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# Physical property differences of source rocks in the Lucaogou Formation and adsorption ability variation under their influence

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Shale gas exploration is very important for the China in the further. There are many findings of shale gas in Northwest China, especially in Junggar Basin Permian Lucaogou Formation of the Xinjiang Uygur Autonomous Region. However the basic geological characteristics of Lucaogou Formation is not still clear. In order to solve this problems, Quanzijie shale gas collected from the southern margin of the Junggar Basin is taken as the study area, and shale gas resources are investigated and evaluated using the organic-rich shale section. On this basis, the investigation and evaluation of shale gas resources in the Lucaogou Formation is performed. The reservoir performance, preservation conditions, compressibility, and gas-bearing properties of gas-bearing shale intervals are identified. The results are as follows: The vertical and horizontal lithological and lithofacies variations of the source rocks of the Lucaogou Formation are clarified. This layer is characterized by the upward thinning of grain size and the upward darkening of color in the vertical direction. The layer can be divided into upper and middle-lower sections, with organic shale being most developed in the upper section, and sandstone and dolomite being more developed in the middle and lower area. The main hydrocarbon source rocks are developed in the upper area of the Lucaogou Formation. The total thickness of the hydrocarbon source rocks is 78.12-513.73 m. The lithology is mainly gray-black shale, silty mudstone, and oil shale. The average organic carbon content is 4.86%. The maturity of the organic matter is mainly 0.7-0.8%. The shale reservoirs have a high content of brittle minerals, but the physical properties both of the shale and sandstone reservoirs are poor. These are generally ultra-low porosity and ultra-low permeability reservoirs, in which the oil and gas reservoir space is mainly fractures.

#### KEYWORDS

Junggar Basin, Lucaogou Formation, shale reservoirs, porosity-permeability, hydrocarbon source rocks

### Introduction

In the past few decades, the large-scale commercial development of shale oil and gas has changed the global energy supply structure (Chen et al., 2011; Aljamaan et al., 2017; Tang et al., 2021). The "shale oil and gas revolution" in the United States has turned it from an energy-importing country into an energy-exporting country. China is rich in shale oil and gas resources (Wang et al., 2014; Li et al., 2018; Ju et al., 2019). Continental shale, marine shale, and sea–land transitional shale are widely developed in oil- and gas-bearing basins in the northwest, north, and northeast of China, as well as in Qinghai Tibet and southern regions. This is an important field of unconventional oil and gas exploration and development in China (Gale et al., 2007; Gou., 2017; Gong et al., 2021).

Shale reservoirs are characterized by low porosity and ultralow permeability (Hou et al., 2018; Javed et al., 2019; Yuan and Rezaee., 2019). Natural fractures are widely developed in shale, which is an important reservoir space and seepage channel of shale reservoirs (Dong et al., 2010; Habina et al., 2017; Guo et al., 2019). This can significantly improve the storage and seepage capacity of shale, affecting the enrichment, preservation conditions, single-well productivity, fracturing, and other development effects of shale oil and gas, and it is an important geological factor to be considered for shale oil and gas development (Yang et al., 2016; Yan et al., 2017; Li et al., 2018; Zhang et al., 2020).

Sets of hydrocarbon source rocks are developed in the southeast margin of Junggar Basin, including Permian, Triassic, Jurassic, Cretaceous, and Paleogene. Among them, the Permian Lucaogou Formation is one of the most important hydrocarbon source rocks, and it is also the stratum with the most potential for oil and gas exploration (Jiao et al., 2014; Hu et al., 2016; Jiang et al., 2019). Until now, several industrial oil wells have been drilled in the Lucaogou Formation. Some scholars have studied the oilbearing properties, geochemical characteristics, hydrocarbon generation potential, and the mechanisms of the source rocks in this formation. The results show that this set of source rocks is a set of saline and semi-saline lacustrine sediments, with complex and diverse lithology, strong heterogeneity, wide distribution range, and large thickness, which shows good prospects for oil and gas exploration (Li et al., 2015; Liu et al., 2018).

In conclusion, the physical properties of shale reservoirs have become the main controlling factor for shale oil and gas production, and numerous research achievements have been made. However, research on the basic geological (sedimentary facies, reservoir physical properties, etc.) characteristics of shale oil and gas in the southern margin of the Junggar Basin is still insufficient, which seriously restricts the exploration and development of shale oil and gas resources in this area.

