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Assessment of landslide susceptibility, exposure, vulnerability, and risk in shahpur valley, eastern hindu kush

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This study assessed landslide susceptibility in Shahpur valley, situated in the eastern Hindu Kush. Here, landslides are recurrent phenomena that disrupt the natural environment, and almost every year, they cause huge property damages and human losses. These damages are expected to escalate in the study area due to the high rate of deforestation in the region, population growth, agricultural expansion, and infrastructural development on the slopes. Landslide susceptibility was assessed by applying “weight of evidence” (WoE) and “information value” (IV) models. For this, the past landslide areas were identified and mapped on the SPOT5 satellite image and were verified from frequent field visits to remove the ambiguities from the initial inventory. Seven landslide contributing factors including surface geology, fault lines, slope aspect and gradient, land use, and proximity to roads and streams were identified based on indigenous knowledge and studied scientific literature. The relationship of landslide occurrence with contributing factors was calculated using WoE and IV models. The susceptibility maps were generated based on both the WoE and IV models. The results showed that the very high susceptible zone covered an area of 14.49% and 12.84% according to the WoE and IV models, respectively. Finally, the resultant maps were validated using the success and prediction rate curves, seed cell area index (SCAI), and R-index approaches. The success rate curve validated the results at 80.34% for WoE and 80.13% for the IV model. The calculated prediction rate for both WoE and IV was 83.34 and 85.13%, respectively. The SCAI results showed similar performance of both models in landslide susceptibility mapping. The result shows that the R-index value for the very high LS zone was 29.64% in the WoE model, and it was 31.21% for the IV

model. Based on the elements at risk, a landslide vulnerability map was prepared that showed high vulnerability to landslide hazards in the lower parts of the valley. Similarly, the hazard and vulnerability maps were combined, and the risk map of the study area was generated. According to the landslide risk map, 5.5% of the study area was under high risk, while 2% of the area was in a very high-risk zone. It was found from the analysis that for assessing landslide susceptibility, both the models are suitable and applicable in the Hindu Kush region.

KEYWORDS

landslide, weight of evidence, information value, susceptibility, risk

Introduction

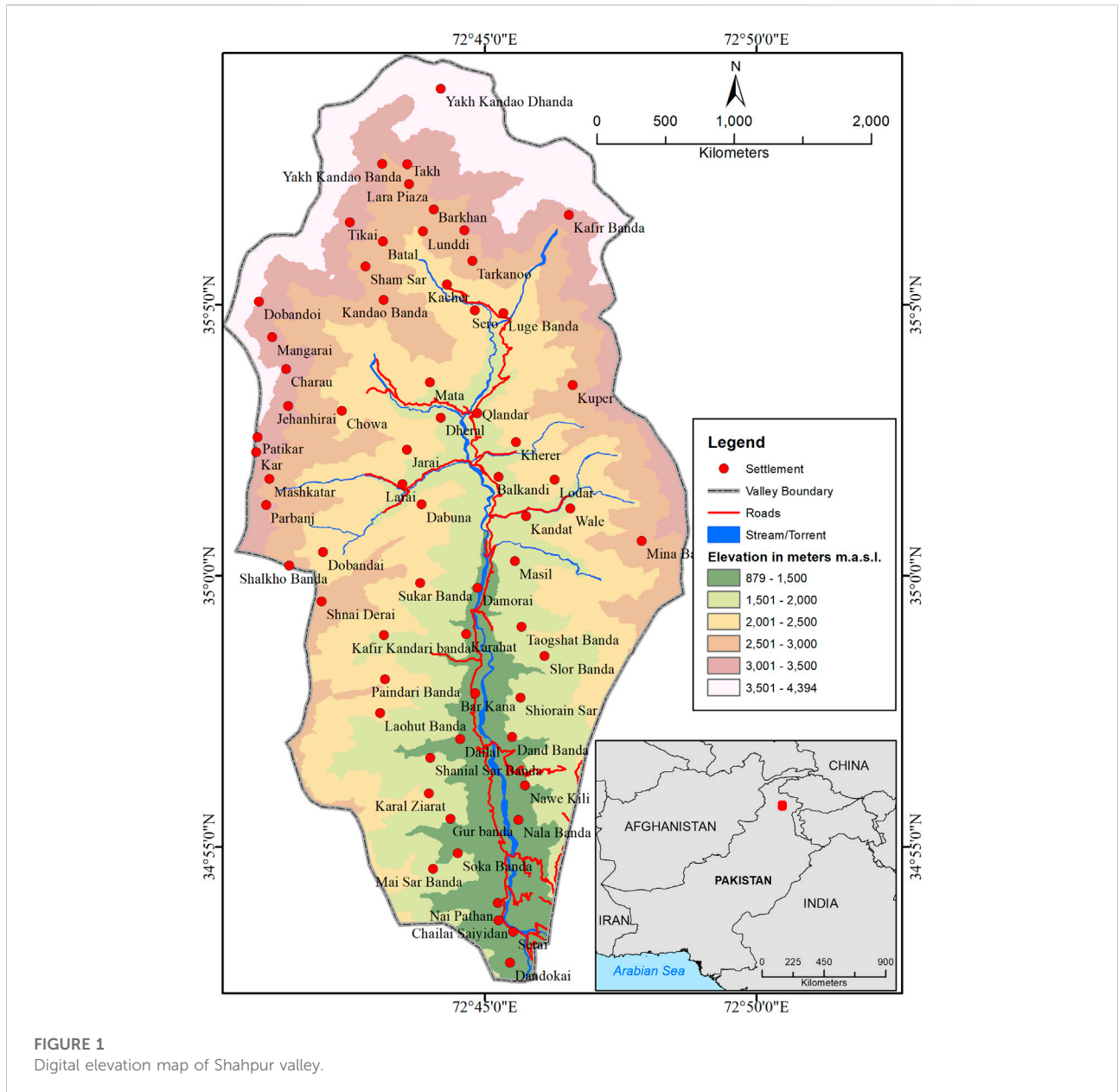
Globally, the frequency of geo-physical and hydro-meteorological disasters has increased in the last two decades with its devastating consequences (Rahman et al., 2017). Landslides are among the geo-physical hazards that cause huge damage to human lives, property, and infrastructure (Wubalem, 2021; Shano et al., 2020). The Hindu Kush Himalayan (HKH) region is a young mountain system where landslides, snow avalanches, floods, and earthquakes are very common (Rahman et al., 2017). In this region, landslides are recurrent phenomena and are often initiated either by seismic activity, prolonged rainfall, and/or human-induced factors (Regmi et al., 2014; Santangelo et al., 2021; Li, 2022a). The frequent landslide events have been causing damages to property as well as infrastructure and sometimes led to human losses. In the literature, it is reported that the global share of landslides was five percent among all the natural hazards during 1990–2005 and tend to increase in the future because of seismic activities, increasing rainfall intensity due to climate change, and anthropogenic activities on the fragile slope (Kanungo et al., 2009; Cui et al., 2021). In 2018, the cumulative share of flooding and landslides in global catastrophic events was 46% among all natural hazards. Landslide is a natural geomorphic process (Allen et al., 2011; Nandi and Shakoor, 2010) mainly governed by the local area surface geology, seismicity, slope gradient/aspect, drainage density/pattern, land use type, and precipitation (Sudmeier-Rieux et al., 2012). In the mountainous area, prolonged rainfall increases the probability of landslide. Similarly, seismic activities and surface lithology are other important contributing factors that govern and initiate slope instability (Van Westen et al., 2010; Li et al., 2021b). Additionally, human intervention in the form of road construction, expansion of human settlement, deforestation, and agricultural activities on fragile slopes further intensifies the probability of landslide occurrence (Rahman et al., 2017; Zhou J. et al., 2021; Li, 2022b).

Many researchers have conducted studies to explore the landslide impacts including human losses, property damage, and infrastructure damages (Davies, 2022; Spegel and Ek, 2022). Very little work has been carried out yet on the

impacts of landslides on the natural environment and ecological system. The environmental effects caused by landslides are changes in agricultural activities, changes in river morphology, and changes to the natural ecosystems (Nakamura et al., 2000). Other effects included sedimentation in river channels and flash floods due to breaching of landslide dams (Tien et al., 2021). Landsliding disturbs the natural habitat of certain endangered species in the susceptible zone and also affects the biodiversity of the affected area; therefore, strict forest preservation measures are highly required to reduce environmental damages (Geertsema and Pojar, 2007).

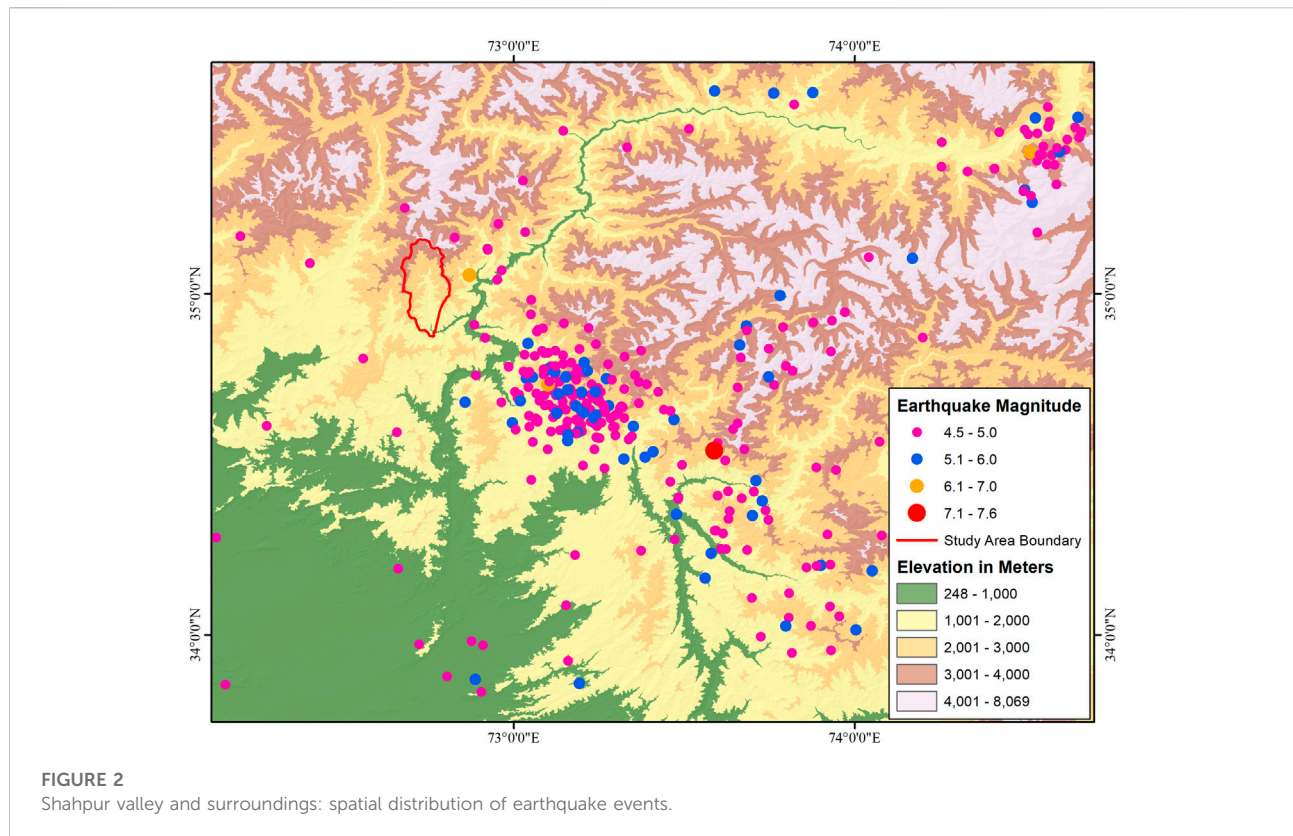
The geo-spatial probability of the slope failure is termed landslide susceptibility. The occurrence of slope failure depends on the presence of some geo-environmental elements (Guzzetti et al., 2005; Li et al., 2021a). During the last decade, numerous scientific studies, including those by Chen and Wang (2007), Goetz et al. (2015), Rahman et al. (2017), Wang et al. (2019), and Ali et al. (2021), have been conducted on fragile mountains and subsequently used a wide range of empirical approaches for analyzing landslide susceptibility to identify the extent of potentially susceptible zones. Quantitative, semi-quantitative, and qualitative techniques have been extensively applied in many studies to map landslide hazard zones (Van Westen et al., 2008; Abdi et al., 2021). In landslide indices, semi-quantitative, quantitative, and qualitative methods are adopted for the identification of landslide-prone areas that have similar geological and geomorphological characteristics (Kouli et al., 2010). In qualitative methodologies, the rating procedure is totally based on indigenous knowledge, whereas semi-quantitative methods are based on weighted procedures (Abdo, 2022). However, quantitative methods use statistical techniques to find out the relationship between factors initiating landslides and landslide events (Rahman et al., 2019).

