



# Changes of Divergence and Confluence Relationship Between Dongting Lake and the Yangtze River After the Operation of the Three Gorges Project and Its Impact on the Waterway Depth

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Yang Y, Liu W, Zhang J, Yang L, Jia M and Zhu L (2022) Changes of Divergence and Confluence Relationship Between Dongting Lake and the Yangtze River After the Operation of the Three Gorges Project and Its Impact on the Waterway Depth. Front. Earth Sci. 10:829669. doi: 10.3389/feart.2022.829669 Human activities and climate change have profoundly changed the hydrological and geomorphological evolutionary processes and trends of rivers and lakes, which, in turn, has affected the utilization of their waterway depth. This study chose the Dongting Lake-Jingjiang confluence of the Changjiang River (Yangtze Rvier), since the operation of the Three Gorges Project, the split ratio of water and the sediment distribution of the three outlets of Dongting Lake-that is, Songzikou, Taipigkou, and Ouchikou-have continued to decrease. Along with the decreasing flow of Dongting Lake, the relative increase in the runoff of the Jingjiang reach and the increase in riverbed erosion intensity have increased the relative height difference between Jingjiang and Dongting Lake, intensifying the reduction of the split ratio of water and the sediment distribution of the three outlets. The riverbed erosion of the Jingjiang reach has created a promising foundation for an increase in the waterway dimensions. However, the length of the reach not meeting the requirements of 4.5× 200 m (water depth × width) is 18.4 km, of which the length in the river-lake confluence is 12.6 km, accounting for 68.35% of the total length of obstructed navigation. Furthermore, at the reach (Zhicheng-Dabujie reach) affected by the first outlet (Songzikou), a 4.5-m flume has been formed, but its width is less than 200 m. Moreover, the "steep slope and rapid current" phenomenon is clear, and it affects the safety of ships. At the reach (Taipingkou reach) affected by the second outlet, the shrinkage of the bottomland, and the intersecting braided river channels make the lowflow routes unstable, the water depth being less than 4.5 m. At the reach (Tianxingzhou-Ouchikou reach) affected by the third outlet, the shoreline of the reach has collapsed and retreated significantly. Additionally, beach erosion and the downstream movement of cut banks have caused the downstream waterway width at the reach to be reduced to less than 200 m or formed shoals with a water depth of less than 4.5 m. Owing to the jacking effect of the lake's outflow backwater, the flow rate of the reach

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(Xiongjiazhou-Chenglingji reach) affected by the confluence of Dongting Lake decreased, creating a shoal with a water depth of less than 4.5 m in the navigation waterway. The results of this study have guiding significance for understanding the changes of channel conditions and the improvement of channel scale in the intersection area of rivers and lakes.

Keywords: Changjiang River waterway, Dongting Lake, deep-water resources, navigation safety, shipping, river-lake confluence, obstructed navigation

### INTRODUCTION

Against the background of intense human activities and global climate change, the hydrology, sedimentation, and landform processes of most lakes globally have changed tremendously (Yu et al., 2018; Mammides, 2020; Grant et al., 2021). Dongting Lake is the second-largest freshwater lake in China and an important lake in the middle reaches of the Changjiang River. Incoming water from the upper reaches of the Changjiang River enters Dongting Lake through Songzikou, Taipingkou, and Ouchikou. Meanwhile, Dongting Lake also receives the "four rivers"-that is, the Xiangjiang River, Zishui River, Yuanjiang River, and Lishui River-which flow from Chenglingji to the Changjiang River, forming a complex river-lake relationship with the Changjiang River (Huang et al., 2012), It affects the disaster prevention and mitigation of river basin flood (Duan et al., 2016), water quality safety (Duan et al., 2018) and economic and social development (Duan et al., 2016). Owing to the impact of climate and human activities, the relationship between Dongting Lake and the Changjiang River has undergone many changes since the beginning of the 20th century (Han et al., 2016; Zheng, 2016; Dai et al., 2017). The opening and operation of the Three Gorges Project in June 2003 meant that the river-lake relationship had entered a new period of change (Hu et al., 2013; Lai et al., 2014; Fang et al., 2014; Zhan et al., 2015; Li et al., 2021a).

Over the past 20 years, researchers have been studying the impact of the Three Gorges Reservoir operation on the river-lake relationship in terms of the spatial and temporal evolution of water and sediment in the Dongting Lake (Nakayama and Watanabe, 2008; Zhang et al., 2016), the water exchange capacity of the Changjiang River and Dongting Lake (Li et al., 2013; Ou et al., 2014; Zhu et al., 2016; Lai and Wang, 2017; Mao et al., 2017), floods (Jiang et al., 2007; Huang et al., 2014; Lai and Wang, 2017; Wu et al., 2019; Geng et al., 2021), the volumes of the river and lake (Liao et al., 2018; Lu et al., 2018), and ecological and environmental factors (Zhang et al., 2016; Lu et al., 2018; Zhang et al., 2020). The Three Gorges Reservoir operation has affected the flow of the Changjiang River (Wang et al., 2016), causing the early dry out of Dongting Lake and the volume of water entering the lake through the three outlets to decline (Zhou and Zhang, 2018). From 2003 to 2008, natural inflow, reservoir regulation, and riverbed erosion contributed 76.19%, 10.32%, and 13.49% to the process of reducing the flow through the three outlets of Dongting Lake. Moreover, from 2009 to 2015, the respective figures were 66.05%, 19.75%, and 14.20% (Chen et al., 2018). The Three Gorges Reservoir's impoundment during the flood

recession period further reduced the water level of Dongting Lake. Combined with the upstream cascade reservoirs, the impact on Dongting Lake increased (Fang et al., 2014).

