Check for updates

OPEN ACCESS

EDITED BY Cong Zhou, East China University of Technology, China

REVIEWED BY Octavio Castillo Reyes, Barcelona Supercomputing Center, Spain Ronghua Peng, China University of Geosciences Wuhan, China

*CORRESPONDENCE Dajun Li, ⊠ lidajun@jlju.edu.cn

RECEIVED 16 April 2023 ACCEPTED 26 June 2023 PUBLISHED 13 July 2023

CITATION

Li D, Wang Z, Li Y and Jin L (2023), A general forward solver for 3D CSEMs with multitype sources and operating environments. *Front. Earth Sci.* 11:1206784. doi: 10.3389/feart.2023.1206784

COPYRIGHT

© 2023 Li, Wang, Li and Jin. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) and the copyright owner(s) are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.

A general forward solver for 3D CSEMs with multitype sources and operating environments

Dajun Li^{1*}, Zhiqiang Wang¹, Yabin Li² and Liubiao Jin¹

¹College of Surveying and Prospecting Engineering, Jilin Jianzhu University, Changchun, China, ²College of GeoExploration Science and Technology, Jilin University, Changchun, China

To determine the electromagnetic (EM) fields of different three-dimensional (3D) controlled-source electromagnetic methods (CSEMs) using the same parameters of the forward solution, by explicitly considering the commonalities, we present a general 3D forward modeling solver for CSEMs with multitype sources and operating environments. The commonality of the solver is reflected in two aspects. First, the solver is based on a frequency-domain (FD) vector Helmholtz equation for determining the scattered electric field. The different types of sources are imposed on the right-hand term of the equation, expressed as background Green's function. Second, sources of any CSEM can be composed of electric dipole (ED) or magnetic dipole (MD) superposition. Thus, the focus of the 3D forward modeling of CSEMs is reduced to determining the EM fields of ED or MD sources for the background medium. The quasi-minimal residual (QMR) method is used to solve the large sparse complex linear system. Once the FD EM fields have been calculated, the time-domain (TD) response can be obtained using the cosine/sine transformation. The numerical results show that the relative error is less than 5% between the 3D numerical and analytical solutions, which verifies the accuracy of the solver. We further study the difference between the real (bent) and theoretical (straight) wires. We suggest that the shape of the source must be considered for TD and FD CSEMs with a wire source during data processing and inversion. The last example investigated the characteristics of FD EM fields from a finite-length wire and TD EM fields from a rectangular fixed loop on the same conductive tilted disk model buried in resistive sediments. According to the numerical results, we recommend FD CSEMs with a wire source for detecting deep anomalies.

KEYWORDS

CSEMs, 3D forward modeling solver, multitype sources, operating environments, frequency/time domain

1 Introduction

Controlled-source electromagnetic methods (CSEMs) are a group of geophysical exploration methods that transmit an electromagnetic (EM) signal using an artificial source (Goldstein and Strangway, 1975; Zonge and Hughes, 1991; Constable and Srnka, 2007; Di et al., 2020). CSEMs exhibit various classifications based on different factors, such as the type of source (e.g., wire, loop, electric, and magnetic dipole) and the operating environment (e.g., land, marine, airborne, and borehole), including marine frequency-/ time-domain EM methods (mFD/TDCSEMs) (Edwards, 2005; Um and Alumbaugh, 2007; Connell and Key, 2013), long- and short-offset transient EM methods (L/SOTEMs) (Commer and Newman, 2004; Xue, 2018), controlled-source audio-frequency

magnetotelluric (CSAMT) methods (Weng et al., 2012; Yang and Oldenburg, 2016), and land/airborne transient EM methods (TEMs) (Yin et al., 2016; Li et al., 2018; Zhang et al., 2018). By examining and analyzing the distribution patterns of the EM fields of CSEMs associated with variations in the resistivity of underground media, geophysicists can explore mineral and hydrocarbon resources (Hu et al., 2013; Streich, 2016; Schaller et al., 2018; Castillo-Reyes et al., 2022) and address geological and environmental engineering challenges (Everett, 2009; Chave et al., 2017; Malovichko et al., 2019; Castillo-Reyes et al., 2022).

Forward modeling is an effective way to study the EM field laws of CSEMs, and it is also the premise and basis of inversion methods (Avdeev, 2005; Börner, 2010). Over the past decades, innovations in numerical calculations have driven remarkable progress in the threedimensional (3D) forward modeling of CSEMs, achieving successful breakthroughs (Constable, 2010; Ansari and Farquharson, 2013; Börner et al., 2015; Oldenburg et al., 2020; Werthmüller et al., 2021). The common methods used for developing the 3D forward modeling of CSEMs include finite element techniques (Tonti, 2002; Um et al., 2012; Cai et al., 2014; Rochlitz et al., 2019; Castillo-Reyes et al., 2022), finite volume methods (Haber and Ascher, 2001; Ren et al., 2017; Peng et al., 2018), finite difference methods (Newman and Alumbaugh, 1995; Weiland, 1996; Li et al., 2022), integral equation techniques (Hursán and Zhdanov, 2002; Tang et al., 2018), and spectral element methods (Huang et al., 2019; Xu and Tang, 2022). Furthermore, numerous scholars have dedicated research efforts to enhance the efficiency of numerical solutions in 3D forward modeling for CSEMs. They have employed parallel programming and numerical computation platforms to accelerate the computation speed (Unno et al., 2012; Koldan et al., 2014; Castillo-Reyes et al., 2018; Castillo-Reyes et al., 2019; Castillo-Reyes et al., 2022; Liu et al., 2023).

By employing forward modeling of different CSEMs, geophysical companies, scientific institutions, and individuals can compare the characteristics of EM responses and determine which CSEM offers the strongest resolution for specific geological targets while optimizing the survey parameter design before conducting field studies. However, the majority of forward modeling codes have been developed by various scientific institutions using different programming languages and focusing on a single CSEM with a specific source type and operating environment. The code of each CSEM is based on different essential mathematical procedures and utilizes different parameters for forward solutions, such as grids, interpolations, stations, numerical solution methods, and staggering schemes. These variations in parameterization can introduce biases during resolution analysis. Meanwhile, the expensive cost of secondary development for those codes poses practical challenges. The application effectiveness of a CSEM is influenced by the complexity and risk associated with the forward modeling theory. Technology implementation is the underlying support for practical applications. Therefore, it is necessary to decouple practical applications and concrete mathematical technology to advance geophysical science.

To simplify the forward theory and analyze the EM characteristics of different CSEMs for the same targets using the same parameters of the forward solution and mathematical procedures, we integrated forward modeling technologies and developed a forward solver for CSEMs with multitype sources

and operating environments. In this study, our approach involves decomposing the total EM fields into the primary/background fields and secondary/scattered fields. This decomposition helps eliminate the singularity of the EM fields near the source. The primary field can be calculated by using the frequency-domain (FD) and full-space Green's function of the source in a one-dimensional (1D) layered medium. The secondary field can be obtained by solving an FD vector Helmholtz equation for the scattered electric field, which is the reusable design part of the forward modeling for CSEMs. The procedure is the same regardless of the source type used to generate this field and regardless of operating in land, marine, airborne, or borehole environments.