A number of methodologies and models are in practice for prediction and susceptibility mapping of landslides globally. The spatial probability of landslides can be predicted by applying various quantitative methodologies such as frequency ratio (FR), weight of evidence (WoE), information value (IV), fuzzy neural network (FNN), analytical hierarchy (AH), index of entropy (IoE), analytical hierarchy process (AHP), logistic regression (LR), and fuzzy logic (FL) (Mohammady et al., 2012; Talaei,



2014; Shahabi and Hashim, 2015; Moazzam et al., 2020). However, WoE and IV models have been frequently used for areas with a similar environmental set-up, and therefore both models are applied in this study (Bacha et al., 2018; Mersha and Meten, 2020; Singh et al., 2021). The HKH is an active seismic region, and hence most of the landslides have been initiated by seismic activities and torrential rainfall during monsoon (Kamp et al., 2010). In the HKH region, developmental work over the fragile slopes frequently encourages landslides. This study is an integrated approach that assesses landslide susceptibility, exposure, vulnerability, and risk in the Shahpur valley, eastern Hindu Kush, and is a basic part of disaster risk management

planning. Worldwide extensive work has been carried out recently on landslide susceptibility assessment and mapping (Myronidis et al., 2016; Stanley and Kirschbaum, 2017; Bui et al., 2020; Wang et al., 2020; Chen and Chen, 2021; Ngo et al., 2021; Youssef and Pourghasemi, 2021) and in the Hindu Kush Himalayan region (Ahmed et al., 2014; Rahman et al., 2017; Rahim et al., 2018; Khan et al., 2019; Gautam et al., 2021), but limited research has been carried out on landslide vulnerability and risk (Ram and Gupta, 2022). It is, therefore, necessary to map the landslide vulnerable and risk areas which will help in minimizing the landslide risk and will provide a base for management. In this regard, the present study has



investigated landslide-prone areas using GIS-based information value and weight of evidence models. Consequently, the resultant maps were validated through success and prediction rate curves, seed cell area index (SCAI), and R-index approaches. The unavailability of temporal landslide records is the major limitation of this study, and therefore the susceptibility map was then considered a hazard map for vulnerability assessment. The vulnerability to landslide hazards was calculated based on the exposure of elements at risk to landslide hazard/susceptibility. Based on indigenous knowledge, the probability of loss to each element at risk was calculated for vulnerability mapping. Then, based on the function of hazard and vulnerability, the landslide risk map was generated for the study area.

The study area

The study area, Shahpur valley, lies in the Hindu Raj Mountains. These mountains are considered the offshoot of the Hindu Kush ranges (Rahman et al., 2019). Moving from the north to south, the height of these mountains tends to decrease. The latitudinal extent of the valley is $34^{\circ} 52' 31''$ to $35^{\circ} 9' 35''$ north, while the longitudinal extent is $72^{\circ} 40' 10''$ to $72^{\circ} 48' 44''$ east (Figure 1). Climatically, Shahpur valley is part of the

moist temperate zone. The valley receives ample amount of rainfall from monsoon in summer, while in winter, heavy snowfall occurs at higher altitudes (Rahman et al., 2019). The average annual temperature of the valley is 18°C , and it receives 1,200 mm of rainfall annually (GOP, 1998). In winter, the rainfall occurs due to western disturbances which are usually accompanied by snowfall, while in summer, rainfall occurs in the months of July and August from monsoon. Both summer and winter rainfall causes landslides in the study region. The total area of the district is 258.96 km^2 , out of which 16.82% is cultivated, while the rest (83.18%) is uncultivated including 40.14% under forest cover dominated by softwood coniferous forest. In the study area, agriculture is a major revenue-generating activity carried out in terraced fields which largely depends on rainfall, whereas only 2% of the cultivated area is irrigated (GOP, 1998). The study area consists of rugged topographic terrain where the elevation varies from 879 m to 4,394 m. According to the 2017 population census, the total population of the valley was 115,711 persons, with an average household size of 8.3 persons. The population density in the valley was 446.8 person/sq. km according to the 2017 census.

The study area falls in the active seismic zone of the HKH region. The USGS earthquake catalog reported 499 earthquakes of ≥ 4.5 magnitude during 1973–2019 in the HKH region of Pakistan (Figure 2). The study area consists of young mountain

TABLE 1 Dataset used in the study and source of data.

Data type	Acquired from	Acquisition date	Spatial resolution
Land use and existing landslide layers were obtained from SPOT 5 satellite image	SUPARCO Pakistan	4/11/2013	10 m multispectral and 2.5 m panchromatic
Surface geology and fault lines were digitized from the geological map of north Pakistan	Geological Survey of Pakistan	3/15/2013	—
Slope gradient, slope aspect, and river/stream were derived from ASTER Global Digital Elevation Model V003	https://search.earthdata.nasa.gov/	3/12/2013	30 m
Settlement and administrative unit data	Survey of Pakistan	5/5/2013	—
Road network data	Communication and Works Department, Peshawar, Khyber Pakhtunkhwa	5/10/2013	—

systems that have immature geology (Jehan and Ahmad, 2006) which are prone to landslides and often result in considerable property damage and human losses. The probabilities of these losses are expected to increase due to increasing population pressure on these prone slopes, anthropogenic activities, deforestation, and increasing rainfall intensity. Population growth has posed more pressure on fragile slopes and has made them more vulnerable to landslides.

Materials and methods

Data

The data were collected through primary and secondary sources to meet the required objectives of the study. As a mandatory part of the study, the sites of previous landslides were identified and mapped on the SPOT5 image with a 2.5 m resolution in April 2013, and to reconfirm the ambiguous sites, consecutive field visits were also conducted to validate the landslide sites on the ground. The landslide contributing factors were selected based on previous literature combined with local community knowledge regarding the factors that contribute to or cause landslides. Thus, seven landslide contributing factors, namely, surface geology, proximity to fault lines, slope gradient and aspect, land use, and proximity to roads and streams were selected. Surface geology and tectonic formation determine the rock permeability and strength that ultimately influence slope stability. Similarly, the slope gradient has a direct relationship with landslides as more landslides occur at higher slopes, while the slope aspect has an indirect relationship with landslide occurrence (Khan et al., 2011). Thick vegetation bounces back the initiation of landslides, while vegetation-free land leads to erosion of soil and slope instability. The construction of roads over fragile slopes and stream lateral erosion also encourage slope instability (Zhou Z. et al., 2021); therefore, all these seven landslide contributing factors were selected for landslide susceptibility assessment.

Spatial data of the aforementioned contributing factors were acquired from various sources. The surface geology and tectonics data were traced and extracted from the Geological map of north Pakistan. The vector geodatabase of settlement and administrative units was obtained from the Survey of Pakistan. The shapefiles of the road network were acquired from the Communication and Works Department, Peshawar (Table 1). Similarly, the land use map was derived from a SPOT satellite image through a supervised classification algorithm in ArcGIS 10.2 (Table 1). The stream network, slope gradient, and slope aspect maps were derived from the ASTER 30 m Global Digital Elevation Model (GDEM).

Inventory of landslides

The past landslide sites were mapped on a multispectral satellite image of SPOT (Figure 3A). These sites were verified through a series of field visits. About 300 landslides of varying sizes were marked on the satellite image, and almost 50% of the total landslides were randomly verified through field investigation. Among these landslides, 50% were debris flow and debris slides, while the remaining 40% consisted of mudflow and mudslide (Figure 3B). Only 10% among these landslides were rockfall and rockslide. The debris flow and debris slide mostly occur in the rainy season along the terrace agricultural field, unstable slopes along roads, and streams flowing on steep slopes in the study area. The debris flow and debris slides were also observed in the high snow-melting areas. Rockfall and rockslides were observed along fault lines in steep slope areas. According to the local people, rockfall and rockslides mostly occur during earthquakes in the study region. The mapped landslides were randomly divided into 75 and 25% as training and validation landslides, respectively (Figure 3A). These landslides were then rasterized to find out the pixel number in all classes of each contributing factor map to calculate the weight values for IV and WoE models.

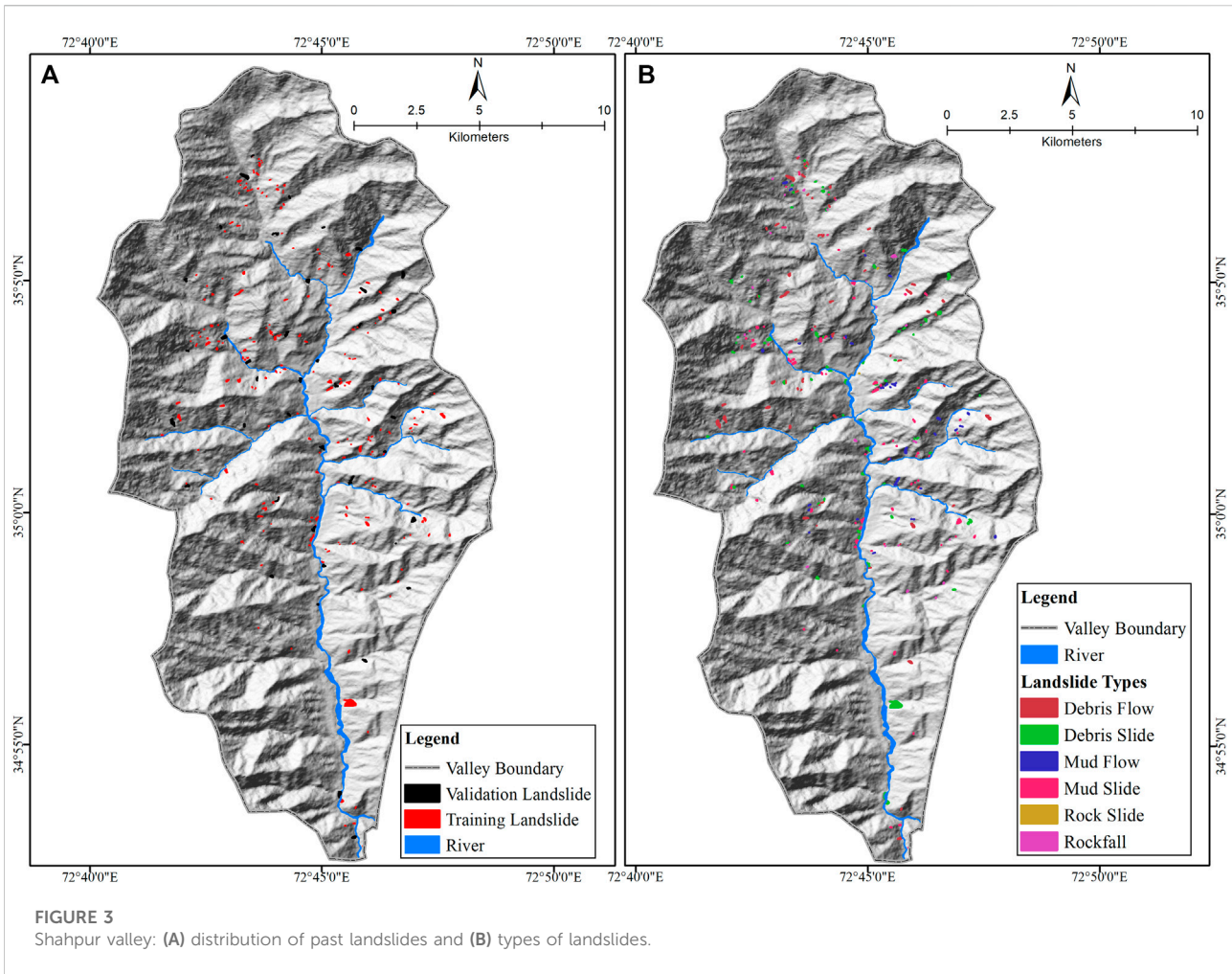


FIGURE 3 Shahpur valley: (A) distribution of past landslides and (B) types of landslides.