The Jingjiang Cut-off Project was the main reason for the decrease in the volume of sediment coming out of Dongting Lake from 1971 to 2002. With the operation of the Three Gorges Reservoir, siltation in the lake area changed to erosion (Zhou et al., 2016; Zhu et al., 2016; Lu et al., 2018) and bottomland shrinkage (Yu et al., 2019). For example, a mathematical model predicts that after the Three Gorges Project has been in operation for 20 years, the decrease in the middle and high-water levels of the Jingjiang reach will be limited, but the riverbed elevation of the flood channels at the three outlets will decrease significantly, although the diversion ratio of the three outlets will not decrease (Li et al., 2009). Under a scenario in which the lake volume remains constant, the reduction in the volume of water outflow from Dongting Lake is approximately 23% of the water storage capacity of the Three Gorges Reservoir (Zhang et al., 2019); under a scenario with different flow-water levels, the jacking strength between floods of the river and lake are both weakened (Duan et al., 2019; Liu et al., 2019); under a scenario with the same flow from the Zhicheng Hydrological Station of the Changjiang River, the flow of the Jianli Hydrological Station is weakened by the jacking of the Dongting Lake outflow, which is also weakened by the jacking of the discharge water of Jingjiang reach (Yu et al., 2017).

Since the operation of the Three Gorges Reservoir, the erosion of the Dongting Lake area has caused the low water level to increase and the depth of the waterway to increase (Zhang et al., 2015). The phase 1 waterway regulation project has been implemented in the Jingjiang reach, and the waterway depth has been increased from 2.9 m in 2002 to 3.5-3.8 m in 2020 (Yan et al., 2019; Yang et al., 2019). An inspection of waterway conditions found that the depth of the waterways from Yichang to Wuhan and Wuhan to Anging had favorable conditions, increasing from 3.5 m and 4.5 m to 4.5 m and 6.0 m, respectively (Yang et al., 2019). Meanwhile, the complex relationship between the hydrodynamic, sediment discharge, and bottomland evolution in the confluence area of Dongting Lake and the Changjiang River has been noted, a key reach for waterway management and dredging maintenance (Dong et al., 2020). Existing studies have also focused on changes in riverbed erosion and silting (Xia et al., 2016; Xia et al., 2017; Yang et al., 2017; Yang Y. et al., 2018; Yang et al., 2018b; Yang et al., 2018c), bottomland changes (Zhao et al., 2020; Yang et al., 2021), diversion relationships (Liu et al., 2017; Hu et al., 2020; Yang et al., 2021), and navigation obstruction characteristics (Yan et al.,



FIGURE 1 | Overview of the middle reaches of the Changjiang River and Dongting Lake; (A) Changjiang River Basin; (B) Relationship between Dongting Lake and Jingjiang reach; (C) Jingjiang reach.

2012; Yang et al., 2019; Li et al., 2021b) in the confluence area of Dongting Lake and Changjiang River since the operation of the Three Gorges Reservoir. The existing studies have paid little attention to the influence of the adjustment of diversion relationship of Dongting Lake on channel conditions in the intersection area of river and lake, especially the influence on the development and utilization of water depth resources in the main channel of the Changjiang River.

The shipping function of the river supports economic and social development along the river. Until 2020, the water discharge volume of the Changjiang River has been 30.6  $\times$ 10<sup>8</sup> t, accounting for 78.6% of that of China's inland rivers. However, previous studies have rarely examined changes in waterway depth in the confluence areas of rivers and lakes. Consequently, using the Jingjiang reach, which is affected by the confluence of the Changjiang River and Dongting Lake, as the research object, the water, sediment, and contour data of the Changjiang River-Dongting Lake region from 1955 to 2020 were used to study the hydrological effects of Dongting Lake and the impact of the overbank evolution of the Jingjiang reach on the waterway conditions at the river-lake confluence area, as well as to analyze the influence of the changing river-lake relationship on the waterway conditions of the Jingjiang reach of the Changjiang River.

### **RESEARCH AREA**

# Overview of Dongting Lake and Changjiang River

The Jingjiang reach (Zhicheng–Chenglingji reach) is located in the middle reaches of the Changjiang River (**Figure 1A**), with a total length of 375 km. Dongting Lake is located on the south bank of the Changjiang River. Water diverges and enters Dongting Lake from Songzikou, Taipingkou, and Ouchikou, merging into the Changjiang River at Chenglingji (**Figure 1B**). The hydrological stations along the Jingjiang reach are Zhicheng, Shashi, Jianli, and Luoshan Hydrological Stations. The hydrological Stations at Songzikou of Dongting Lake are Xinjiangkou Station and Shadaoguan Station, at Taipingkou the Hydrological Station is Mituosi Station, at Ouchikou the hydrological stations are Tangjiagang Station and Guanjiapu Station, and at the confluence into the Changjiang River, the hydrological station is Chenglingji Station (**Figure 1C**).