On the other hand, any source type can be viewed as a combination of electric dipoles (EDs) or magnetic dipoles (MDs), each of which can be further decomposed into two horizontal EDs or MDs along the x and y directions, and one vertical ED or MD along the z direction. Thus, the focus of 3D forward modeling of CSEMs is reduced to solving EM fields for the background medium for ED or MD sources. By employing this approach, we can analyze the EM characteristics of different CSEMs using the same parameters for the forward solution, thus enhancing the comparability and understanding of the results.

The forward solver defines the overall structures and the main responsibilities of each module, thereby reducing the difficulty of solving EM fields of CSEMs and simplifying the implementation process of the theory. By avoiding the need to repeatedly construct the underlying mathematical logic, this approach enables geophysicists to focus on their unique application innovation. This paper first gives a brief overview of the mathematical methodology for 3D FD modeling of CSEMs. Once the FD EM fields have been calculated, the time-domain (TD) response can be obtained using the cosine/sine transformation. Then, the accuracy of the solver is verified by comparing the 3D modeling results with reference results obtained from 1D and 3D numerical solutions. Finally, we present two 3D modeling examples and discuss the effects of source shape and type on the EM fields of CSEMs over a 3D conductive Earth model.

2 Methodology

2.1 Maxwell's equations

Assuming harmonic time dependence $e^{-i\omega t}$, Maxwell's equations for the EM fields in the FD can be written as (Streich, 2009)

$$\nabla \times \mathbf{E} = -\mathrm{i}\omega\mu_{0}\mathbf{H},\tag{1}$$

$$\nabla \times \mathbf{H} = \mathbf{J} + \sigma^* \mathbf{E},\tag{2}$$

where $i^2 = -1$, $\omega = 2\pi f$ is the angular frequency, *f* is the frequency, **E** and **H** denote the total electric and magnetic fields, μ_0 within the Earth is assumed to be constant and set to that of free space, **J** is the source current distribution, $\sigma^* = \sigma + i\omega\varepsilon$ is the ohmic conduction term, σ is the electric conductivity, ε represents the air dielectric constant, and σ^* E describes the induced currents inside the Earth. According to Eqs 1, 2, the expression of the FD vector Helmholtz equation for the scattered electric field is as follows (Alumbaugh et al., 1996):



The electric or magnetic dipole is used as the unit of any CSEM sources (A). Finite-length wire source of LOTEM (B); A and B are the endpoints of the source, the solid line represents the real layout of the source, and the dotted line represents the source that is discretized by electric dipoles. Vector decomposition of an electric dipole (C).

$$\nabla \times \nabla \times \mathbf{E}^{\mathrm{s}} - k^{2} \mathbf{E}^{\mathrm{s}} = \left(k^{2} - k_{p}^{2}\right) \mathbf{E}^{\mathrm{p}},\tag{3}$$

where $k^2 = -i\omega\mu_0 (\sigma + i\omega\varepsilon)$, the superscripts "p" and "s" represent the primary and scattered fields, respectively, and E^s and E^p denote the scattered and primary electrical fields, respectively. Eq. 3 forms the basis of the presented modeling scheme and is well suited for simulations in CSEMs. Eq. 3 can overcome the difficulty of source description and the influence of source singularity on the numerical stability. Regardless of the type of source used to create the field, the procedure for calculating the scattered field remains the same. Then, we obtain the following PDE for the component \hat{i} :

$$\frac{\partial^2 E_y^s}{\partial x \partial y} - \frac{\partial^2 E_x^s}{\partial y^2} - \frac{\partial^2 E_x^s}{\partial z^2} + \frac{\partial^2 E_z^s}{\partial x \partial z} - k^2 E_x = \left(k^2 - k_p^2\right) E_x^p.$$
(4)

2.2 The multitype source

In this paper, we adopt the ED or MD as the basic composition unit for any complex geometry source (Figure 1A). For example, the signal emission source of the LOTEM can be decomposed into many EDs (Figure 1B), and vector decomposition is applied to each ED (Figure 1C). The EM fields at any given position can be obtained by superimposing the EM fields generated by the ED component along the *x*, *y*, and *z* directions. Therefore, the primary objective of 3D forward modeling in CSEMs, regardless of the specific source type used, is to solve the electric field of an electric dipole or a magnetic dipole within the background models.



FIGURE 2

Homogeneous stratified medium with an EM transmitting source embedded in an intermediate layer. *N* is the number of layers, ρ_0 , ρ_1 , ρ_2 , ..., ρ_N are individual layer resistivities, h_0 , h_1 , h_2 , ..., h_N are layer thicknesses, and z_1 , z_2 , ..., z_N denote the depth of the layer. The solid black circle denotes that the source is located at z_{ls} in the *ls*th layer, and the dashed line denotes that a virtual interface is added at the source location parallel to the layer boundaries.

2.3 Background models for primary fields

The primary field, as described by Weng et al. (2016), employs a virtual interface technique (Das and De Hoop, 1995) to solve the whole-space EM fields in a 1D layer model for different types of sources (Figure 2). These sources include vertical and horizontal electric dipoles (VED and HED) as well as vertical and horizontal magnetic dipoles (VMD and HMD).

By separating the partial wave solutions of the Helmholtz equations into upward and downward waves within certain boundaries, the potentials for Green's function are obtained. Starting from the source level, the amplitudes of the potentials in each layer are derived recursively based on the initial amplitudes. For different types of sources, only the initial terms that are associated with the transmitting sources need to be modified, and the kernel connected to the layered media remains the same. Hence, the aforementioned scheme can be easily applied to EM transmitting sources with slight modifications.

