



Comparative Assessment and Future Prediction Using CMIP6 and CMIP5 for Annual Precipitation and Extreme Precipitation Simulation

Jingjing Li[†], Ran Huo[†], Hua Chen^{*}, Ying Zhao and Tianhui Zhao

State Key Laboratory of Water Resources and Hydropower Engineering Science, Wuhan University, Wuhan, China

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*Correspondence:

Hua Chen
chua@whu.edu.cn

[†]These authors have contributed
equally to this work

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This study assesses the improvement of the latest Coupled Model Intercomparison Project Phase 6 (CMIP6) over Coupled Model Intercomparison Project Phase 5 (CMIP5) for precipitation simulation. Precipitation simulations under different future climate scenarios are also compared in this work. The results show that: 1) CMIP6 has no overall advantage over CMIP5 in simulating total precipitation (PRCPTOT) and maximum consecutive dry days (CDD). The performance of CMIP6 increases or decreases regionally in PRCPTOT and consecutive dry days. But it is slightly worse than CMIP5 in simulating very wet days (R95pTOT). 2) Comparing the trend test results of CMIP5 and CMIP6 in the future, there are more areas with significant trend based on Mann–Kendall test in CMIP6 compared with that of CMIP5. The differences in PRCPTOT are mainly found in Amazon Basin and Western Africa. The differences between the R95pTOT trends mainly noticeable in South America and Western Africa, and the differences in CDD are mainly reflected in Central Asia, Sahara Desert and central South America. 3) In Southern South America and Western North America, the PRCPTOT changing rate of CMIP6 in the future under various scenarios is always greater than that of CMIP5; in Alaska, Western Africa, Southern Africa, the PRCPTOT changing rate of CMIP6 in the future under various scenarios is always less than that of CMIP5. In Southern South America, the R95pTOT changing rate of CMIP6 in the future under various scenarios is always greater than that of CMIP5; in Alaska, East Asia, North Asia, the R95pTOT changing rate of CMIP6 in the future under various scenarios is always less than that of CMIP5. In almost half of the regions, the CDD changing rate of CMIP6 is less than that of CMIP5 under all scenarios, namely Australia, Amazon Basin, Southern South America, Central America, Western North America, Central North America, Eastern North America, Central Asia, Tibet.

Keywords: CMIP6, precipitation, worldwide, general circulation models, regions

INTRODUCTION

To analyze the impact of future climate change, the World Climate Research Program (WCRP) initiated the Coupled Model Intercomparison Project (CMIP) in 1995 (Meehl et al., 2000). With scientific development and the gradually improved understanding of climate change mechanisms, the CMIP has evolved from CMIP Phase 1 (CMIP1), Phase 2 (CMIP2), Phase 3 (CMIP3), and Phase 5 (CMIP5) to the latest Phase 6 model (CMIP6). CMIP formulates climate model test standards, and

the mechanism for sharing simulated climate data promotes model development and improvement. Therefore, it has gradually developed and become an indispensable element of climate science. Many studies based on CMIP data provided support for the Intergovernmental Panel on Climate Change (IPCC) assessment reports (IPCC 2007; IPCC 2013). CMIP6 will provide data support for climate change research in the next five to ten years.

CMIP6 newly proposed shared socioeconomic pathways (SSPs) (Moss et al., 2010), which describe different socioeconomic reference assumptions. The SSPs and RCPs (representative concentration pathways) were combined to provide an integrated scenario in CMIP6 (Van Vuuren et al., 2014). In addition to the new future scenario, a new generation of general circulation models (GCMs) is used in CMIP6 (Veronika et al., 2016). The GCM is an indispensable tool for exploring the interaction of climate systems, future climate change predictions, and climate change research. It has been proven to behave well in simulating climate change and can be used to predict future precipitation under climate change (Li et al., 2019; Khan et al., 2020).

Climate change significantly affects the water cycle (Madsen et al., 2014; Zobel et al., 2018; Ortiz-Gómez et al., 2020). Obtaining high-quality water resource forecast data under climate change is critical for sustainable water resource development. In addition to the total amount of water resources, the temporal and spatial distribution of precipitation, especially extreme precipitation events, is worthy of attention. Extreme precipitation can lead to drought and flood events (Ulbrich et al., 2003; Wang et al., 2016; Funk et al., 2018; Funk et al., 2019; Lal et al., 2020), which in turn will cause enormous societal, economic, and ecological losses (Kundzewicz et al., 2014; Gao et al., 2019). To manage the hidden dangers of water security caused by global climate change and deploy adaptive countermeasures in advance, it is crucial to study the total precipitation and extreme precipitation events under climate change (Birkmann 2011; Forestieri et al., 2018; Rahmani and Harrington 2019; Hosseinzadehtalaei et al., 2020; Ukkola et al., 2020). CMIP result data provide a meaningful way to analyze the temporal and spatial distribution of future precipitation.

