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## Application of potential vorticity tendency diagnosis method to high-resolution simulation of tropical cyclones

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As the grid spacing in the numerical simulation decreases to ~1 km, the potential vorticity (PV) structure of the simulated tropical cyclone (TC) is an annular tower of high PV with low PV in the eye and the resulting reversal of the radial PV gradient in the inner core is subject to dynamic instability, leading to complicated small-scale features in the PV field. While the PV tendency (PVT) method has been successfully used to diagnose TC motion in numerical simulations with relatively coarse resolution (~10 km), it has been found that the PVT diagnosis method fails in the TC simulation with grid spacing of ~1 km. This study reveals that the failure of the PVT diagnosis method in the high-resolution simulation with grid spacing of ~1 km arises from the induced small-scale features in the PV field. The high localized PV features do not affect TC motion, but make it difficult to calculate the gradient of azimuthal mean PV when the time interval of the model output is ~1 h. It is suggested that the PVT method can be applied to high-resolution simulations by increasing the time interval of the model output and/or smoothing the model output to reduce the influence of small-scale PV features.

### KEYWORDS

tropical cyclone motion, nurmerical simulation, potential vorticity (PV), small-scale features, model resolution

### Introduction

The potential vorticity (PV) tendency (PVT) method proposed by Wu and Wang (2000) treats a tropical cyclone (TC) as a positive PV anomaly relative to its environment and TC motion is essentially the propagation of the PV anomaly. The PVT method has been successfully used to diagnose the influences of various physical processes on TC motion in numerical simulations, such as the effects of diabatic heating, land surface friction, river deltas, coastal lines, mountains, islands, cloud-radiative processes, and sea surface pressure gradients (e.g., Wu and Wang 2001a, Wu and Wang 2001b; Wong and Chan, 2006; Yu et al., 2007; Fovell et al., 2010; Hsu et al., 2013; Wang et al., 2013; Choi

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et al., 2013). Over the past two decades, numerical models for TC simulation have been greatly improved and the horizontal grid spacing has been decreased from ~10 to ~1 km, even to 100 m or less (Zhu 2008; Rotunno et al., 2009; Zhu et al., 2013; Green and Zhang 2015; Wu et al., 2018, Wu et al., 2019). However, it has been found that the PVT method cannot be applied to the high-resolution (1 km or less) simulations of TCs (Zhao et al., 2019).

Over the past two decades, the PVT method was used to diagnose the influences of various physical processes on TC motion (e.g., Wu and Wang 2001a, Wu and Wang 2001b; Wong and Chan, 2006; Yu et al., 2007; Fovell et al., 2010; Hsu et al., 2013; Wang et al., 2013; Choi et al., 2013). Using the hourly model output with the 9-km horizontal resolution, Wu and Chen (2016) demonstrated that the conventional steering calculated over a certain radius from the TC center in the horizontal and a deep pressure layer in the vertical plays a dominant role in TC motion since the coherent structure of TC circulation makes the contributions of other processes largely cancelled out, but the instantaneous motion and the trochoidal motion around a mean track can considerably deviate from the conventional steering (Chan et al., 2002) also found that the horizontal PV advection played a dominant role when there were no significant changes in the speed and direction of the TC motion. Using the hourly model output with the 9-km horizontal resolution, (Liang and Wu 2015) investigated the sudden northward track turning of the western North Pacific TCs embedded in a monsoon gyre, which is a low-frequency, cyclonic vortex in the lower troposphere with a diameter of about 2500 km (Lander 1994; Carr and Elsberry 1995; Wu et al., 2013). They showed that the sudden northward track turning can be generally accounted for by changes in the synopticscale steering flow resulted from the interaction between the TC and the monsoon gyre. Note that the PVT method used in these studies were based on simulation data with relatively coarse resolution (~10 km), although some numerical simulations include nested domains with the horizontal resolution (~1 km).

As the grid spacing decreases to ~1 km or less, the PV structure of the simulated TC is an annular tower of high PV with low PV in the inner-core region, rather than a tower of high PV (e.g., Möller and Smith 1994; Yang et al., 2020). The reversal of the radial PV gradient in the eyewall is subject to dynamic instability (Montgomery and Shapiro 1995; Menelaou et al., 2018), leading to a complicated PV structure due to breaking and mixing of the PV distribution (Schubert et al., 1999). With the increase of the model resolution, the simulated PV structure in the TC inner core is increasingly complicated due to the presence of small-scale structures (Ito et al., 2017; Wu et al., 2018, Wu et al., 2019). Given that the tendency and horizontal gradient in the PVT method are calculated using the difference method, it is necessary to reevaluate the PVT method in the high-resolution (~1 km) simulation, which is an important approach to understand TC intensity and structure change.