2.4 FD discretization

To calculate the scattered electric field in the medium, the geoelectric model is discretized by the cuboid cell (including the air layer) (Figure 3A), each denoted by a subscript (i, j, and k). The resistivity of each cell is represented by the symbol ρ (i, j, and k). i, j, and k represent the mesh indices in the *x*, *y*, and *z* directions, ranging from 1 to Nx, Ny, and Nz, respectively. Nx, Ny, and Nz denote the numbers of cells in the *x*, *y*, and *z* directions, respectively. Therefore, Eq. 4 can be discretized using the 3D staggered-grid finite-difference method, and the discretized expression is given in Eq. 5. The electric field components are defined on cell edges, while the magnetic field components naturally correspond to the cell faces (Figure 3B). The



discrete finite difference equations of the \hat{j} and \hat{k} components are given in Appendix A.

$$\begin{cases} \left[\frac{1}{\Delta x_{i}} \left(E_{y_{i+1,j+1/2,k}}^{s} - E_{y_{i,j+1/2,k}}^{s} \right) - \frac{1}{\Delta y_{j}} \left(E_{x_{i+1/2,j+1,k}}^{s} - E_{x_{i+1/2,j,k}}^{s} \right) \right] \\ - \left[\frac{1}{\Delta x_{i}} \left(E_{y_{i+1,j-1/2,k}}^{s} - E_{y_{i,j-1/2,k}}^{s} \right) - \frac{1}{\Delta y_{j-1}} \left(E_{x_{i+1/2,j,k}}^{s} - E_{x_{i+1/2,j-1,k}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \overline{y}_{j}} \\ + \left\{ \left[\frac{1}{\Delta x_{i}} \left(E_{z_{i+1,j,k+1/2}}^{s} - E_{z_{i,j,k+1/2}}^{s} \right) - \frac{1}{\Delta z_{k}} \left(E_{x_{i+1/2,j,k}}^{s} - E_{x_{i+1/2,j,k}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \overline{z}_{k}} \\ - \left[\frac{1}{\Delta x_{i}} \left(E_{z_{i+1,j,k-1/2}}^{s} - E_{z_{i,j,k-1/2}}^{s} \right) - \frac{1}{\Delta z_{k-1}} \left(E_{x_{i+1/2,j,k}}^{s} - E_{x_{i+1/2,j,k-1}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \overline{z}_{k}} \\ + i\omega \mu_{p} \hat{y}_{p} E_{x_{i+1/2,j,k}}^{s} \\ = -i\omega \mu_{p} \left(\hat{y}_{i+1/2,j,k} - \hat{y}_{p} \right) E_{x_{i+1/2,j,k}}^{p}, \tag{5}$$

where $\hat{y} = \sigma + i\omega\varepsilon$, w = x, y, z, and l = i, j, k denote the length of the *l*th cell in the *w* direction, and $\Delta \overline{w_l}$ is the distance between the center of the *l*th and *l*1th cells in the *w* direction.

2.5 Preconditioning and numerical implementation

Equation 5 can be assembled into the following system:

$$\mathbf{K}\mathbf{E}^{\mathrm{s}} = \mathbf{b},\tag{6}$$

where \mathbf{K} is the coefficient matrix, and \mathbf{b} in the right-hand side consists of the dot products between the primary/background field and conductivity abnormalities, as well as the appropriate boundary conditions.

Generally, the coefficient matrix **K** for Eq. 6 is a large, sparse, and ill-conditioned matrix that is difficult to solve. To reduce the condition number of the coefficient matrix **K**, Eq. 6 can be written in the preprocessing form as follows:

$$(\mathbf{K}\mathbf{M}^{-1}) \cdot \mathbf{y} = \mathbf{b}. \tag{7}$$

Here, $\mathbf{y} = \mathbf{M}\mathbf{E}^{s}$ is the modified unknown vector, where the matrix \mathbf{M} is called the (right-hand side) preprocessor, and the matrix $\mathbf{K}\mathbf{M}^{-1}$ is considered very close to the identity matrix. Once Eq. 7 is solved approximately, we can obtain \mathbf{E}^{s} using the relationship between \mathbf{y} and \mathbf{E}^{s} . In this study, an incomplete Cholesky decomposition is

used as the precondition to accelerate the convergence and improve the accuracy of the iterative solution. Eq. 7 can be solved by using the quasi-minimal residual method (QMR) (Mackie et al., 1994).

2.6 Boundary conditions and grid generation

Due to the significant difference between the solution region and the anomalous bodies, the secondary field will decay to zero at the boundary far from the anomalous body. As a result, we applied typical homogeneous Dirichlet's boundary conditions. With **n** being the normal vector on the domain boundary $\delta\Omega$, it is defined as:

$$\mathbf{n} \times \mathbf{E}^{\mathbf{s}} = \mathbf{0}. \tag{8}$$

The grid generation is centered on the area of interest and is divided into finer grids near the anomaly bodies. As depicted in Figure 3A, the grid size gradually increases as the distance from the area of interest expands. The air layer is set to seven layers, with the uppermost layer having a thickness of 30 km. In our paper, the mini software CSEM mesh is used to generate the mesh, which was developed by teachers and students in our research group. It greatly reduces the cost of generating meshes.

2.7 Divergence correction

The divergence of the secondary electric field \mathbf{E}^s is zero at any point except the source (Shen, 2003; Chen et al., 2011). Due to the accuracy of the numerical calculation, the divergence of \mathbf{E}^s does not disappear during an iteration solution process which can be calculated using $\varphi = \nabla \cdot \mathbf{E}^s$, considering the additional electrical field generated by the source φ . To calculate the additional electrical field, it is necessary to use the Neumann boundary conditions to solve the potential obtained by Poisson's equation as follows:

$$\nabla \cdot \nabla \psi = \varphi. \tag{9}$$

When Eq. 6 is solved, the E^s values required for divergence correction can be expressed as follows:

$$\mathbf{E}_{\text{correct}}^{\text{s}} = \mathbf{E}_{\text{old}}^{\text{s}} - \nabla \psi, \qquad (10)$$

specifically, to satisfy the following equation:

$$\nabla \cdot \mathbf{E}_{\text{correct}}^{\text{s}} = \nabla \cdot \mathbf{E}_{\text{old}}^{\text{s}} - \nabla \cdot \nabla \psi = \varphi - \varphi = 0.$$
(11)

The preconditioned conjugate gradient (PCG) algorithm is used to solve the divergence in Eq. 11. The iterative solution's convergence rate is greatly improved, particularly at low frequencies (Smith, 1996a; Smith, 1996b).

2.8 FD EM field interpolation to receiver positions

Once E^s on the center of the grid edge is obtained by Eq. 6, the total electric field E can be estimated by adding a primary field E^p , that is,

$$\mathbf{E} = \mathbf{E}^{\mathrm{p}} + \mathbf{E}^{\mathrm{s}},\tag{12}$$

while the total magnetic field normal to the surface confined by the grid edge at the center can be approximated by Faraday's law:

$$-i\omega \mathbf{B} = \nabla \times \mathbf{E},\tag{13}$$

from the estimated EM fields. Then, the EM fields at the position of interest can be interpolated by

$$\mathbf{E} = \mathbf{L}_e \mathbf{E} \tag{14}$$

and

$$\mathbf{H} = \mathbf{L}_{\mathrm{h}} \mathbf{H},\tag{15}$$

where L_e and L_h represent bilinear splines from the 3D grid nodes and edges, respectively, to the data sites.

