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A practical approach to develop a proper anisotropic rock physics model for media comprising of multiple fracture sets in the absence of sophisticated laboratory/wireline data

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For economical production from a fractured reservoir, a characteristic analysis of the fracture parameters like its density and orientation within the reservoir is essential to improve the fluid flow during extraction. This study deals with the development of a proper anisotropic rock physics model for a media with multiple fracture sets to study the spatial distribution of important fracture parameters i.e., fracture density and orientation in the absence of sophisticated laboratory/wireline and pre-stack seismic data. The crest of hydrocarbon producing fault-bounded Balkassar Anticline in Northern Potwar, Upper Indus Basin, Pakistan is selected as a case study representing a potential zone for development of fractures at reservoir level (Sakesar Limestone). The methodology consists of the interpretation of 3D post-stack seismic and conventional wireline log data to demarcate the reservoir containing fractures. The Ant-tracking discrete fracture network (DFN) attribute is applied on 3D post-stack seismic data to obtain an initial estimate about the presence of fracture corridors and their orientations. Based on this initial estimate, a proper rock physics model has been developed utilizing inverse Gassmann relations, T-matrix approximation, and Brown and Korringa relations. The output from the developed rock physics model has been displayed in the form of 13 effective independent elastic stiffness constants (monoclinic symmetry-representing media comprising of multiple fracture sets) as a function of fracture densities and azimuthal fracture orientations. A clear decreasing trend in effective elastic stiffness constants with increasing fracture densities can be observed. Similarly, a periodic trend of effective elastic stiffness constants with fracture orientations can be observed. These trends are more or less expected, but they would have been difficult to quantify without a proper rock physics model. The use of independent effective elastic constants for the generation of synthetic seismic amplitude versus angle and azimuth (AVAZ) data and its correlation with observed seismic AVAZ data in a geostatistical sense has been discussed.

KEYWORDS

fractured reservoirs, ant-track attribute, anisotropic rock physics model, seismic AVAZ data, bayesian inversion, fracture density, fracture orientation

1 Introduction

Fractures are commonly found in carbonate reservoir rocks (Zhiqian et al., 2016; Bagni et al., 2020) and hold the key for the extraction of hydrocarbons from the reservoir. Prolific management of the fractured reservoirs involves effectively defining fluid flow pathways during production for which effective characteristic analysis of fracture systems is essential (Ali and Jakobsen, 2011a; Ali and Jakobsen, 2011b). Furthermore, fracture properties such as fracture density and azimuthal fracture orientation, when studied spatially, can assist in the optimization of fractured reservoir production (Sayers, 2009; Ali et al., 2015). However, despite the advancement in methodologies pertaining to subsurface data acquisition and processing, identification and adequate characteristic analysis of fractures crucial to production remain a task yet to be accomplished (Gholipour et al., 2016; Bagni et al., 2020).

The investigation of fractures within a carbonate reservoir can be carried out efficiently through sophisticated subsurface borehole data (Bagni et al., 2020). However, the fracture corridors within the carbonate reservoirs are typically extended over tens to hundreds of meters in terms of height and width while being regionally spread over a kilometer (Huang et al., 2017). Furthermore, the well-based petrophysical analysis used in the characterization of fracture properties (like fracture density) cannot be available throughout the reservoir (Jadoon et al., 2006; Bagni et al., 2020). Therefore, characterizing a fractured reservoir using borehole data may require drilling more wells, which is practically not feasible. In order to tackle this problem, a simple yet practical approach is to develop a proper rock physics model which can be utilized to generate synthetic (calculated) seismic amplitude versus angle and azimuth (AVAZ) data. This synthetic seismic AVAZ data can be correlated with observed seismic AVAZ data, if available, in a geostatistical sense to obtain the spatial distribution of important fracture parameters required to enhance fluid flow.

Rock physics modeling is a vital tool for studying the seismicbased characterization of a reservoir (Misaghi et al., 2010). The focus of rock physics concepts is to link the composition of a rock with its macroscopic properties (Wang et al., 2018). The use of an appropriate rock physics model can, therefore, delineate characteristics of a rock effectively (Wang et al., 2018). In terms of seismic fracture characterization, rock physics modelling offers a reduced number of unknown parameters, thus offering a relative advantage during seismic inversion (Ali and Jakobsen, 2014). Furthermore, a proper rock physics model delivers a better physical description of the subsurface provided that the model is able to relate the fracture parameters with the effective elastic stiffness constants (Ali and Jakobsen, 2014).