With the further development of CMIP, studying the difference between CMIP5 and CMIP6 has received critical attention. Several studies have assessed CMIP6 by comparing the historical simulation data of CMIP5 and CMIP6 with the observed data (Gusain et al., 2020; Yuanhai et al., 2020; Zamani et al., 2020; Zhu et al., 2020; Zhu and Yang 2020). When CMIP5 and CMIP6 are compared, the multi-model mean result (Xin et al., 2020; Yuanhai et al., 2020) rather than a specific GCM (Wyser et al., 2020) is used in most cases. Gusain et al. (2020) compared CMIP6 and CMIP5 when simulating Indian summer monsoon precipitation. The results showed that CMIP6 performed better than CMIP5, as the model deviation in CMIP6 was significantly reduced. Thus, CMIP6 could better simulate the Indian monsoon characteristics. Zhu et al. (2020) compared CMIP5 and CMIP6 in China based on a daily observational dataset. They found that CMIP6 behaved better in the simulation of total precipitation, precipitation intensity, heavy precipitation, and extremely heavy rain days, except for consecutive dry days (CDD). For precipitation, the simulation of precipitation intensity was better in CMIP6, although the simulation of precipitation frequency was comparable to that in

CMIP5. Overall, most studies have focused on a specific region (Bracegirdle et al., 2020; Davy and Outten 2020; Jiang et al., 2020) rather than the world when assessing historical precipitation in CMIP6 (Fan et al., 2020; Ukkola et al., 2020). Compared to research that evaluated historical data in CMIP5 and CMIP6, those with further analysis of future climate change in CMIP5 and CMIP6 (Chen et al., 2020; Ukkola et al., 2020) are sparse. Ukkola et al. (2020) found that drought changes in the future (2051–2100) in CMIP6 are more considerable and more consistent than those in CMIP5. The results also showed that mean precipitation and variability influenced drought duration and frequency. Chen et al. (2020) studied precipitation and temperature extremes in the future after a historical data assessment. The work highlighted the changing rate of climate indicators in the last 2 decades of this century compared with the historical record. The results suggested that extreme climate changes are more pronounced in CMIP6 than in CMIP5. Davy and Outten (2020) investigated the model distributions of surface air temperature and sea ice extent and volume in the Arctic until 2100 and compared CMIP6 with CMIP5. The results showed that under SSP126, the Arctic climate will stabilize by 2060. Most research only consider several decades at the end of the 21st century when analyzing the future, without a detailed analysis of the near future. Very little work has been devoted to the analysis of future precipitation projected by CMIP6 worldwide after comparing it with CMIP5.

More work is needed to analyze worldwide precipitation based on CMIP6 and CMIP5 in detail. This study has two objectives: 1) to compare the historical precipitation simulation capability between CMIP6 and CMIP5; and 2) analyze the differences in the temporal and spatial distributions of precipitation in the future in CMIP6 and CMIP5. In the middle of this century and at the end of this century, the impact of climate change will be very different. Therefore, when analyzing future changes in precipitation, this study divides the future into near-future and far-future periods. Three indicators are used in this study: total precipitation (PRCPTOT), very wet days (R95pTOT) and CDD.

The remainder of this paper is organized as follows. *Methods and Materials* describes the methods and materials used in this study. *Results and Discussions* presents the study results, focusing on four key themes. *Evaluation of CMIP5 and CMIP6 Historical Precipitation Simulation Capabilities* presents the evaluation of historical precipitation simulation capabilities. Then, trend analysis and comparison of long series precipitation are presented in *Trend analysis and comparison of long series precipitation in CMIP5 and CMIP6*, followed by a comparison of the changing rates of CMIP5 and CMIP6 precipitation in the future compared with history in *Changing rate comparison of precipitation in the future over the history*. Finally, the model uncertainties in CMIP5 and CMIP6 are compared in *Comparison of Model uncertainty in CMIP5 and CMIP6 precipitation*. Conclusions are drawn in *Conclusion*.

METHODS AND MATERIALS

Data

In this study, the Global Precipitation Climatology Center (GPCC) V 2018 grid daily precipitation data were used as the

TABLE 1 | List of general circulation models (GCMs) in the Coupled Model Intercomparison Project phases 5 and 6 (CMIP5 and CMIP6).

CMIP6		CMIP5		Institution
Model	Resolution (°) Lat*Lon	Model	Resolution (°) Lat*Lon	
BCC-CSM2-MR	1.12 × 1.13	BCC-CSM1.1	2.79 × 2.81	BCC
CESM2	0.94 × 1.25	CCSM4	0.94 × 1.25	NCAR
FGOALS-g3	2.0 × 2.81	FGOALS-g2	2.78 × 2.81	CAS
GFDL-CM4	1.00 × 1.25	GFDL-CM3	2.0 × 2.5	NOAA-GFDL
GFDL-ESM4	1.00 × 1.25	GFDL-ESM-2G	2.02 × 2.0	NOAA-GFDL
INM-CM4-8	1.5 × 2.0	INM-CM4	1.5 × 2.0	INM
IPSL-CM6A-LR	1.27 × 2.5	IPSL-CM5A-LR	1.89 × 3.75	IPSL
MRI-ESM2-0	1.12 × 1.13	MRI-ESM1	1.12 × 1.13	MRI

TABLE 2 | Indicators used in this study.