2.9 TD EM responses

According to Weng et al. (2017), the TD EM signal h(t) through

$$h(t) = \frac{1}{2\pi} \int_0^\infty I(\omega) H(\omega) e^{-i\omega t} d\omega$$
(16)

is related to the FD response $H(\omega)$ from a source excited by current $I(\omega)$ over a conductive model; in the aforementioned equation, $i^2 = -1$, $\omega = 2\pi f$ is the angular frequency, f is the frequency, and t is the time. If a step-off current is assumed, using the Euler formula, Eq. 16 can be further split into a sine transformation as follows:

$$dB/dt = \mu_0 dh/dt = -\mu_0 \frac{2}{\pi} \int_0^\infty \operatorname{Im}\left[H(\omega)\right] \sin(\omega t) d\omega, \quad (17)$$

where B is the magnetic flux density and Im [·] denotes the imaginary part of the FD EM fields. We sample the exact six points per decade FD data between 10^{-3} Hz and 10^8 Hz and obtain 67 frequencies (Liu et al., 2016). After obtaining the FD results, the TD EM responses for the step wave are calculated using Eq. 17. Based on the aforementioned theory, we have programmed a 3D forward modeling code using Fortran 90 for FD/TD CSEMs with multitype sources.

3 Model verification

To test the correctness and reliability of the aforementioned forward modeling solver for different types of sources, we initially design a 100 Ωm uniform half-space model for the airborne EM method (AEM) with a vertical magnetic dipole (VMD) source and a unit current (Liu and Yin, 2013). The flight altitude is set to 30 m, and the receiver is positioned 2 m away from the transmitter (Figure 4A). The model is divided into $20 \times 20 \times 25$ prisms of dimension 10 m \times 10 m \times 10 m. The background model with 50 Ωm is used for computing the primary field. Figure 4B illustrates the comparisons between the 3D solution and the 1D results using a step-off current waveform. The results of the 3D solution agree well with the 1D result. The overall relative errors are less than 5%. Figure 4C shows a plot of the error versus iteration for the QMR solver for the uniform half-space model for the data with the four frequencies. We can clearly see that the QMR solver is stable and converges quickly to the values of 10⁻⁷. We obtained similar results for the other models presented in this paper. The total memory required to solve this model was 73.24 MB. It took approximately 4.3 min per frequency to solve this model on a personal computer with an Intel[®] Core[™] i5-2320 processor and 8 GB memory.

Second, we take the RFLTEM as an example, and a loop source with dimensions of 10 m \times 10 m can be decomposed into many EDs. A 3D conductive body (50 m \times 50 m \times 50 m) of 20 Ωm is embedded in a twolayer Earth model in Ji et al. (2017) (Figure 5A). In the numerical simulation, the origin of the coordinate system is at the surface with the z-axis downward through the center of the abnormal body. The model is divided into $51 \times 51 \times 30$ prisms of dimension $10 \text{ m} \times 10 \text{ m} \times 10$ m. The FD EM fields are calculated at the measurement points Rx1 and Rx2 in Figure 5A. Then, dB_z/dt is obtained using Eq. (17), and it was normalized by the square of the single-turn receiving coil. The solution of the forward modeling solver agrees well with that in the study by Ji et al. (2017) (Figure 5B). These examples serve as strong evidence for the accuracy and reliability of our forward modeling code. Figure 5C shows a plot of the error versus iteration for the QMR solver for a two-layer model for the data with four frequencies. The error of the QMR solver converges quickly to the values of 10⁻⁷. The total memory required to solve this model was 114.26 MB. It took approximately 6.8 min per frequency.

4 Applications

4.1 The real and theoretical source

In real field surveys, surface obstacles and topography can preclude laying out the source in a theoretical shape such as a straight line, rectangle, or circle. The aforementioned scheme can segment a source with an arbitrarily complex shape into a large number of EDs and MDs. To illustrate this, we conducted a study taking the LOTEM and CSAMT as examples, focusing on the differences between the TD and FD EM fields for both straight and non-straight wire sources. We computed the TD and FD responses for the bent and straight wires. Both wires centered at (0, 0) have the same grounding points; the straight wires of 1 km have a 1 A current on the surface, and the bent wire is the real source consisting of 22 segments (Figure 6A). The 3D model of a 100 Ω m homogeneous half-space is shown in Figure 6B.



FIGURE 4

A uniform half-space is used to verify the numerical accuracy of 3D modeling for AEM with a VMD source. (A) 3D model. (B) Comparison of our 3D result against the 1D numerical solution for the TD EM fields (dB/dt). (C) Error curve of the QMR solver.



FIGURE 5

A two-layer model is used to verify the numerical accuracy of 3D modeling for the RFLTEM. (A) 3D model. (B) Comparison between the field solutions obtained in this paper and those of Ji et al. (2017) for the receiver positions denoted by Rx_1 (105, 0, 0) and Rx_2 (-155, 0, 0). (C) Error curve of the QMR solver.



In Figure 7, we display the TD magnetic field dB/dt for a bent wire compared to dB/dt for a straight wire at different times: 0.065 ms, 0.9 ms, 4 ms, and 10 ms. As time increases, the difference in dB/dt between the bent and straight wires decreases. dB/dt

generated by the straight wire exhibits a symmetrical distribution around the source center and is concentrated near the surface (Figure 7A). At the early stage, there is a significant disparity in the dB/dt between the bent and straight wires (Figure 7C). Therefore,



when analyzing early-time TEM data, it is crucial to consider the shape of the source. With a time delay, the center of dB/dt propagates downward and outward, and the strength of the field gradually decays. At the later stage, dB/dt continues to propagate and becomes more uniform, and the relative differences in dB/dt between the bent and straight wires are minimal (Figure 7C). Spatially, the closer the distance to the source, the greater the difference in dB/dt between the bent and straight wires. Conversely, the farther the distance to the source, the smaller the difference in dB/dt between the bent and straight wires. Therefore, when processing data of TEMs observed close to the source, such as the SOTEM, the source's shape must be taken into account.

Figure 8 shows the FD electric field component Ey for the bent wire relative to Ey for a straight wire at frequencies of 0.1 Hz, 2 Hz, 16 Hz, and 128 Hz. The relative differences in Ey between the bent and straight wires increase with increasing frequency and decrease with increasing distance from the source. At low frequencies, the difference in Ey between the bent and straight wires was very small, and it was mainly concentrated near the source and both sides along the emitting source direction. At high frequencies, the difference in Ey was most pronounced in the survey area. These results highlight the significant influence of the grounding points' locations and the entire wire layout on the EM fields at these frequencies. Therefore, during the inversion of the high-frequency data from a wire source, such as the CSAMT, it is impossible to ignore the impact of the source's shape on the measured data.

The comparison between the TD and FD EM methods shows that the FD CSEM appears to be more significantly influenced by the entire wire layout than the TD CSEM. This difference can be attributed to the nature of the measurements. In the TD CSEMs, observations are made after the current is turned off, capturing the induced eddy field or secondary field. In this case, the shape of the emission source only affects the early data. On the other hand, in FD CSEMs, observations are made under the harmonic excitation of the source, representing the total field. The impact of the source's shape on the field becomes more noticeable, particularly at higher frequencies. At low frequencies, the electric field resembles a potential field that occurs in the static limit. Here, the only factors affecting the electric field are the locations of the grounding points.