## High-resolution simulation of TCs

The simulation data used in this study are from two numerical experiments, which were conducted in previous studies (Chen et al., 2011; Wu and Chen 2016). While the first simulation mainly provides the 1-h outputs with different horizontal resolutions, the second simulation provides the 5-min output in the 1-km resolution domain.

The first numerical experiment is a semi-idealized numerical simulation conducted with the version 3.2.1 of the WRF model. The details of the experimental design can be found in Wu and Chen (2016). It was designed with four two-way interactive domains embedded in the 27-km resolution domain. The grid spacing decreases by a factor of 3 for the nested domains. The corresponding horizontal grid sizes of the nested domains are 9 km, 3 km, 1 km, 1/3 km (333 m) and the number of their grid meshes is 230  $\times$  210, 432  $\times$  399, 333  $\times$  333, 501  $\times$  501, and 720  $\times$ 720, respectively. The three innermost domains move with the simulated TC. The model includes 75 vertical levels with a model top of 50 hPa. The Kain-Fritsch cumulus parameterization scheme and the WRF single-moment 3-class scheme are used in the outermost domain (Kain and Fritch 1993), and in the four nested domains the WRF 6-class microphysics scheme is selected (Hong et al., 2006). The Yonsei University scheme was adopted for PBL parameterization in the outer domains (Noh et al., 2003) and the LES simulation was used in the sub-kilometer domains (Mirocha et al., 2010). In this study we focus only on the evaluation of the PVT method and the evolution of the TC structure simulated with the different grid sizes is not discussed. The multi-domain datasets can be used to examine the influence of the small-scale PV features in the inner core of the simulated TC.

Following (Wu and Chen 2016), the low-frequency background was from that of Typhoon Matsa (2005) from 0000 UTC on 5 August to 0000 UTC on 9 August 2005 (Duchon 1979). A symmetric vortex was first spun up for 18h on an f-plane and then put in the low-frequency background at the center of Matsa (25.4°N, 123.0°E). The experiment was run for 72 h over an open ocean with a constant sea surface temperature of 29°C. Considering the possible adjustment of the vortex structure, we used the output from 24 to 72 h. Figure 1 shows the 48-h track and intensity of the simulated TC in the semiidealized experiment at 1-h intervals. The simulated storm takes a generally recurving track. The minimum sea-level pressure fluctuates around 950 hPa, while the 10-m maximum wind generally decreases during the 48-h period. The simulated TC maintains the strength of categories 2 and 3.

The other experiment is the prediction of Hurricane Wilma (2005) from (Chen et al., 2011). The initial and lateral boundary conditions were from then-operational GFDL model data. The details about the setup of the WRF model and the corresponding verification can be found in (Chen et al., 2011). The experiment included four interactive domains with the horizontal grid



pressure (SLP) (black, hPa) and 10-m maximum wind speed (red, ms<sup>-1</sup>).



spacing of 27 km, 9 km, 3 km and 1 km, respectively, with 55 vertical levels. The output of the innermost domain at 5min intervals is used in this study. The 72-h experiment was initialized at 0000 UTC 18 October 2005 and terminated at 0000 UTC 21 October 2005. Figure 2 shows the predicted track and intensity of Wilma after the first 18-h spin-up. The predicted storm generally takes a northwestward track. Consistent with the observation, the predicted storm experiences an initial spin-up, the rapid intensification from 18 to 36 h and the intensity change associated with the eyewall replacement during the last 36 h.

## PVT method and TC center detection

The PV tendency in the coordinates moving with a TC is written as (Wu and Wang 2000)

$$\left(\frac{\partial P}{\partial t}\right)_{1m} = \left(\frac{\partial P}{\partial t}\right)_{1f} + \vec{C} \cdot \nabla P_s,\tag{1}$$

where subscript 1 represents the wavenumber-1 component of the PV tendency and subscripts m and f denote the moving and fixed reference frames relative to the TC center, respectively.  $\vec{C}$  is the translation velocity of the TC, and  $P_s$  is the symmetric component of PV with respect to the TC center. In (1), the Ertel PV (P) is written as

$$P = \frac{1}{\rho} \vec{\eta} \cdot \nabla \theta_{\nu}, \tag{2}$$

where the absolute vorticity  $\vec{\eta} = 2\vec{\Omega} + \nabla \times \vec{V}$ , and  $\vec{\Omega}$  and  $\vec{V}$  are the angular velocity of rotation of Earth and three-dimensional velocity, respectively. The total air density ( $\rho$ ) includes the densities of dry air, water vapor, condensed water, and hydrometeors.  $\theta_{\nu}$  is the virtual potential temperature.