Indicator	Acronym	Definition	Units
Total precipitation	PRCPTOT	Let R_{wj} be the daily precipitation amount for day w of period j . Then the total precipitation in period j is $\text{PRCPTOT}_j = \sum_{w=1}^W R_{wj}$	mm
Very wet days	R95pTOT	Let R_{wj} be the daily precipitation amount for wet day ($R > 1$ mm) of period j and $R95p$ the 95th percentile of precipitation for wet days in the specified period. Then $R95pTOT$ is determined as $R95pTOT_j = \sum_{w=1}^W R_{wj}, R_{wj} > R95p$	mm
Maximum consecutive dry days	CDD	Let R_{ij} be the daily precipitation amount for day i of period j . Then count the largest number of consecutive days where $R_{ij} < 1$ mm	day

observation data (1982–2005) (Fuchs et al., 2009), with a resolution of $0.5^\circ \times 0.5^\circ$. Eight GCMs of CMIP6 and eight corresponding GCMs in CMIP5 were adopted for the modeled data. The specific GCM information is presented in **Table 1**. We selected paired GCMs in CMIP5 and CMIP6. For example, the FGOALS-g3 in CMIP6 and FGOALS-g2 in CMIP5. It can reflect the difference between CMIP6 and CMIP5.

Three corresponding future scenarios from CMIP6 and CMIP5 were used to explore how forcing level and socioeconomic development affect the temporal and spatial distribution of precipitation. Those scenarios are SSP126 (RCP26), SSP245 (RCP45), and SSP585 (RCP85), where radiative forcing stabilizes at approximately 2.6 W/m^2 , 4.5 W/m^2 , and 8.5 W/m^2 in 2100, respectively. Due to historical simulation data limitations, the historical period was from 1982 to 2005 (24 years). The future is divided into near-future period (2037–2060, 24 years), and far-future period (2077–2100, 24 years) in this study.

Because the resolution of each GCM is different, all GCM and observation data were interpolated into $1^\circ \times 1^\circ$ grid data sets using the bilinear interpolation method.

Indicators

Three precipitation indices were selected to describe the precipitation characteristics from the 27 climate change indicators recommended by the World Meteorological Organization (WMO) and the Expert Team on Climate Change Detection, Monitoring Indices (ETCCDMI) (Zhang et al., 2011): PRCPTOT, R95pTOT, and CDD (**Table 2**). It helps to understand the total amount of water resources and the risks of drought and flooding.

In this study, the root mean square error (RMSE) was used as the criterion for evaluating modeled precipitation.

$$\text{RMSE} = \sqrt{\frac{1}{n} \sum_{i=1}^n (X_i - Y_i)^2} \quad (1)$$

where X represents the GCM precipitation, Y represents the observation value, and n represents the series length, i represents the number of the variable.

The Mann–Kendall (MK) Trend Test Method

MK trend test method is used to analyze hydrological series trends. The MK trend test method is widely used because of its simplicity (Mann 1945; Kendall 1990; Hamed 2008). When the standard normal statistics Z is beyond the threshold of $[-2.32, 2.32]$, the hydrological series has a significance changing rate at the 99% confidence level. The standard normal statistics Z is calculated as follow:

$$Z = \begin{cases} (S-1)/\sqrt{\text{Var}(S)}, & S > 0 \\ 0, & S = 0 \\ (S+1)/\sqrt{\text{Var}(S)}, & S < 0 \end{cases} \quad (2)$$

where S is a statistic,

$$S = \sum_{i=1}^{n-1} \sum_{j=i+1}^n \text{sgn}[x(j) - x(i)] \quad (3)$$

where n represents the series length, i, j represents the number of the variable, sgn is calculated as:

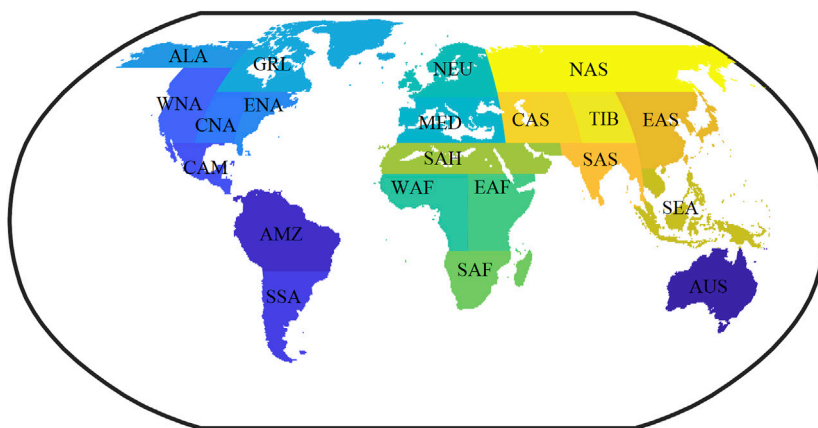


FIGURE 1 | Regions used in this study.