According to the aforementioned analysis, we suggest that the shape of the source must be considered for TD and FD CSEMs with a wire source during data processing and inversion. However, in regard to the influence of the source shape, we recommend TD CSEMs when processing data based on the theory of an ideal straight wire or electric dipole source for the same target.

4.2 Wire and loop sources

4.2.1 Model parameters

3D FD forward modeling and TD forward modeling of CSEMs with wire and loop sources were conducted using a conductive



geological model in which a conductive tilted sheet is embedded in a uniform half-space with a strike of 525 m and two high- and low-resistivity bodies on the surface (Figure 9A). In the numerical simulation, the origin of the coordinate system was set at the center of the model at the ground, with the *z*-axis pointing downward. The model space is subdivided into $40 \times 40 \times 20$ prisms of size 25 m \times 25 m \times 25 m. To ensure that Dirichlet's boundary condition was satisfied, boundary cells in the five layers outside the model were expanded by a factor of 1.5 (Figures 9B–D). In the following solution calculation, all the parameters remained consistent for the 3D forward modeling of CSEMs with both loop and wire sources.

4.2.2 TD CSEM with a loop source

In the TD CSEM, a larger rectangular fixed loop is commonly used as a source, encompassing the target area. The vertical TD magnetic field dB/dt is then measured both inside and outside the loop, simulating the RLFTEM. In this paper, the RLFTEM uses a fixed loop of dimension 300 m × 300 m with a 1 A current on the Earth's surface (Figure 9A). Figure 10 shows the TD scattered field for two specific sections (Y = -270 m and Y = 270 m). According to Faraday's law, the induced eddy currents within the anomalous body are excited when the transmitter switch is abruptly turned off, thus preventing the internal magnetic field of the anomalous body from weakening. At early times, the presence of surface resistive and conductive bodies distorts the induced eddy currents. With a time delay, the induced eddy currents diffuse downward and outward underground, leading to a gradual decay in the strength of the magnetic field. As observed from dB/dt, the RLFTEM has a higher sensitivity to conductive bodies than resistive anomalies.

4.2.3 FD CSEM with a wire source

The wire source is generally used in land and marine CSEMs, such as the CSAMT, LOTEM, and GATEM, and the transmitter is a finitelength wire with complex geometry. In the numerical implementation, we take the CSAMT as an example. A finite-length wire with a length of 1 km and a current of 1 A was laid on the ground along the x direction as a signal source. The nearest measuring point is located at a distance of 10 km from the center of the source (Figure 9A). For the CSAMT, the electric field component Ex is generally used as the observed data, as in the wide-field EM method (He, 2010). The proposed forward modeling solver computes Ex in the whole space. The results of the calculation for two sections (Y = -270 m and Y = 270 m) are shown in Figure 11. The FD EM fields from the wire source are similar to those of a plane wave in the measurement area. Among these frequencies, f = 16 Hz is the most sensitive frequency for the tilted target body. The low-frequency data demonstrate a stronger resolution ability to the deep abnormal body. Moreover, high-frequency EM fields have higher energies and stronger abilities to resolve shallow abnormal bodies. However, due to their shorter wavelength and lower penetration, their overall ability to resolve deep anomalous bodies is relatively limited.

The RFLTEM and CSAMT utilize loop and wire sources, representing the magnetic and electric sources, respectively. By comparing the forward results of the RFLTEM and CSAMT, we



FIGURE 9

3D conductive geological model and the grid used for the model. (A) View of the model used for 3D FD and TD forward modeling of CSEMs. The solid circles represent the position of the receivers. The yellow rectangle and the black line indicate the locations of the loop and wire sources, respectively; (B) plan view and rectangular mesh with Z = 120 m; (C) and (D) vertical sections and rectangular mesh with Y = -270 m and Y = 270 m, respectively.



TD scattered field at t = 0.013 ms, 0.87 ms, 3.6 ms, and 11 ms: (A) Y = -270 m; (B) Y = 270 m. The yellow solid lines denote the position of the abnormal bodies. The black solid lines represent the contour line.



can see that both methods are sensitive to shallow conductive anomalies. However, in terms of resolving shallow resistive anomalies, the CSAMT demonstrates a better resolution than the RFLTEM. The observed data distortion, primarily caused by deep conductivity anomalies, is more pronounced in the CSAMT than in the RLFTEM. Furthermore, the resolution of the CSAMT is stronger than that of the RLFTEM for spatial information, such as the occurrence of deep tilted sheets.

Based on the aforementioned analysis, TD CSEMs with loop sources (central loop, coincident loop, and dipole–dipole) are recommended for detecting shallow conductivity anomalies due to their convenience and higher efficiency. On the other hand, for detecting deep anomalies, FD CSEMs with a wire source are sensitive to both resistive and conductive abnormal bodies. The 3D forward modeling solver of CSEMs provides valuable assistance to geophysicists in selecting the most suitable exploration method based on the same forward modeling parameters.

5 Conclusion

In this paper, we have considered the significant commonalities of 3D forward modeling for TD and FD CSEMs with different types of sources and operating environments and developed a general forward solver. Using the same format, the source term of CSEMs is imposed on the right-hand term of an FD vector Helmholtz equation for the scattered electric field. Any complex geometry source can be decomposed into EDs or MDs used as the basic composition unit, each of which can be further decomposed into two horizontal EDs or MDs along the *x* and *y* directions and one vertical ED or MD along the *z* direction. Through this solver, geophysicists can compare the EM field

characteristics of different CSEMs for the specific geological targets using the same parameters of the forward solution.

Based on the numerical experimental results of the real and theoretical sources, we suggest that the shape of the source must be considered for TD and FD CSEMs. According to the numerical experiment results of wire and loop sources, we recommend TD CSEMs with the loop source for detecting shallow conductivity abnormal bodies. Moreover, we recommend FD CSEMs with a wire source for detecting deep anomalies.

The solver proposed in this paper leads us to clearly define the basic target objects and methods needed to solve the 3D TD and FD forward problem of the general CSEMs. In the future, we need to further optimize the code computational efficiency, improve the speed of program running, reduce the memory requirements, and lay the foundation for the geophysical interpretation of TD and FD EM data.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary Material; further inquiries can be directed to the corresponding author.

Author contributions

DL implemented the algorithms and wrote the manuscript. YL and ZW conceived the idea of the algorithm, supervised the study, performed the analyses, and edited the manuscript. LJ helped in processing the data and prepared the figures. All authors contributed to the article and approved the submitted version.

Funding

This research was supported by the Natural Science Foundation of Jilin Province, China (Grant Number YDZJ202301ZYTS222) and Jilin Province Education Department Science and Technology Research Project (Grant Number JJKH20230347KJ).

