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Numerical stability assessment of a mining slope using the synthetic rock mass modeling approach and strength reduction technique

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Structurally controlled slope failure in open-pit mining occurs when the shear stress acting on the geological structure exceeds its shear strength. Mining slope stability is an extremely important topic from the ramifications of safety, social, economic, environmental and regulatory factors. This study reports the engineering geological setting of a bedded mining slope in China, and evaluates its stability via a numerical approach. First, a slope profile model is constructed using a synthetic rock mass (SRM) modeling approach. More specifically, the mechanical behavior of colluvium, intact rock and discontinuities are represented by linear contact model, bonded particle model and smooth joint model, respectively. Then, the factor of safety (FOS) and instability process are investigated by integrating the discrete fracture network (DFN)-distinct element method (DEM) and strength reduction technique (SRT). In addition, shear stress analyses of colluvium and bedrock are conducted for revealing the potential failure mechanism. Finally, the well-established limit equilibrium (LEM) and finite element method (FEM) are adopted for simulation results comparison and validation.

KEYWORDS

opencast slope, stability assessment, numerical simulation, distinct element method, strength reduction technique, factor of safety

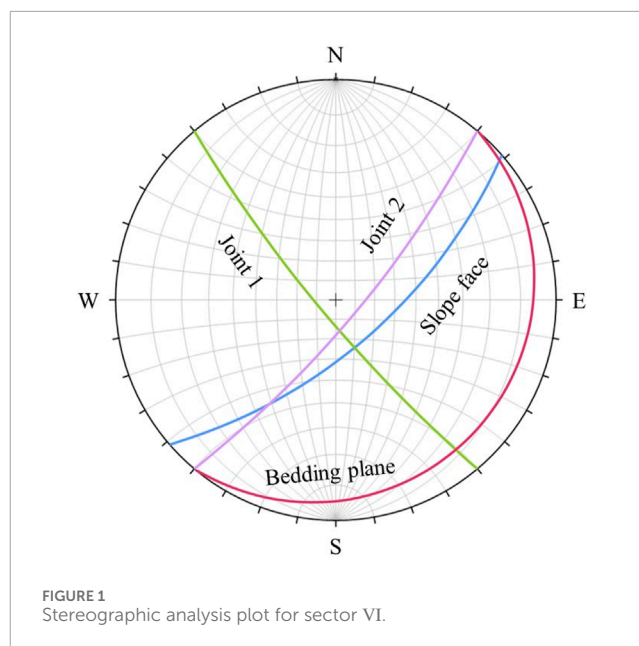
1 Introduction

The uncontrolled rock slope failure in open-pit mining operations and quarries may cause many adverse consequences, such as loss of life, operation disruption and closure consideration (Read and Stacey, 2009). Rock slope stability is mainly governed by the discontinuities within the rock mass, e.g., fault, bedding plane and joint. An unfavorable orientation combination of major structural planes and bedding planes commonly yields a structurally controlled failure, i.e., planar, wedge or toppling (Hoek and Bray, 1981; Stead and Wolter, 2015; Gong et al., 2023). As a result, the geometrical features and mechanical behavior of discontinuities should be of great concern when aiming to evaluate rock slope stability.

The available approaches for opencast slope stability assessment include kinematic analysis, rock mass rating, analytical method, numerical method and probabilistic method (Stead and Wolter, 2015; Raghuvanshi, 2019). Kinematic method aims to investigate the potential failure modes of rock slope based on the angular relationship between discontinuities and slope surfaces, without considering the forces that cause the sliding mass to move (Hoek and Bray, 1981; Yan et al., 2022a). For this reason, it is a pure geometric method. Empirical rock mass classification systems for slope stability condition rating have been well developed (Jaiswal et al., 2024). They are effective tools, but the accuracy is strongly limited by the user's geological experiences. The commonly mentioned analytical method is the limit equilibrium method (LEM). It calculates the factor of safety (FOS) of a generalized rock slope for both moment and force equilibrium using a searching algorithm. Different slip surfaces can be tested by varying the assumptions on the interslice forces (Zhou et al., 2019). Another analytical method is the block theory (Goodman and Shi, 1985). It aims to search key blocks from an excavation slope by integrating geometric topology and simplified LEM. The analytical methods fail to consider the complex stress-strain relationship in the rock mass and usually ignore the resistant shear forces that act on the lateral release surface (Jiang and Zhou, 2017).

To address the limitations in the LEM, numerical simulation methods are developed with the improvement of computation efficiency in the recent decade years. Continuum modelling approaches treat the extremely jointed rock mass as an equivalent medium which is meshed using triangular or quadrilateral elements in 2d (Stead and Wolter, 2015). A constitutive model allows the variety of rock mass mechanical behavior from elastic to elasto-plastic. The most common continuum numerical methods are finite element method (FEM) and finite difference method (FDM) (Bao et al., 2019a; Ren et al., 2020; Chand and Koner, 2024). Discontinuum modelling approaches focus on the simulation of joint media and block separation behavior using force-based contact models. According to the deformability and interpenetrability of the contact objects, distinct element method (DEM) and discontinuous deformation analysis (DDA) method are differentiated (Yan et al., 2020; Ma et al., 2020; Zheng et al., 2021; Cui et al., 2021). Nowadays, discrete fracture network (DFN)-DEM or coupled numerical methods (such as the finite-DEM, FDEM) provide more precise solutions on the simulation of non-persistent joint and crack propagation (Elmo et al., 2022; Li C. et al., 2022; Li X. et al., 2022). The probabilistic methods facilitate to incorporate the variety and uncertainty of rock mass properties by integrating kinematic analysis (Obregon and Mitri, 2019; Yan et al., 2022b), LEM (Du et al., 2022), block theory (Zheng et al., 2015; Xue et al., 2023) or numerical methods (Basahel and Mitri, 2019; Xing et al., 2023; Zhang and Yang, 2023). They express failure criteria in probabilistic terms and provide the probability distribution of the FOS values.

This study describes a bedded opencast slope in China, and evaluates its stability for landslide disaster mitigation. First, a rigorous numerical slope model is constructed using the synthetic rock mass (SRM) modeling approach. Then, its FOS, instability process and failure mechanism are investigated using the DFN-DEM and strength reduction technique (SRT). Finally, the slope stability simulation results are verified by comparing with the LEM and FEM.



