a (p < 0.05). WYs are high in the south, low in the north, and high in the northwest. Except for the southeast, WYs remained stable or increased in other regions. (2) The relative importance of the main influencing factors of WYs in AEGC successively were precipitation (1.57), evapotranspiration (1.29), temperature (1.12), population density (1.10), NPP (1.05) and land use intensity (1.02). (3) The variation trend of precipitation is consistent with that of WYs, which is the most important factor affecting the overall WYs in AEGC. In areas with low WYs, evapotranspiration is several times higher than WYs, so WYs is not sensitive to changes in evapotranspiration. Increased precipitation over the Tibet Plateau caused by rising temperatures promoted WYs in the west. Human activities in densely populated areas greatly changed the surface energy budget and hydrological cycle (Georgescu, 2015; Jiang et al., 2021), and the increase of land use intensity significantly inhibited WYs in water conservation areas. At the same time, ecological restoration projects can promote WYs. The results show that: (1) Large-scale regional nature conditions are the primary force driving change in WYs, while in small-scale regions (Feng et al., 2021), human activities and land use are the primary drivers of WYs. (2) The effects of the influencing factors on WYs showed significant spatiotemporal heterogeneity, and the responses of WYs to the changes of the influencing factors were different in different regions.

Data availability statement

The original contributions presented in the study are included in the article/supplementary material, further inquiries can be directed to the corresponding author.

Author contributions

JL and JW designed research. GL performed research. WM, YaL, and YuL modified paper. JL and GL wrote the paper. All

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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