Acknowledgments

Thanks are given to the two reviewers and editors for their helpful suggestions and for improving the manuscript. We would like to thank Prof. Aihua Weng, who modified the whole paper.

References

Alumbaugh, D. L., Newman, G. A., Prevost, L., and Shadid, J. N. (1996). Threedimensional wideband electromagnetic modeling on massively parallel computers. *Radio Sci.* 31 (1), 1–23. doi:10.1029/95RS02815

Ansari, S., and Farquharson, C. G. (2013). 3D finite-element forward modeling of electromagnetic data using vector and scalar potentials and unstructured grids. *Geophysics* 79 (4), E149–E165. doi:10.1190/GEO2013-0172.1

Avdeev, D. B. (2005). Three-dimensional electromagnetic modelling and inversion from theory to application. *Surv. Geophys.* 26 (6), 767–799. doi:10.1007/s10712-005-1836-x

Börner, R. U., Ernst, O. G., and Güttel, S. (2015). Three-dimensional transient electromagnetic modelling using Rational Krylov methods. *Geophys. J. Int.* 202 (3), 2025–2043. doi:10.1093/gji/ggv224

Börner, R. U. (2010). Numerical modelling in geo-electromagnetics: Advances and challenges. *Surv. Geophys.* 31 (2), 225–245. doi:10.1007/s10712-009-9087-x

Cai, H., Xiong, B., Han, M., and Zhdanov, M. (2014). 3D controlled-source electromagnetic modeling in anisotropic medium using edge-based finite element method. *Comput. Geosciences* 73, 164–176. doi:10.1016/j.cageo.2014.09.008

Castillo-Reyes, O., Amor-Martin, A., Botella, A., Anquez, P., and García-Castillo, L. E. (2022). Tailored meshing for parallel 3D electromagnetic modeling using high-order edge elements. *J. Comput. Sci.* 63, 101813. doi:10.1016/j.jocs.2022.101813

Castillo-Reyes, O., de la Puente, J., and Cela, J. M. (2022). HPC geophysical electromagnetics: A synthetic vti model with complex bathymetry. *Energies* 15 (4), 1272. doi:10.3390/en15041272

Castillo-Reyes, O., de la Puente, J., and Cela, J. M. (2018). Petgem: A parallel code for 3D CSEM forward modeling using edge finite elements. *Comput. Geosciences* 119, 123–136. doi:10.1016/j.cageo.2018.07.005

Castillo-Reyes, O., de la Puente, J., García-Castillo, L. E., and Cela, J. M. (2019). Parallel 3-D marine controlled-source electromagnetic modelling using high-order tetrahedral Nédélec elements. *Geophys. J. Int.* 219 (219), 39–65. doi:10.1093/gji/ggz285

Castillo-Reyes, O., Modesto, D., Queralt, P., Marcuello, A., Ledo, J., Amor-Martin, A., et al. (2022). 3D magnetotelluric modeling using high-order tetrahedral Nédélec elements on massively parallel computing platforms. *Comput. Geosciences* 160, 105030. doi:10.1016/j.cageo.2021.105030

Castillo-Reyes, O., Queralt, P., Marcuello, A., and Ledo, J. (2022). Land CSEM simulations and experimental test using metallic casing in a geothermal exploration context: Vallès basin (NE Spain) case study. *IEEE Trans. Geoscience Remote Sens.* 60, 1–13. doi:10.1109/TGRS.2021.3069042

Chave, A. D., Everett, M. E., Mattsson, J., Boon, J., and Midgley, J. (2017). On the physics of frequency-domain controlled source electromagnetics in shallow water. 1: Isotropic conductivity. *Geophys. J. Int.* 208 (2), 1026–1042. doi:10.1093/gji/ggw435

Chen, H., Deng, J. Z., Tan, H. D., and Yang, H. Y. (2011). Study on divergence correction method in three-dimensional magnetotelluric modeling with staggered-grid finite difference method. *Chin. J. Geophys. (in Chinese)* 54 (6), 1649–1659. doi:10.3969/j. issn.0001-5733.2011.06.025

Commer, M., and Newman, G. (2004). A parallel finite-difference approach for 3D transient electromagnetic modeling with galvanic sources. *Geophysics* 69 (5), 1192–1202. doi:10.1190/1.1801936

Connell, D., and Key, K. (2013). A numerical comparison of time and frequencydomain marine electromagnetic methods for hydrocarbon exploration in shallow water. *Geophysical Prospecting* 61 (1), 187–199. doi:10.1111/j.1365-2478.2012.01037.x

Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Publisher's note

All claims expressed in this article are solely those of the authors and do not necessarily represent those of their affiliated organizations, or those of the publisher, the editors, and the reviewers. Any product that may be evaluated in this article, or claim that may be made by its manufacturer, is not guaranteed or endorsed by the publisher.

Constable, S., and Srnka, L. J. (2007). An introduction to marine controlled-source electromagnetic methods for hydrocarbon exploration. *Geophysics* 72 (2), WA3–WA12. doi:10.1190/1.2432483

Constable, S. (2010). Ten years of marine CSEM for hydrocarbon exploration. *Geophysics* 75 (5), 75A67–75A81. doi:10.1190/1.3483451

Das, U. C., and De Hoop, A. T. (1995). Efficient computation of apparent resistivity curves for depth profiling of a layered Earth. *Geophysics* 60 (6), 1691–1697. doi:10.1190/1.1443901

Di, Q., Xue, G., Yin, C., and Li, X. (2020). New methods of controlled-source electromagnetic detection in China. *Science China Earth Sciences* 63 (9), 1268–1277. doi:10.1007/s11430-019-9583-9

Edwards, N. (2005). Marine controlled source electromagnetics: Principles, methodologies, future commercial applications. *Surveys in Geophysics* 26 (6), 675–700. doi:10.1007/s10712-005-1830-3

Everett, M. E. (2009). Transient electromagnetic response of a loop source over a rough geological medium. *Geophysical Journal International* 177 (2), 421–429. doi:10. 1111/j.1365-246X.2008.04011.x

Goldstein, M. A., and Strangway, D. W. (1975). Audio-frequency magnetotellurics with a grounded electric dipole source. *Geophysics* 40 (4), 669–683. doi:10.1190/1. 1440558

Haber, E., and Ascher, U. M. (2001). Fast finite volume simulation of 3D electromagnetic problems with highly discontinuous coefficients. *SIAM Journal of Scientific Computing* 22 (6), 1943–1961. doi:10.1137/S1064827599360741

He, J. S., Yang, Y., Di-quan, L., and Jing-bo, W. (2010). Wide field electromagnetic sounding methods. *Journal of Central South University (Science and Technology)* 41 (3), 1065–1072. doi:10.4133/sageep.28-047

Hu, X., Peng, R., Wu, G., Wang, W., Huo, G., and Han, B. (2013). Mineral exploration using CSAMT data: Application to Longmen region metallogenic belt, Guangdong Province, China. *Geophysics* 78 (3), B111–B119. doi:10.1190/geo2012-0115.1