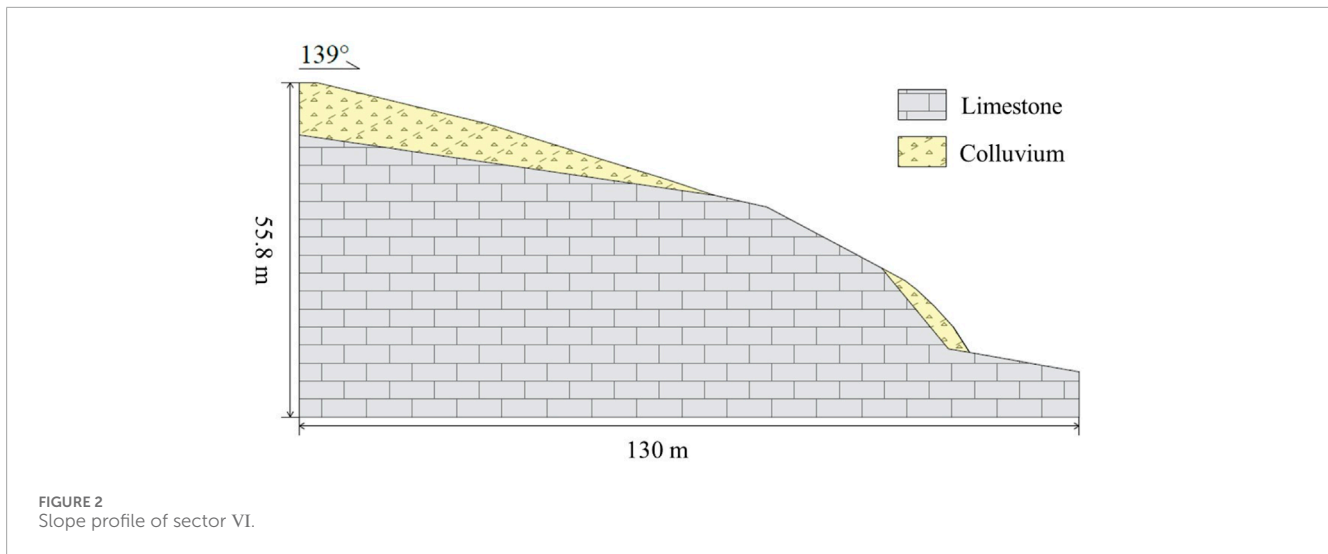
2 Study area

2.1 Engineering geology setting

The studied open-pit mining area is located in the Hubei province, China. It is in a hill region that covered well with plants, and the relative relief height is 120 m. About 50% of the annual precipitation is concentrated from June to September, and the maximum monthly precipitation is 1,600 mm. The tectonic setting is relatively stable and no large faults cut through the mining area. The main lithology of the mining area is Ordovician limestone with a bedding orientation of $130^{\circ}\angle 8^{\circ}$, and karrens are observed in the excavation face. Additionally, two steeply inclined discontinuity sets are developed in the rock mass. Their orientation are $230^{\circ}\angle 81^{\circ}$ and $130^{\circ}\angle 78^{\circ}$, respectively. Since there are no known major structures (e.g., faults) within the immediate mining area, the stability of opencast side slope is mainly controlled by the orientation combination of minor structures (Obregon and Mitri, 2019), such as bedding planes and joints.

2.2 Slope failure mode analysis

Nine sectors are differentiated according to the slope dip direction. They are numbered from I to IX. Stereographic analyses are conducted for all sectors, and one plot indicates that sector VI is a consequent bedding rock slope (Figure 1). Of these, the high steep excavation face is the free surface; the gentle dip bedding plane is the potential sliding surface; and the steeply inclined joints are lateral release surfaces. Under the combination of given orientations, block instabilities or even cataclinal rock slides may be triggered by torrential rainfall or human activities (Wang F. et al., 2022). For this reason, the stability of the sector VI should be carefully investigated.



2.3 Slope material properties

A typical pit slope profile of sector VI is shown in Figure 2. The slope height is 55.8 m. Colluvium deposits with an average thickness of 4.5 m are distributed on the upper and lower slope. These loose materials are a mixture of gravel (25%) and silty clay (75%). The gravel is composed of non-rounded limestone block with a diameter mainly distributed between 5 and 15 cm. The internal friction angle (ϕ_c) and cohesion (C_c) of the colluvium deposits refer to a laboratory direct shear test presented by She et al. (2018). They are 30° and 87.8 kPa, respectively. The colluvium deposits are in unconformable contact with the underlying heavy-layer limestone of the Dawan Formation. The spacing of the bedding plane is about 4 m. The strength and deformation parameters of the bedrock refer to a laboratory uniaxial compression test presented by Wang X. et al. (2022). The uniaxial compressive strength (UCS), elastic modulus (E) and Poisson ratio (ν) are 64.5 MPa, 25.1 GPa and 0.24, respectively. The shear strength parameters of discontinuities refer to the empirical values suggested by the Standard for Engineering Classification of Rock Mass. The corresponding internal friction angle (ϕ_j) and cohesion (C_j) are 42° and 0.4 MPa, respectively.

3 Methodologies

3.1 Distinct element method

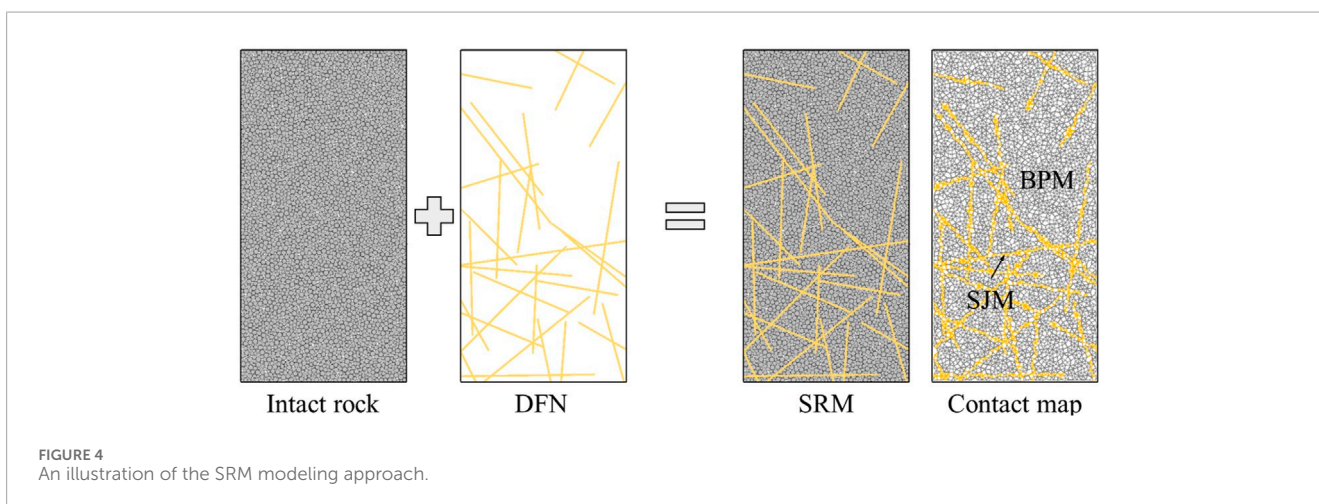
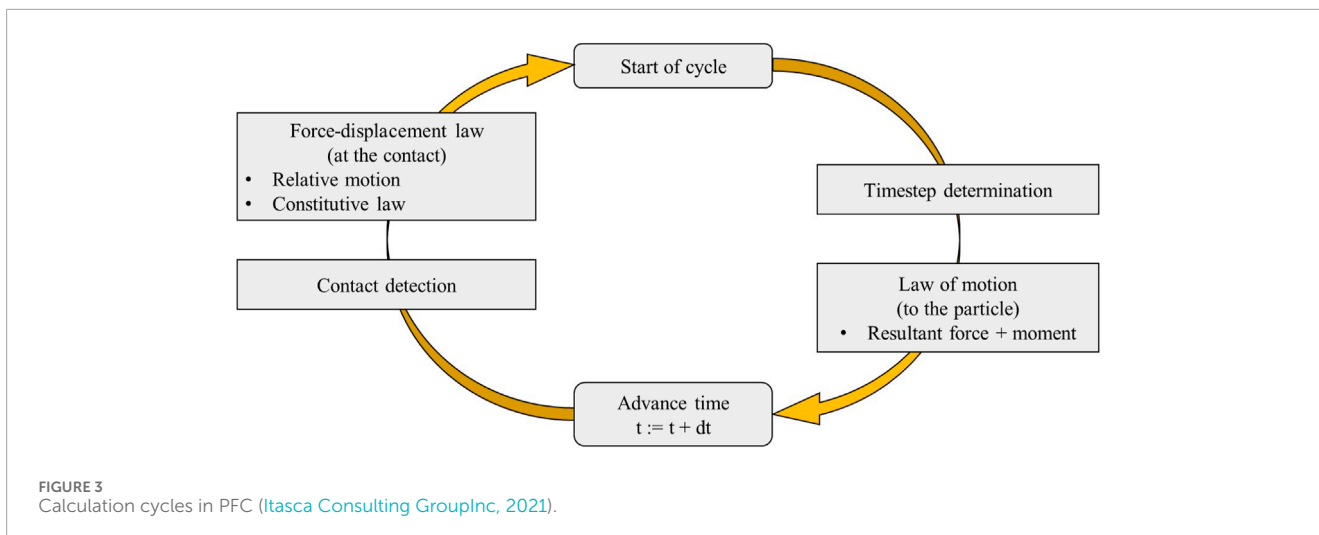
The distinct element method (DEM) introduced by Cundall (1971) was originally used for the analysis of rock mechanics problems. It was later applied to granular material by Cundall and Strack (1979). The Particle Flow Code (PFC) can be viewed as a simplified implementation of the DEM because of the strict assumption on non-deformable particles (Itasca Consulting Group Inc, 2021). The dynamic behavior is represented numerically by a timestepping algorithm, and the solution scheme is identical to the explicit finite-difference method in continuum analysis. The calculation cycles performed in the PFC

alternate between the application of Newton's second law to the particles and a force-displacement law at the contacts (Figure 3). Following this, the velocity and displacement of a particle within each timestep can be obtained. In this study, the two dimensional PFC (PFC2D) is used for pit slope stability assessment.