Weight of evidence model

Bonham-Carter (1989) applied the weight of evidence model for the first time using Eq. 1 and Eq. 2:

$$W^+ = \ln \frac{P(\frac{B}{D})}{P(\frac{\bar{B}}{\bar{D}})} \tag{1}$$

$$W^- = \ln \frac{P(\frac{\bar{B}}{\bar{D}})}{P(\frac{B}{D})} \tag{2}$$

In the abovementioned equations, the probability is denoted by P , while \ln denotes natural log. B and \bar{B} , respectively, represent the existence or non-existence of potential landslide evidence factors. Likewise, D and \bar{D} denote landslide presence and absence, respectively. Eqs 1 and 2 were modified to calculate the weight of the classes of each causal factor contributing to landslide occurrence and statistically derived using Eq. 3 and Eq. 4.

$$W^+ = \ln \left\{ \left(\frac{[Npix1]}{[Npix1] + [Npix2]} \right) / \left(\frac{[Npix3]}{[Npix3] + [Npix4]} \right) \right\} \tag{3}$$

$$W^- = \ln \left\{ \left(\frac{[Npix3]}{[Npix1] + [Npix2]} \right) / \left(\frac{[Npix4]}{[Npix3] + [Npix4]} \right) \right\} \tag{4}$$

$Npix1$ are the pixels that express the existence of landslides and contributing factors; $Npix2$ denotes the existence of landslides and absence of contributing factors, while $Npix3$ denotes the existence of contributing factors and non-existence of landslides. Similarly, $Npix4$ shows the pixels where both landslides and landslide contributing factors are not found. W^c is the final weight which is calculated using Eq. 5:

$$W^c = (W^+) - (W^-) \tag{5}$$

W^c represents the spatial association of landslides and contributing factors.

Information value (IV) model

Measuring the effect of landslide contributing factors on landslide occurrence was also examined using the IV model. The information value model is based on Bayes' theorem, and information values are derived from each class of contributing factors. The IV model calculates landslide susceptibility based on past landslide events and its correlation with landslide contributing factors. The positive information value of a class shows a stronger correlation with landslide occurrence, while the negative value denotes a negative correlation of that class with landslides. In the IV model, the statistical values are calculated for all classes of each contributing factor using Eq. 6:

$$IV = \log \left(\frac{N_{pix(Si)} / N_{pix(Ni)}}{\sum N_{pix(Si)} / \sum N_{pix(Ni)}} \right), \quad (6)$$

where $N_{pix(Si)}$ denotes the landslide pixels in class i , whereas $N_{pix(Ni)}$ is the total pixels of class i . Similarly, $\sum N_{pix(Si)}$ is the landslide pixels in the entire study region, whereas $\sum N_{pix(Ni)}$ denotes the total pixels in the study region.

Landslide susceptibility index (LSI)

The LSI for both IV and WoE models was calculated using the sum of the landslide causal/contributing factors in the GIS based on the W^c and IV values for overlay analysis in ArcGIS using the following equations:

$$LSI_{WoE} = \sum W^c, \quad (7)$$

$$LSI_{IV} = \sum IV, \quad (8)$$

where $\sum W^c$ denotes the total derived weight of WoE and $\sum IV$ is the derived weight of the information value model.

Validation techniques

The success rate curve, prediction rate curve, R-index, and seed cell area index (SCAI) were used to validate the LSI maps prepared based on WoE and IV models. The success rate curve validation is based on training landslides which is 75% of the total landslides, while the prediction rate curve is based on validation landslides which is 25% of the total landslides. Both curves elucidate the accuracy of WoE and IV for selected causal factors to landslide occurrences. The validation curves for WoE and IV were calculated using the LSI values overlaid

with the existing layer of landslide area using the geo-statistical tool in the GIS. Cumulative percentages for both susceptibility class and landslide area were calculated, and susceptibility class was plotted on the x -axis and landslide area on the y -axis to generate both success and prediction rate curves. The landslide model performance was assessed using specificity, sensitivity, and accuracy. Sensitivity is the proportion of pixels that have been accurately classified as occurrences of landslides, whereas the proportion of pixels of non-landslide pixels that are accurately classified as non-landslides is the specificity (Abedi Gheshlaghi and Feizizadeh, 2021) and accuracy is the proportion of landslides and non-landslide pixels that are accurately classified (Chen et al., 2018).

$$Sensitivity = \frac{TP}{TN + FN}. \quad (9)$$

$$Specificity = \frac{TN}{TN + FP}. \quad (10)$$

$$Accuracy = \frac{TP + TN}{TP + TN + FP + FN}. \quad (11)$$

In the equations, TN is the true-negative, TP is the true-positive, FN is the false-negative, and FP is the false-positive (Abedi Gheshlaghi and Feizizadeh, 2021).

The SCAI technique was applied on both WoE and IV models as first tested by Szen and Doyuran (2004). SCAI is the ratio of the percentage of the pixels of a landslide susceptibility zone to the percentage of existing landslide pixels in that susceptibility zone. The model is considered excellent if the value of the SCAI decreases from a very low to a very high LS class (Arabameri et al., 2020; Pawluszek-Filipiak et al., 2020). The SCAI is calculated using the following equation:

$$SCAI = \frac{Ni\%}{ni\%}. \quad (12)$$

In addition, the relative landslide density (R-index) was also used to validate the LSI maps. R-index is calculated using the following equation:

$$R - Index = \left(\frac{ni}{Ni} / \sum \frac{ni}{Ni} \right) \times 100, \quad (13)$$

where ni is the number of landslides in a LS zone and Ni is the total number of pixels in a LS zone.

Landslide vulnerability assessment

Vulnerability is the essential element of risk assessment which refers to the potential degree of loss to the element at risk (Ram and Gupta, 2022). According to the definition of UNISDR (2009), vulnerability is "the characteristics and circumstances of a community, system, or asset that make it susceptible to the damaging effects of a hazard." The calculation of social, economic, and environmental vulnerability is still a challenge in landslide studies

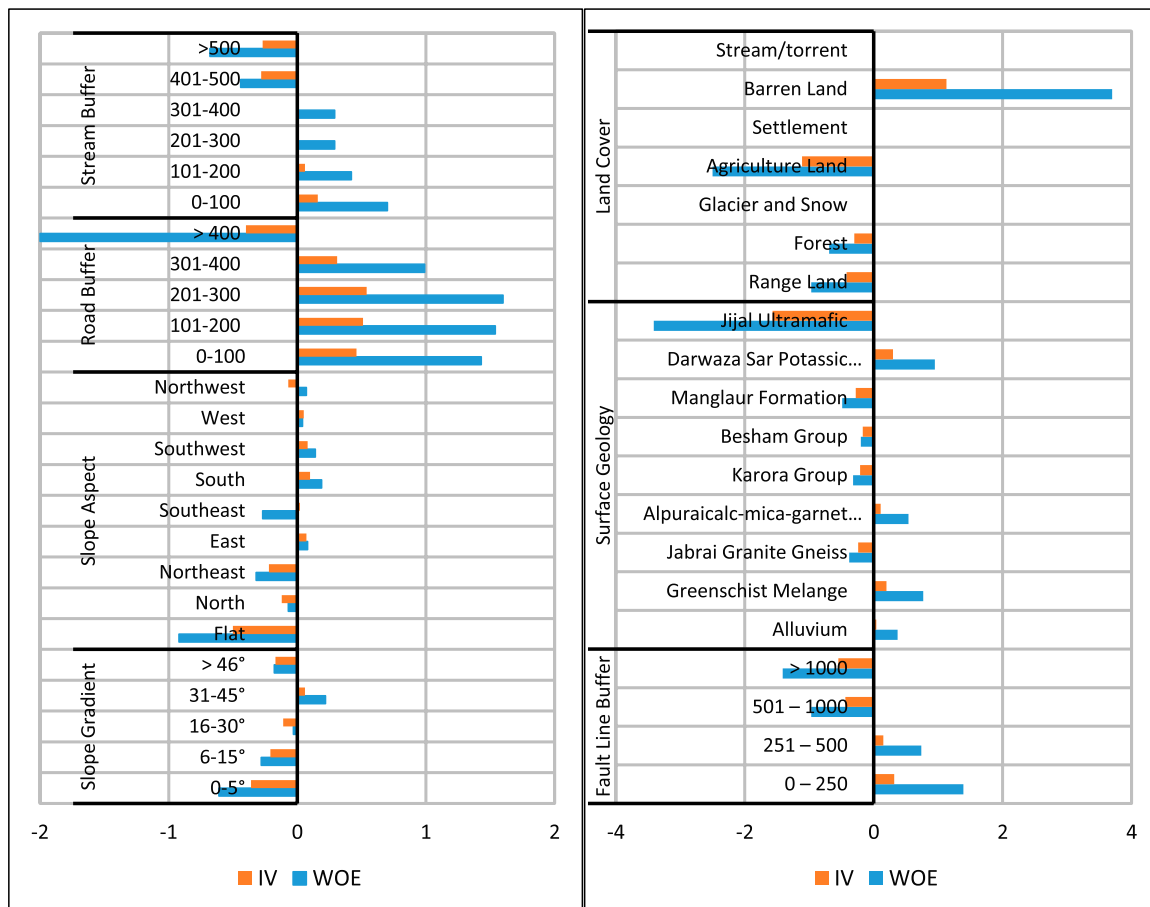


FIGURE 4
Shahpur valley, the calculated weight of each class of causal factors.

(Fu et al., 2020); therefore, the physical vulnerability was assessed based on elements at risk exposure level and the probability of loss due to landslide. Based on indigenous knowledge and expert opinions, the agricultural land, settlement, rangeland, forest, schools, roads, and bridges were selected as elements at risk for vulnerability assessment. Based on the element at risk, the exposure maps were prepared using the landslide susceptibility map to measure the level of exposure of these elements at risk. These exposure maps were used to calculate to prepare the vulnerability map of the area using a geo-statistical approach based on Eq. 14.

$$V = P(D), \tag{14}$$

where P is the probability of loss due to landslides and D is the element at risk. From the knowledge gained during fieldwork, the value of each element at risk was then used for calculation of the probability of loss which is expressed on a scale of 0 to 1, where zero means no loss while 1 is the total loss. The final map was reclassified into five zones ranging from very low- to very high-vulnerability zones.