Huang, X., Yin, C., Farquharson, C. G., Cao, X., Zhang, B., Huang, W., et al. (2019). Spectral-element method with arbitrary hexahedron meshes for time-domain 3D airborne electromagnetic forward modeling. *Geophysics* 84 (1), E37–E46. doi:10. 1190/geo2018-0231.1

Hursán, G., and Zhdanov, M. S. (2002). Contraction integral equation method in three-dimensional electromagnetic modeling. *Radio Science*, 37(6), 1–13. doi:10.1029/2001rs002513

Ji, Y., Hu, Y., and Imamura, N. (2017). Three-dimensional transient electromagnetic modeling based on fictitious wave domain methods. *Pure and Applied Geophysics* 174 (5), 2077–2088. doi:10.1007/s00024-017-1528-8

Koldan, J., Puzyrev, V., De la Puente, J., Houzeaux, G., and Cela, J. M. (2014). Algebraic multigrid preconditioning within parallel finite-element solvers for 3-D electromagnetic modelling problems in geophysics. *Geophysical Journal International* 197 (3), 1442–1458. doi:10.1093/gji/ggu086

Li, D., Shan, X., and Weng, A. (2022). Applying three-dimensional inversion to the frequency-domain response converted from transient electromagnetic data for a rectangular fixed loop. *Journal of Applied Geophysics* 196, 104489. doi:10.1016/j. jappgeo.2021.104489

Li, J., Lu, X., Farquharson, C. G., and Hu, X. (2018). A finite-element time-domain forward solver for electromagnetic methods with complex-shaped loop sources. *Geophysics* 83 (3), E117–E132. doi:10.1190/GEO2017-0216.1

Liu, Y. H., Yin, C. C., Ren, X. Y., and Qiu, C. K. (2016). 3D parallel inversion of time-domain airborne EM data. *Applied Geophysics* 13 (4), 701–711. doi:10.1007/s11770-016-0581-x

Liu, Z., Ren, Z., Yao, H., Tang, J., Lu, X., and Farquharson, C. (2023). A parallel adaptive finite-element approach for 3-D realistic controlled-source electromagnetic problems using hierarchical tetrahedral grids. *Geophysical Journal International* 232 (3), 1866–1885. doi:10.1093/gji/ggac419

Mackie, R. L., Smith, J. T., and Madden, T. R. (1994). Three-dimensional electromagnetic modeling using finite difference equations: The magnetotelluric example. *Radio Science* 29 (4), 923–935. doi:10.1029/94RS00326

Malovichko, M., Tarasov, A. V., Yavich, N., and Zhdanov, M. S. (2019). Mineral exploration with 3-D controlled-source electromagnetic method: A synthetic study of sukhoi log gold deposit. *Geophysical Journal International* 219 (3), 1698–1716. doi:10. 1093/gji/ggz390

Newman, G. A., and Alumbaugh, D. L. (1995). Frequency-domain modelling of airborne electromagnetic responses using staggered finite differences. *Geophysical Prospecting* 43 (8), 1021–1042. doi:10.1111/j.1365-2478.1995.tb00294.x

Oldenburg, D. W., Heagy, L. J., Kang, S., and Cockett, R. (2020). 3D electromagnetic modelling and inversion: A case for open source. *Exploration Geophysics* 51 (1), 25–37. doi:10.1080/08123985.2019.1580118

Peng, R., Hu, X., Li, J., Hu, Z., and Yang, H. (2018). 3-D finite-volume forward modeling of wide-field EM using scattered potentials. *Chinese Journal of Geophysics (in Chinese)* 61 (10), 4160–4170. doi:10.6038/cjg2018L0363

Ren, X., Yin, C., Liu, Y., Cai, J., Wang, C., and Ben, F. (2017). Efficient modeling of time-domain AEM using finite-volume method. *Journal of Environmental and Engineering Geophysics* 22 (3), 267–278. doi:10.2113/JEEG22.3.267

Rochlitz, R., Skibbe, N., and Günther, T. (2019). custEM: Customizable finite-element simulation of complex controlled-source electromagnetic data. *Geophysics* 84 (2), F17–F33. doi:10.1190/geo2018-0208.1

Schaller, A., Streich, R., Drijkoningen, G., Ritter, O., and Slob, E. (2018). A land-based controlled-source electromagnetic method for oil field exploration: An example from the Schoonebeek oil field. *Geophysics* 83 (2), WB1–WB17. WB1–WB17. doi:10.1190/geo2017-0022.1

Shen, J. (2003). Modeling of 3-D electromagnetic responses in frequency domain by using the staggeredgrid finite difference method. *Acta Geophysica Sinica* 46 (2), 396–408. doi:10.1002/cjg2.355

Smith, J. T. (1996a). Conservative modeling of 3-D electromagnetic fields, Part I: Properties and error analysis. *Geophysics* 61, 1308–1318. doi:10.1190/1.1444054

Smith, J. T. (1996b). Conservative modeling of 3-D electromagnetic fields, Part II: Biconjugate gradient solution and an accelerator. *Geophysics* 61, 1319–1324. doi:10. 1190/1.1444055

Streich, R. (2009). 3D finite-difference frequency-domain modeling of controlledsource electromagnetic data: Direct solution and optimization for high accuracy. *Geophysics* 74 (5), F95–F105. doi:10.1190/1.3196241

Streich, R. (2016). Controlled-source electromagnetic approaches for hydrocarbon exploration and monitoring on land. *Surveys in Geophysics* 37 (1), 47–80. doi:10.1007/s10712-015-9336-0

Tang, J., Zhou, F., Ren, Z., Xiao, X., Qiu, L., Chen, C., et al. (2018). Three-dimensional forward modeling of the controlled-source electromagnetic problem based on the

integral equation method with an unstructured grid. Acta Geophysica Sinica 61 (4), 1549–1562. doi:10.6038/cjg2018L0121

Tonti, E. (2002). Finite formulation of electromagnetic field. *IEEE Transactions on Magnetics* 38 (2 I), 333–336. doi:10.1109/20.996090

Um, E. S., and Alumbaugh, D. L. (2007). On the physics of the marine controlledsource electromagnetic method. *Geophysics* 72 (2), WA13–WA26. doi:10.1190/1. 2432482

Um, E. S., Harris, J. M., and Alumbaugh, D. L. (2012). An iterative finite element timedomain method for simulating three-dimensional electromagnetic diffusion in Earth. *Geophysical Journal International* 190 (2), 871–886. doi:10.1111/j.1365-246X.2012. 05540.x

Unno, M., Aono, S., and Asai, H. (2012). GPU-based massively parallel 3-D HIE-FDTD method for high-speed electromagnetic field simulation. *IEEE Transactions on Electromagnetic Compatibility* 54 (4), 912–921. doi:10.1109/TEMC.2011.2173938