### Waterway Regulation Project

Songzikou is the first diversion area from the Changjiang River to Dongting Lake, covering the reach from Zhicheng to Changmenxi, the distributed central bars being Guanzhou and Lujiahe, the waterway regulation project (**Figure 2A**) being



implemented in the Zhicheng–Changmenxi reach from 2013 to 2017. The purpose was to stabilize the diversion relationship and have a good overbank shape. The control target was  $3.5 \times 150 \times 1,050$  m (water depth × width × bending radius).

Taipingkou is the second diversion area from the Changjiang River to Dongting Lake, covering the Taipingkou reach, the distributed central bars being Taipingkou and Sanbatan, and the beaches being Lalinzhou and Yanglinji. From 2005 to 2012, two phases of waterway regulation projects were implemented at Sanbatan, mainly to stabilize it, and from 2010 to 2017, two phases of Lalinzhou protection projects were implemented to stabilize the waterway boundaries (**Figure 2B**).

Ouchikou is the third diversion area from the Changjiang River to Dongting Lake, covering the Tianxingzhou–Ouchikou reach, with the Xinchang Beach, Tianxingzhou, Tuoyangshu Beach, Ouchikou central bars, Doukouyao central bars, and Xiangjiazhou Beach. The bars most directly affected by the diversion of Ouchikou are the Tianxingzhou and Tuoyangshu Beaches. From 2010 to 2013, the protection of the high bar on the left of Tianxingzhou, the lower section of the Tuoyangshu Beach, and the bank protection project was implemented, and from 2013 to 2017, the Dakouyao central bars, Tuoyangshu Beach, and Xinchang Beach protection projects were implemented to stabilize the waterway boundary and increase the waterway water depth to 3.5 m and above (**Figure 2C**).

# RESEARCH DATA, METHODS, AND MODELS

### Data

The water level, water, and sediment data of Zhicheng, Shashi, Jianli, and Luoshan Hydrological Stations from 1955 to 2020 were collected to analyze changes in the water and sediment conditions in the Jingjiang reach. Data on the water and sediment volume of Songzikou (Xinjiangkou, Shadaoguan Hydrological Station), Taipingkou (Mituosi Hydrological Station), Ouchikou (Guanjiapu, Tangjiagang Hydrological Stations), and Chenglingji Hydrological Station of Dongting Lake from 1955 to 2020 were collected to analyze the water and sediment conditions of the Dongting Lake area, the contour data of the riverbed of the Jingjiang reach (Zhicheng–Chenglingji reach) being collected to analyze the waterway conditions, bottomland deformation, and cross-section changes in the Dongting Lake–Jingjiang confluence area (**Table 1**).

### Calculation Method of Branch Diversion Ratio and Waterway Dimension

The navigational datum is the water level H<sub>0</sub> corresponding to the position of a section, usually the water surface line along the river course with a 98% guarantee calculated using the synthetic duration curve method (Figure 3A) (Changjiang Waterway Bureau, 2014). In addition, the difference between the water level of the navigational datum and the elevation of the waterway bottom is the water depth (h). The dimensions of a waterway include its water depth (H), width (B), bend radius (R), and navigation clearance height  $(H_{max})$ . If the water depth corresponding to the actual lowest navigable water level (h) is less than the target navigation depth (H), a break will appear in the depth contour corresponding to H-a navigation obstacle due to insufficient water depth (Figure 3B). If a location on the waterway has h greater than H-the depth contour at H is not broken-but a width less than B, the location is a navigation obstacle caused by insufficient navigable width. Similarly, if R is too small for safe passage, route adjustments may lead to insufficient waterway width and/or depth.

The calculation of the split ratios of water and sediment ratios (**Figure 3C**) is as follows: firstly, the total inflow of the braided reach Q and S is obtained by measuring the runoff at the cross-section of its inlet. If the runoff flowing into each branch is  $Q_i$  and  $S_i$  (i = 1, 2, ..., n where n is the number of branches), the split ratios of water and sediment distribution ratios  $\eta_i$  and  $\xi_i$  of each branch is given by:

#### TABLE 1 | Data characteristics and sources.

	•								
Number	Profile name	Data type	Time	Source					
1	Zhicheng, Shashi, Jianli, Luoshan Hydrological Stations	Runoff,	1955–2020	Changjiang Waterway Bureau; Changjiang Hydrology Bureau of					
2	Xinjiangkou, Shadaoguan, Nituosi、Guanjiapu,	sediment		Changjiang Water Resources Committee; Changjiang Waterway					
	Tangjiagang, Chenglingji Hydrological Stations			Planning, Design and Research Institute					
3	Jiangjiang reach	Terrain	2003–2020						



$$\eta_i = \frac{Q_i}{Q_1 + Q_2 \dots Q_n} \times 100\% = \frac{Q_i}{Q} \times 100\%; i = 1, 2, \dots, n$$
  
$$\xi_i = \frac{S_i}{S_1 + S_2 \dots S_n} \times 100\% = \frac{S_i}{S} \times 100\%; i = 1, 2, \dots, n$$

# $(S_{\text{Mituosi}}, 10^8 \text{ t/y})$ ; the hydrological control stations at Ouchikou are Guanjiapu $(S_{\text{Guanjiapu}}, 10^8 \text{ t/y})$ and Tangjiagang $(S_{\text{Tangjiagang}}, 10^8 \text{ t/y})$ . Consequently, the formula for sediment erosion and deposition can be expressed as follows:

$$T_{Jing jiang reach} = S_{Zhicheng-SLuoshan} - S_{Xinjiangkou} - S_{Shadaoguan}$$

$$-S_{\rm Mituosi} - S_{\rm Guanjiapu} - S_{\rm Tangjiagang}$$
(2)