Landslide risk assessment

The function of hazard (likelihood of landslide occurrence) and damage potential is called risk (Anbalagan and Singh, 1996; Sujatha and Rajamanickam, 2015). Risk is defined as the potential for adverse consequences or the chances of loss (Lee and Jones, 2004). UNISDR (2009) defines risk as “the combination of the probability of an event and its negative consequences.” In this study, risk assessment is calculated based on damage potential to the elements at risk which we have named vulnerability (Ram and Gupta, 2022).

In this study, landslide risk is calculated using the following equation:

$$R = H \times V, \tag{15}$$

where R is the potential risk, H is the hazard probability (landslide susceptibility), and V is the calculated vulnerability (damage potential to the elements at risk). In the study region, the unavailability of temporal records of landslides is the limitation of the study; therefore, we have used the likelihood

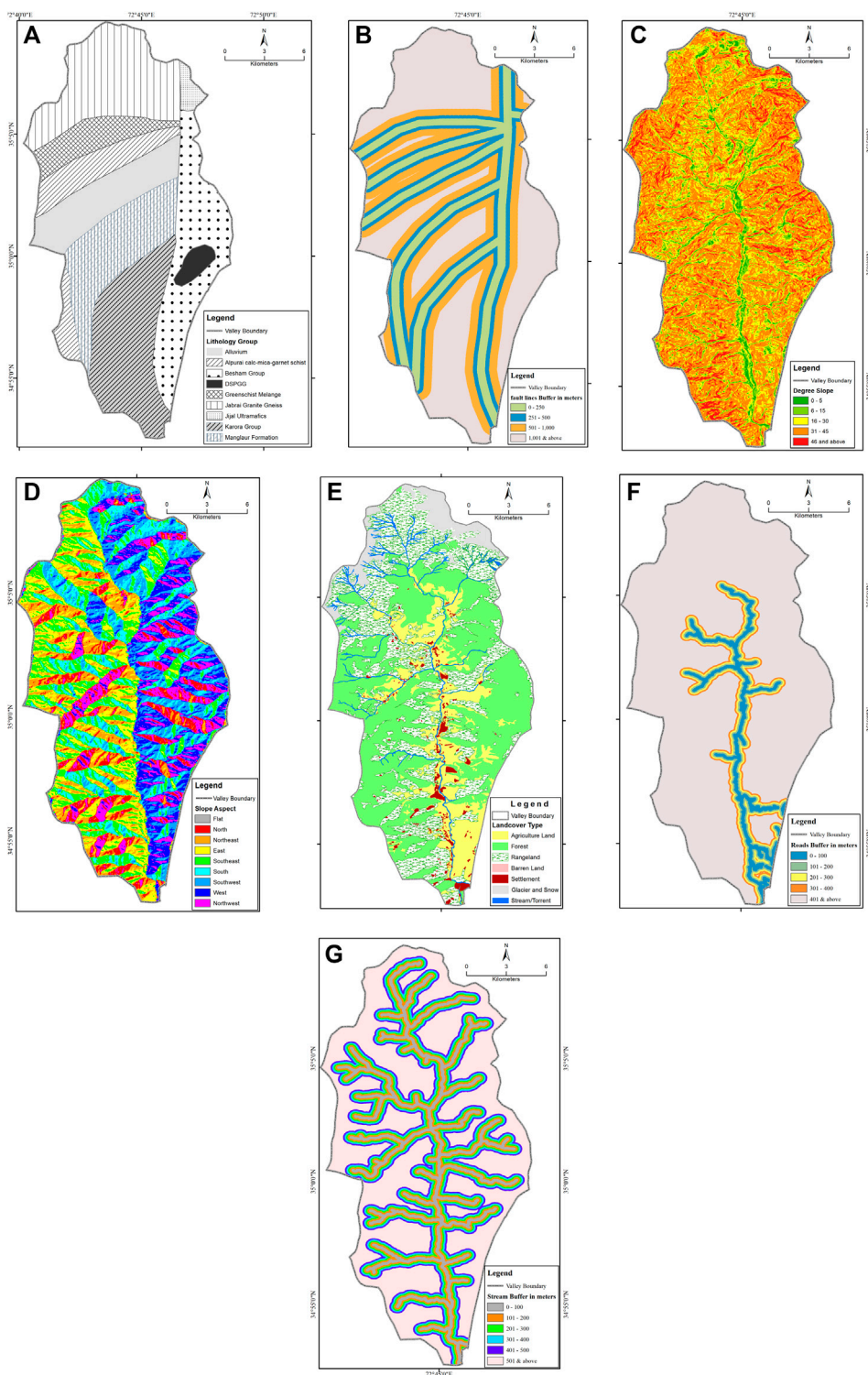


FIGURE 5
 Shahpur valley: (A) surface geology; (B) proximity to fault lines; (C) slope gradient; (D) slope aspect; (E) land use/land cover; (F) proximity to road; (G) proximity to stream/torrent (Rahman et al., 2017).

TABLE 2 Shahpur valley: the calculated weight of each class of causal factors.

Class	Npix (Si)	% age of Npix (Si)	Npix (Ni)	% age of Npix (Ni)	W ⁺	W ⁻	W ^c	IV (W)
Surface geology								
Alluvium	995	15.43	290137	11.20	0.32	-0.05	0.37	0.04
Greenschist melange	826	12.81	165892	6.40	0.70	-0.07	0.77	0.20
Jabrai granite gneiss	903	14.00	497979	19.22	-0.32	0.06	-0.38	-0.24
Alpuraicalc-mica-garnet schist	945	14.65	235014	9.07	0.48	-0.06	0.54	0.11
Karora group	957	14.84	501955	19.37	-0.27	0.05	-0.32	-0.21
Besham group	933	14.47	441986	17.06	-0.17	0.03	-0.20	-0.17
Manglaur formation	615	9.53	378895	14.62	-0.43	0.06	-0.49	-0.28
Darwaza Sar potassic granite gneiss	273	4.23	43693	1.69	0.92	-0.03	0.95	0.30
Jijal ultramafic	3	0.05	35939	1.39	-3.40	0.01	-3.41	-1.57
Fault line buffer (m)								
0–250	2933	45.47	448304	17.30	0.97	-0.42	1.39	0.32
251–500	1820	28.22	409420	15.80	0.58	-0.16	0.74	0.15
501–1000	763	11.83	676634	26.11	-0.79	0.18	-0.97	-0.44
>1000	934	14.48	1057133	40.79	-1.04	0.37	-1.41	-0.55
Slope gradient								
0–5°	93	1.44	67722	2.61	-0.60	0.01	-0.61	-0.36
6–15°	506	7.84	261492	10.09	-0.25	0.02	-0.28	-0.21
16–30°	1624	25.18	668931	25.81	-0.02	0.01	-0.03	-0.11
31–45°	3747	58.09	1366442	52.73	0.10	-0.12	0.22	0.06
>46°	480	7.44	226903	8.76	-0.16	0.01	-0.18	-0.17
Slope aspect								
Flat	1	0.02	1004	0.04	-0.92	0.00	-0.92	-0.50
North	503	7.80	214667	8.28	-0.06	0.01	-0.07	-0.12
Northeast	531	8.23	284530	10.98	-0.29	0.03	-0.32	-0.22
East	1029	15.95	387999	14.97	0.06	-0.01	0.08	0.07
Southeast	781	12.11	395492	15.26	-0.23	0.04	-0.27	0.02
South	1075	16.67	366954	14.16	0.16	-0.03	0.19	0.10
Southwest	1003	15.55	356943	13.77	0.12	-0.02	0.14	0.08
West	819	12.70	317383	12.25	0.04	-0.01	0.04	0.05
Northwest	708	10.98	266520	10.28	0.07	-0.01	0.07	-0.07
Land cover								
Range land	1007	15.61	847632	32.71	-0.74	0.23	-0.97	-0.42
Forest	1621	25.13	1036194	39.98	-0.47	0.22	-0.69	-0.30
Glacier and snow	0	0.00	111086	4.29	0.00	0.00	0.00	0.00
Agriculture land	100	1.55	416925	16.09	-2.34	0.16	-2.50	-1.11
Settlement	0	0.00	37521	1.45	0.00	0.00	0.00	0.00
Barren land	3722	57.71	87880	3.39	2.87	-0.83	3.70	1.13
Stream/torrent	0	0.00	54252	2.09	0.00	0.02	0.00	0.00

(Continued on following page)

TABLE 2 (Continued) Shahpur valley: the calculated weight of each class of causal factors.

Class	Npix (Si)	% age of Npix (Si)	Npix (Ni)	% age of Npix (Ni)	W^+	W^-	W^c	IV (W)
Road buffer (m)								
0–100	1169	18.12	130869	5.05	1.28	–0.15	1.43	0.46
101–200	1041	16.14	103117	3.98	1.41	–0.14	1.54	0.51
201–300	991	15.36	92441	3.57	1.47	–0.13	1.60	0.54
301–400	541	8.39	85731	3.31	0.93	–0.05	0.99	0.31
>400	2708	41.98	2179333	84.10	–0.70	1.30	–2.00	–0.40
Stream buffer (m)								
0–100	1318	20.43	294902	11.38	0.59	–0.11	0.70	0.16
101–200	955	14.81	265711	10.25	0.37	–0.05	0.42	0.06
201–300	821	12.73	255277	9.85	0.26	–0.03	0.29	0.01
301–400	799	12.39	247979	9.57	0.26	–0.03	0.29	0.01
401–500	395	6.12	238952	9.22	–0.41	0.03	–0.44	–0.28
>500	2162	33.52	1288669	49.73	–0.40	0.28	–0.68	–0.27

of landslide occurrence instead of landslide hazard which might possess some degree of uncertainty in risk assessment.

Results

Landslide causative factors and landslide occurrence

Surface geology

The relationship between surface geology and landslide occurrence was assessed using WoE and IV models. The types of surface geology are shown in Figure 4A. The highest positive W^c weight was found in Darwaza Sar Potassic Granite Gneiss (0.95) and Greenschist Melange (0.77) (Figure 5 and Table 2). Both classes have a very positive correlation with landslides as assessed in WoE model results. In the study area, Darwaza Sar Potassic Gneiss and Greenschist Melange are metamorphic rocks located in a seismo-tectonically active zone and therefore bear a high potential for slope failure. Similar results were found in the IV model results. In both WoE and IV models, the highest negative correlation was found in Jijal Ultramafics, where the W^c value was –3.41, while the IV(W) was –1.57 (Table 2). Ultramafic are igneous rocks having low silica content and heavily comprise mafic minerals and therefore resist weathering and registered a highly negative correlation to landslide occurrence in the results of both the models.