The reason for the difficulty in the prediction of subsurface fracture systems from seismic data is the lack of resolution at

which these fractures can be mapped (Souque et al., 2019). The accurate prediction of the fracture systems requires sophisticated laboratory/wireline and pre-stack seismic data, including core cuttings, full bore formation micro imager (FMI) logs, and seismic amplitude versus angle and azimuth (AVAZ) data at the reservoir level. However, this data is not easily available. Therefore, it is a challenging task to develop a proper anisotropic rock physics model utilizing conventional wireline log and 3D post-stack seismic data.

This study proposes a practical approach for the development of a proper rock physics model in the case of a reservoir (Sakesar Formation in this study) containing multiple fracture sets in the absence of sophisticated laboratory/wireline data. The crest of the Balkassar anticline has been taken as the case study as it provides a geological advantage of high probability of presence of multiple fracture sets. Balkassar oil field is present in Potwar basin, a north-western part of Himalayan Mountain ranges (Gee and Gee, 1989; Kemal et al., 1991) as depicted in Figure 1. A compressional regime acting on the area led to the development of anticlinal structural traps bounded by reverse faults. Patala Shale and Pre-Cambrian Salt Range Formation are considered to be the main source rocks in the locality (Bender et al., 1995; Masood et al., 2017). The main reservoirs are the fractured carbonates of the Sakesar and Chorgali formations in the Balkassar oil field with the Chorgali Formation also acting as the seal rock (Khan et al., 1986).

The input data for developing the rock physics model is obtained from available conventional wireline log data (three tracks-lithology, resistivity, and porosity) of Well OXY-01 and 3D seismic post-stack data acquired at the Balkassar oil field. The workflow followed for the generation of such a rock physics model in the case of multiple fracture sets is displayed in Figure 2. section 2 of this paper focuses of the methodology followed to achieve the objective of developing a proper rock physics model. section 3 discusses the results obtained from each applied method including the output of the proposed rock physics model. Section 4 deals with the discussion and implication of the applied method, while section 5 concludes the study.

2 Methodology

2.1 Demarcation of fractured reservoir (Sakesar Limestone)

Interpretation of the 3D seismic line focuses on the extension of the formation as well as the demarcation of regional structures. In this study, 3D seismic data of Balkassar area with a dominant



frequency of 35 Hz has been interpreted for the purpose of developing the discrete fracture network (DFN) model at the crest of Balkassar anticline (Figure 3). The important horizons including Chorgali, Sakesar, Patala, and Lockhart formations have been marked. The demarcation has been done with the help of synthetic seismogram (seismic to well tie tool) generated from the well log data of OXY-01 and wavelet extracted from the seismic traces nearest to the well. Faults have been marked on the basis of seismic discontinuities and background geological knowledge (Figure 3).

2.2 Discrete fracture network modeling

Model development of complete 3D characteristic profiling of fracture networks present within naturally fractured reservoirs (NFR) is essential for improving the supervision of a reservoir and consequently, the rate of recovery (Narr et al., 2006; Bisdom et al., 2014). DFN modeling is a proficient tool used for the depiction of the spatial distribution of fractures as well as the details of properties like location, size, density, orientation, and conductivity of individual fractures within an NFR (Tran et al., 2006). It takes into account the geometry of fractures, their conductive ability, and interconnection to construct a fracture network (Aydin and Akin, 2019). Furthermore, DFN ensures the contribution of each fracture as a separate entity unlike the continuum approach and dual-porosity modeling. It, therefore, establishes a more realistic display of fracture systems and paves a path to study response to the flow impact from properties of individual fractures (Aydin and Akin, 2019). This study makes use of the DFN model for Sakesar Formation in the Balkassar area to ensure the type of fractures present in the subsurface reservoir before proceeding to develop a rock physics model based on this priori information.