3.2 Slope numerical model construction

Ball and wall are two fundamental body types in the PFC. Of these, ball is a rigid disk with unit thickness in 2D, while wall is a manifold surface composed of lines segments in 2D termed facets (Itasca Consulting Group Inc, 2021). Using the two elements, ball-wall and ball-ball models are available for construction of a slope numerical model (Li et al., 2012; Wei et al., 2019). The latter model is adopted mainly because that the slip surface has not been determined in this case.

Regarding the slope materials, colluvium deposits are simulated by an assembly of balls. The large and small ball elements represent soil and gravel, respectively. Note that the mass percentage within two size ranges are in accordance with onsite investigation. The mechanical behavior of colluvium deposits is modeled using the linear contact model. The bedrock is simulated using the synthetic rock mass (SRM) modeling approach (Esmaili et al., 2010; Mas Ivars et al., 2011). To facilitate reading, an illustration of this approach is shown in Figure 4. Note that the superiority of SRM in characterizing fractured rock mass has been examined by some recent published studies (Bester et al., 2021; Zhang et al., 2021; Xing et al., 2023), thus this study follows their modelling procedures. First, the bonded particle model (BPM) proposed by Potyondy and Cundall (2004) is used to represent intact rock. Then, discrete fractures network (DFN) are generated in the intact rock to represent the *in-situ* joint network. Finally, by assigning the smooth-joint contact model (SJM) to the ball-ball contact that intersects with fractures, the mechanical behavior of jointed rock mass can be simulated. Regarding the boundary conditions, virtual and boundary walls are separated for model construction (Bao et al., 2019b). Of these, the former is used to create the shape of the



slope and partition material zones where particles can be generated; and the latter is used to fix the displacement of the bottom particles.

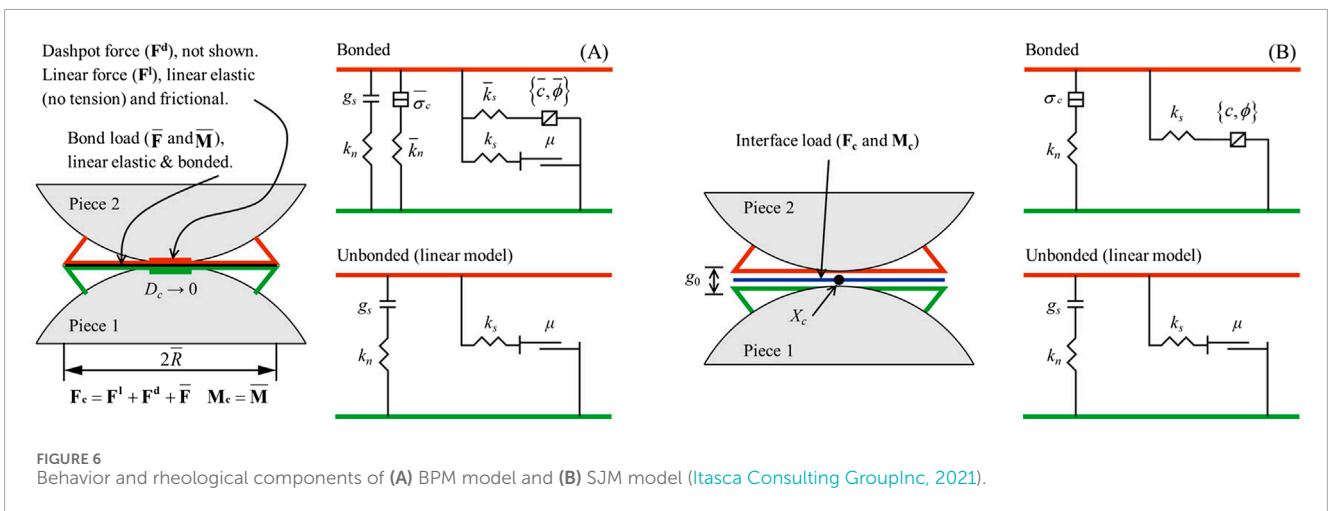
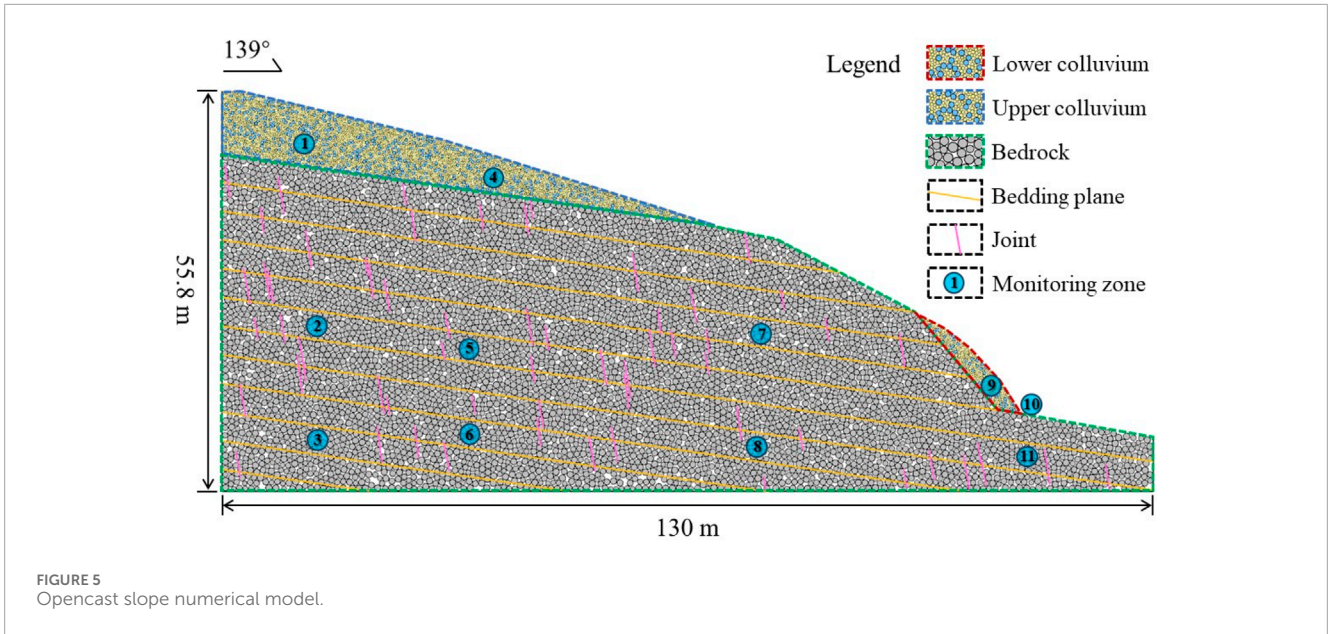
The area of colluvium deposits in the slope profile is 413.9 m², and is filled by 6067 balls with radius between 0.1 and 0.15 m for soil and 0.2 and 0.25 m for gravel (yellow and blue particles in Figure 5). Upper and lower colluvium herein are differentiated by their elevations. The area of bedrock is 4,202.4 m² where an assembly of 5914 balls with radius between 0.35 and 0.5 m are generated (gray particles in Figure 5). The number of bedding planes and joints are defined as 7 (orange lines in Figure 5) and 46 (pink lines in Figure 5), respectively. Their size and position are determined based on a detailed field investigation. Considering the significant differences in the stochasticity and persistence of the two types of geological structures, multiple geometrical parameters are assigned. More specifically, the position of bedding planes is in accordance with the field rock slope exposure, and full-persistence is assumed for such geological structure. For comparison, the joints in the slope numerical model follows a uniform distribution, and non-persistence is assumed for them. Furthermore, interval values (79–83° and 2.5–5.5 m) are assigned to

simulate the variety of orientation and trace length involved in the dominant joint sets.