$$\text{sgn}[x(i) - x(j)] = \begin{cases} 1, & x(i) > x(j) \\ 0, & x(i) = x(j) \\ -1, & x(i) < x(j) \end{cases} \quad (4)$$

Sen's Slope Estimator Test

We use Sen's slope estimator (Swain 2014; Verma et al., 2016) to detect the linear trend of precipitation in this study. The size of the trend is given as:

$$\beta = \text{Median}\left(\frac{x_i - x_j}{i - j}\right), \quad \forall j < i \quad (5)$$

where β is the size of the trend of data series. Median means the median of the data series.

Regions

There are considerable differences in the climate and geographical characteristics across the world. Therefore, regional assessments are more meaningful for practical guidance. The following factors should be considered when choosing regions (Giorgi and Francisco 2000): 1) the shape of the regions should be simple; 2) within a region, the climate and geographical characteristics should be similar; and 3) the regions should cover all the land except Antarctica. Based on the above standards, the world can be divided into 21 regions (Giorgi and Francisco 2000) (Figure 1). The 21 regions include Australia (AUS), Amazon Basin (AMZ), Southern South America (SSA), Central America (CAM), Western North America (WNA), Central North America (CNA), Eastern North America (ENA), Alaska (ALA), Greenland (GRL), Mediterranean Basin (MED), Northern Europe (NEU), Western Africa (WAF), Eastern Africa (EAF), Southern Africa (SAF), Sahara (SAH), Southeast Asia (SEA), East Asia (EAS), South Asia (SAS), Central Asia (CAS), Tibet (TIB), and North Asia (NAS). Table 3 provides detailed information on the 21 regions.

TABLE 3 | List of regions used in this study. Only land grid points are used in the analysis.

Name	Acronym	Latitude (°)	Longitude (°)
Australia	AUS	45 ± 11 S	110 E ± 155 E
Amazon Basin	AMZ	20 S ± 12 N	82 ± 34 W
Southern South America	SSA	56 ± 20 S	76 ± 40 W
Central America	CAM	10 ± 30 N	116 ± 83 W
Western North America	WNA	30 ± 60 N	130 ± 103 W
Central North America	CNA	30 ± 50 N	103 ± 85 W
Eastern North America	ENA	25 ± 50 N	85 ± 60 W
Alaska	ALA	60 ± 72 N	170 ± 103 W
Greenland	GRL	50 ± 85 N	103 ± 10 W
Mediterranean Basin	MED	30 ± 48 N	10 W ± 40 E
Northern Europe	NEU	48 ± 75 N	10 W ± 40 E
Western Africa	WAF	12 S ± 18 N	20 W ± 22 E
Eastern Africa	EAF	12 S ± 18 N	22 ± 52 E
Southern Africa	SAF	35 ± 12 S	10 W ± 52 E
Sahara	SAH	18 ± 30 N	20 W ± 65 E
Southeast Asia	SEA	11 S ± 20 N	95 ± 155 E
East Asia	EAS	20 ± 50 N	100 ± 145 E
South Asia	SAS	5 ± 30 N	65 ± 100 E
Central Asia	CAS	30 ± 50 N	40 ± 75 E
Tibet	TIB	30 ± 50 N	75 ± 100 E
North Asia	NAS	50 ± 70 N	40 ± 180 E

RESULTS AND DISCUSSIONS

Evaluation of CMIP5 and CMIP6 Historical Precipitation Simulation Capabilities

PRCPTOT, R95pTOT, and CDD were selected to evaluate the agreement between the simulated GCMs and the observed data in the historical period. The RMSE was used to characterize the simulation capability; the lower the RMSE, the better the simulation capability. The global distribution maps of the multi-model mean RMSE of the three indicators in the historical period for CMIP5 and CMIP6 are shown in **Supplementary Figures S1–S3**. Global distribution maps of the RMSE difference between CMIP5 and CMIP6 are shown in **Figure 2**, calculated by the RMSE of CMIP5 minus the RMSE of CMIP6.