The Ant-Tracking algorithm has been utilized for the construction of the DFN model. The ant-tracking algorithm simulates a fracture pattern analogous to ants making pathways optimal for searching food (Hu et al., 2014). During their search for food, ants leave pheromones along the pathways to inform other ants about the possibly shortest passage to any located food (Silva et al., 2005). The same concept is used in the Ant-tracking algorithm for 3D seismic volume (Pedersen et al., 2002). Ant-tracking distributes a vast number of electronic "ants" over a seismic volume with each "ant" moving through a fractured surface-emitting "pheromone" for the rest to follow (Hu et al., 2014). Hence a large number of "ants" will trace planar surfaces like faults and fractures in a 3D seismic volume marked by the "pheromones" that the "ants" left compared to the nonstructure element i.e., noise, which will be marked by a smaller number of "pheromones" (Cox and Seitz, 2007; Hu et al., 2014). Furthermore, deviation up to only 15° is allowed to "ants" from their original direction in order to improve and reinforce the tracing of planar structure (faults and fractures) rather than seismic irregularity due to noise (Fang et al., 2017).

Ant-tracking process uses a series of geophysical attributes which are inserted sequentially in a background program for the detection of faults and fractures on a 3D seismic volume (Ouenes, 2000). The general workflow to run

the Ant-track volume is given in Figure 4. The final output is an attribute volume that shows faults and fracture zones in detail. In this workflow, amplitude contrast with dip guide has been used as an edge detection attribute. Two attributes have been computed in series from the result of the edge detection i.e., 3D edge enhancement and structural smoothing (Figure 5). Output from edge detection is subtracted from the input seismic data. The edge evidence attribute is computed on the output of edge enhancement which further enhances the discontinuities of interest. The final fault volume is obtained by multiplying the output of edge enhancement (A), with the output of edge evidence (B), by using the following syntax in the seismic calculator: if (A>0, A, 0)B. On Fault Cube, Ant track attribute slice is generated to extract regional faults visible on seismic data. The results obtained from the Ant track attribute applied to the seismic data are given in Figures 5A,B.

The application of DFN Ant-track algorithm on 3D seismic post-stack data at the crest of Balkassar anticline gives us an initial estimate of fracture distribution. This initial estimate helps in defining the conceptual model for fracture distribution at the reservoir level, within the Sakesar Formation, for which the effective anisotropic elastic properties can be obtained utilizing rock physics algorithms.



2.3 Rock physics modeling

In carbonates rocks, fractures play imperative role in providing a suitable pathway for flow of the hydrocarbons. The primary/ secondary migration and alignment of these fractures to preferred orientation lead us towards the anisotropic behavior of subsurface rock formation (Ali and Jakobsen, 2011a; Ali and Jakobsen, 2011b). Therefore, in order to study the characteristics of fractures in fractured media, an anisotropic rock physics model has to be developed.

Multiple fracture sets in Sakesar Formation have been identified by means of a seismic Ant-tracking algorithm. On the basis of this observation, a monoclinic anisotropic rock physics model, taking fracture density and fracture orientation as the variables under observation has been developed. The workflow for rock physics modeling is shown in Figure 6. The monoclinic anisotropic system gives 13 independent effective elastic stiffness constants which have been plotted against the fracture densities and fracture orientations of the fractures sets. These effective elastic constants are required for the computation of AVAZ data at the reservoir level.

2.3.1 Mineral properties

The Sakesar Formation acting as a reservoir in the study area consists of compact limestone as dominant lithology indicated by the petrophysical analysis of Well OXY-01, with shale volume of 4%, effective porosity of 5%, and water saturation of 27.7% (Table 1; Figure 7). The standard values of bulk and shear modulus of mineral (calcite) along with density is taken from (Mavko et al., 2009) listed in Table 1.