In this paper, Quanzijie shale gas collected from the southern margin of the Junggar Basin is taken as the study area, and shale gas resources are investigated and evaluated by using the organicrich shale section. First, a set of strata is divided into upper and middle-lower sections, the basic characteristics of the source rocks are described in detail, and the shale facies type is identified by using rock type and TOC content. Second, the spatial distribution differences of the hydrocarbon source rocks are studied. Third, the type of organic matter, reservoir physical properties, and pore structure of different rock types and layers are compared, and differences in the organic geochemical characteristics are analyzed. On this basis, the physical property differences of source rocks in the Lucaogou Formation and dynamic variation of free surface energy under their influence is studied. These results can provide a theoretical basis for the efficient development of shale oil and gas in the southern margin of the Junggar Basin.

# Geological setting and sample preparation

### Geological setting

The tectonic environment of the study area in this paper is located in the eastern segment of the Dahuangshan–Dalongkou reversed turn. Owing to the influence of the Bogda Mountain front thrust belt, a series of tectonic deformations dominated by compressive stress has developed in this study area. Crumpled deformation formed by south to north napping can be seen in the Lucaogou Formation, where hydrocarbon source rocks are developed. This area is located in the east section of the south Junggar thrust belt. Regional and secondary faults are relatively developed. Generally, there are two large folds along the back, namely, the Dalongkou Syncline and the West Dalongkou Anticline. The hydrocarbon source rock horizon in this area is mainly the Lucaogou Formation, which can be mainly divided into the upper and middle-lower sections (Zhao et al., 2017; Qiao et al., 2020).

The upper section of this formation is composed of grayish black carbonaceous shale, silty mudstone, oil shale, and multilayer dolomitic limestone and marlstone. The lithology of the upper section of the north wing of the Dalongkou Syncline in the area changes continuously, but the sand content in the upper section of the formation is significantly increased near the core of the Dalongkou Syncline and the Bogda Mountain front area in the south wing. The lithology is black-gray argillaceous siltstone, siltstone, and even fine sandstone, and the shale content is lower (Figure 1).

### Sample and experimental methods

Shale and sandstone samples were collected from the upper and middle-lower sections of the Lucaogou



Formation for experimental testing. The organic geochemical characteristics and mineral composition of all the samples were analyzed. Then, the morphological characteristics of all the samples were observed by FE-SEM and quantitatively characterized by PCAS software. Subsequently, pore volume, specific surface area, and pore size distribution of the pores were quantitatively studied by  $N_2$  adsorption.

#### **TOC** experiments

All the samples were manually selected, manually ground to <60 meshes, and weighed 10 g. According to the national standard (GB/T 19145-2003), they were tested in the Jiangsu Institute of Geology and Mineral Resources Design. In the XRD experiment, the clay test sample was artificially broken into 60–80 meshes, weighing 200 g, and the whole rock test sample was broken into small pieces weighing 30 g, ground completely to <200 meshes, which was completed in Sinopec East China Branch's test center. The instrument used was Ultima IV, and the experiments were conducted according to the industry standard SY/T 5163-2010.

#### **FE-SEM** experiment

Small samples were selected after argon ion polishing, after which a Quanta250 scanning electron microscope was used for testing and analysis. The FE-SEM photos taken were quantitatively analyzed by PCAS software, which is a professional software developed by Liu et al. (2011) to identify the micro pore structure in high-resolution SEM photos. This technique has been successfully applied to the quantitative identification of clay mineral micro-fractures and organic-rich shale micro pores.

#### Isothermal adsorption test

The instrument used in this experiment was a IS-100 high-pressure gas isothermal adsorption and desorption instrument. In strict accordance with the GB/T 19560-2008 Test Method for High Pressure Isothermal Adsorption of Coal, the sample was ground to 60-80 meshes, and a 100-g sample was taken and put it into the vacuum drying oven. The coal sample was dried at 373K for 6 h. Then, the sample was cooled to room temperature and put into the coal sample tank for evacuation. Methane adsorption experiments were carried out at ambient temperature (30°C) and an equilibrium of 0–15 MPa. pressure range The adsorption equilibrium time at each pressure point was greater than 12 h.