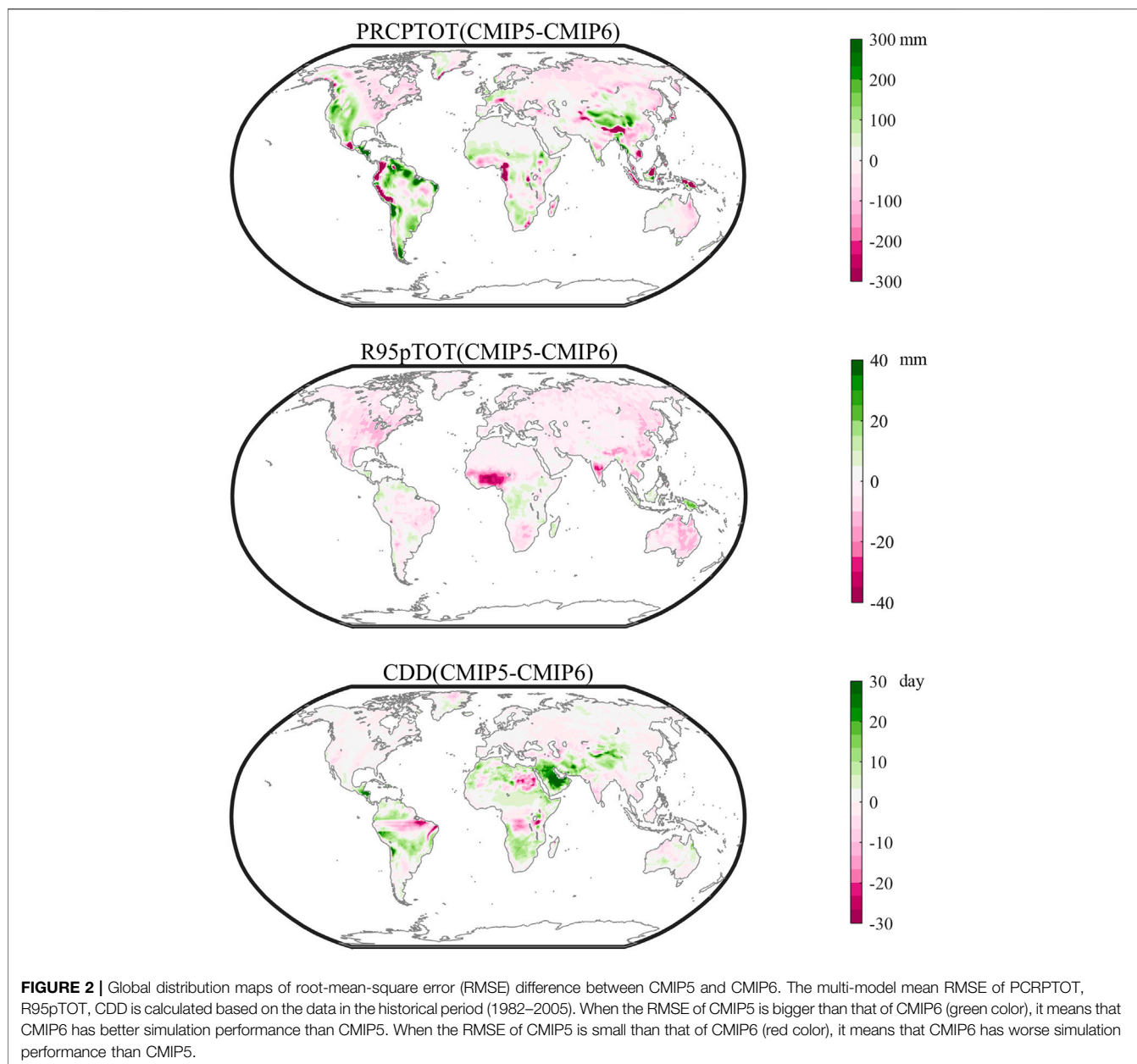
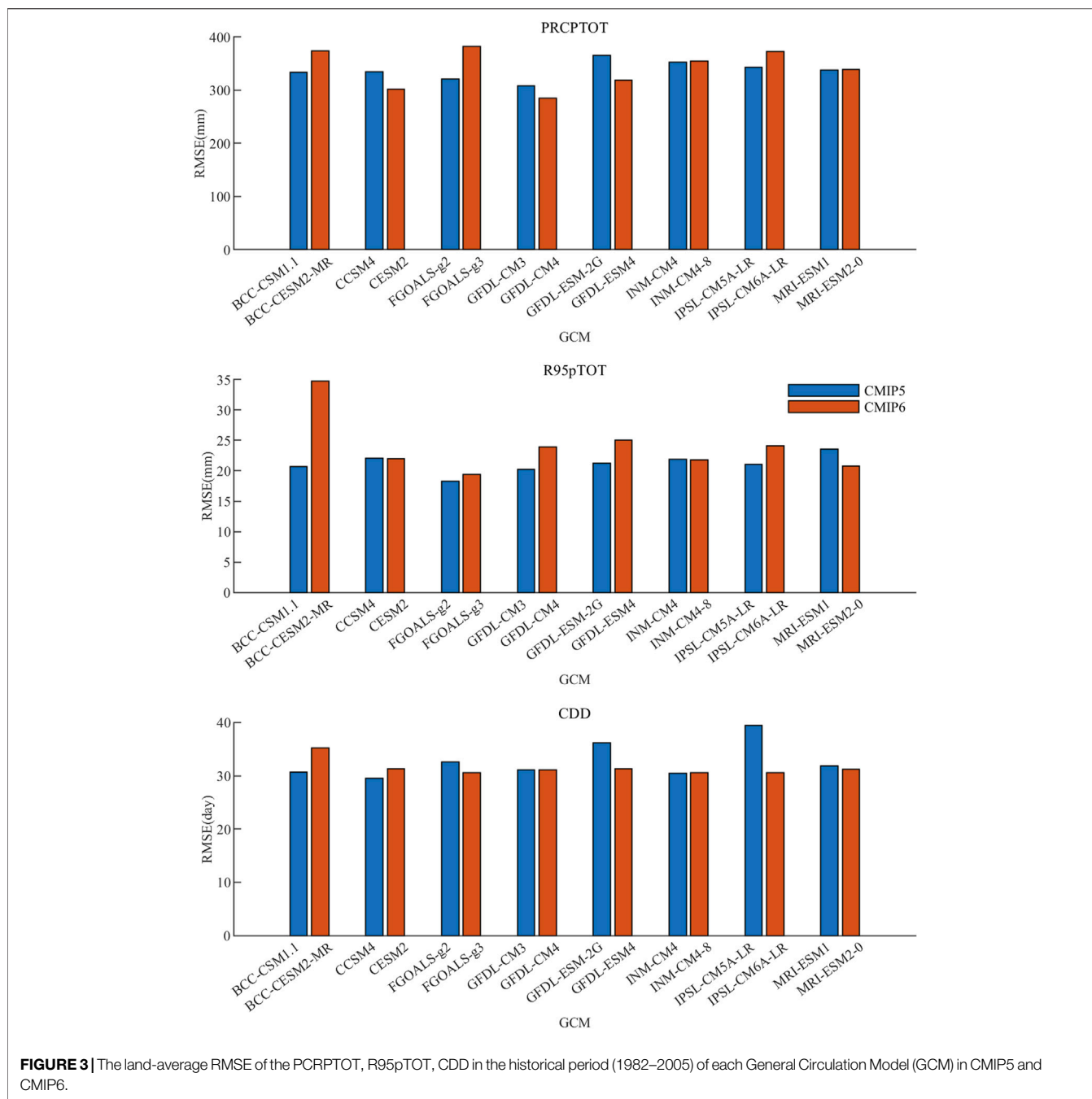