2.3.2 Dry rock matrix properties (without fractures)

The bulk modulus for the frame (K_{dry}) of the reservoir can be measured from laboratory experimentation, empirical relation,





or formulation using wireline log data (Kumar, 2006). K_{dry} can be determined from wireline log data using rephrased Gassmann equation for K_{dry} (Zhu and McMechan, 1990; Smith et al., 2003; Kumar, 2006; Mavko et al., 2009) given in Eq. 1 as:

$$K_{dry} = \frac{K_{sat} \left(\frac{\varnothing K_{mat}}{K_{fl}} + 1 - \varnothing\right) - K_{mat}}{\frac{\varnothing K_{mat}}{K_{fl}} + \frac{K_{sat}}{K_{mat}} - 1 - \varnothing} .$$
(1)

Here, K_{dry} is bulk modulus of dry porous rock, K_{sat} is the *insitu* value of the bulk modulus of saturated rock, K_{mat} is matrix (mineral) bulk modulus, K_{fl} is effective fluid bulk modulus and \emptyset is porosity. The K_{sat} is found using the density (ρ), P-wave (V_p), and S-wave (V_s) logs using the relation given in Eq. 2 as:

$$K_{sat} = V_P^2 * \rho - \frac{4}{3} \mu_{sat} \tag{2}$$



where, μ_{sat} is the *insitu* shear modulus which is equal to ρV_S^2 . V_P and V_S (in m/s) are estimated by taking inverse of compressional slowness and shear slowness respectively. K_{fl} is calculated for a fractured media saturated with oil and water using Wood's equation for homogenous saturation (Wood, 1955) given in Eq. 3 as:

$$\frac{1}{K_{fl}} = \frac{S_w}{K_w} + \frac{S_o}{Ko}$$
(3)

where, K_w and K_o are bulk moduli of water and oil, respectively, whereas S_w and S_o are saturation of water and oil respectively. The S_w has been estimated using resistivity and porosity logs of Well OXY-01. The S_o can be estimated by subtracting the S_w from 1, since the total sum of saturation of the two fluids is equal to 1. In particular, the bulk moduli for fluid phases (water and oil in this case) have been computed using (Batzle and Wang, 1992) relations. The detailed TABLE 1 Input data for rock physics algorithms obtained from petrophysical analysis of OXY-01 well at reservoir (Sakesar Formation) level (Figure 7). The bulk and shear modulus of mineral (calcite) along with density is taken from (Mavko et al., 2009). The *insitu* velocities and density obtained from wireline log data helps to obtain saturated bulk modulus (K_{sat}) at reservoir level for computation of dry rock properties.

No.	Reservoir properties	Values
1	Bulk Modulus of mineral (calcite)	72 (GPa)
2	Shear Modulus of mineral (calcite)	45 (GPa)
3	Density of mineral (calcite)	2700 (Kg/m ³)
4	Effective porosity	5 (%)
5	Volume of Shale	4.1 (%)
6	Density (in-situ)	2670 (Kg/m ³)
7	Vp (in-situ)	5.86 (km/sec)
8	Vs. (in-situ)	3.88 (km/sec)
9	Water saturation	27.7 (%)
10	Oil saturation	72.3 (%)
11	Bulk Modulus of water	2.328 (GPa)
12	Bulk Modulus of Oil	1.349 (GPa)

workflow for computing the bulk moduli of fluid phases is provided by Kumar (2006). The values used as input for the determination of K_{fl} and K_{dry} are given in Table 1.

2.3.3 T-matrix for incorporating the effect of fractures in poroelastic monoclinic model

For the dry case of the fractured porous medium, the dry effective stiffness tensor C_d^* has been estimated using the T-matrix approach shown in Eq. 4, given by (Jakobsen et al., 2003a; Jakobsen et al., 2003b; Ali and Jakobsen, 2011a; Ali and Jakobsen, 2011b; Ali et al., 2015).

$$C_d^* = C^{(0)} + C_1: (I_4 + C_1^{-1}: C_2)^{-1},$$
(4)

where, ': 'implies double scalar product (Auld, 1990), $C^{(0)}$ adds the matrix properties input into the equation and is termed as the background stiffness tensor, I_4 is the identity matrix for the fourth-rank tensors, and C_1 is the fourth rank tensor representing first-order corrections in order to incorporate the effects of isolated fractures. C_2 is the second-order correction which incorporates the fracture-fracture interaction effect. C_1 and C_2 are computed using Eqs 5, 7.