### **Results and discussion**

# Detailed description and difference comparison of lithofacies types

Outcrop observation and drilling revealed that the typical sedimentary rocks of the Lucaogou Formation  $(P_3l)$  in this area are gray shale, oil shale, siltstone with dolomite, limestone lens, and a small amount of sandstone lens. According to the differences in the lithology, sedimentary structure, contents, mineral components, and logging curve of the sediments, the



FIGURE 2

Thin section photos of black oil shale samples (A), Asphaltic concretion; (B-D) Muddy structure and layered structure can be seen under the microscope

oil shale of the Lucaogou Formation in the study area is divided into five lithologic facies: black oil shale/shale facies; gray-black/dolomitic mudstone facies; gray-white argillaceous dolomite/limestone facies; gray-white micritic dolomite/micritic limestone facies; and granular limestone/ dolomite facies. Among them, black oil shale/shale facies and gray-black calcareous/dolomitic mudstone facies are the main lithofacies types.

#### Black oil shale/shale facies

Black oil shale/shale facies are the most important lithologic facies in this study area. The main lithology is black/black-brown oil shale and gray-black shale. Many shales are intercalated with siltstone laminae or bands (mm). The siltstone can be seen in the liquefaction deformation structure, which may be turbidite sandstone at the far end of the semi-deep lake. The most intuitive features of this phase are the black color and prominent page structure.

The weathered color of the oil shale is mainly light bluegray, gray-white, etc. The fresh surface is mainly black. After weathering, it presents an obvious book-like page texture. The fresh rock has thick foliation and common horizontal bedding. It contains fish (ancient cod, complete or fragmented), abundant phyllopods, bivalves, ostracods, and plant fossils preserved in situ. There are cubic or globular pseudomorphs of luteite with punctate ferritization and gypsum snowflake crystals (secondary changes) and bitumen-bearing nodules

(Figure 2A). In this section, microscopically, the argillaceous structure and the layered structure can be seen (Figure 2).

#### Lime/dolomitic mudstone facies

Compared with the black oil shale/shale lithologic facies, the most intuitive sign of this facies is grayish black with undeveloped foliation, mostly in a transitional relationship with the black shale facies. The lithology of this lithofacies includes calcareous mudstone, dolomitic mudstone, and silty mudstone. It is mostly mixed with siltstone laminations or bands.

The weathered color of the mudstone is mostly grayish white, brown-gray, and other colors. The fresh surface is grayish black, with a massive structure, and the section is mostly shell-shaped. Lime shales mixed with thin siltstone laminations or bands are common, and liquefaction deformation structures are found in the siltstone (Figure 3).

### Spatial distribution differences of the hydrocarbon source rocks

Figure 5 shows that the hydrocarbon source rocks of  $P_3l$  in the study area are distributed in the upper section of the Lucaogou Formation vertically. The thickness of the upper section is 185.22-522.72 m. The lithology of the hydrocarbon



#### FIGURE 3

Thin section photo of grayish black silty banded dolomitic mudstone rock (A), Lime mudstone; (B), Dolomitic mudstone; (C), Silty mudstone; (D), mudstone rock.









source rocks is gray-black shale, silty mudstone, and black oil shale. The total thickness of the upper section is 78.12–513.73 m. The proportion of hydrocarbon source rocks in the upper section of the formation is 36%–98%, with an average of 85%. The lithology of the non-hydrocarbon source rocks in the upper section is gray dolomite, argillaceous

dolomite, argillaceous siltstone, and a small amount of argillaceous silty limestone, with a total thickness of 2.92–136.46 m, ranging from 1 to 19 layers. The middle-lower section is composed of gray-white silty fine sandstone and dolomitic siltstone, mixed with a thin layer of dark mudstone, and the source rock is not developed.



syncline are exposed to the surface, and their thickness is centered around the Xiaolongkou area, showing a trend of gradual thinning around. The core and south wing of the syncline are covered by the Quaternary system. The thickness of the hydrocarbon source rocks controlled by drilling is much smaller than that of the north wing of the syncline. The thickness of the hydrocarbon source rocks changes: in the south wing of the syncline it is less thick than in the core of the syncline, while it is at its thickest in the north wing of the syncline.