# Calculation Method of Riverbed Erosion and Deposition

The amount of sediment erosion and silting in the Jingjiang reach can be calculated using the sediment discharge method. The upstream control station is the Zhicheng Station ( $S_{Zhicheng}$  10<sup>8</sup> t/ y), and the downstream control station is the Luoshan Station ( $S_{Luoshan}$ , 10<sup>8</sup> t/y). Sediment diverges from Songzikou, Taipingkou, and Ouchikou into Dongting Lake—that is, ( $S_{three}$ outlets =  $S_{Songzikou+}S_{Taipingkou} + S_{Ouchikou}$ , 10<sup>8</sup> t/y). The difference in sediment discharge volume between the upstream and downstream, plus Dongting Lake, is the amount of sediment erosion and silting in the Jingjiang reach ( $V_{Jingjiang reach}$ , 10<sup>8</sup> t/y), which can be calculated as follows:

$$V_{\text{Jingjiang reach}} = S_{\text{Zhicheng}} - S_{\text{Luoshan}} - S_{\text{Songzikou}} - S_{\text{Taipingkou}} - S_{\text{Ouchikou}}$$
(1)

The hydrological control stations at Songzikou are Xinjiangkou ( $S_{Xinjiangkou}$ ,  $10^8$  t/y) and Shadaoguan ( $S_{Shadaoguan}$ ,  $10^8$  t/y); the hydrological control station at Taipingkou is Mituosi

### **RESEARCH RESULTS**

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### Runoff and Sediment Variation of Changjiang River

From 1955 to 2002 and 2003 to 2020, the runoffs of Zhicheng, Shashi, Jianli, and Luoshan Hydrological Stations were 4,450 × 108 m<sup>3</sup>/y, 3,933 × 108 m<sup>3</sup>/y, 3,567 × 108 m<sup>3</sup>/y, and 6,408 × 108 m<sup>3</sup>/y and 4,282 × 108 m<sup>3</sup>/y, 3,907 × 108 m<sup>3</sup>/y, 3,780 × 108 m<sup>3</sup>/y, and 6,222 × 108 m<sup>3</sup>/y, respectively, the variations being -3.76%, -0.65%, +5.95%, and -2.90% (Figure 4A). From 1955 to 2020, the sediment at the Zhicheng, Shashi, Jianli, and Luoshan Hydrological Stations exhibited a decreasing trend. From 1955 to 2002 and 2003 to 2020, the sediment at the Zhicheng, Shashi, Jianli, and Luoshan Hydrological Stations exhibited a decreasing trend. From 1955 to 2002 and 2003 to 2020, the sediment at the Zhicheng, Shashi, Jianli, and Luoshan Hydrological Stations were  $5.00 \times 10^8$ ,  $4.33 \times 10^8$ ,  $3.70 \times 10^8$ , and  $4.11 \times 10^8$ , and  $0.42 \times 10^8$  t/y,  $0.52 \times 10^8$ ,  $0.68 \times 10^8$ , and 0.84





 $\times$  10  $^8$  t/y, respectively, the variations being –91.58%, –87.95%, –81.53%, and –79.45% (**Figure 4B**).

# Variation of Runoff and Sediment in Dongting Lake

From 1955 to 2002 and 2003 to 2020, the runoffs of Songzikou, Taipingkou, and Ouchikou of Dongting Lake were  $419 \times 108$ ,  $163 \times 108$ , and  $336 \times 108 \text{ m}^3/\text{y}$ , and  $306 \times 108 \text{ m}^3/\text{y}$ ,  $81 \times 108 \text{ m}^3/\text{y}$ y, and  $111 \times 108 \text{ m}^3/\text{y}$  respectively, the extent of the decrease being 26.97%, 50.31%, and 66.96% (Figure 5A). The extent of the decrease between all three outlets was 48.00% (Figure 5C). From 1955 to 2002 and 2003 to 2020, the sediment of Songzikou, Taipingkou, and Ouchikou of Dongting Lake were  $4,652 \times 10^4$ ,  $1887\times10^4,$  and  $5{,}842\times10^4,$  and  $474\times10^4,$   $116\times10^4,$  and  $282\times10^4$  $10^4$  t/y, respectively, the extent of the decrease being 89.81%, 93.85%, and 95.17% (Figure 5B). The extent of the decrease in sediment between all three outlets was 92.94% (Figure 5C). Comparing the data from 2003 to 2020 with that from 1955 to 2002, the runoff of Chenglingji Hydrological Station decreased by 13.84% (Figure 6). From 1955 to 2007, the amount of sediment at Chenglingji Station showed a decreasing trend, whereas from 2007 to 2012 it increased slightly, before

showing a decreasing trend from 2013 to 2020 (**Figure 5D**). Comparing the data from 2003 to 2020 with that from 1955 to 2002, the amount of sediment at Chenglingji Station decreased by 55.8%.