$$C_{1} = \sum_{r=1}^{N} \nu^{(r)} t_{d}^{(r)}, \qquad (5)$$

where, $v^{(r)}$ is the porosity of inclusion, $t_d^{(r)}$ incorporates the effects of individual fractures of elastic stiffness $C^{(0)}$ and is given in Eq. 6 as:

$$\boldsymbol{t}_{d}^{(r)} = -\boldsymbol{C}^{(0)} \colon \left[\boldsymbol{I}_{4} + \boldsymbol{G}^{(r)} \colon \boldsymbol{C}^{(0)} \right]^{-1}, \tag{6}$$



where, $G^{(r)}$ is a fourth-rank tensor integrated over a characteristic spheroid with shape similar to fractures of type r presented by strain green's function (Jakobsen et al., 2003a; Jakobsen et al., 2003b; Ali and Jakobsen, 2011a; Ali and Jakobsen, 2011b; Ali et al., 2015). The second-order correction is given by:

$$C_{2} = \sum_{r=1}^{N} \sum_{s=1}^{N} v^{(r)} t_{d}^{(r)} \colon G_{d}^{(rs)} \colon t_{d}^{(s)} v^{(s)}.$$
 (7)

Here, $G_d^{(rs)}$ is the strain green's function in the form of a fourth-rank tensor which represents the spatial distribution of fractures over an ellipsoid determining the symmetry of the correlation for fractures (Carcione, 1995; Ali et al., 2015). This function estimates the probability of distribution of the two fracture sets namely r and s. The function $G_d^{(rs)}$ is integrated over an ellipsoid with an aspect ratio equal to $p^{(s|r)}(x - x')$, that calculates the probability density for determining a type s inclusion at x' provided a type r inclusion is present at the point x (Jakobsen et al., 2003a; Jakobsen et al., 2003b; Ali and Jakobsen, 2011a; Ali and Jakobsen, 2011b; Ali et al., 2015). The correlation function $p^{(s|r)}(x - x')$ defines how individual fractures are distributed throughout the strata (Nguyen and Nam, 2011; Ali et al., 2015). The symmetry of the correlation function is represented by the aspect ratio. A rough aspect ratio value used in this study was measured to be 1/1,000 from exposed strata of Sakesar Formation in the vicinity of acquired seismic data. The porosity $v^{(r)}$ of type r fractures relate to the fracture density $\varepsilon^{(r)}$ through the equation $\nu^{(r)} = \frac{4}{3}\varepsilon^{(r)}\alpha^{(r)}$ where $\alpha^{(r)}$ represents the type r fracture's aspect ratio. The modified form of $t_d^{(r)}$ for the two sets of fracture parameters can now be written as given in Eq. 8.

$$\boldsymbol{t}_{d}^{(r)} = \boldsymbol{t}_{d}^{(r)} \left(\boldsymbol{\psi}_{1}, \boldsymbol{\psi}_{2}, \boldsymbol{\varepsilon}_{1}, \boldsymbol{\varepsilon}_{2} \right), \tag{8}$$

where, ψ_1 , ψ_2 denotes the azimuthal fracture orientations for the two fracture sets while ε_1 , ε_2 are the fracture densities for the respective fracture set. In general, T-matrix gives the initial output in form of VTI (Vertically transversely isotropic) model. This output matrix upon 90° rotation gives HTI (Horizontally transversely isotropic) symmetry. An arbitrary rotation through bond transformation converts the output matrix into monoclinic symmetry representing the multiple fracture sets (Figure 6). The effect of fluid in this anisotropic poroelastic model has been incorporated using an anisotropic form of Gassmann's equation called Brown-Korringa relation (Brown and Korringa, 1975) which can be written in the symbolic matrix notation given in Eq. 9 as:

$$S^{*} = S_{d}^{*} + \frac{(S_{d}^{*} - S_{m}): (I_{2} \otimes I_{2}): (S_{d}^{*} - S_{m})}{\phi^{0} (I_{2}.S_{m}.I_{2} - \frac{1}{K_{fl}}) - I_{2}.(S_{d}^{*} - S_{m}).I_{2}}$$
(9)

Here \otimes symbolizes the dyadic product of tensor (Auld, 1990), I_2 is second rank tensor identity, S_m represents the compliance tensor for solid mineral constituent (calcite), S_d^* denotes the effective compliance tensor for dry fractured rock and S^* inculcates the effective compliance tensor for the saturated fractured porous medium into the equation. Total porosity is represented by ϕ^0 which is the sum of storage porosity of the homogenous matrix and the fracture porosity.