In the study area, the thickness variation of the hydrocarbon source rocks in the upper section is larger. The thickness of the hydrocarbon source rocks in the area of Dalongkou–Gangou in the west decreases gradually toward the surrounding area, while the hydrocarbon source rocks in the area of the Quancan 1 Well are



The results show that the hydrocarbon source rocks in the upper section are basically distributed around the Dalongkou Syncline. The hydrocarbon source rocks in the north wing of the thinner, and thickness of the hydrocarbon source rocks in the area of Well Qi1 increases gradually toward the northeast (Figure 4).



#### FIGURE 10

Fracture type of all samples: (A) dissolution fissure in the upper section; (B) straight gap between mineral particles of sandstone in the lower section; (C) micro-fracture of sandstone in the upper section; (D) micro-fracture of sandstone in the upper section.



## Comparison of shale rock density and organic carbon content in the upper part of the Lucaogou Formation.

# Difference comparison of organic geochemical characteristics

#### Type of organic matter

The organic macerals of kerogen are mainly sapropel amorphous and humic amorphous. The sapropel formation is



composed of sapropel amorphous and algal bodies, with a relative abundance of 0%–86%, with an average of 49.91%. Most of the non-resinous bodies are cuticle and chitin, with a relative abundance of 0%–87%, with an average of 18.91%. The vitrinite consists of structural vitrinite and unstructured vitrinite, with a relative abundance of 5%–79%, with an average of 26.7%.



Distribution frequency of the specific surface area in different intervals of the upper shale reservoirs.

sapropel humic type. Vertically, hydrocarbon protoliths of the upper section are type  $II_1$ , type  $II_2$ , and a small amount of type III, while the middle–lower section is dominated by type  $II_1$ .

#### Organic matter content (TOC)

The organic carbon content of the source rocks is generally higher. The surface and drilling revealed that the TOC content is 0.5%-27.48%, with an average of 4.31%. Based on 22 samples of non-hydrocarbon source rocks, the TOC content was found to be 0.06%-0.48%, with an average of 0.33%, and the TOC content of source rock was mainly distributed between 0.6% and 7.0% (Figure 5). According to distribution of the TOC content (Figure 8), the higher TOC value area of the hydrocarbon basically distributed source rocks is around the basin-mountain junction in the north of the study area,



surface area vs mineral content)

The relative abundance of the inert group is 2%-53%, with an average of 10.48%. The kerogen is light-yellow-brown, with no weak fluorescence, and the kerogen type index is -82 to 81.

The results show that the types of kerogen in the source rocks are type II<sub>1</sub>, type II<sub>2</sub>, and type III, with few type I. Type II is the main type, accounting for 89% in total. This shows that the main type of kerogen in the survey area is sapropel type, followed by

indicating a decreasing trend from north to south. The higher organic carbon content area is concentrated in the Dalongkou–Gangou area in the west, with an average organic carbon content of more than 6%. The average organic carbon content in the area from Heijiawan to Dongdalongkou is 3%–4%, while the average organic carbon content in the area from Quancan Well 1 to Zhunye Well 4 in the south of the survey area is 0.5%–2%.





Relationship between the porosity and permeability of shale

reservoirs in the middle–lower section. The blue line is porosity and red line is permeability.



In conclusion, the hydrocarbon source rock abundance in the north of this study area is better than that in the south. From the tectonic location of the source rocks, the hydrocarbon source rock abundance in the north wing of Dalongkou Syncline is higher than that in the core and south wing of the syncline. This is because the northern area was a deep-water area of the lake basin during the deposition, and the water body gradually became shallow toward the south. Bogda Mountain provided the main material source for the formation of hydrocarbon source rocks in the area during the Lucaogou Formation of the Late Permian, and the TOC content was mainly affected and controlled by the sedimentary environment. The deeper the sediment body, the higher its TOC content will be.

#### Maturity

The results show that the  $R_{o, max}$  of the hydrocarbon source rocks is 0.53%–1.24%, with an average of 0.78%. It generally follows normal distribution characteristics. The main body belongs to the mature stage, the second belongs to the lower-maturity stage, and a small amount is distributed in the immature stage, with a low degree of thermal evolution (Figure 6).