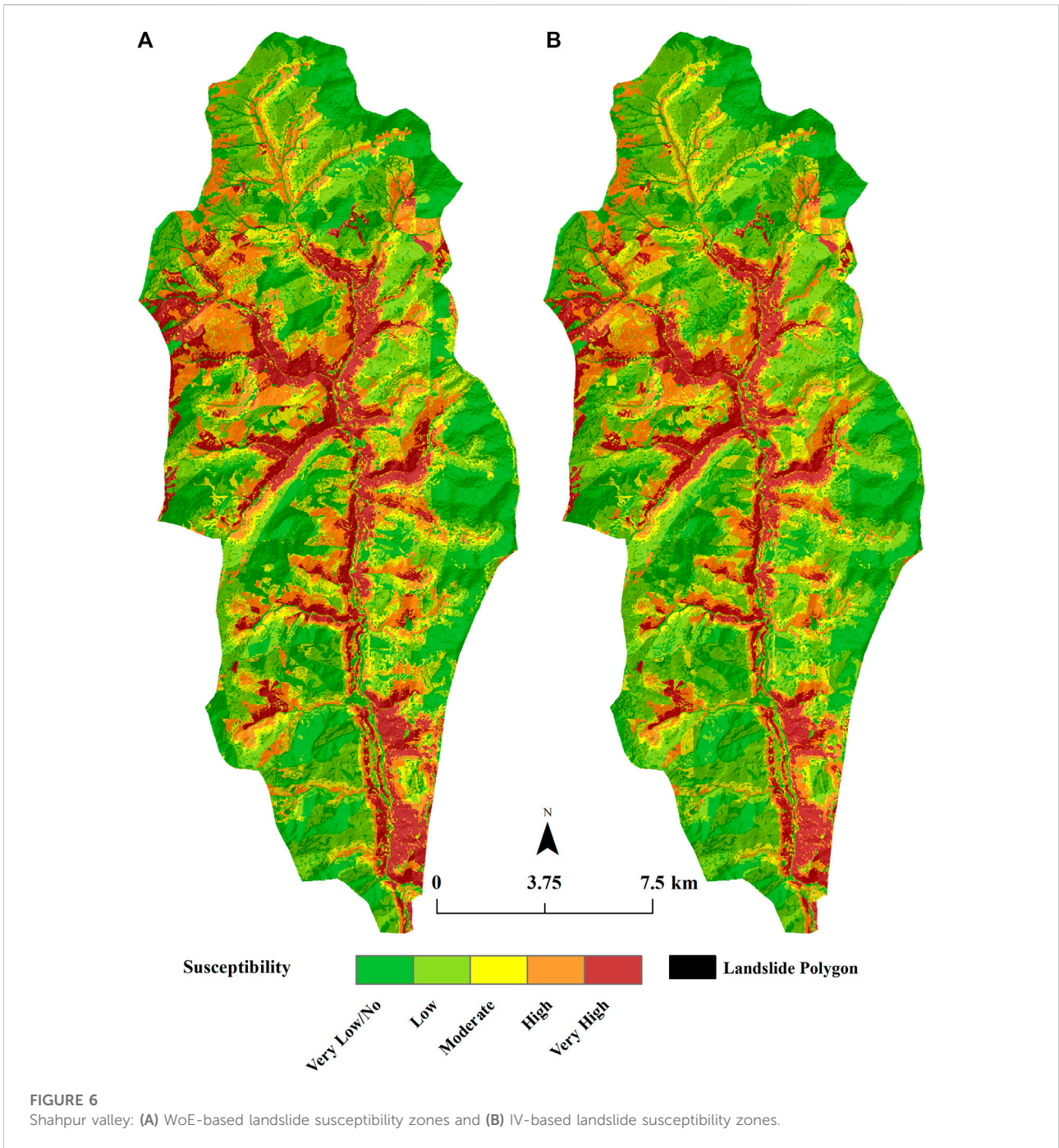
Proximity to fault line

The occurrence of landslides has a strong correlation with fault lines (Korup, 2004; Rahman et al., 2019). The existence of

fault lines at high slope gradients provides favorable settings for the occurrence of slope failure. There is a complex tectonic structure in the study area which is considered a causal factor for slope instability. It is evident from the analysis that the tectonic structures have a major role in landslide occurrence. The highest positive W^c value (1.39) was in the buffer zone of 0–250 m followed by the 251–500 m buffer zone according to the WoE model. Similarly, the IV model also showed the highest landslide probability in the buffer zone 0–250 m followed by the 251–500 m buffer zone. The buffer zones 501–1,000 and >1,000 m showed a negative correlation with landslide occurrence in both WoE and IV models (Figure 5 and Table 2). Thus, it is evident from the results of both models that landslides occur in the vicinity of fault lines.

Slope gradient

The slope gradient affects the population distribution, their activities, and natural resource distribution. Likewise, slope gradient has a close association with landslide distribution, and it acts as a governing factor in slope stability as the chances of landslide incidence escalate with the increase in slope gradient. It was observed during field visits that a high number of landslides occurred on the slope along the road and stream where the lateral cutting was a dominant factor. The map of the slope gradient for the study area was generated in the GIS from ASTER-GDEM having a 30 m spatial resolution (Figure 4C). The analysis of both WoE and IV shows that the slope failure is higher in 31–45° slopes as the highest W^c value (0.22) and IV value (0.06) was found in this class of slope gradient (Figure 5 and Table 2), while the slope gradient 0–5° and 6–15°



class showed the highest negative correlation with landslides in both model results as both the slope ranges, to some extent, consist of flat land where a low possibility of slope failure is expected.

Slope aspect

Slope direction governed sunlight intensity and duration, determining the amount of rainfall, moisture-holding capacity of

land, and distribution of vegetation. The analysis revealed that the south-facing slope has the highest value of W^c (0.19) and IV (0.10) followed by the southwest facing slope W^c (0.14) and IV (0.08) as shown in Figure 5 and Table 2. The high value of W^c and IV in south-facing slopes denotes that there is a high number of landslides that occurs, which might be because of exposure to more sunlight and ample amount of rainfall. High amounts and more intense rainfall in the summer monsoon causes more rain-

TABLE 3 Areal coverage of landslide susceptibility zones.

Susceptibility class	WoE		IV	
	Area (Km ²)	Percentage	Area (Km ²)	Percentage
Very low/no	80.53	31.21	71.35	27.65
Low	61.41	23.80	85.77	33.24
Moderate	32.65	12.66	30.16	11.69
High	46.03	17.84	37.58	14.57
Very high	37.38	14.49	33.14	12.84

induced land sliding. Similarly, during the winter season, the topsoil and porous rocks absorb more water from the accumulated snow on these slopes. Therefore, it often accelerates the saturation of ground and thus causes landsliding. This may be attributed to the low vegetation on south and southwest-facing slopes as compared to north and northeastern-facing slopes where vegetation cover was found to be high during field visits.

Land use/land cover

The forest cover protects the mountainous slope from weathering and mass wasting processes as the roots hold the underneath soil and keep the slope stable. Increasing population growth has increased the demand for wood and land for food, which has disturbed the slope of almost all the mountainous regions of the world and has led to slope instability. The land use map of the study area shows seven different land cover types, that is, rangeland, forest, glacier and snow, agricultural land, settlement, barren land, and stream/torrent (Figure 4E). The statistical weight for land use classes was derived using WoE and IV models. The highest weight of both WoE ($W^c = 3.70$) and IV (1.13) was found in the barren land class (Figure 5 and Table 2). As barren land is exposed to a variety of erosional and mass wasting processes, it is highly prone to the likelihood of landslide occurrence. The highest negative landslide correlation was found in the agricultural land where the WoE value was -2.50 and the IV value was -1.11 ; although in the study area, terrace agriculture is practised, the local community has constructed protection walls around the field.

Proximity to road

Road constructions often disturb the slopes and expedite the weathering and mass wasting processes, thus increasing the probability of landslide occurrences. It also provides means of accessibility and accelerates the process of deforestation in mountainous regions. In the current study, proximity to the road is used as a contributing factor of landslide hazard. The results show a high positive correlation with proximity to the road, especially the distance from the road up to 300 m. The highest W^c value (1.60) and IV (0.54) were found in 201–300 m of road proximity (Figure 5 and Table 2). Similarly, the area that ranges from 101 to 200 m away from the roads has the second highest positive correlation toward landslide probability, while the areas >400 m away from the roads recorded the highest negative correlation with landslide occurrence. The construction of roads in mountainous areas usually destabilizes the nearby fragile slope (Khan et al., 2011). Often, shallow landslides occur along the roads mainly in rainy seasons, and these landslides are initiated due to edge cutting and the use of explosive material during road construction on these slopes (Pourghasemi et al., 2020). Therefore, the density of landslides was found to be high

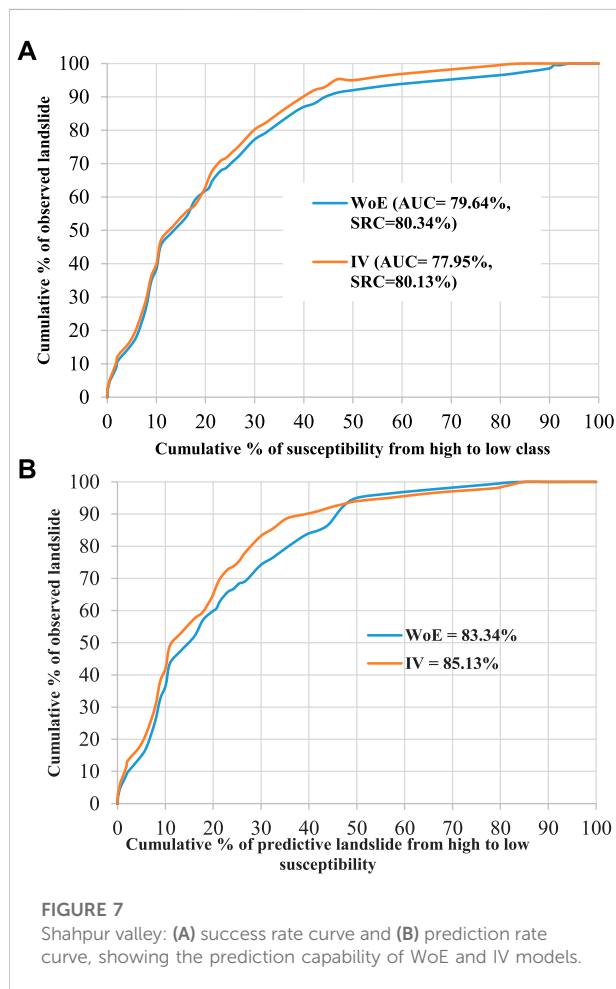


TABLE 4 Model performance using the training dataset.

Evaluation parameter	Model	
	WoE	IV
TP	114	118
TN	102	103
FP	24	25
FN	18	12
Sensitivity	86.36%	90.77%
Specificity	80.95%	80.47%
Accuracy	83.72%	85.66%

near the roads, and the density of landslide decreases as the distance from roads increases.

Proximity to stream/torrent

The relationship of stream/torrent with landslide was also examined through WoE and IV statistical models. It was found from the analysis that both WoE and IV have higher values near the stream, which indicates high landslide probability in this region. The highest W^c (0.70) and IV values (0.16) were found in the proximity of 0–100 m (Figure 5 and Table 2). The results show that the region up to 400 m in proximity to the stream recorded a positive correlation toward landslide probability. The highest negative correlation was found in the buffer zone greater than 500 m of stream in both models' results. Steep slopes along the stream and consistent lateral erosion destabilize the slopes and cause landslides (Shano et al., 2021). In the study area, the stream undercut the banks that result in sloughing and stimulate slope failures. Therefore, the highest values in W^c and IV were found in the proximity of the streams.

Landslide susceptibility zonation

Landslide is a common menace to property, human lives, and infrastructure in Shahpur valley. For its mitigation, the first important step is to identify susceptible landslide areas. In the LSZ map, all the landslide causative factors were integrated into the GIS, and the region was divided into very low- to very high-susceptible zones based on their susceptibility level. To minimize the subjectivity, quantitative weight to each class of factor maps was applied based on WoE and IV models for the preparation of the LSZ map of Shahpur valley. The LSZ map was created based on both WoE and IV models by integrating all the relative weights of all classes of factor maps.

$\sum W^c$ denotes the sum of the W^c of all factor maps for the WoE model, while $\sum IV$ is the sum of the IV model value of all factor maps. In both models, the higher LSI means a high probability of landslide incidents. Based on LSI, Figure 6

TABLE 5 Model performance using a validation dataset.