The final product of rock physics modeling for two mesoscopic fracture sets with different fracture orientations and fracture densities for monoclinic symmetry is 13 independent effective elastic stiffness constants. These 13 independent constants in terms of Voigt index notation are C_{11} , C_{22} , C_{33} , C_{12} , C_{13} , C_{23} , C_{44} , C_{55} , C_{66} , C_{16} , C_{26} , C_{36} , and C_{45} . The 13 independent constants have been varied with fracture orientations of the two fracture sets while keeping their fracture densities constant and vice versa. The Voigt condensed notation in the matrix from for the case of monoclinic assumption having 13 independent stiffness constants is given in Eq. 10 as:

$$C_{ij} = \begin{bmatrix} C_{11} & C_{12} & C_{13} & 0 & C_{15} & 0 \\ C_{12} & C_{22} & C_{23} & 0 & C_{25} & 0 \\ C_{13} & C_{23} & C_{33} & 0 & C_{35} & 0 \\ 0 & 0 & 0 & C_{44} & 0 & C_{46} \\ C_{15} & C_{25} & C_{35} & 0 & C_{55} & 0 \\ 0 & 0 & 0 & C_{46} & 0 & C_{66} \end{bmatrix}$$
(10)

3 Results

3.1 Spatial extension of Sakesar Formation

A synthetic seismogram is generated using sonic and density logs of OXY-01 for the purpose of seismic-to-well-tie. With the help of this, the top and bottom of Sakesar Formation have been marked at 1.37 and 1.475 s respectively, by correlating the seismic (time domain) data to well (depth domain) data. The interpretation of seismic Inline 235 (Figure 3) shows a clear anticlinal structure which is called the Balkassar anticline. Reverse faults formed as a result of slippage caused by compressional regime bounding the anticline have also been interpreted (Figure 3).

3.2 Ant-tracking DFN model analysis

The DFN model effectively delineates the fractures present within the carbonate reservoir. Greater values displayed by Anttracking volume shows high fractured zone while low values correspond to the less fractured strata which can be related directly with secondary porosity evolution in carbonates (Figure 5A). Fractures are concentrated more towards the south within the anticline while in general, the highly populated fractures towards the east corresponds to the faulting phenomena towards the east of the anticline (Figure 5A). Two dominant fracture sets, almost orthogonal





to one another, oriented in NE-SW and NW-SE direction can be observed in the formation (Figure 5B).

3.3 Rock physics model for monoclinic symmetry

A rock physics model for monoclinic symmetry has been developed for Sakesar Limestone reservoir in Balkassar locality. 13 elastic stiffness constants have been computed and displayed against fracture orientations (ψ_1 , ψ_2) and fracture densities (ε_1 , ε_2) of two mesoscopic fracture sets observed in the formation using Ant-tracking DFN model (Figure 5B). Elastic stiffness constants as a function of ε_1 and ε_2 have been plotted in Figure 8 keeping ψ_1 and ψ_2 constant at 35° and 45° respectively. From Figure 8, it is clear that there is a decrease in elastic stiffness constants with increasing fracture densities showing an inverse linear relation.

In Figure 9, the 13 elastic stiffness constants have been plotted as a function of ψ_1 and ψ_2 for the two fracture sets with constant ε_1 and ε_2 at 0.03 and 0.05, respectively. There is a periodic trend of each elastic stiffness constant with ψ_1 and ψ_2 , thus following the transformation law of the fourth ranked tensor given by (Auld, 1990). There will be no periodic trend if fractures are not present within the formation.