# Adsorption of reservoirs in the Lucaogou Formation

#### Petromineralogical characteristics Clay minerals

The maximum clay mineral content (59%) of the shale reservoir samples in the upper section is slightly higher than the maximum clay content (38%) in the middle and lower sections, while the minimum clay content (4%) is equivalent to that of the samples in the middle and lower sections (3%), and the average clay content (28%) in the upper section is higher than that of the samples in the middle and lower sections (16%). The clay mineral composition of the shale reservoirs in the upper and middle-lower sections is mainly illite, followed by the illite/ montmorillonite interbeds, while kaolinite, chlorite, and green/montmorillonite interbeds are relatively small. In the clay composition of the upper section, the average illite content (62%) is significantly higher than average illite content (41%) of the lower-middle section (Figure 7). In both



the upper sample and the middle-lower sample, the illite/ montmorillonite interlayer in the clay mineral components decreases with the increase of illite content.

#### Brittle minerals

The brittle minerals of the upper and lower sections are mainly quartz, feldspar, calcite, and dolomite. Among them, the average content of quartz and feldspar in the upper section (60%) is close to that in the middle–lower section (62%), the average content of carbonate rock (10%) is lower than that in middle–lower section (20%), and the average content of brittle minerals (71%) is lower than that in the middle–lower sample (84%) (Figure 8).

#### Type of reservoir space

#### The combination of macro and micro

Based on the core observation and scanning electron microscope results, it can be seen that the reservoir space types of this formation are diverse, mainly including macro-scale fractures and pores, as well as micro-scale pores and fractures.

#### Pores

The pores are mainly intergranular pores, intragranular pores, intergranular gaps, and mold holes at the micro-scale

(Figure 9). All kinds of pores can form good self-sourced oil and gas reservoir spaces.

The fractures include macro-scale fractures, dissolution fractures, sutures and micro-scale fractures, lamina fractures, dissolution fractures, intergranular fractures, and internal fractures of clay minerals (Figure 10). For fluids, fractures play a dual role, which can be both a fluid migration channel and an effective reservoir space. It can not only destroy the oil and gas reservoir, resulting in the loss of oil and gas, but can also improve the reservoir pore connectivity and permeability.

#### Reservoir physical properties and pore structure Shale reservoirs in the upper section

The rock density of the shale reservoir has a significant negative correlation with organic carbon content (TOC), that is, low TOC content represents high rock density (Figure 11). Therefore, it can be concluded that rock density logging data can be used to identify carbonaceous shale intervals during drilling.

As shown in Figure 12, permeability tends to increase with the increase of porosity, but the correlation between porosity and permeability is not obvious. On the contrary, some low porosity samples have high permeability values, indicating that their permeability is obviously affected by the micro-fractures. Some



0.6 Sample 1 Sample 2 Total pore volume (cm<sup>3</sup>.g<sup>-1</sup>) 70 F Sample 3 Sample 4 Sample 5 4 12 16 20 Pressure (MPa) 24 28 32 8 FIGURE 20 Isothermal adsorption simulation experiment curve for Zhunye Well 4

higher porosity samples have lower permeability values, indicating poor pore connectivity.

The specific surface area test results show that the minimum value of the specific surface area in the upper section is  $0.12 \text{ m}^2/\text{g}$ , the maximum value is  $14.2 \text{ m}^2/\text{g}$ , and the average value is  $4.08 \text{ m}^2/\text{g}$ , mainly distributed in the range of  $0-6 \text{ m}^2/\text{g}$  (Figure 13).

With the increase of the specific surface area, the porosity generally has an upward trend, but the correlation between them is general. The specific surface area is negatively correlated with the percentage of TOC content, and it is approximately positively correlated with the percentage of clay minerals (Figure 14).

# Sandstone reservoirs in the upper part of the Lucaogou Formation

Figure 15A shows that permeability tends to increase with the increase of porosity, but the correlation between porosity and permeability is not obvious. On the contrary, some low porosity samples have high permeability values, indicating that their permeability is obviously affected by the microfractures in the samples. Figure 15B shows that the minimum and maximum specific surface areas of the sandstone reservoirs in the upper part of Lucaogou Formation are 0.951 and 5.75 m<sup>2</sup>/g, respectively, with a main distribution range of 3–6 m<sup>2</sup>/g. As the specific surface area increases, the porosity has an overall upward trend.