Figure 2 shows the following: 1) for PRCPTOT, the simulation ability of CMIP6 improved in CAM, WNA, SAH, TIB, the east coast of AMZ, the coastal areas of SAF and SSA. However, the simulation effect of CMIP6 in SAS the west coast of AMZ and WAF is slightly worse than in CMIP5. Overall, CMIP6 has no advantage over CMIP5. 2) For R95pTOT, the overall simulation effect of CMIP6 decreases, especially in WAF. 3) For CDD, the simulation ability of CMIP6 improved in eastern SAH, SAF, and central South America. However, the simulation ability of CMIP6 is weakened in Northeastern Brazil, central Africa, and central SAH. The areas with large differences between CMIP5 and

CMIP6 are mainly concentrated at low latitudes. Same as PRCPTOT, CMIP6 has no advantage over CMIP5 in CDD simulation.

The land-average RMSE values of the sixteen GCMs are listed in **Supplementary Table S1** and drawn in **Figure 3**. Comparing the land-average RMSE of the GCM, which participates in both CMIP5 and CMIP6, the precipitation indicator simulation capability of a single GCM in CMIP6 is not enhanced compared to that of CMIP5.

In general, CMIP6 has no overall advantage over CMIP5 in the simulation of PRCPTOT and CDD. The performance of CMIP6



increases or decreases regionally in PRCPTOT and CDD. But it is slightly worse than CMIP5 in the simulation of R95pTOT.

Trend Analysis and Comparison of Long Series Precipitation in CMIP5 and CMIP6

The Mann—Kendall (MK) trend test method is used to analyze precipitation series trends. The spatial distributions of the Sen's slope estimator test (β) for multi-model mean PRCPTOT, R95pTOT, and CDD are shown in **Figures 4–6**, which show β

values of the long series precipitation (1982–2100) under different future scenarios in CMIP5 and CMIP6. And the positive β values mean the increasing trend of precipitation indices and vice versa. In addition, the points in **Figures 4–6** represent non-significant trend at the 99% confidence level based on the MK test. Besides, the global distribution maps of β without showing the significant areas are shown in **Supplementary Figures S4–S6**. And global distribution maps of β difference between CMIP5 and CMIP6 are also shown in **Supplementary Figures S7–S9**.