Evaluation parameter	Model	
	WoE	IV
TP	46	45
TN	39	41
FP	7	8
FN	6	4
Sensitivity	88.46%	91.84%
Specificity	84.78%	83.67%
Accuracy	86.73%	87.76%

shows the landslide susceptibility level in the study area. The WoE model predicted more areas as high (17.84%) and very high (14.49%) landslide susceptible to landslide hazard as compared to the IV model where the high landslide susceptible class covers 14.57 and 12.84% of the area in very high susceptibility zones (Table 3). The landslide spatial distribution and the derived susceptibility maps show that the selected causative factors are relevant, and both the models accurately depicted the landslide-susceptible areas. The existing landslide distribution in different landslide susceptibility zones in Figure 6 shows more coherent results in the WoE-derived susceptibility map as compared to the landslide susceptibility map derived through the IV model (Figure 6).

Validation of the landslide susceptibility map

The result of both success and prediction rate curves has a steep slope, which indicates significantly good results for both WoE and IV models. The model validation from the success rate curve shows 80.34% accuracy for WoE and 80.13% for the IV model (Figure 7A). The area under the curve (AUC) was calculated using a success rate curve based on the training dataset. An AUC of 79.64% was recorded for WoE and 77.95% for the IV model (Figure 7A). The model's prediction rate curve was also calculated based on the validation dataset and shown in Figure 6B. The calculated prediction rate for both WoE and IV was 83.34 and 85.13%, respectively.

Based on the training dataset, the performance of the models, that is, WoE and IV models has been assessed as shown in Table 4. Both the models showed very good performance in the classification of landslide pixels where the sensitivity for the WoE was 86.36%, while it was 90.77% for the IV model (Table 4). Similarly, the models showed the highest performance in the classification of non-landslide pixels, where the WoE model recorded 80.95% specificity, whereas the specificity of the IV model was 80.47%. The highest accuracy is for the IV model

TABLE 6 WoE and IV model result validation based on seed cell area index (SCAI).

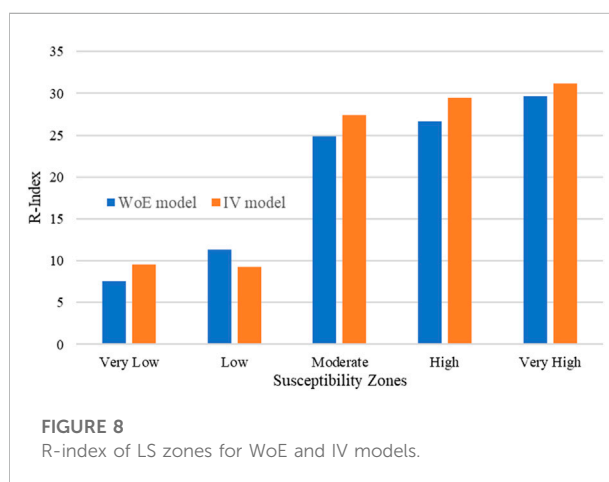
Model	Susceptibility zone	No. of pixels in the susceptibility zone	%	No. of landslide pixel	%	SCAI
WoE	Very Low	515414	31.20	754	15.60	2.00
	Low	393019	23.80	630	13.00	1.82
	Moderate	208960	12.70	746	15.50	0.82
	High	294565	17.80	1070	22.20	0.80
	Very High	239247	14.50	1628	33.70	0.43
IV	Very Low	456638	27.65	669	13.86	2.00
	Low	548936	33.24	1046	21.67	1.53
	Moderate	193009	11.69	713	14.77	0.79
	High	240528	14.57	943	19.53	0.75
	Very High	212094	12.84	1457	30.18	0.43

(85.66%), whereas the accuracy for the training dataset of the WoE model is 83.72% (Table 4).

Likewise, in the training dataset, the performance of the models was assessed using the validation dataset. The WoE and IV models showed excellent performance in the classification of landslide pixels. The sensitivity results for the WoE model are 88.46%, whereas it is 91.84% for the IV model (Table 5). Similarly, both the models showed the highest performance for non-landslide pixels classification, which recorded 84.78% specificity for WoE and 83.67% for the IV model. The accuracy of the WoE model is 86.73%, while the IV model's accuracy is 87.76% (Table 5).

The results of the SCAI validation index are shown in Table 6. It is evident from the previous literature that the SCAI value should be lower in the higher LS zone and higher in the lower LS zone, whereas the greater difference in SCAI values of both LS zones (lower and higher LS zone) indicates a good fit of the model (Arabameri et al., 2020; Saha and Saha, 2020). In the current study, the SCAI results for very high and low landslide susceptibility classes in both WoE and IV models show the same value range where the very low LS zone has 2 and the very high LS zone has 0.43 SCAI value (Table 6). This means that both models have almost similar performance for landslide susceptibility mapping in the Hindu Kush region.

In this research article, the susceptibility model results were also validated through the R-index which is considered one of the important validation techniques in landslide studies. The result shows that the R-index value for high and very high LS zones was 26.64 and 29.64%, respectively, in the WoE model, and it was 29.48 and 31.21% for the IV model, respectively (Figure 8). Both landslide models have high accuracy and are suitable for similar studies in the Hindu Kush region; however, based on the validation



results of all mentioned techniques, it was elucidated that the WoE model gives slightly better results than the IV model.

Elements at risk and landslide vulnerability

In this study, elements at risk were assessed based on indigenous knowledge and expert opinions, and these included agriculture land, range land, settlements, road network, bridges, and schools. The elements at risk are studied in relation to landslide hazard (susceptibility) zones, and their degree of exposure is explored through the GIS. Each element at risk that falls in the high-susceptibility zone was categorized into the highly exposed group, and similarly, those lying in the low-susceptibility zone are categorized as low exposure to landslide as one of the core characteristics of

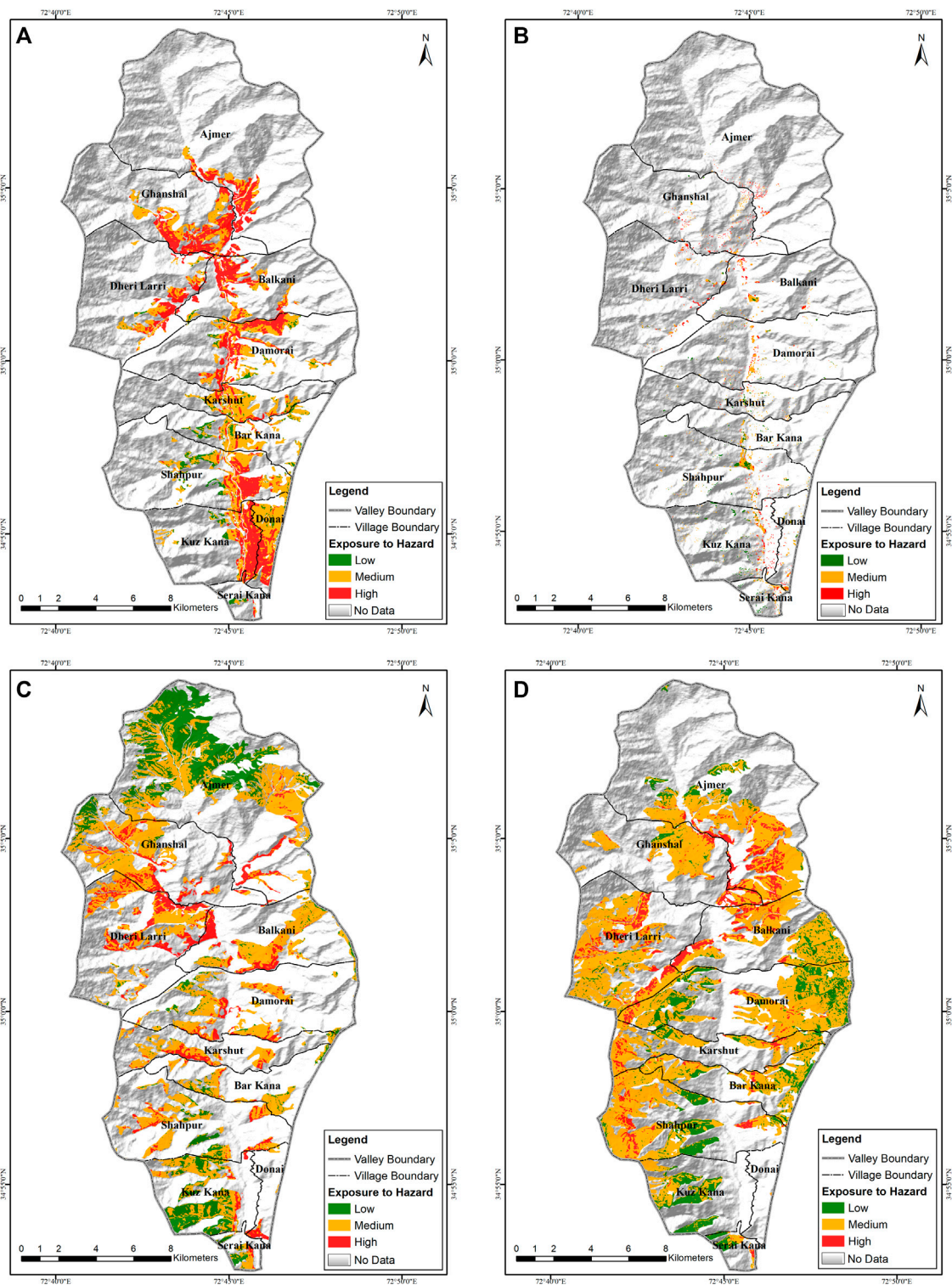
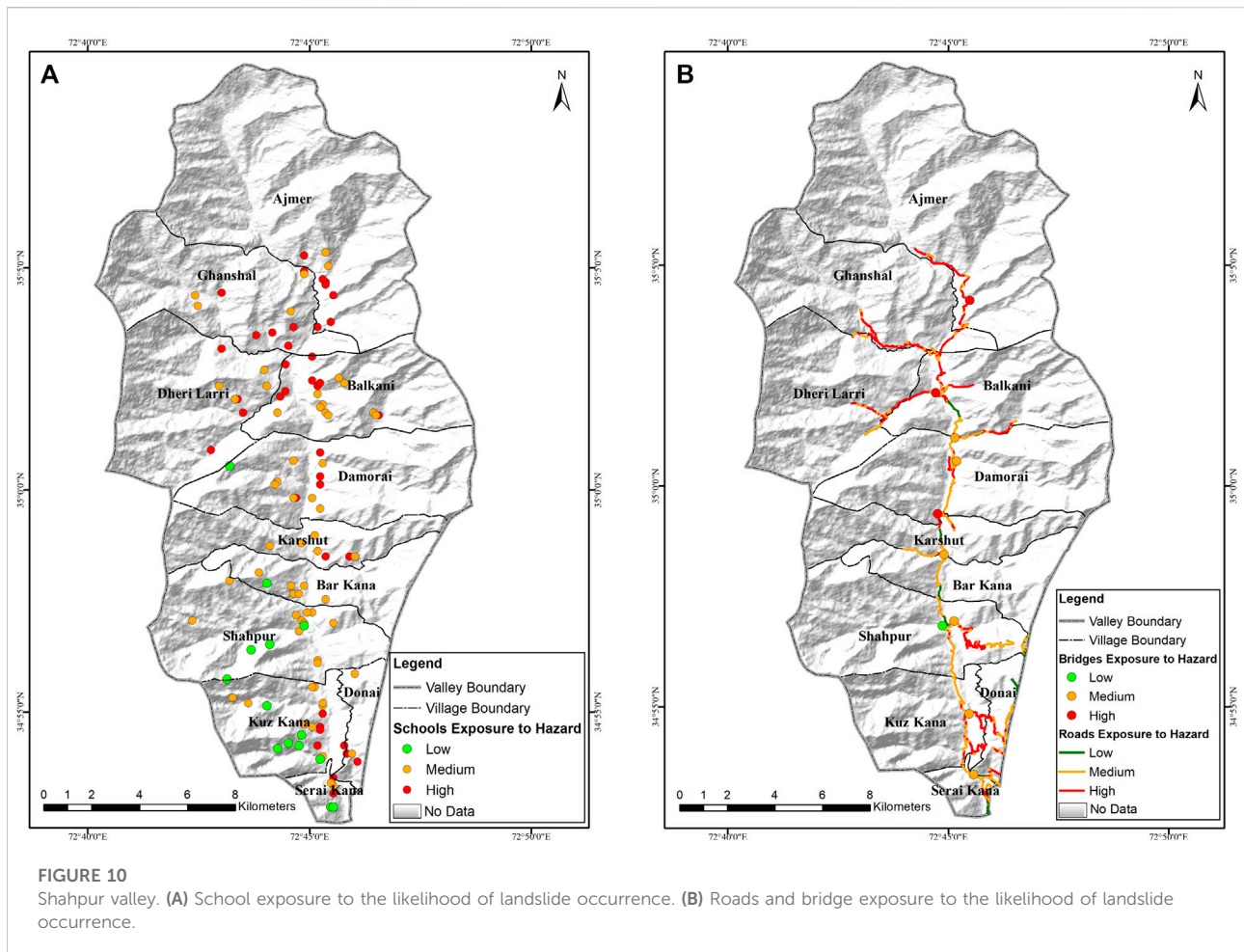


FIGURE 9 Shahpur valley. (A) Agriculture land exposure to the likelihood of landslide occurrence. (B) Settlement exposure to the likelihood of landslide occurrence. (C) Rangeland exposures to the likelihood of landslide occurrence. (D) Forest cover exposures to the likelihood of landslide occurrence.



vulnerability is the ability to see the conditions of exposed elements or communities at risk.