3.3 Contact models and micro-parameters

Contacts are created/deleted automatically during cycling. They are detected via termed piece based on the principles of different contact models. The contact types of piece are ball-ball and ball-facet. The aforementioned BPM and SJM are two built-in contact models in the PFC.

The BPM model provides the behavior of two interfaces: a linear model and a parallel bond (Potyondy and Cundall, 2004) (Figure 6A). In the linear model, rotation is allowed, and slip is adapted to the Coulomb limit on the shear surface. Linear contact is active if the surface gap is less than or equal to zero, and can be activated again when the criterion is satisfied. The micro-parameters of linear group are mainly: friction coefficient (μ), normal stiffness (k_n) and shear stiffness (k_s). The parallel-bond component acts in parallel with the linear component and establishes an elastic interaction between the pieces. The parallel bonds are broken when the threshold bond strength is exceeded,



and cannot be reactivated. The parallel-bond group includes the following micro-parameters: bond normal stiffness (\bar{k}_n), bond shear stiffness (\bar{k}_s), tensile strength ($\bar{\sigma}_c$), cohesion (\bar{c}) and friction angle ($\bar{\theta}$).

In the SJM model, smooth joints are created by removing the bond between particles and applying a set of elastic spring uniformly over a rectangular cross section (Mas Ivars et al., 2011) (Figure 6B). Particle pairs intersected by a smooth joint may overlap and pass through each other rather than forced to move around one another (Bahaaddini et al., 2015). The smooth joints remain active while there is a nonzero overlap between particles. The main micro-parameters in the SJM model are normal stiffness per unit area (k_n), shear stiffness per unit area (k_s), friction coefficient (μ), tensile strength (σ_c), cohesion (C^*) and joint friction angle (θ^*).

PFC derives macro-scale material properties from the interactions among micro-scale particles (Lu et al., 2014; Yan et al., 2022a). Unfortunately, the universal or accurate analytical equations between micro- and macro-parameters have not been established

(Tang et al., 2009; Lu et al., 2014; Lin; Lin and Lin, 2015; Wei et al., 2019). Alternatively, the trial test was used to determine the appropriate micro-parameters. It calibrates micro-parameters from the macroscopic response of a particle assembly by repeating numerical tests until the modeled macro-parameters approach to the laboratory values (Bao et al., 2020; Yan et al., 2022b). For linear contact model and SJM, their micro-parameters can be obtained after several rounds of failed attempts. However, a rational strategy should be adopted for calibrating the micro-parameters of the BPM. Fortunately, some previous studies suggested a straightforward flowchart (Zhang et al., 2012; Lu et al., 2014; Bahaaddini et al., 2015). First, E is matched by adjusting k_n, \bar{k}_n , particle normal/shear stiffness ratio k_n/k_s and \bar{k}_n/\bar{k}_s and parallel bond normal/shear stiffness ratio (\bar{k}_n/\bar{k}_s). This is followed by calibration of the ν which is influenced by k_n/k_s and \bar{k}_n/\bar{k}_s . Finally, the best fit of UCS is found by adjusting $\bar{\sigma}_c, \bar{c}$ and $\bar{\theta}$.

Using the trial and error test, numerical direct shear tests of colluvium and joint and numerical uniaxial compression test of

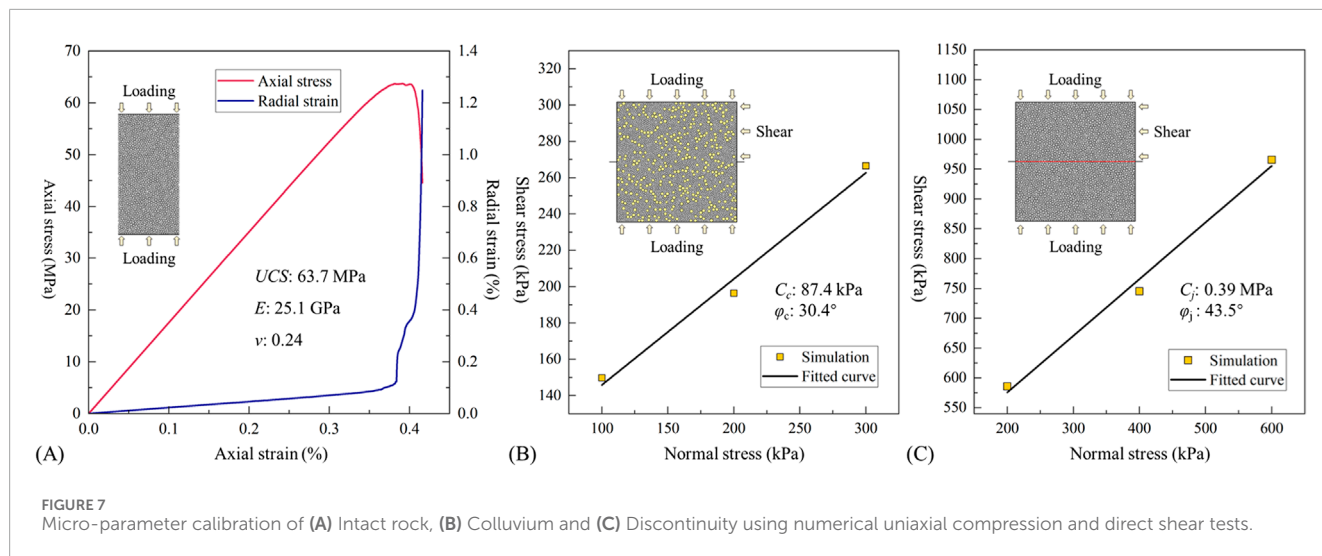


TABLE 1 Micro-parameters of slope materials.

Material	Model name	Micro-parameters							
Colluvium	Linear	μ (-)		k_n (N·m ⁻³)		k_s (N·m ⁻³)			
		0.4		$1 \times 10^7 - 2 \times 10^8$		$1 \times 10^7 - 1 \times 10^8$			
Limestone	BMP	Linear group			Parallel-bond group				
		μ (-)	k_n (N·m ⁻³)	k_s (N·m ⁻³)	\bar{k}_n (N·m ⁻³)	\bar{k}_s (N·m ⁻³)	$\bar{\sigma}_c$ (MPa)	\bar{c} (MPa)	$\bar{\theta}$ (°)
		0.6	2×10^8	1×10^8	1×10^9	1×10^9	50	30	30
Fracture	SJM	k_n (N·m ⁻³)		k_s (N·m ⁻³)		μ (-)	σ_c (MPa)	C^* (MPa)	θ^* (°)
		2×10^{14}		2×10^{14}		0.7	0	0	0

intact rock are repeated by adjusting micro-parameters until the simulated macro-parameters are similar to the reference values (Figure 7). The calibrated micro-parameters of slope materials are listed in Table 1. Then, a balanced numerical slope model is established using these micro-parameters, and a contact map is shown in Figure 8.