Since 1955, affected by the cut-off of the Lower Jingjiang reach, the construction of the Gezhouba Dam, and the Tree Gorges Dam, the river regime change at the gate section of the three outlets, and other factors, the riverbed of the Jingjiang reach has been eroded. Furthermore, at the same flow rate, the water level has dropped and the riverbeds of the splitting channels of the three outlets have been silted up-that is, the split ratio of water and the sediment distribution capacities of the three outlets have been in a state of attenuation (Table 2). From 1955 to 1966, the split ratios of water and sediment of the three outlets of Dongting Lake were 29.8% and 35.4%, respectively; during the systematic cut-off period of the Lower Jingjiang reach from 1967 to 1972, the split ratio water and the sediment distribution ratio of the three outlets were 23.7% and 28.3%, respectively, 6.1% and 7.1% less than before the cutting off. From 1973 to 1980, the decay rate of the split ratio of water and the sediment distribution capacity of the three outlets increased, with the average distribution ratios being 18.8% and 21.6%, respectively. After the Gezhouba Dam was built in 1981, the decay rate of the diversion ratios of the three



**TABLE 2** | Runoff and sediment quantity, split ratio of water and the sedimentof the three outlets of Dongting Lake.

Parameter/Time		1955–1966	1967–1972	1973–1980	1981–1998	1999–2002	2003-2020
Zhicheng	Runoff (10 <sup>8</sup> m <sup>3</sup> /y)	4,539	4,302	4,441	4,442	4,454	4,283
	Sediment (10 <sup>4</sup> t/y)	55,640	50,359	51,246	49,112	34,550	4,215
Songzikou	Runoff (10 <sup>8</sup> m <sup>3</sup> /y)	490	445	428	377	345	306
	Split ratio of sediment (%)	10.8	10.3	9.6	8.5	7.7	7.1
	Sediment (10 <sup>4</sup> t/y)	5,433	4,846	4,711	4,442	2,846	474
	Split ratio of sediment (%)	9.8	9.6	9.2	9.0	8.2	11.2
Taipingkou	Runoff (10 <sup>8</sup> m <sup>3</sup> /y)	210	186	160	133	126	81
	Split ratio of water (%)	4.6	4.3	3.6	3.0	2.8	1.9
	Sediment (10 <sup>4</sup> t/y)	2,397	2,202	1935	1,644	1,019	116
	Split ratio of sediment (%)	4.3	4.4	3.8	3.3	2.9	2.7
Ouchikou	Runoff (10 <sup>8</sup> m <sup>3</sup> /y)	652	390	247	189	155	111
	Split ratio of water (%)	14.4	9.1	5.6	4.2	3.5	2.6
	Sediment (10 <sup>4</sup> t/y)	11,842	7,220	4,430	3,244	1795	283
	Split ratio of sediment (%)	21.3	14.3	8.6	6.6	5.2	6.7
Three outlets	Runoff (10 <sup>8</sup> m <sup>3</sup> /y)	1,352	1,022	834	699	625	498
	Split ratio of water (%)	29.8	23.7	18.8	15.7	14	11.6
	Sediment (10 <sup>4</sup> t/y)	19,672	14,268	11,076	9,330	5,660	873
	Split ratio of sediment (%)	35.4	28.3	21.6	19	16.4	20.7

outlets slowed down from 1981 to 1998. The split ratio of water and the sediment distribution ratio of the three outlets were 14.0% and 16.4% from 1999 to 2002 and 11.6% and 20.7% from 2003 to 2020, respectively. In other words, with the operation of the Three Gorges Project, the split ratio of water of the three outlets decreased by 2.4%, and the sediment distribution ratio increased by 4.3%. From 1955 to 2002, the sediment distribution ratio of the three outlets of Dongting Lake was higher than the



split ratio—the split effect of the three outlets caused the sediment content of the Jingjiang reach to trend downwards. From 2003 to 2020, the sediment distribution ratio was higher than the split ratio of water by approximately 9.1%, which was higher than the 2.4%–5.6% difference in the period 1955–2002, before the impoundment of the Three Gorges Project, further increasing the erosion intensity of the Jingjiang reach.

### Sediment Deposition in Jingjiang Reach

From 1955 to 2002, the riverbed of the Jingjiang reach was alternately eroded and silted. From 2003 to 2020, it was in a state of continuous scouring, the cumulative erosion total being  $5.98 \times 108$  t, and the annual erosion total being  $0.33 \times 108$  t/y (Figure 6A). From 2003 to 2020, 39.4% of the sediment volume at Luoshan Hydrological Station was from the erosion of the Jingjiang reach, 10.7% was due to the exchange interaction with Dongting Lake, and 49.9% was from the river basin upstream of Zhicheng Station. Before the operation of the Three Gorges Project, the sediment content of the three outlets was higher than that of the Jingjiang reach during the medium flood years. After the operation of the Three Gorges Project, the sand content of the three outlets was higher than that of the Jingjiang reach. In other words, the changes in Dongting Lake's sediment distribution from 2003 to 2020 have caused the Jingjiang reach to erode (Figure 6B); the sediment content of Chenglingji Station, where Dongting Lake convergences into the Changjiang River, is lower than that of Luoshan Hydrological Station on the Changjiang River, its confluence effect aggravating the sediment erosion in the Jingjiang reach (Figure 6C).