In (Wu and Wang 2000), the PV tendency in the moving reference frame was neglected since it is much smaller than the one in the fixed reference frame in their study. In this study, we retain the term and calculate it with the time difference method. The PV tendency in the fixed reference frame can be calculated with the time difference or from the PV tendency equation. The gradient of the symmetric component of PV are calculated with the space difference method. The least square method is used to solve the zonal and meridional components of the translation velocity within a radius of 90 km (Wu and Wang 2000). In fact, when the calculation area covers the inner-core region, the results are insensitive to the selection of the radius since the TC translation velocity is weighed by the gradient of the symmetric component of PV.

When quantifying the contributions of individual physical processes, we use the PV tendency equation in the fixed reference as follows

$$\left(\frac{\partial P}{\partial t}\right)_{1f} = \wedge_1 \left( -\overrightarrow{V_h} \cdot \nabla_h P - w \frac{\partial P}{\partial z} + \frac{\overrightarrow{\eta}}{\rho} \cdot \nabla \frac{d\theta_v}{dt} + R \right), \quad (3)$$

where  $\overrightarrow{V_h}$  and w are the horizontal and vertical components of the wind velocity. The right-hand side of (3) includes the horizontal advection (HA), vertical advection (VA), diabatic heating (DH) terms. R includes the friction term and the effects related to momentum flux. In this study, R is calculated as a residual of (3) and thus also includes the calculation error. The operator  $\wedge_1$  is used to obtain the wavenumber one component. In addition to the PV tendency in the moving coordinates (PVTm), as indicated in (1) and (3), the TC translation velocity calculated with the PVT method can result from five physical processes: the development of the wavenumber-one component (PVTm), HA, VA, DH and R.

As indicated in (1), the symmetric and wavenumber one components with respect to the TC center are calculated in the PVT method. The accuracy of the decomposed symmetric and asymmetric components of the TC circulation depends critically on the TC center position due to strong wind speed and the associated strong radial gradient in the inner-core region (Willoughby 1992; Yang et al., 2020). Focusing on the resulting symmetric and asymmetric inner-core structures, Yang et al. (2020) evaluated four center-detecting methods that are often used in TC simulations by comparing the evolution of the small-scale track oscillation and vortex tilt. The centers detected with the four methods are the pressure centroid center (PCC), the PV centroid center (PVC), the maximum tangential wind center (MTC), and the minimum pressure variance center (MVC). They found that the maximun symmetric inner-core structures can be obtained with the MVC and MTC, while the resulting track oscillations and vortex tilt are smoother with the MVC than with the MTC. It is suggested that the MVC and MTC can be used in high-resolution simulations.

The method for detecting the MVC was described in Braun et al. (2006) and Yang et al. (2020). That is, the azimuthallyaveraged variance of the pressure field in the inner core is calculated with an assumed TC center in the inner-core region until the minimum azimuthally-averaged variance is reached. The radius for calculating pressure variance is 90 km in this study. The method for detecting the MTC was described in Marks et al. (1992), Wu et al. (2006) and Reasor et al. (2013). The MTC is associated with the maximum azimuthal mean tangential



Comparisons of the zonal **(A)** and meridional **(B)** TC translation speed based on the MVC (black) and MTC (red) at the 5-km altitude. The 1-h model output in the 9-km domain is used from the semi-idealized experiment.



wind. To reduce abrupt changes in the track oscillations, we detect the maximum azimuthal mean tangential wind in an annulus, rather than a single circle. The annulus is defined between 0.7 and 1.3 times the radius of maximum wind (RMW).