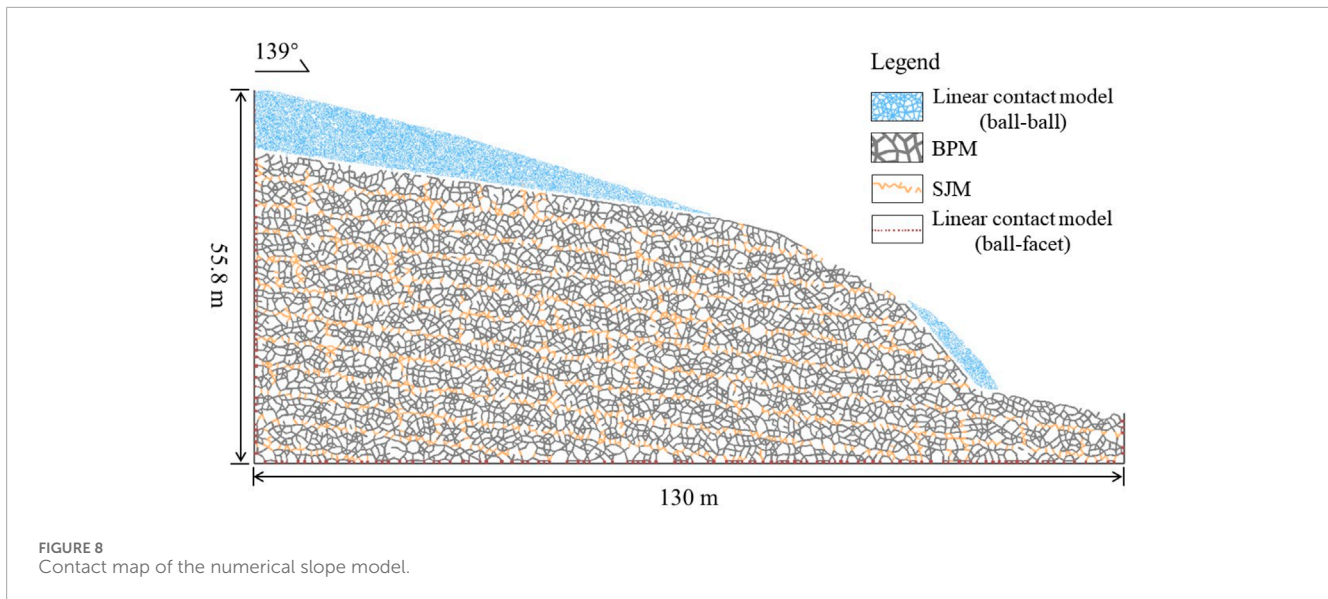
3.4 Strength reduction technique

Factor of safety (FOS) is the most widely used quantitative index in slope stability assessment. The concept of the FOS is derived from the limit equilibrium method (LEM), and can be expressed as Equation 1. To combine FOS with numerical approach, strength reduction technique (SRT) was first introduced by Zienkiewicz et al. (1975). Following this, SRT has gradually become a well-recognized technique for FOS calculation in the finite element (Jiang et al., 2015; Bao et al., 2019a; Jia et al., 2024) or distinct element (Bao et al., 2019b; Su et al., 2019; Li Y. et al., 2022; Xing et al., 2023) analyses of slope stability.

$$FOS = \frac{\int (C + \sigma \tan \phi) dA}{\int \tau dA} \tag{1}$$

where C and ϕ are cohesion and internal friction angle, respectively; σ is the normal stress; τ is the shear strength and A is the area of sliding surface.

In the finite element analysis, the SRT defines two new shear strength parameters (C_R and ϕ_R) by dividing the actual shear strength parameters (C and ϕ) to a strength reduction factor (SRF). By modifying SRF, a set of C_R and ϕ_R are generated for finite element calculations. Of these, the FOS is equal to the SRF when the slope is in a limit failure state. Regarding the distinct element analysis, the procedures are the same as finite element analysis except that the macro-scale parameters need to be replaced by the micro-parameters. Considering that the micro-parameters in the BPM and SJM are more than twenty, the reduction parameters can hardly be determined. For this reason, an alternative approach named “gravity increase method” reported by Li et al. (2009) is adopted. In this method, the FOS is defined as the ratio between the gravitational



acceleration at the critical slope instability state and the natural gravitational acceleration.

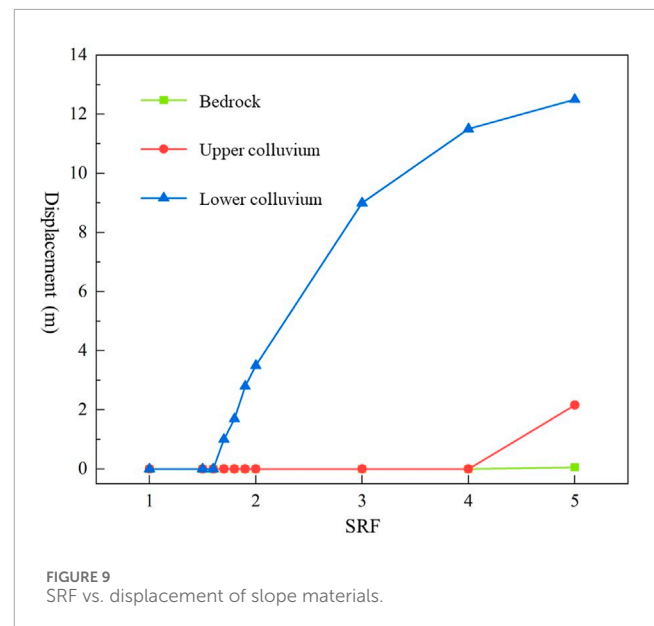
Numerical non-convergence criterion, plastic yield zone connection criterion and displacement mutation criterion are commonly used for judging the limit failure state of a slope (Jiang et al., 2015). Of these, the displacement-based criterion is more clear in the distinct element analysis (Li C. et al., 2022; Xing et al., 2023) and is adopted in this study. Using the SRT, the instability dynamic process of the studied rock slope can be simulated and the FOS can also be obtained.

4 Results

4.1 Stability of the opencast slope

As mentioned above, FOS is equal to SRF when the slope reaches a limit failure state in the SRT. From this, SRF should be preliminarily set as an interval value for seeking the interest value corresponding to the critical stability state. In this study, SRF refers to previous papers (Bao et al., 2019a; Su et al., 2019; Li X. et al., 2022), as well as considers the current geological model. The trial SRF values are tentatively created from 1.0 to 5.0 with an increment of 1.0. The relationship between SRF and the displacement of different slope materials is plotted in Figure 9. The modeled slope is stable when the SRF is equal to 1; and that exhibits significant failure when the SRF is equal to 5. From this, the pre-defined interval value is reasonable.