In Shahpur valley, agriculture is the mainstay of the economy, a source of foodstuff, and a major economic activity in the low-lying areas. In Shahpur valley, agriculture land occupied 42.22 sq km, which is 16.3% of the total study area. The local people practice agriculture in the low-lying areas, specifically in the vicinity of rivers or streams, and mostly gentle slopes are dominated by terrace farming. It was found from the results that 37.6% of the total agriculture land is highly exposed to the likelihood of landslide occurrence (Figure 9A). Landslide is a major threat to the agriculture sector and to the people engaged in agriculture activities. About 52.2% of agriculture land is moderately exposed, and only 10.2% of agriculture land is in the category of low exposure to the likelihood of landslide occurrence.

In the Hindu Kush region, the pace of population growth is higher than the national average (Rahman and Shaw, 2015). This tremendous population growth has increased pressure on the available natural and land resources. With the population growth, demand for food, shelter, communication network, and other

requirements also increases. In the areas like Shahpur valley, the people have no choice other than engulfing forest and rangeland to construct shelter for their living and clear the area for agriculture to gain food self-sufficiency. Such situations have accelerated the likelihood of landslide occurrence and vulnerability in the whole Shahpur valley. Settlements taken into consideration as elements at risk include housing units and commercial areas such as shops. People living in these settlements and their vulnerability depend on the location of settlements: either they are situated in hazard-free zone or in hazard-prone areas. Based on the analysis results, about 18.3% of the area under settlements is highly exposed to the likelihood of landslide occurrence, and these settlements are highly vulnerable to any unusual landslide event (Figure 9B). About 66.2% of settlements are moderately exposed to the likelihood of landslide occurrence and have comparatively low vulnerability, while 15.5% of settlements have low exposure to the likelihood of landslide occurrence.

In Shahpur valley, rangelands are present at higher altitudes and are mainly dominated by grasses, shrubs, and bushes. Such lands are very suitable for watershed management and mainly utilized for

grazing and livestock ranching. These rangelands are the main pastureland in Shahpur valley. The vegetation cover in these rangelands protects soil erosion and discourages landslide phenomena. In Shahpur valley, out of the total rangeland area, 13.6% was found highly exposed to the likelihood of landslide occurrence and 25.67% area has low exposure, while the rangeland moderately exposed to the likelihood of landslide occurrence is 60.73% of the total area under rangeland (Figure 9C).

Forest is an important resource for the local people and contributes equally to the country's economy. Forest provides timber for house construction and wood for cooking and heating in winter to protect from harsh cold winter. Forest cover is a bioengineering solution for landslide hazards and soil erosion. Forest helps in binding soil and protecting it from slope failure. The study area has thick vegetation cover, particularly over the higher altitudes and the slopes which receive an ample amount of precipitation. At higher altitudes, the coniferous forest is predominant. In landslide vulnerability analysis, it is also an important indicator like agriculture land and human settlement, but it works as a landslide mitigation strategy and its importance cannot be denied. About 10% of the area under forest cover is highly exposed to the likelihood of landslide occurrence (Figure 9D). Mostly, the highly exposed areas are on the slope which is destabilized either by a natural intervention such as earthquake, geological structure, and lateral erosion by rivers and streams during floods or by human intervention such as deforestation, clearing area for settlement construction and agricultural activities, road construction, and quarrying. Forest cover is basically a capacity to cope with the landslide hazard phenomena, but unfortunately, due to the lack of forest conservation, this precious resource is vanishing and, as a result, vulnerability to landslide hazard is increasing.

Likewise, agriculture, settlements, forest and rangeland, and the infrastructure including schools, bridges, and roads were also analyzed for exposure to landslide hazard. Road and bridge data were obtained from the Communication and Works Department, Khyber Pakhtunkhwa, while the school data were obtained from the Elementary and Secondary Education Department. The school data consist of primary, middle, and high schools for both boys and girls. In Shahpur Valley, there are a total of 120 schools. Out of total schools, 42 were found in very high and high hazard zones, which make up 35% of the total schools and have been categorized as highly exposed to landslide hazard (Figure 10A).

In Shahpur valley, there is a total of 75 km road network in terms of length, out of which 38 km falls in highly exposed landslide hazard class (Figure 10B). Most of the highly exposed road section is in the upper part of the valley. Roads in Shahpur valley mostly follow the riverside, so the river lateral erosion and road construction itself pose a serious threat, whereas human intervention over the natural slope increases the vulnerability to landslide hazards. About 25 km road is moderately exposed to landslide hazard. Likewise, bridges were also analyzed for landslide exposure. According to the Communication and Works Department, there are 10 bridges in the valley, out of which three bridges are highly exposed to landslide hazard. Similarly,

six bridges are located in a moderately exposed area, and the remaining one bridge is in the zone of low exposure to landslide hazard (Figure 10B).

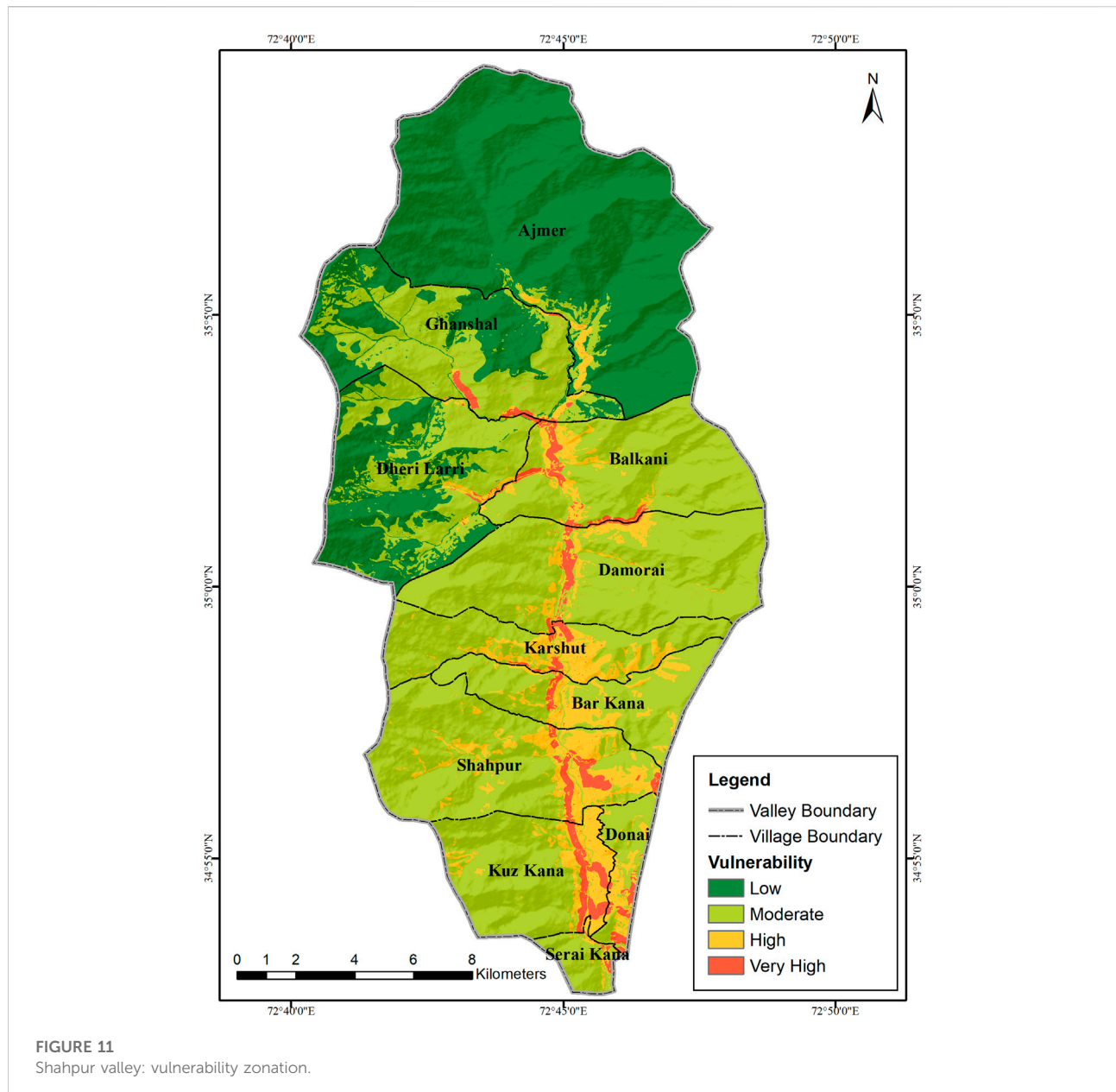
In the study area, agriculture and settlements have high importance and were given a high vulnerability score based on expert opinion, while the rangeland which is not that important in terms of monetary benefits, given a moderate vulnerability score. Thus, vulnerability zonation maps were produced through weighted overlay analysis in the ArcGIS environment using the element at risk layers. Figure 11 shows the different zones of vulnerability for landslide hazard in Shahpur valley. The high and very high vulnerability is mostly found in the lower parts of the valley, while the upper parts of the valley mostly have low vulnerability. The vulnerability variation in the upper and lower parts of the valley is characterized by the variation in the population inhabiting the area and agriculture activity carried out in the valley. Villages Kuz Kana, Shahpur, Bar Kana, and Karshut have high vulnerability as compared to other villages such as Ajmir, Ganshal, and Dheri Larri.

Landslide Risk

The likelihood of landslide occurrence and damage potential to the elements at risk maps (vulnerability) were combined through overlay analysis, and the final landslide risk map was prepared. The risk map depicts the areas of high and low risk. The areas that have low or moderate risk means that there is only likelihood of landslide occurrence or damage potential to elements at risk that exists, while both likelihoods of landslide occurrence and damage potential to elements at risk are high, categorized as high and very high landslide risk zones. According to Figure 12, in Shahpur valley, 41.5% of the total area falls in low-risk areas, while the majority (51%) of the area is in a moderate landslide risk zone. The remaining 5.5% area is at high risk, and 2% is in a very high landslide risk zone. In Shahpur valley, while comparing Figure 12 with the land use, road network, and other infrastructure, it was found that most of the agriculture land, settlements, and road network are located in high and very high landslide risk zones. It was found that 10% of the total agriculture land is in a very high landslide risk area and another 22.75% is at high risk of landslide hazard. This is an alarming issue for the agriculture sector, and similarly, in Shahpur valley, enormous damages could occur to this sector if proper mitigation measures are not taken into consideration.