Weiland, T. (1996). Time domain electromagnetic field computation with finite difference methods. *International Journal of Numerical Modelling Electronic Networks, Devices and Fields* 9 (4), 295–319. doi:10.1002/(sici)1099-1204(199607)9:4<295::aid-jnm240>3.0.co;2-8

Weng, A. H., Liu, Y. H., Yin, C. C., and Jia, D. Y. (2016). Singularity-free Green's function for EM sources embedded in a stratified medium. *Applied Geophysics* 13 (1), 25–36. doi:10.1007/s11770-016-0549-x

Weng, A., Li, D., Yang, Y., Li, S., Li, J., and Li, S. (2017). Transforming a timedomain electromagnetic signal to a frequency-domain electromagnetic response using regularization inversion. *Geophysics* 82 (5), E287–E295. doi:10.1190/ GEO2016-0505.1

Weng, A., Liu, Y., Jia, D., Liao, X., and Yin, C. (2012). Three-dimensional controlled source electromagnetic inversion using non-linear conjugate gradients. *Chinese Journal of Geophysics (in Chinese)* 10, 3506. doi:10.6038/j.issn.0001-5733.2012.10.034

Werthmüller, D., Rochlitz, R., Castillo-Reyes, O., and Heagy, L. (2021). Towards an open-source landscape for 3-D CSEM modelling. *Geophysical Journal International* 227 (1), 644–659. doi:10.1093/gji/ggab238

Xu, J., and Tang, J. (2022). Spectral element method for 3D wide field electromagnetic method forward modeling. *Chinese Journal of Geophysics (in Chinese)* 65 (4), 1461–1471. doi:10.6038/cjg2022P0067

Xue, G. Q. (2018). The development of near-source electromagnetic methods in China. *Journal of Environmental and Engineering Geophysics* 23 (1), 115–124. doi:10. 2113/JEEG23.1.115

Yang, D., and Oldenburg, D. W. (2016). Survey decomposition: A scalable framework for 3D controlled-source electromagnetic inversion. *Geophysics* 81 (2), E69–E87. doi:10. 1190/GEO2015-0217.1

Yin, C., Qi, Y., and Liu, Y. (2016). 3D time-domain airborne EM modeling for an arbitrarily anisotropic Earth. *Journal of Applied Geophysics* 131, 163–178. doi:10.1016/j. jappgeo.2016.05.013

Zhang, B., Yin, C., Ren, X., Liu, Y., and Qi, Y. (2018). Adaptive finite element for 3D time-domain airborne electromagnetic modeling based on hybrid posterior error estimation. *Geophysics* 83 (2), WB71–WB79. doi:10.1190/geo2017-0544.1

Zonge, K. L., and Hughes, L. J. (1991). "Controlled source audio-frequency magnetotellurics," in *Electromagnetic methods in applied geophysics: Volume 2, application, parts A and B* Editor M. N. Nabighian (United States: Society of Exploration Geophysicists).

Appendix A: Finite difference equations

According to Eq. 5, the expressions of \hat{j} and \hat{k} components are, respectively, as follows:

$$\begin{split} &\left\{ \left[\frac{1}{\Delta y_{j}} \left(E_{z_{i,j+1,k+1/2}}^{s} - E_{z_{i,j,k+1/2}}^{s} \right) - \frac{1}{\Delta z_{k}} \left(E_{y_{i,j+1/2,k+1}}^{s} - E_{y_{i,j+1/2,k-1}}^{s} \right) \right] \right. \\ &\left. - \left[\frac{1}{\Delta y_{j}} \left(E_{z_{i,j+1,k-1/2}}^{s} - E_{z_{i,j,k-1/2}}^{s} \right) - \frac{1}{\Delta z_{k-1}} \left(E_{y_{i,j+1/2,k}}^{s} - E_{y_{i,j+1/2,k-1}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \bar{z}_{k}} \\ &\left. + \left\{ \left[\frac{1}{\Delta y_{j}} \left(E_{x_{i+1/2,j+1,k}}^{s} - E_{x_{i+1/2,j,k}}^{s} \right) - \frac{1}{\Delta x_{i}} \left(E_{y_{i,j+1/2,k}}^{s} - E_{y_{i,j+1/2,k}}^{s} \right) \right] \right\} \\ &\left. - \left[\frac{1}{\Delta y_{j}} \left(E_{x_{i-1/2,j+1,k}}^{s} - E_{x_{i-1/2,j,k}}^{s} \right) - \frac{1}{\Delta x_{i-1}} \left(E_{y_{i,j+1/2,k}}^{s} - E_{x_{i-1,j+1/2,k}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \bar{x}_{i}} \\ &\left. + i\omega \mu_{p} \hat{y}_{p} E_{y_{i,j+1/2,k}}^{s} \\ &= -i\omega \mu_{p} \left(\hat{y}_{i,j+1/2,k} - \hat{y}_{p} \right) E_{y_{i,j+1/2,k}}^{p} \end{split}$$
 (A - 1)

and

$$\begin{split} &\left\{ \left[\frac{1}{\Delta z_{k}} \left(E_{x_{i+1/2,jk+1}}^{s} - E_{z_{i+1/2,jk}}^{s} \right) - \frac{1}{\Delta x_{i}} \left(E_{z_{i+1,jk+1/2}}^{s} - E_{z_{i,jk+1/2}}^{s} \right) \right] \right. \\ &\left. - \left[\frac{1}{\Delta z_{k}} \left(E_{x_{i-1/2,jk+1}}^{s} - E_{x_{i-1/2,jk}}^{s} \right) - \frac{1}{\Delta x_{i-1}} \left(E_{z_{i,jk+1/2}}^{s} - E_{y_{i-1,jk+1/2}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \bar{x}_{i}} \\ &\left. + \left\{ \left[\frac{1}{\Delta z_{k}} \left(E_{y_{i,j+1/2,k+1}}^{s} - E_{y_{i,j+1/2,k}}^{s} \right) - \frac{1}{\Delta y_{j}} \left(E_{z_{i,jk+1/2}}^{s} - E_{z_{i,jk+1/2}}^{s} \right) \right] \right] \\ &\left. - \left[\frac{1}{\Delta z_{k}} \left(E_{y_{i,j-1/2,k+1}}^{s} - E_{y_{i,j-1/2,k}}^{s} \right) - \frac{1}{\Delta y_{j-1}} \left(E_{z_{i,jk+1/2}}^{s} - E_{z_{i,j-1,k+1/2}}^{s} \right) \right] \right\} \cdot \frac{1}{\Delta \bar{y}_{j}} \\ &\left. + i \omega \mu_{p} \hat{y}_{p} E_{z_{i,jk+1/2}}^{s} \\ &= -i \omega \mu_{p} \left(\hat{y}_{i,j,k+1/2} - \hat{y}_{p} \right) E_{z_{i,jk+1/2}}^{p} \end{split} \right.$$
 (A - 2)

Li et al.