#### Shale reservoirs in the middle--lower section

Figure 16 shows that the permeability tends to increase with the increase of porosity, but the correlation between porosity and permeability is general. This indicates that the minimum specific

surface area of the shale reservoirs in the middle–lower section is  $0.162 \text{ m}^2/\text{g}$ , and the main distribution range is  $0-1 \text{ m}^2/\text{g}$  (Figure 17).

#### Sandstone reservoirs in the middle-lower section

Figure 18A shows that there is an upward trend with permeability, but the correlation between porosity and permeability is not obvious. On the contrary, lower porosity samples have high permeability values, which indicates that permeability is obviously affected by micro-fractures in the samples. The minimum and maximum specific surface areas of sandstone reservoirs are 5.75 and 2.056 m<sup>2</sup>/g, respectively, with a main distribution range of  $0-2 \text{ m}^2/\text{g}$  (Figure 18B). Figure 18C shows that porosity has an overall upward trend with the increase of the specific surface area. Figure 18D shows that the pore diameter interval distributions in the samples which reflect the statics data are effective.

#### Physical properties of reservoir rocks

Figure 19A shows that the porosity of the shale reservoirs in the upper section is greater than that in the middle-lower section. The porosity of the upper sandstone reservoirs is less than that of lower sandstone reservoirs. The porosity of the sandstone reservoirs is greater than that of shale, and the porosity of the sandstone reservoirs in the middle-lower section is greater than that of mudstone. Figure 19B shows that the permeability of the shale and sandstone reservoirs in the upper section is higher than that in the middle-lower section. The permeability of the sandstone reservoirs is higher than that of shale. Figure 19C shows that the specific surface areas of sandstone and shale are similar, and the specific surface areas in the upper section are larger than those in the middle-lower section. According to the scanning electron microscope statistical analysis results (Figure 19D), the average pore length and diameter of sandstone rocks are larger than that of shale.

# Adsorption of hydrocarbon source rocks in the Lucaogou Formation

In this paper, high-temperature isothermal adsorption tests were carried out on 13 samples to characterize the adsorption of organic-rich shale in the study area (Figure 20). The test conditions comprised a laboratory temperature of 30°C, a helium concentration of 99.999%, and a methane concentration of 99.99%. These results show that the average Langmuir volume  $V_{\rm L}$  is 2.93 m<sup>3</sup>/t and the average Langmuir pressure  $P_{\rm L}$  is 7.83 MPa. In general, the larger the Langmuir volume, the stronger the adsorption capacity of mudstone. The average Langmuir volume  $V_{\rm L}$  of siltstone fine sandstone is 0.89 m<sup>3</sup>/t, which is much smaller than that of mudstone, and the average Langmuir pressure  $P_{\rm L}$  is 14.8 MPa.

## Conclusion

In this paper, Quanzijie shale gas collected from the southern margin of the Junggar Basin was taken as the study area, and shale gas resources were investigated and evaluated by using the organic-rich shale section. On this basis, the investigation and evaluation of shale gas resources in the Lucaogou Formation was performed. The reservoir performance, preservation conditions, compressibility, and gas-bearing properties of gas-bearing shale intervals were identified. The results are as follows.

The vertical and horizontal lithological and lithofacies variations of the source rocks of the Lucaogou Formation were clarified. This layer is characterized by the upward thinning of grain size and the upward darkening of color in the vertical direction. The layer can be divided into upper and middle–lower sections; organic shale is most developed in the upper section, and sandstone and dolomite are more developed in the middle–lower section.

The main hydrocarbon source rocks are developed in the upper section of the Lucaogou Formation. The total thickness of the hydrocarbon source rocks is 78.12–513.73 m. The lithology is mainly gray-black shale, silty mudstone, and oil shale. The average organic carbon content is 4.86%. The maturity of the organic matter is mainly 0.7–0.8%. The shale reservoirs have a high content of brittle minerals, but the physical properties both of the shale and sandstone reservoirs are poor. They are generally ultra-low porosity and ultra-low permeability reservoirs, in which the oil and gas reservoir space is mainly fractures.

### Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

## Author contributions

DG is responsible for writing. DL designed the ideas. PY is responsible for drawing. SW is responsible for analyzing the data. JL is responsible for analyzing the data. GL is responsible for analyzing the data.

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## **Conflict of interest**

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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