Discussion

This study focused on landslide susceptibility using WoE and IV models, assessed elements at risk exposure to landslide hazard (susceptibility), assessed vulnerability based on exposure to hazard and probability of loss, and prepared the landslide risk map based on the function of hazard and



vulnerability. Due to unavailability of spatio-temporal landslide inventory in Pakistan and especially in the study region, the inventory of the existing landslide was prepared using high-resolution remote sensing images and verified with extensive field visits. The existing landslides were mapped on the image using polygons. In landslide susceptibility and hazard mapping, polygon data are more effective and reliable than point data (Pourghasemi et al., 2020). In the current study, seven landslide causative factors were selected based on indigenous people's knowledge and studied literature (Khan et al., 2011; Khan et al., 2019; Rahman et al., 2019; Moazzam et al., 2020; Abdi et al., 2021). The relationship between landslide causative factors and landslide occurrences was investigated based on WoE and IV

models as shown in Table 2. It is evident from the results of this study that land cover, geology, proximity to roads, and proximity to fault lines have high Wc value and IV values. It is, therefore, concluded that these factors are more contributing in landslide occurrences, and the same has been observed in other studies (Moazzam et al., 2017; Bacha et al., 2018; Rahim et al., 2018). Moreover, the land cover parameter, the barren land, was found highly susceptible to landslide occurrence as this class is more exposed to erosional and mass wasting processes (Mohammady et al., 2012). Similarly, the geology parameters, Darwaza Sar Potassic Gneiss and Greenschist Melange, are found in the seismic-tectonically active zone and therefore bear a high potential for slope failure.

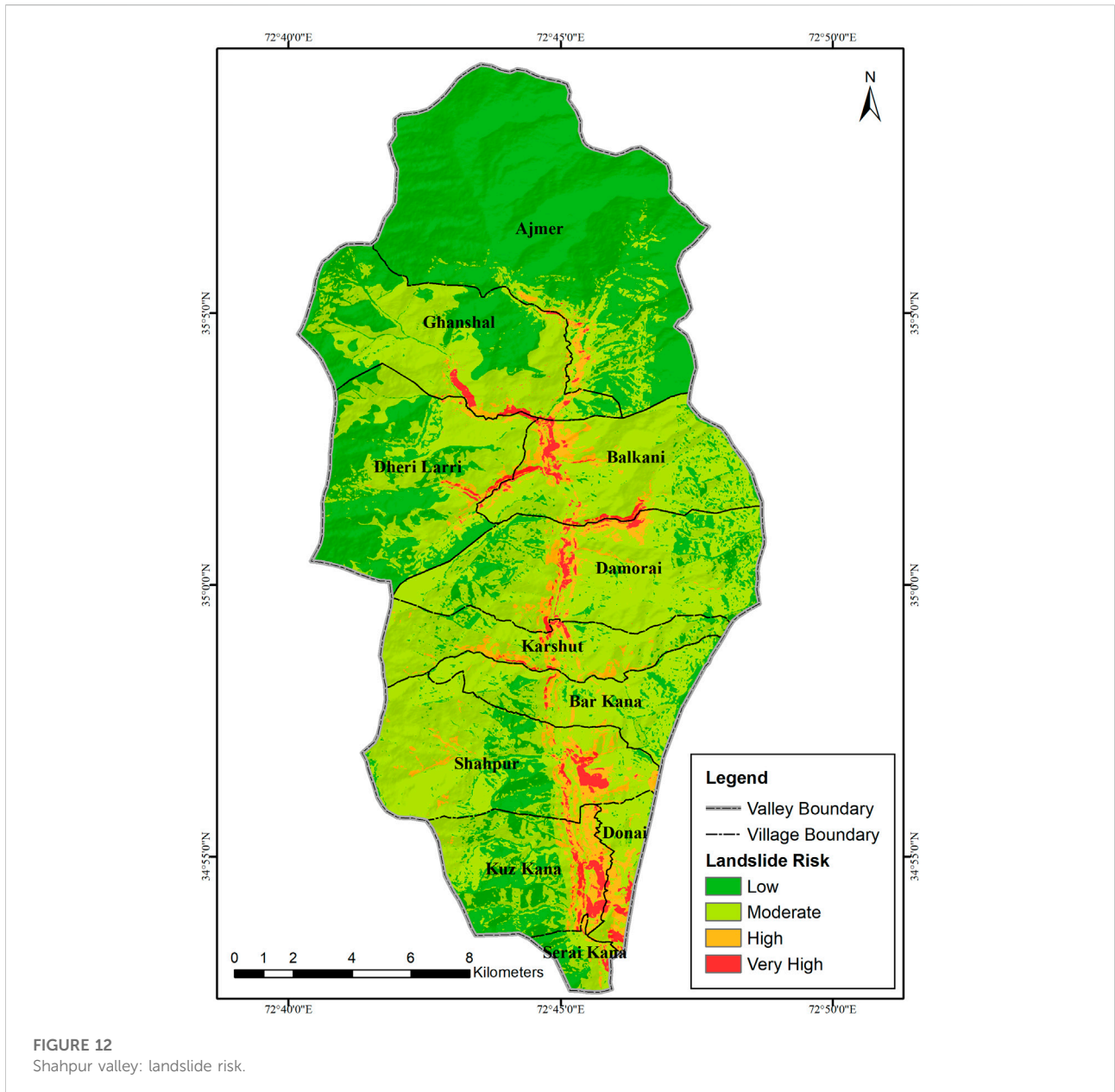


FIGURE 12
Shahpur valley: landslide risk.

Moreover, it is observed from Table 2 that the influence of landslide occurrence decreases with an increase in distance from the roads and streams, and the same has been observed in other similar studies (Regmi et al., 2014; Rahman et al., 2017; Bacha et al., 2018). The construction of roads and stream lateral erosion destabilized the nearby slopes and caused landslides. Earthquake and rainfall are the major landslide causative factors, but due to the unavailability of landslide data associated with both these parameters, we could not use these variables in landslide susceptibility mapping.

The derived susceptibility maps were validated using success and prediction rate curves using training and validation landslides

for both curves, respectively. Both curves showed high accuracy for both WoE and IV models, and the results are in line with other studies (Pradhan et al., 2010; Regmi et al., 2014). Both models (WoE and IV) showed excellent performance in sensitivity, specificity, and accuracy results based on both training and validation datasets as shown in Tables 4 and 5, and similar results have been observed in some previous studies (Abdi et al., 2021; Abedi Gheshlaghi and Feizizadeh, 2021). Furthermore, the susceptibility maps derived based on WoE and IV models were validated using the SCAI validation index, and the results showed a good fit for the models, which are in line with other studies (Arabameri et al., 2020; Saha and Saha, 2020).

In this study, exposure to landslide hazard (susceptibility) was assessed using agricultural land, range land, human settlements, roads, bridges, and school data as elements at risk which were selected based on indigenous people's knowledge and expert opinions. Exposure of elements at risk is one of the core characteristics of vulnerability assessment (Shahabi and Hashim, 2015; Singh et al., 2021). It is difficult to assess social, economic, and environmental vulnerability in a region of scarce data availability; therefore, we just focused on the physical vulnerability of the elements at risk in this study. In the Hindu Kush region, the pace of population growth is higher than the national average (Rahman and Shaw, 2015). This tremendous population growth has increased pressure on the available natural and land resources, which ultimately increased the demand for food, shelter, communication network, and other necessities. In the areas like Shahpur valley, the people have no choice other than engulfing forest and range land to construct shelter for their living and clear the area for agriculture to gain food self-sufficiency. Such situations have accelerated the landslide hazard and vulnerability in the whole Shahpur valley. In the valley, 37.6% of the total agriculture land is highly exposed to landslide hazards (Figure 9A), while 18.3% of settlements are highly exposed to landslides (Figure 9B). Landslides deposit debris on prime agricultural land or remove the productive soil (Nguyen and Kim, 2021). From the fieldwork, it was deduced that most of the agriculture is practiced on greater than 30° slopes due to limited land availability in lower areas. Furthermore, heavy landsliding was observed in heavily grazed and cultivated areas. The forest and natural vegetation clearance for agricultural practices as well as for wood used as a fuel and construction have increased vulnerability in the study region. More than 50% of the road sections in the study area are facing the continuous threat of landslide (Figure 10B). The areas with high population and agricultural practices were found in the highly vulnerable zone (Figure 11). Out of the total study area, 7.5% of the area was found in a high to very high-risk zone where 32% of agriculture is in practice. Evading and reducing the impacts of landslide hazards is possible with education and awareness of people of the community, looking to the history of the hazard to avoid such prone areas. Land use planning, policies, and regulation implementation through local government can reduce the risk of such hazards, and government intervention in this regard has been observed by Almeida et al. (2017).

Conclusion

The WoE and IV models were applied to develop landslide susceptibility maps in this study. Initially, past landslides were marked on the SPOT5 satellite image, validated with consecutive field visits, and then plotted on a map. Landslide causal factors identified through literature review and from indigenous people's knowledge were surface geology, fault lines, land cover, slope gradient and aspect, and distance from streams and roads. The maps of these factors were prepared for susceptibility analysis. The roles of each class of these

factor maps in landslide occurrence were analyzed, and assigned weights were calculated by implementing Bayesian probability models, that is, the weight of evidence and information value. The required susceptibility maps were generated using $\sum W^c$ and $\sum IV$ values through overlay analysis in the GIS.

The LSI maps were prepared based on both models and then validated using the success and prediction rate curves, SCAI, and R-index techniques. The results concluded that the very high landslide susceptibility zone covers 14.49% of the area according to the WoE model and 12.84% of the area in the IV model. Similarly, the area covered by the high-susceptibility zone was 17.84 and 14.57% in WoE and IV models, respectively. The validation results of all applied techniques showed better performance of both the models in the study region. Based on elements at risk, the landslide vulnerability map was generated which showed a high vulnerability to the likelihood of landslide occurrence in the lower parts of the valley. Similarly, the likelihood of landslide occurrence and vulnerability maps were combined, and the risk map of the study area was generated. According to the landslide risk map, 5.5% of the study area was under high risk while 2% of the area was in the very high-risk zone. The availability of temporal landslide records in the study region is the main hindrance in the preparation of landslide hazard and risk mapping. Extensive fieldwork in similar mountainous terrain is another hindrance in the preparation of social and economic landslide vulnerability maps. This study provides a base for disaster management authorities to develop location-specific mitigation measures for landslide hazards to avoid human losses and infrastructural damages in the future. The study concludes that landslide hazards in the region may have negative impacts on agricultural activities, natural ecosystems, river morphology, human lives, and infrastructure in future. In this regard, proper land use planning and strict forest preservation measures are highly required to reduce the environmental and socio-economic damages from landslide hazards.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary Material; further inquiries can be directed to the corresponding author.

Author contributions

GR, MU, AB, SM, and AR proposed the main concept and are highly involved in the write-up. AB, MU, SM, GR, MA, HA, AA, MA-M, and HA assisted in data analysis and preparation of the spatial map. GR, MU, AR, SM, MA-M, AA, HA, MA, and HA are involved in the write-up and review. MM, AB, AA, SM, AR, HA, MA, MA-R, and HA are involved in the review, editing and review, and English grammar correction.

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Conflict of interest

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