### FIGURE 5

The zonal (left) and meridional (right) components of the PVT translation velocity based on the symmetric PV gradient from the 9-km domain and the wavenumber-one component of PV tendency from the 1/3-km domain (red), the symmetric PV gradient from the 1/3-km domain and the wave-number-one component of PV tendency from the 9-km domain (blue), and the symmetric PV gradient from the 9-km domain and the wave-number-one component of PV tendency from the 9-km domain (black) at **(A,B)** 8 km, **(C,D)** 5 km and **(E,F)** 1 km in the semi-idealized experiment. The TC speed is based on the MVC.



### FIGURE 6

The zonal (left) and meridional (right) components of the PVT translation velocity based on the symmetric PV gradient from the 9-km domain and the wave-number-one component of PV tendency from the 1/3-km domain (red), the symmetric PV gradient from the 1/3-km domain and the wave-number-one component of PV tendency from the 9-km domain (blue), and the symmetric PV gradient from the 9-km domain and the wavenumber-one component of PV tendency from the 9-km domain (black) at (A,B) 8 km, (C,D) 5 km and (E,F) 1 km in the semi-idealized experiment. The TC speed is based on the MTC. In the semi-idealized experiment, the RMW is about 30 km, and the width of the annulus is about 18 km. Figure 3 shows the comparisons of the zonal and meridional components of the translation velocity of the center at the 5-km altitude based on the data from the 9-km domain at 1-h intervals in the semi-idealized experiment. Despite the small differences, Figure 3 indicates that the TC velocity can be calculated based on both the MVC and MTC.

## Failure of the PVT diagnostic method in the high-resolution simulation

We first demonstrate the failure of the PVT diagnostic method in the high-resolution simulation when the hourly model output is used. The TC translation velocity (hereafter PVT velocity) is retrieved by using the 1-h instantaneous outputs from the various model domains in the semi-idealized experiment. Note that the TC translation velocity calculated with the center position (hereafter TC velocity) at 1-h intervals are nearly identical in these domains despite the varying grid sizes (figure not shown). In this section, the MVC is used as the TC center to demonstrate that the PVT velocity can be very different when the grid sizes become smaller than 9 km.

Figure 4 shows the time series of the zonal and meridional components of the TC translation velocity retrieved with the PVT method at the 5-km altitude. The PVT tendency terms were calculated with the time difference method. As shown in Figure 4, while the PVT velocity in the 9-km domain is in good agreement with the TC velocity, the PVT method fails to retrieve the TC velocity in the 3-km, 1-km and 1/3-km domains. Moreover, as the grid size becomes finer, the PVT velocity is closer to zero.

With the increase of the numerical model resolution, the simulated PV structure contains complicated small-scale structures (Ito et al., 2017; Wu et al., 2018, 2019). To understand the influence of the complicated small-scale features on the PVT method, we first conducted two experiments, in which we replaced the wave-number-one component of PVT and horizontal gradient of the symmetric PV component with those in the 1/3-km domain, respectively. By doing this we can pinpoint the influence of the complicated small-scale features. Figures 5, 6 show the comparisons of the resulting zonal and meridional components of the PVT velocity for the three selected levels, which represent the TC motion at the lower, middle and upper levels. When the wavenumber-one component of PV tendency is replaced with that in the 1/3km domain, the TC motion can be well retrieved with the MVC and MTC. However, when the horizontal gradient of the symmetric PV component in (1) is replaced with that in the 1/3-km domain (blue), the TC motion is poorly retrieved. Both the zonal and meridional components are close to zero. It is suggested that the failure of the PVT method in the 1/3-km

domain arises from the calculation of the horizontal gradient of the symmetric PV component.

Why cannot we use the 1-h instantaneous output to represent the horizontal gradient of PV over the 1-h period? Figure 7 shows the radius-altitude cross section of the azimuthal mean PV in the 9-km and 1/3-km domains at 48 h. In the 9-km domain, as shown in Figure 7A, the simulated TC exhibits a bowl-shaped PV anomaly within the radius of ~30 km, extending to 12 km. The maximum of PV is more than 35 PVU (1 PVU =  $10^{-6}$  kg<sup>-1</sup> K m<sup>2</sup> s<sup>-1</sup>) in the eye, with large radial PV gradient in the inner side of the eyewall. The simulated PV structure agrees well with previous studies (e.g., Chen and Yau 2001; Martinez et al., 2019). As the horizontal grid spacing decreases to 1/3 km (Figure 7B), there are many small-scale features although the bowl-shaped structure can be generally identified. A significant difference is localized PV maxima in the inner side of the eyewall. Figure 8 further shows the comparison of the horizontal cross section of PV in the two domains and their difference at the 5-km altitude. Comparing with the PV in the 9-km domain (Figure 8A), there are small-scale features outside of the high-PV core (Figure 8B). The small-scale structures are clearly seen in their difference (Figure 8C). It is suggested that the presence of the small-scale features can make difference in the calculated horizontal gradient of the symmetric PV component.