4 Discussion

4.1 Seismic modeling using rock physics model for monoclinic symmetry

Once the saturated effective elastic properties of the porous fracture media have been calculated, the multiple fracture sets can then be characterized on the basis of seismic modeling. The variation of fracture based elastic properties within a rock can be detected based on sophisticated seismic attributes like seismic AVAZ, azimuthal variation of velocity in fractured interval, and shear wave birefringence analysis (Crampin et al., 1980; Lynn et al., 1995; Lynn et al., 1999; Zhu et al., 2004; Will et al., 2005). In this study, a practical method is presented for the development of rock physics model in case of a media containing multiple fracture sets utilizing conventional wireline log and 3D seismic post-stack data. The output from this rock physics model in terms of 13 independent effective elastic stiffness constants can be utilized for determination of monoclinic reflectivity given by Schoenberg and Protazio (1992):

$$\mathbf{R} = \begin{bmatrix} R_{PP} & R_{SP} & R_{TP} \\ R_{PS} & R_{SS} & R_{TS} \\ R_{PT} & R_{ST} & R_{TT} \end{bmatrix}$$
(11)

$$\mathbf{X} = \begin{bmatrix} e_{P1} & e_{S1} & e_{T1} \\ e_{P2} & e_{S2} & e_{T2} \\ \{ - (C_{13}e_{P1} + C_{36}e_{P2})s_1 & \{ - (C_{13}e_{S1} + C_{36}e_{S2})s_1 & \{ - (C_{13}e_{T1} + C_{36}e_{T2})s_1 \\ - (C_{23}e_{P2} + C_{36}e_{P1})s_2 & - (C_{23}e_{S2} + C_{36}e_{S1})s_2 & - (C_{23}e_{T2} + C_{36}e_{T1})s_2 \\ - C_{33}e_{P3}s_{P3} & - C_{33}e_{33}s_{33} \} & - C_{33}e_{T3}s_{37} \end{bmatrix}$$

$$(12)$$

$$\mathbf{Y} = \begin{bmatrix} \{ - (C_{55}s_1 + C_{45}s_2)e_{P3} & \{ - (C_{55}s_1 + C_{45}s_2)e_{S3} & \{ - (C_{55}s_1 + C_{45}s_2)e_{T3} \\ - (C_{55}e_{P1} + C_{45}e_{P2})s_{3P} \} & - (C_{55}e_{S1} + C_{45}e_{S2})s_{33} \} & - (C_{55}e_{T1} + C_{45}e_{T2})s_{3T} \} \\ \{ - (C_{45}s_1 + C_{44}s_2)e_{P3} & \{ - (C_{45}s_1 + C_{44}s_2)e_{S3} & \{ - (C_{45}s_1 + C_{44}s_2)e_{T3} \\ - (C_{45}e_{P1} + C_{44}e_{P2})s_{3P} \} & - (C_{45}e_{S1} + C_{44}e_{S2})s_{33} \} & - (C_{45}e_{T1} + C_{44}e_{T2})s_{3T} \} \\ \\ \{ - (C_{45}e_{P1} + C_{44}e_{P2})s_{3P} \} & - (C_{45}e_{S1} + C_{44}e_{S2})s_{33} \} & - (C_{45}e_{T1} + C_{44}e_{T2})s_{3T} \} \\ \\ e_{P3} & e_{S1} & e_{T1} \end{bmatrix}$$

Here, **R** is reflection matrix; s_1 and s_2 defines the phase slowness vector (horizontal components), the associated eigenvectors are denoted by e_P , e_S , and e_T derived from the Christoffel equations (Mavko et al., 2009), and C_{ij} denotes the elastic stiffness constants. For the case, when **X** and **Y** both are not singular and $(X^{-1}X' + Y^{-1}Y')$ can be mathematically inverted, the reflection matrix is given as:

$$\mathbf{R} = \left(\mathbf{X}^{-1}\mathbf{X}' - \mathbf{Y}^{-1}\mathbf{Y}'\right) \left(\mathbf{X}^{-1}\mathbf{X}' + \mathbf{Y}^{-1}\mathbf{Y}'\right)^{-1}$$
(14)

Here, X' and Y' are similar to X and Y with only difference being replacement of unprimed parameters (incidence medium) by primed parameters (transmission medium). Once, reflectivity from above Eqs. (11-14) is obtained, it can be convolved with the seismic wavelet to obtain the synthetic (calculated) seismic AVAZ data at reservoir level.