### **Verification of Waterway Conditions**

The water depths of the Jingjiang reach waterway were tallied based on the river contour that was surveyed in October 2020 (**Figure 7**). Given a waterway width of 200 m, it was found that there were 14 waterwar with water depths less than 4.5 m in the Jingjiang reach, including the Guanzhou, Lujiahe, Zhijiang, Jiangkou, Dabujie, Yuanshi, Taipingkou, Wakouzi, Zhougongdi, Jianli, Damazhou, Tiepu, Chibakou, and Guanyinzhou waterways. The minimum water depths of the 19 remaining waterways were all greater than 4.5 m.

After drawing a 4.5 m depth contour through the Jingjiang reach, it was found that there were 13 waterways with widths less than 200 m, that is, the Zhicheng, Guanzhou, Lujiahe, Zhijiang, Jiangkou, Dabujie, Yuanshi, Taipingkou, Wakouzi, Zhougongdi, Jianli, Damazhou, and Guanyinzhou waterways. The other 20 channels had widths greater than 200 m on their 4.5 m depth contours. Given a waterway scale of  $4.5 \times 200$  m, the Jingjiang reach was either insufficiently wide or deep in the Guanzhou, Lujiahe, Zhicheng, Jiangkou, Dabujie, Yuanshi, Taipingkou, Wakouzi, Zhougongdi, Jianli, Damazhou, Chibakou, and waterways. These navigation Guanyinzhou hindering waterways accounted for 5.3% of the Jingjiang reach's total length (18.4 km). The Songzikou, Taipingkou, Ouchikou of Dongting Lake, and confluence-affected reaches have a navigation obstruction length of 12.6 km, accounting for 68.35% of the total navigation obstruction length. In other words, the confluence area of Dongting Lake exhibited a higher degree of navigation obstruction than other mainstream reaches.





FIGURE 9 | Hydrodynamic and waterway changes in shunt area of Songzikou; (A) Velocity distribution; (B) Water depth; (C) Maximum velocity; (D) Maximum gradient.

### DISCUSSION

# Waterway Conditions of Songzikou Diversion Area

In 1955–2002 and 2003–2020, the split ratio of water of Songzikou was 419  $\times$  108 m<sup>3</sup>/y and 305.8  $\times$  108 m<sup>3</sup>/y, respectively, the split ratios of water being 9.5% and 7.1%. The

right branch of the Lujiahe has a sandy riverbed, and the left branch has a sandy pebble riverbed. The left branch has higher erosion resistance than the right branch (Li Y. et al., 2021). Since the decrease in the split of Songzikou means the relative increase in runoff of the Jingjiang reach, the left branch of the Lujiahe reach was eroded and developed due to weak eroded resistance, which caused a relative decrease in the split ratio of water of the



left branch of the Lujiahe reach (**Figure 8A**). On the relationship between the flow of Zhicheng Hydrological Station and the split ratio of water of the left branch of the Lujiahe reach, 2008–2017 exhibited a decreasing trend compared to 2003–2007 (**Figure 8B**). After a comprehensive analysis, it is believed that the decrease in the split ratio of water of Songzikou is an important factor causing the instability of the Lujiahe branch. From 2003 to 2017, phase 1 of the channel regulation project of the Jingjiang reach was implemented, and the split ratio of water of the left branch of the Lujiahe increased by 11.1%. The implementation of the project played a positive role in suppressing the eroded of the right branch and the reduction of the left branch split ratio of water.

The planned waterway size of the Jingjiang reach is  $4.5 \times$  $200 \text{ m} \times 1,050 \text{ m}$ , the weight of navigable ships being 5,000 t. According to the Changjiang River trunk line, the maximum speed of ascending rapids for the existing 5000-ton ships is 2.8 m/ s and the maximum slope is 8500. From 2005 to 2018, the maximum flow velocity of the Zhicheng-Changmenxi reach increased significantly, reaching a maximum of 3.50 m/s. In 2020, the maximum flow velocity was reduced to 3.28 m/s (Figure 9A, Figure 9C), which was still not conducive to the safety of ships; the vertical slope of the Zhicheng-Chenerkou reach relatively small, the slope was of the Chenerkou-Changmenxi reach having a significant increasing trend. Affected by the decrease in the split ratio of flow of the Songzikou, the increase in the split ratio of water of the right branch of the Lujiahe reach caused significant erosion of the branch, resulting in a weak hydrodynamic force on the left branch, resulting in the waterway water depth less than 4.5 m or the width being less than 200 m (Figure 9B).

The erosion resistance of the riverbed of the right branch was higher than that of the downstream sandy riverbed, which meant the headward transmission effect due to the downstream water level drop was significant. Since 2007, the slope of the Changmenxi–Dabujie reach has been higher than 85000 (Figure 9D), creating the phenomenon of "steep slopes and rapid currents," which is not conducive to the safety of ships.

# Waterway Conditions of Taipingkou Diversion Area

## Relationship Between Branch Diversion and Main Waterway

During the dry season of 1975–2000, the flow diversion ratio of the north branch of the Sanbatan Central bar was higher than that of the south branch; thus, the main navigable waterway was located in the north branch (**Figure 10**). The flow diversion ratio of the north branch decreased during 2000–2003, after which a multi-phase waterway regulation project was implemented. From 2004 to 2012, the flow diversion ratio of the north branch exhibited an increasing trend, the desired flow diversion ratio being achieved through the waterway regulation project. Since 2013, the enlargement of the Yanglinji Beach, through silting, significantly narrowed the inlet of the north branch, rapidly reducing its flow diversion ratio from more than 60% in 2012 to approximately 15% in.