### FIGURE 7

The radius-altitude cross section of the azimuthal mean PV (shaded, unit:  $10^{-6}$  m<sup>2</sup>s<sup>-1</sup> Kkg<sup>-1</sup>) and the azimuthal mean wind (contour, unit: ms<sup>-1</sup>) at 48 h in the semi-idealized experiment. The PV is calculated with the outputs from (A) the 9-km domain and (B) 1/3-km domain, respectively.



### FIGURE 8

The 5-km PV ( $10^{-6}$  m<sup>2</sup>s<sup>-1</sup> Kkg<sup>-1</sup>) field calculated from (A,D) the 9-km domain and (B,E) the 1/3-km domain and (C,F) their difference at 48 h (left) and 57 h (right) in the semi-idealized experiment. The plus symbol denotes the position of the TC center.

Figure 9 shows the comparisons of the zonal and meridional gradients of the azimuthal mean PV calculated from the 9-km domain and the 1/3-km domain at 48 h at the altitude of 5-km in the semi-idealized experiment. Compared with the 9-km domain, the PV gradient in the 1/3-km domain exhibits complicated radial structures, which result from the azimuthal mean of the small-scale features characterized by localized high PV maxima. In other words, the symmetric PV component is distorted by the localized PV maxima when making the azimuthal average.

# Improvement of the PVT method for high-resolution simulations

In the last section, we show that the presence of the smallscale PV features can distort the resulting symmetric PV component. If the simulation with the high spatial resolution can resolve the small-scale features with very high PV, the evolution of the small-scale features must be represented in the calculation of the PV tendency. On the other hand, the



small-scale features must be smoothed out as the PVT method is applied to the data at longer time intervals. Thus we can have two approached to improve the PVT method for high-resolution simulations.

### Increasing time intervals

Since the PV structure cannot be well resolved with the 1-h model output, we first calculated the PV tendency and TC velocities with the 2-h model output of the semi-idealized experiment at 1-min intervals. Figure 10 shows the comparison between the PVT velocity and TC velocity based on the MTC and MVC, respectively. We can see that the TC motion at 1-min intervals can be retrieved with the PVT method, the correlation coefficients of zonal and meridional

velocities are 0.95 (0.80) and 0.91 (0,72) for the MTC (MVC), respectively. The fluctuations of PVT velocity with the MVC are smaller than that with the MTC. Note that the TC translation velocity from the 1-min output can also have larger errors than that from 1-h output.

We also examined the retrieved PVT velocity using the prediction data of Hurricane Wilma (2005). Using the 5-min data, the TC velocity can be well retrieved with both the MTC and MVC although there are differences of the PVT velocity between the MVC and the MTC (Figure 11). For clarity, the 1-h running average is applied to the time series. Taking the MTC as an example, the zonal and meridional correlation coefficients between the TC and PVT velocities are 0.94 and 0.98, respectively. It is suggested that the TC velocity in the high-resolution simulation can be retrieved with the PVT method using the minute-interval output.



#### -TC Zonal Speed (ms<sup>-1</sup>) ms\_ Zonal Speed Aeridional Speed (ms<sup>-1</sup>) Speed (ms<sup>-1</sup>) D тс 4 Meridional 24 36 48 Time (h) 72 24 36 60 60 48 72 Time (h) FIGURE 11 Comparisons of the zonal (A,B) and meridional (C,D) TC translation velocity derived from the TC center position (black) and PV tendency (red) based on the MTC (A,C) and MVC (B,D) of the 5-min output in the 1-km domain at 5-km altitude in the prediction of Hurricane Wilma. The 1-h running average is applied to the time series.

# Reducing influence of small-scale PV features

(Zhao et al., 2019) suggested that the PVT method can be improved by smoothing the meteorological fields such as wind speed and temperature. However, the smoothing method and time were subjectively selected in (Zhao et al., 2019). In this study, a horizontal-thinning method is introduced to smooth the model output. The model output is horizontally thinned to the resolution of ~10 km for the calculation of the PVT velocity at the interval of ~1 h.