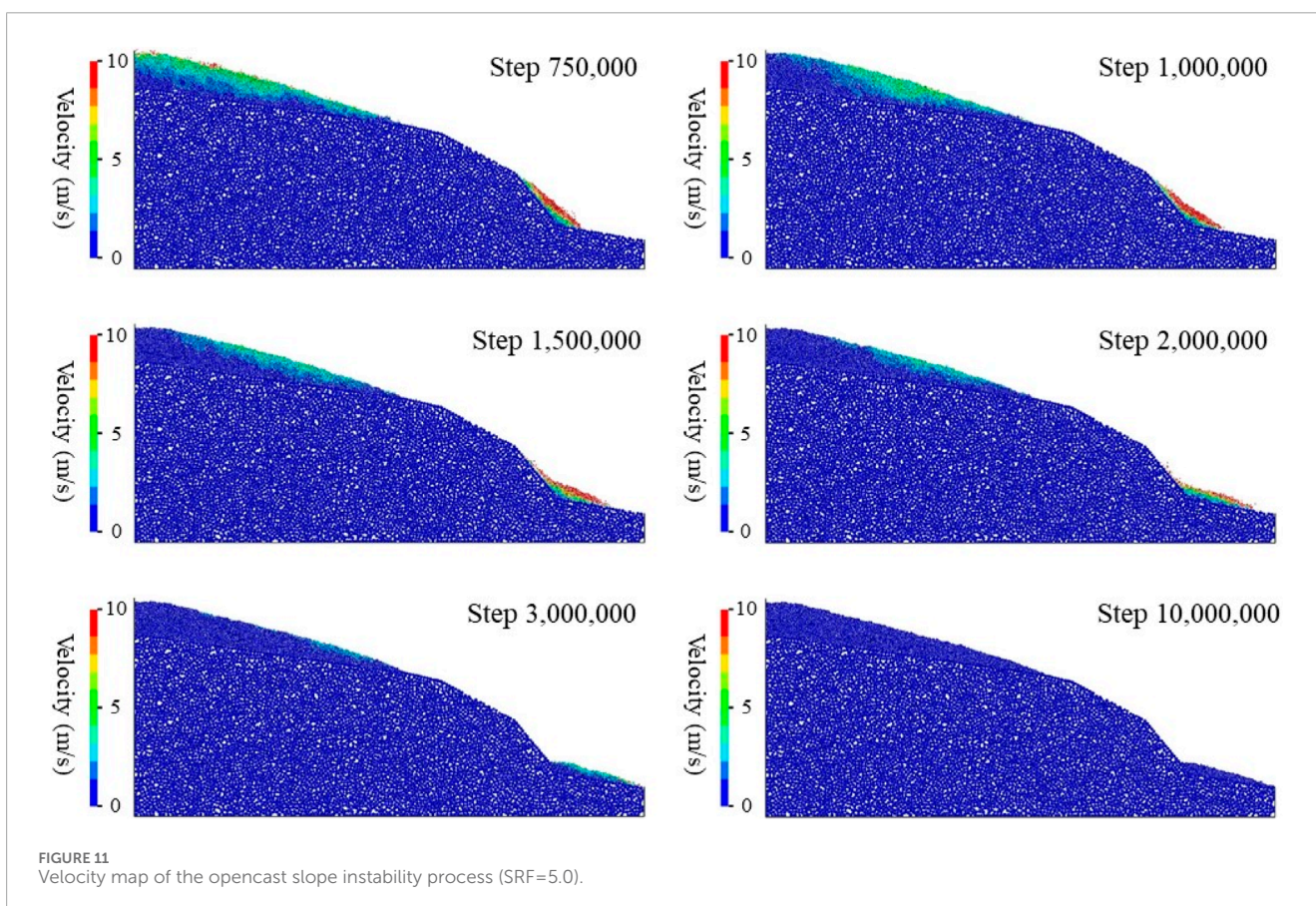
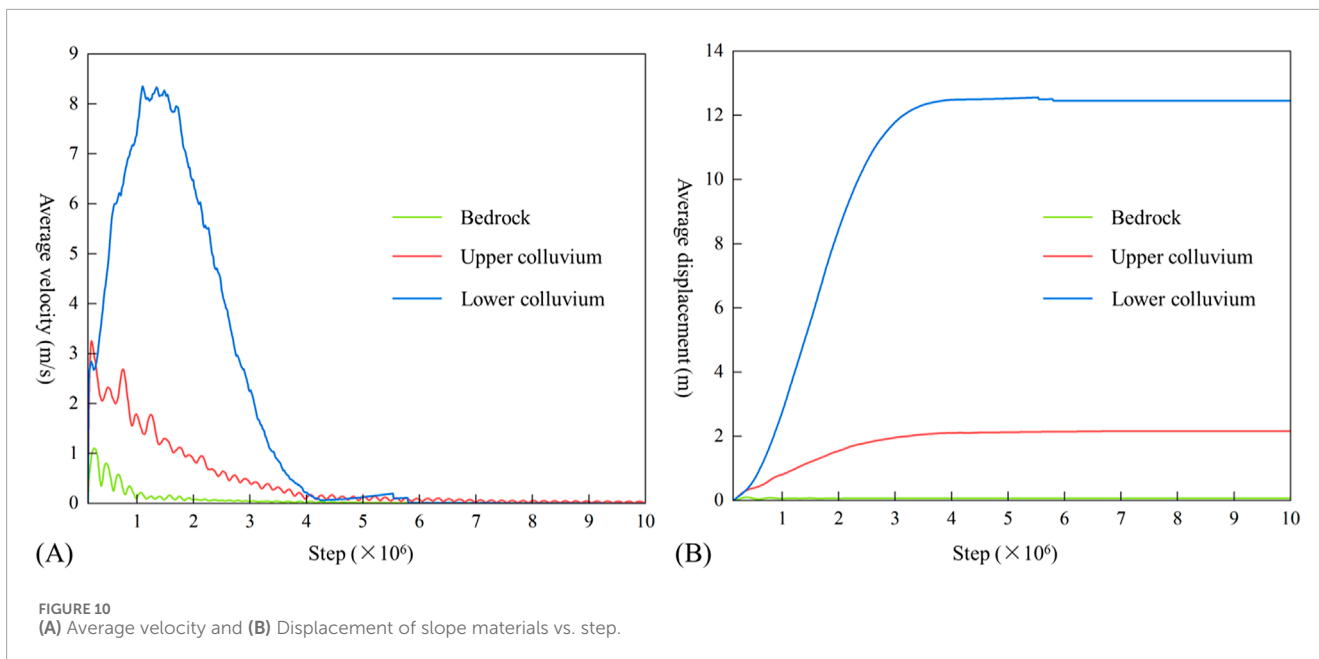
More specifically, the average displacement of the lower colluvium is gradually greater with increasing the SRF values, but the corresponding displacement increment reduces gradually (Figure 9). In contrast, the displacement of the upper colluvium and bedrock is not sensitive to the SRF values. The FOSs of the upper colluvium and bedrock are greater than 4.0 and 5.0, respectively. To obtain a more accurate FOS value for the lower colluvium, a total of four SRF values with a range between 1.5 and 2.0 with an increment of 0.1 are created. The SRF value corresponding



to the displacement mutation point is 1.7. Therefore, the current opencast slope is stable in general.

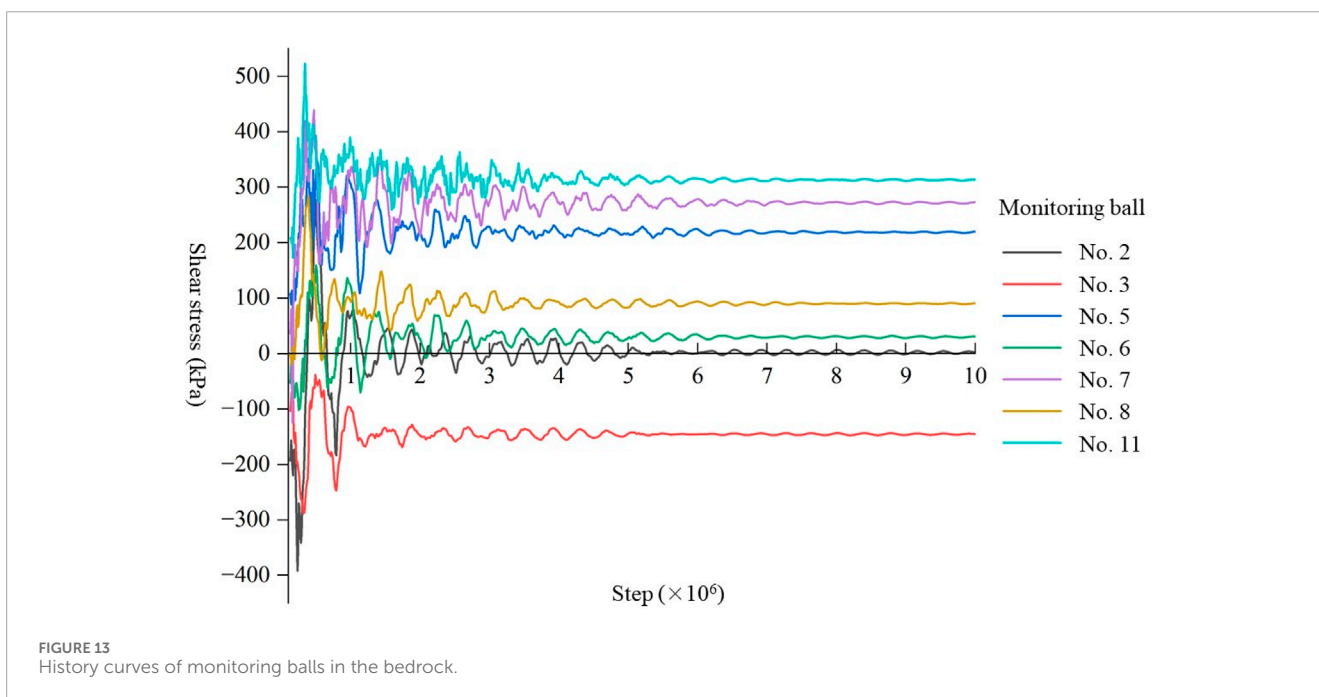
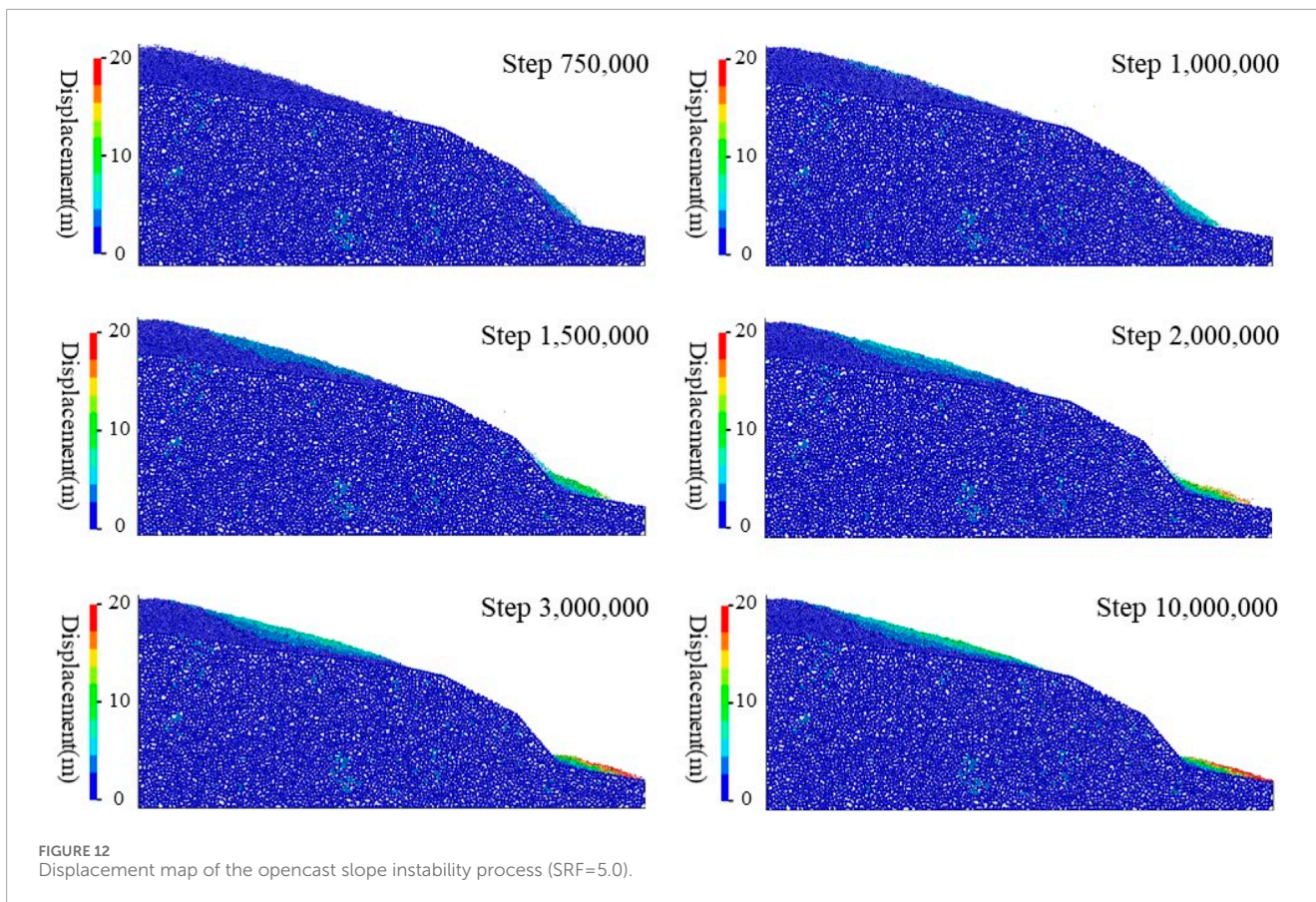
4.2 Slope instability process