4.2 Seismic inversion for porous media with multiple fracture sets

Inverse modeling targets the spatial estimation of fracture parameters, which influence the fluid flow during production (Ali et al., 2015). To obtain the modeled fracture parameters on a reservoir scale, the calculated seismic AVAZ data has to be matched with the field based observed seismic AVAZ data in order to minimize the error before the parameters under observation can be studied. The inverse problem in this particular case can be formulated as given in Eq. 15.

$$G(\mathbf{m}) \approx \mathbf{d}.$$
 (15)

Here, *d* is a vector of observable quantities (acquired AVAZ data), *G* combines the monoclinic rock physics model (as discussed in *Rock Physics Modeling* for Sakesar Formation) with the seismic model (as discussed in *Seismic modeling using rock physics model for monoclinic symmetry*), and *m* inculcates the fracture parameters, selected for the developed monoclinic model (ε_1 , ε_2 and ψ_1 , ψ_2), in the form of a vector. The workflow for inversion of parameters using correlation of synthetic AVAZ data with acquired AVAZ data is given in Figure 10.



Correlation of the developed synthetic data with the prestack 3D seismic AVAZ data will lead to spataial mapping the parameters at reservoir scale, thus immensely assisting in reservoir characterization provided that the synthetic data has been modelled correctly. The plots for these parameters with 13 effective elastic stiffness constants helps to monitor the sensitivity of the forward model. Changing the forward model will change the outcome of the reflectivity. Therefore, the geostatistical correlation helps to select the best possible model to invert for the fracture parameters.

Due to limited availability of borehole data over large areas, a monoclinic rock physics model can be opted to spatially characterize the azimuthal fracture orientations and fracture densities of multiple fracture sets in a fractured reservoir. The challenge however is requirement of accurate priori information regarding petrophysical properties like fluid saturation, porosity, volume of shale etc.; matrix properties like matrix moduli, matrix density, and fracture geometry. The importance of rock physics model lies in its ability to inculcate a large number of input parameters. This however can turn out to be a problem provided an improper priori information is used. It is therefore essential to know the alignment of fracture sets present in a fractured reservoir before a rock physics model can be developed. The Ant-tracking DFN modeling proves to be a good tool for this purpose, however, core data analysis is always preferable. The aspect ratio for fractures used in this study was obtained from outcrop of the studied formation, but core sample analysis for aspect ratio determination is suggested as it affects the results of T-matrix model.

5 Conclusion

In the absence of sophisticated laboratory/wireline data, characterization of fracture parameters (like azimuthal orientation and density) of multiple fracture sets in a fractured media can be achieved using a suitable rock physics model. The accuracy of the developed model depends upon the precision of input parameters and awareness of background geological knowledge. The Ant-tracking DFN model gives a potential initial estimate of fractures, which can be used as *a priori* information for the development of the required rock physics model. Here, in this study, a proper rock physics model for a media comprising of multiple fracture sets (monoclinic symmetry) was developed using T-matrix approach, Wood's equation, and Brown-Korringa relation. The

output in the form of 13 independent effective elastic stiffness coefficients shows an inverse linear trend with ε_1 and ε_2 There is a periodic variation of the effective stiffness coefficients with ψ_1 and ψ_2 confirming the presence of fractures within the formation. These elastic constants can practically be applied to generate a synthetic pre-stack (calculated) seismic AVAZ data using the Schoenberg and Protazio (1992) solution. Fracture properties can effectively be inverted through geostatistical correlation of this synthetic data with pre-stack (observed) 3D seismic AVAZ data. This research facilitates and provides a practical approach to develop a proper rock physics model applicable to fractured reservoirs in order to study the parameters of existing fractures, thus facilitating in increasing the production from these reservoirs.

Author's note

This work is a part of FM's Ph.D thesis.

Data availability statement

The data analyzed in this study is subject to the following licenses/restrictions: The data was officially obtained after applying through the university to the Directorate General of Petroleum Concessions (DGPC) which has jurisdiction over handling the data. Requests to access these datasets should be directed to dgpc@mpnr.gov.pk.

Author contributions

FM implemented the methodology and carried out the research. ZA proof read the manuscript and contributed in finalizing the manuscript. YA provided technical support and

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Conflict of interest

Author ZA is employed by the Digital and Integration, Schlumberger, Pakistan

The remaining authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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