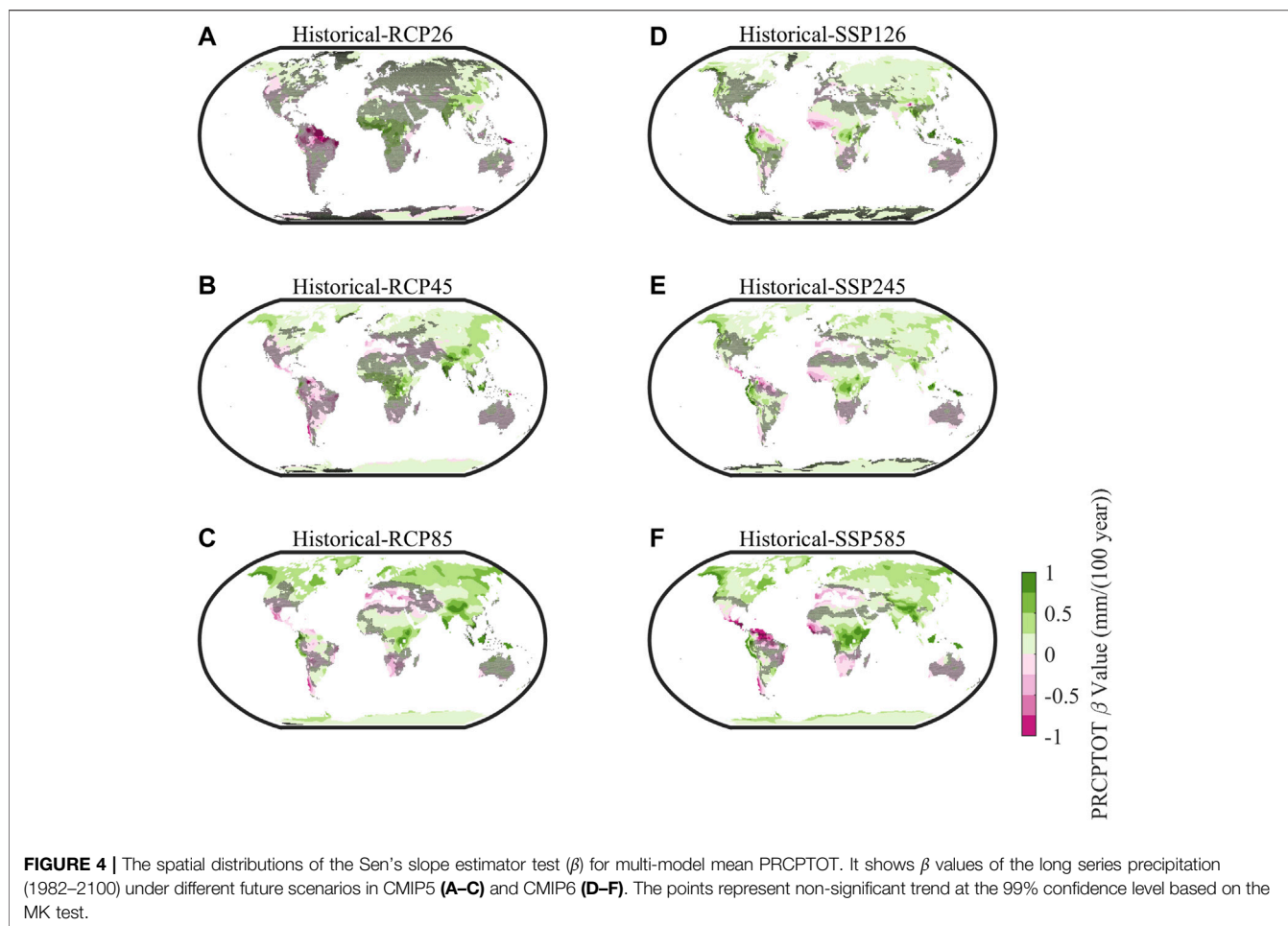
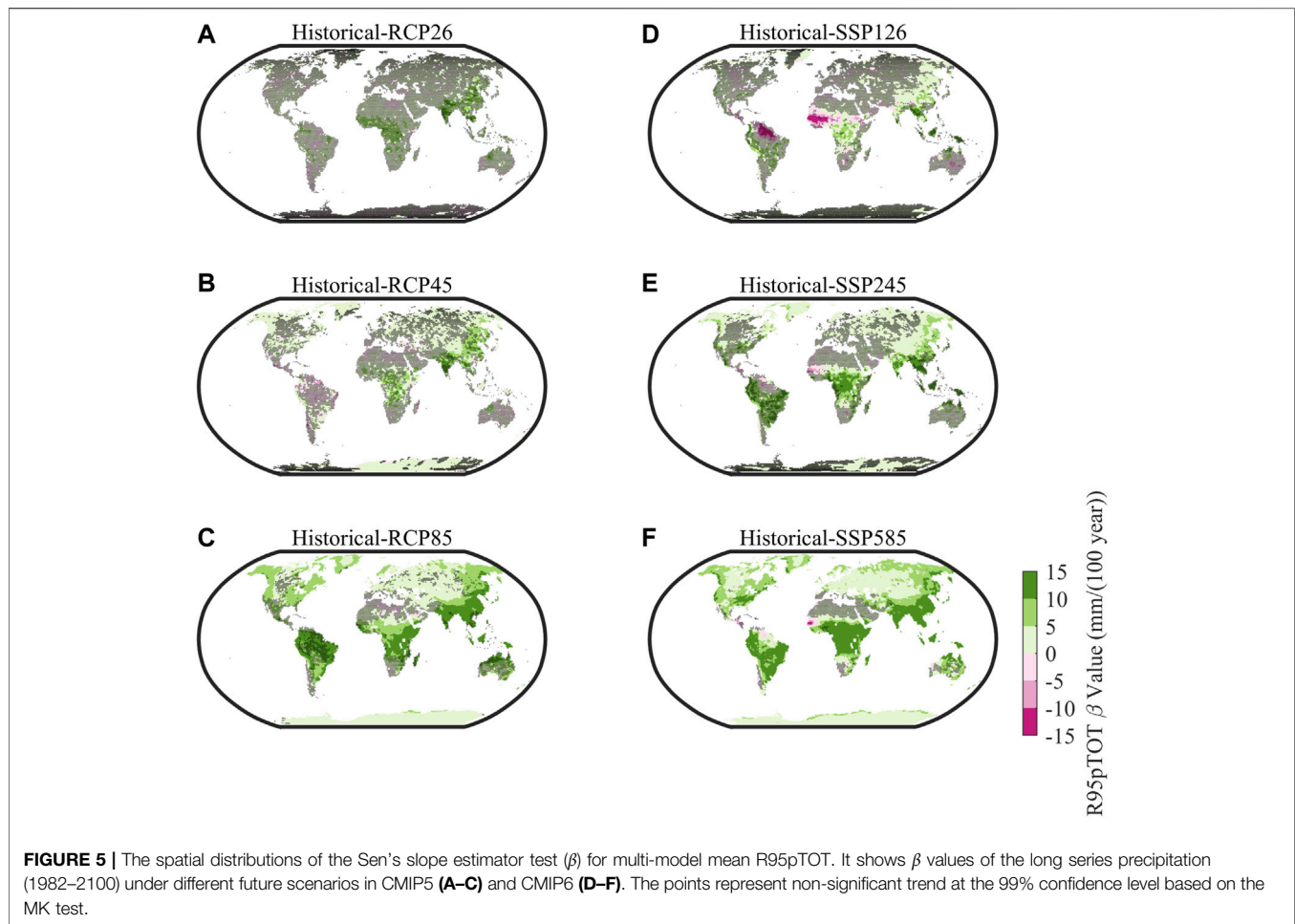