The thinning method can be illustrated by calculating the PVT velocity in the simulation of Hurricane Wilma. Here we use the instantaneous output in the 1-km domain at a time interval of

1 h. First, a 9-point smoothing is applied to the fields required in the PVT method, and then we have the resulting fields with the horizontal resolution of 3 km. Note that we cannot directly smooth the PV field since it is dominated by the positive PV. Second, using the smoothed fields, we repeat the first step and then we have the thinned output with the horizontal resolution of 9 km. Finally, we use the thinned, smoothed model output to calculate the PVT velocity.

Figure 12 shows the comparisons between the PVT velocity and TC velocity based on the instantaneous 1-h output in the 1km domain. The PVT velocity is retrieved with the original output (black dashed) and thinned output (red dashed), respectively. For clarity, a 3-h running average is applied to the time series. While the PVT velocity from the original output



deviates significantly from the TC velocity, the TC velocity is well retrieved from the thinned output. The correlation between the TC and PVT velocities is 0.94 and 0.95 in the zonal and meridional components, respectively. Note that we can improve the performance of the PVT method by increasing time intervals and spatially smoothing model output.

## Diagnosing contributions of physical processes

As shown in the last section, the TC translation velocity in the high-resolution can be well retrieved by increasing the output time interval or/and thinning the model output. In this section, we show that using the PV tendency equation, we can examine the individual contributions of the five physical processes: HA, VA, DH, the development of the wavenumber-one component (PVTm), and R. Since the PVT method can retrieve the TC translation velocity with 1-km resolution predicted data of Hurricane Wilma (2005), here we use the experiment as an example to illustrate the diagnosis of the contributions from various physical processes. Figure 13 shows the contributions of the five processes to the PVT velocity at the altitudes of 1 km, 5 and 8 km at 28 h. Note that the PVT velocity at each level is the vector sum of the contributions of the five processes. Consistent with the TC velocity, the retrieved PVT velocities at the three levels are nearly the same, and direct to the southwest. The PVT velocity is dominated by the contributions of HA, VA, and DH at



5 and 8 km, while R also plays a role in the boundary layer at 1 km. Although the individual contributions vary in the vertical, the TC vortex can maintain its coherent structure in the vertical.

Figure 14 further shows the time series of the contributions of the individual contributions at 5 km. Since the predicted Wilma maintains a relatively symmetric structure, the contribution of PVTm is generally small during the period, while the contribution of R increases since 60 h. As indicated in Wu and Chen (2016), the



contributions of these processes are not physically independent due to the coherent structure of the TC. For example, the correlation between the contributions of VA and HA is -0.53 in the zonal component and the correlation between the contributions of VA and DH is -0.4 in the meridional component. It is suggested that the contributions of these processes to the TC motion can partially cancelled each other.

## Summary

As the model grid spacing decreases to ~1 km, the PV structure of the simulated TC is an annular tower of high PV with relatively low PV in the eye and the resulting reversal of the radial PV gradient in the eyewall is subject to dynamic instability. While the PVT method has been successfully used to understand TC motion in numerical simulations with relatively coarse resolution (~10 km), this study reevaluates the performance of the PVT diagnosis method by using the model output from two high-resolution TC simulations with grid spacing of ~1 km.

It is indicated that the failure of the PVT method in the highresolution simulation arises from the influence of small-scale PV features on the calculation of the horizontal gradient of the symmetric PV component. This study demonstrates that the TC translation velocity in the high-resolution can be well retrieved by increasing the output time interval and thinning the model output. As the grid spacing decreases to ~1 km, the minute-interval model output can be used to calculate the PVT velocity. Alternatively, when the 1-h model output is available, a 9-point smoothing is successively used to thin the fields required in the PVT method, until the grid spacing of the smoothed fields is ~10 km. The thinned model output can be used to calculate the PVT velocity.

In this study, we also demonstrate that the PV tendency can be divided into two parts: one for TC motion and the other for the development of the asymmetric structure of TCs. It is suggested that the PVT method can be useful to understand the changes of TC structure and the associated intensity.

## Data availability statement

Publicly available datasets were analyzed in this study. This data can be found here: https://mygeohub.org/resources/1601.

## Author contributions

LW and TX conducted analysis and writing. TX contributed to figures included in this manuscript. JY offered advices to the writing.

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## Conflict of interest

The reviewer YW is currently organizing a Research Topic with the author LW.

The remaining authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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