The scenario with a SRF equal to 5.0 is used to investigate the potential failure process of the opencast slope. The calculation cycles are terminated when the velocity of slope materials is less 0.01 m/s. The history curves of velocity and displacement of different slope materials are plotted in Figure 10. The major instability process lasts 4×10^6 steps. Of these, the lower colluvium experiences a significant acceleration stage and a deceleration stage. The maximum velocity and displacement are 8.36 m/s and 12.56 m, respectively. The velocity of the upper landslide deposit decreases



rapidly after initiation, and the final displacement is only 2.16 m. The velocity of the bedrock converges to 0.1 m/s after 1×10^6 steps, and the associated displacement curve is a near-horizontal line. It indicates that bedrock remains stable even under a high SRF, thus the risk of planar failure is extremely low.

The velocity and displacement maps of the slope instability process are shown in Figures 11, 12, respectively. Different instability processes are observed between the upper and lower colluvium because of the difference of slope angle. Regarding the upper gentle colluvium, failure occurs in the shallow layer and a circular



slip surface is observed. Maximum deformation of this shallow layer is approximately 10 m, and some surficial particles fall at the toe. For comparison, deformation of the deep layer is not obvious. As for the lower steep colluvium, it slides along the

lithologic interface under gravity. More specifically, it accelerates away from the slope toe in the first 2×10^6 steps. Subsequently, the sliding mass moves slower with decreasing slope angle. Finally, the sliding mass deposits at the slope toe.

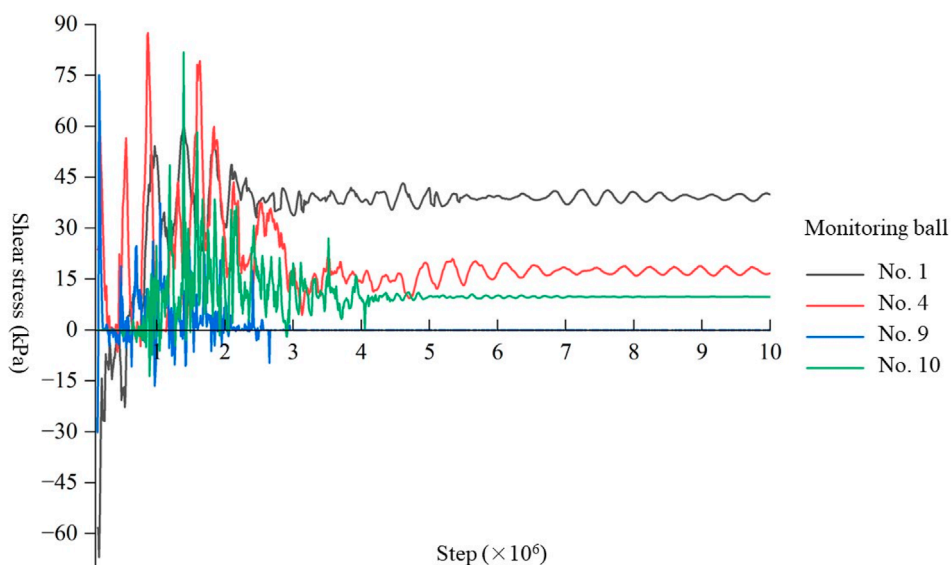


FIGURE 14 History curves of monitoring balls in the colluvium.