Since the implementation of Phase 1 of the Jingjiang Reach Waterway Regulation Project, the flow diversion ratio of the north branch was restored to some extent in 2016–2017, reaching 21.3% at the beginning of 2017, before decreasing to 17.8% by the end of 2017. Since the formation of the Taipingkou Central bar, the dry-season flow diversion ratio of the Taipingkou south branch has increased gradually, exceeding the flow diversion ratio of the north branch since 2004 and maintaining a level above 50% during 2004–2017. As shown by the flow diversion ratios, the Taipingkou south branch has been dominant over the north branch, making it appropriate to select the south branch as the main branch. The flow diversion ratio of the north branch of the Sanbatan Central bar has been higher than that of the south branch in most years; therefore, the north branch was usually selected as the customary main navigable waterway during the



dry season, despite a slight decline in the flow diversion ratio in recent years. From 1955 to 2020, both the split ratios of water and the sediment ratio of Taipingkou were decreasing. Consequently, with the decreasing volume of sediment in the river basin, the south branch of Taipingkou developed a erosion trend. Since 2003, the 4.5 m waterway of the south branch of Taipingkou has been finished (**Figure 10**) and used as the main waterway during the dry season. Affected by sediment excavation activities in the north branch of Taipingkou since 2014 (Zhao et al., 2020; Yang et al., 2021), the north branch of Taipingkou has been greatly erosion and a 4.5 m waterway has been opened since 2018 as a dry-season waterway, forming a two-branch navigable route.

#### Morphological Changes of Beach

From 1955 to 1975, the split ratio flow and split ratio of water of the Taipingkou fluctuated. During this period, the scale of the Lalinzhou Beach was relatively large, and the central bars of Taipingkou had not yet formed (**Figure 11**). The 1975–1993 Taipingkou split ration of water by 29.3% compared with that from 1955 to 1975, which caused a relative increase in the runoff of the Jingjiang reach, causing the upper section of the Lalinzhou Beach to erode and separate, forming the Taipingkou central bars. From 1993 to 2008, the middle and upper sections of Lalinzhou Beach largely collapsed, corresponding to an increase in the area of the Taipingkou central bars, a great decrease in the area of Sanbatan, silting and erode of the upper section. From 2008 to 2019,

the area of the Taipingkou central bars and Sanbatan central bars exhibited a decreasing trend, the area of the upper section of Lalinzhou Beach being stable, and the area of the lower section exhibiting an increasing trend.

#### **Changes in Waterway Conditions**

Since 2017, the water depth of the waterway in the Taipingkou reach has been maintained at 3.5-3.8 m, the planned water depth of the second-phase channel regulation project of the Jingjiang reach being 4.5 m. The water depth conditions of the 4.5 m waterway in the Taipingkou reach are checked based on the actual measured water depth contour from 1975 to 2018 (Figure 12A). The Taipingkou reach consists of four waterways: 1) Waterway 1, from the Taipingkou south branch to the Sanbatan south branch; 2) Waterway 2, from the Taipingkou south branch to the Sanbatan north branch; 3) Waterway 3, from the Taipingkou north branch to the Sanbatan north branch; and 4) Waterway 4, from the Taipingkou north branch to the Sanbatan south branch (Figure 12B). During 1975-2018 (Figure 12C), the 4.5 m water depth conditions gradually improved for waterways 1 and 2. Since 2011, a 4.5 m isobath has run unobstructed throughout Waterway 1, and since 2008 a 4.5 m isobath has run unobstructed throughout Waterway 2.

In terms of the waterway depth, the implemented waterway improvement project has achieved the expected goal, maintaining Waterway 2 as the main navigable waterway during the dry



season. Since the implementation of the Three Gorges Project, the water depth conditions for the north branch have gradually improved as the river branches in the Taipingkou reach have been erode out, creating what is now the existing 4.5 m isobath throughout. Moreover, from the 2018 bathymetric survey map, a 4.5 m isobath has also developed throughout waterways 3 and 4.

# Waterway Conditions of Ouchikou Diversion Area

With the decrease in the split ration of water at Ouchikou, the relative increase in the runoff of the Jingjiang reach has caused the collapse of the right edge of Tianxingzhou Central Bar and the erode of the upper section of the Tuoyangshu Beach, directly affecting the waterway conditions of the downstream reaches (Figure 13A). Since 2003, the right edge of Tianxingzhou has collapsed significantly. Shoreline protection was implemented during the first phase of the channel regulation of the Jingjiang reach, but the unprotected area still exhibited a collapsing trend, collapsing approximately 360 m from 2016 to 2019 (Figure 13B). From 2003 to 2019, the head of the Tuovangshu Beach was headward erode and silted downstream at its lower end, squeezing the navigation channel, making the collapse of the Tokouyao Central Bars even more significant. From 2013 to 2017, the first phase of the channel regulation project of the Jingjiang reach was implemented, and the middle and lower sections of the Tuoyangshu Beach and the Tokouyao Central Bar were protected. However, the upper and middle section of the unprotected Tuoyangshu Beach was still in a state of erode and

receding, scouring sediment downstream which silted up the navigation channel to form a silt-navigational-obstruction body.