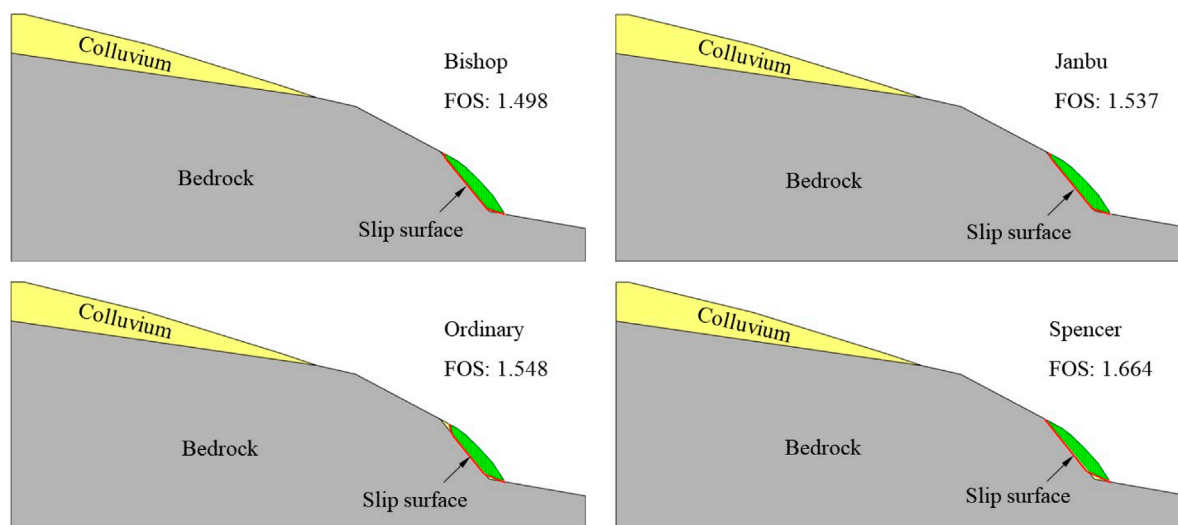


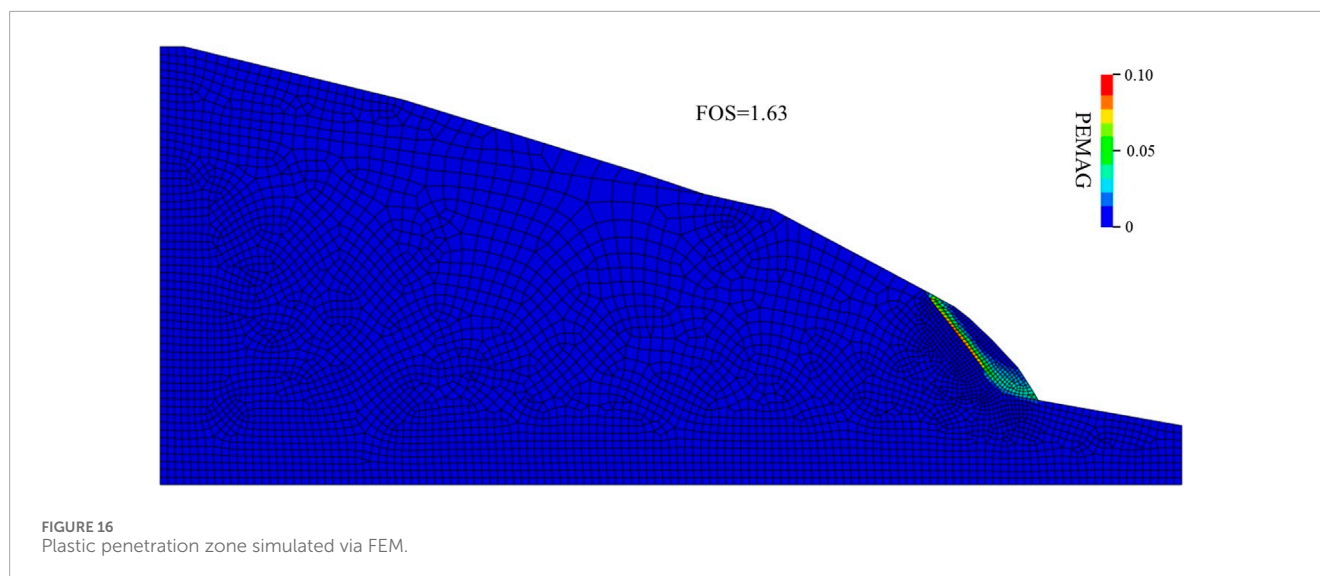
FIGURE 15 Slope stability assessment via LEM.

Affected by the relative larger friction coefficient of the bedrock, high velocity layer is observed in the surficial of the sliding mass, while low velocity layer is observed at the bottom of the sliding mass (Figure 11). Since the acceleration distance is very short, the expansion of the landslide deposits is not obvious. Besides the effect of friction coefficient of slip surface on the particle motion, the front particles also obstruct the latter particle motion. For this reason, the front surficial particles have a relative larger displacement (Figure 12).

4.3 Shear stress analysis for slope instability

To investigate the relationship between the variation of shear stress and slope instability process, a total of eleven monitoring balls are set in the slope (Figure 5). Of these, two for the upper colluvium, two for the lower colluvium and seven for the bedrock.

The shear stress in bedrock fluctuates strongly in the first 1×10^6 steps (Figure 13), which is in accordance with the variation characteristics in the velocity map. After 5×10^6 steps, the shear



stress of each monitoring ball converges to a constant. The general features of the modelled shear stress field are: 1) the shear stress of the monitoring balls near the slope face is greater than that inside the slope; 2) stress concentration is most significant at the slope toe; and 3) shear stress is related to the slope height. From this, bedrock is generally stable because that the shear strength of rock mass is greater than the shear stress. The material at the slope toe is prone to instability, which is consistent with the simulation results (Figures 11, 12).

The history shear stress in the colluvium fluctuates more strongly than that in the bedrock in the first 4×10^6 steps (Figure 14). It is because that instabilities mainly occur in the colluvium layer. Since the lower colluvium has completely slid from its original region, the rear portion of the history curve of the monitoring ball No. 9 is missing. The shear stress in the colluvium is smaller than that in the bedrock due to the differences in the contact models. The lowest shear stress (9.79 kPa) is observed in the lower colluvium because that the landslide deposits reaches a more stable equilibrium state. In general, the modelled shear stress field well explains the mechanism of slope instability process.

5 Discussion

In this section, both LEM and FEM are used to make a result comparison with the DEM. Regarding the LEM, similar slip surfaces are calculated by varying four searching algorithms, i.e., Bishop, Janbu, Ordinary and Spencer (Figure 15). The slight differences in the FOS results are caused by the different pre-defined assumptions on the interslice forces. LEM derives FOS from macro-parameters (e.g., cohesion and internal friction angle), and the complex strain-stress conditions in a slope rock mass is not considered. For comparison, DEM derives FOS from the micro-view of the slope instability process by solving the particle displacement and contact force iteratively. Therefore, different results are obtained from the two methods.

Further comparison is conducted by using the FEM and SRT. Figure 16 shows the simulated plastic penetration

zone of the opencast slope, and the corresponding FOS is 1.63. Both FEM and DEM are numerical methods, and utilize SRT to find the limit failure state. However, since the difference in calculation rationale between continuum-based and discontinuum-based methods (e.g., particle assembly and continuum mesh), as well as the ambiguous relationship between macro- and micro-parameters, different results are also observed from the two methods.

Based on the visible distinction among the methods and results, an argument about “How to model more rationally to make the results more referable” is introduced here. In the authors’ opinion, three aspects should be fully considered, i.e., numerical method, slope numerical model and adequate parameter. For the first aspect, continuum-based or discontinuum-based numerical methods should be determined according to the rock mass structure of the studied slope. In other words, FEM-SRT with macro-parameters is suggested for cataclastic slope, while DEM-SRT is more recommended for fractured rock slope. In this study, the latter is adopted due to the slope mass is mainly composed of bedded rock mass. For the second aspect, a numerical slope model should be consistent with the actual geological conditions. In this study, colluvium deposits are simulated by an assembly of balls, and the linear contact model is assigned. The bedrock is simulated using the SRM modeling approach. Of these, the mechanical behavior of intact rock and discontinuities are represented by BPM and SJM models, respectively. It seems that slope materials are explicitly distinguished by assigning different models, which improve the accuracy of the simulation results. For the third aspect, parameters should refer to laboratory tests or remarkable studies. If conditions permit, comparison with *in-situ* monitoring devices can also verify the simulation results and further make the results more referable.

6 Conclusion

The strength reduction DFN-DEM modelling approach is used to investigate the stability of a bedded mining slope. The major findings of this study are summarized as follows. Using a displacement mutation criterion, the FOS of the mining slope is

equal to 1.7. The possible sliding mass and slip surface are the lower colluvium with a steep angle and the lithologic interface, respectively. Regarding the underlying bedrock, it is still stable even though the SRF is set as 5.0. A shallow failure for the upper colluvium and a translational slide for the lower colluvium are observed in the simulated instability process. They are explained by a shear stress analysis. LEM and strength reduction FEM yield a lower FOS, but the searched slip surfaces are similar. This can be attributed to the differences of model parameters and media assumptions among the three methods. To make a simulation result more referable, numerical method, slope numerical model and adequate parameter should be well considered on the basis of the actual geological conditions of a rock slope.

Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

Author contributions

LT: Data curation, Formal Analysis, Funding acquisition, Methodology, Resources, Software, Validation, Writing—original draft. YH: Conceptualization, Investigation, Project administration, Writing—review and editing. YW: Funding acquisition, Validation, Visualization, Writing—review and editing. CS: Formal Analysis, Visualization, Writing—review and editing. JY: Funding acquisition, Writing—review and editing.

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Conflict of interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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