In 2017–2018, the riverbed scour in the Xinchang–Guchangdi reach was approximately  $792 \times 10^4$  m<sup>3</sup>. The sediment transported downstream was prone to silt in the Ouchikou reach and needed to be dredged and maintained (Fan et al., 2020). With the decrease of the split ratio of water at Ouchikou, the collapse of the shoreline of Ouchikou and the receding of the head of the beaches have been aggravated (Jia et al., 2016). The unstable boundary of the waterways has created navigation obstruction problems, the siltation body formed by erosion causing the downstream water depth to be less than 4.5 m or the width to be less than 200 m.

# Waterway Conditions in Confluence Area of Dongting Lake

From 2012 to 2020, the Xiongjiazhou–Chenglingji reach of the Dongting Lake's confluence area showed "beach erosion on the convex bank and channel siltation on the concave bank" (**Figure 14**), consistent with the distribution characteristics from 2002 to 2012 (Zhu et al., 2016; Zhu et al., 2018). Affected by the operation of the Three Gorges Reservoir, the flow redistribution of the Lower Jingjiang reach during the year ensures that the beaches on the convex bank last longer in the mainstream area than the deep channel on the concave bank, increasing the erosion power of the beaches (Han et al., 2017). The erosion of the beaches on the convex bank of the curved river section causes the bending radius to decrease, making the channel bending radius insufficient and affecting the navigation safety of



FIGURE 13 | Changes of waterway and channel conditions in the confluence area of the Tianxingzhou–Ouchikou reach; (A) Length of 4.5 m isobath in the Tianxingzhou–Ouchikou reach; (B) Transverse section changes.



FIGURE 14 | Water depth condition of the bend channel of the Xiongjiazhou–Chenglingji reach; (A) River bed erosion and deposition distribution; (B) Length of 4.5 m isobath in the Xiongjiazhou–Chenglingji reach.

ships. The Xiongjiazhou–Chenglingji reach has continuous sharp bends, and the Xiongjiazhou, Chibakou, Baxianzhou, and Guanyinzhou reaches are affected by the jacking effect of the outflow of Dongting Lake (Lai et al., 2013; Sun et al., 2017). In 2012, the 4.5 m channel at the confluence of Dongting Lake and Jingjiang reach was disconnected. After the operation of the

Three Gorges Reservoir, the river-lake hydrological jacking in the confluence area weakened (Yu et al., 2017; Liu et al., 2019; Duan et al., 2021), and in 2020, the 4.5 m channel in the confluence area of Dongting Lake and Jingjiang reach opened, only scattered shoals with a water depth of less than 4.5 m causing insufficient navigation width.

### CONCLUSION

This study took Dongting Lake and the confluence area of the middle reaches of the Changjiang River as the research object, analyzed the influence of the confluence relationship between rivers and lakes on the water depth conditions of the main channel of the Changjiang River, and further deepened the understanding of the navigation obstruction characteristics of the main channel of the Changjiang River.

After the operation of the Three Gorges Project (2003-2020), the Jingjiang reach (Zhicheng, Shashi, Jianli, Luoshan), the three outlets of Dongting Lake (Songzikou, Taipigkou, Ouchikou), and the Chenglingji Hydrological Station sediment of decreased by 79.45%-91.58%, 92.94%, and 55.8%, respectively, compared with conditions before the operation of the Three Gorges Project (1955-2002). The extent of runoff changes were -3.76 to 5.95%, -48.00%, and -13.80%, respectively, which has caused new changes in the river-lake relationship. Affected by the decrease in the volume of incoming sediment and the sediment content in the river basin, the Jingjiang reach has been eroded continuously since 2003, which has provided a favorable environment for the improvement of waterway conditions. The length of the Jingjiang reach not meeting the requirements of  $4.5 \times 200$  m was 18.4 km, accounting for 5.3% of the total length of the reach. The Songzikou, Taipingkou, Ouchikou, and outlets of Dongting Lake had a navigation obstruction length of 12.6 km, accounting for 68.35% of the of the Jingjiang reach. length The total reach (Zhicheng-Dabujie reach) affected by the first outlet of Dongting Lake (Songzikou) exhibited the "steep slope and rapid current," phenomenon, affecting the navigation safety of ships. Meanwhile, the downstream riverbed erosionr caused the water level to drop, and the waterway water depth was less than

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4.5 m. The reach affected by the second outlet (Taipingkou reach) exhibited bottomland shrinkage, frequent alternation of branch channels, and extremely unstable dry-season navigation routes. The reach affected by the third outlet (Ouchikou reach) exhibited shoreline collapse, beach erosion, and oscillating movement. Furthermore, the channel boundary was unstable, making the downstream navigation width insufficient or producing shoals with a water depth of less than 4.5 m. The reach affected by the inflow into the Dongting Lake (the Xiongjiazhou–Chenglingji reach) was affected by jacking by the backwater outflow from the lake. The slowing of the flow in this reach caused scattered shoals with a water depth of less than 4.5 m in the navigation waterway.

### **Compliance With Ethical Standards**

The contributing authors have no conflict of interest, including specific financial interests, relationships, or affiliations relevant to the subject matter or materials contained within this manuscript.

### DATA AVAILABILITY STATEMENT

The data and code for this study are available upon request from the corresponding authors.

### AUTHOR CONTRIBUTIONS

Full text written by YY; JZ put forward the research ideas; MJ, LY, WL, LZ, and others contributed to data processing and graph rendering.

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