Figure 4 shows that: 1) PRCPTOT shows an increasing trend from the historical period to the future in many parts of the world under three future scenarios, especially for SSP585 (RCP85) scenarios with significant increase for most areas. Specifically, for SSP245 (RCP45) and SSP585 (RCP85) scenarios, there is a significant increasing trend for PRCPTOT in most regions of ALA, GRL, WNA, ENA, NEU, EAF, and the whole Asia. And the results of MK test show a significant decreasing trend in most parts of CAM and SAF; 2) Comparing the trend from the historical period to the future of CMIP5 and CMIP6, the β values in CMIP5 and CMIP6 show roughly similar spatial distributions in most land areas. There are, however, some exceptions, such as the fact that there is an opposite trend in some areas of AMZ and WAF for CMIP5 and CMIP6. In addition, the differences also exist in the degree of increasing/decreasing trend and coverage area; 3) We can also see more regions with significant trend based on MK test in CMIP6 compared with that of CMIP5. And as the level of radiative forcing increases (RCP26/SSP126---RCP45/SSP245----RCP85/SSP585), there are more regions showing significant change from the historical period to the future.

Figure 5 shows that: 1) all of the CMIP5 and CMIP6 GCMs indicate that R95pTOT will most likely increase across most of World from the historical period to the future under three future scenarios. That is to say, in general R95pTOT increases over wider and decreases over smaller areas than PRCPTOT. For example, under the scenarios of SSP585 (RCP85), except for a few regions in WAF and SAF, the R95pTOT for almost all land areas (with black dots) shows a significant increasing trend. This may lead to more extreme flooding in the world under global warming scenarios based on the CMIP5 and CMIP6 simulations; 2) for the average extreme precipitation indices (R95pTOT) shown in **Figure 4**, we can also see that the trends from the historical period to the future from all CMIP5 and CMIP6 GCMs display same directions (increase or decrease) compared with PRCPTOT for many regions, which is helpful to discern uniform trend for a particular region. However, in some regions of AMZ and SSA where PRCPTOT decreases, there is a statistically significant increase in R95pTOT for all CMIP5 and CMIP6 future scenarios.

Figure 6 shows that: 1) more regions in the world for CDD show a significant decreasing trend. Specifically, CDD shows a



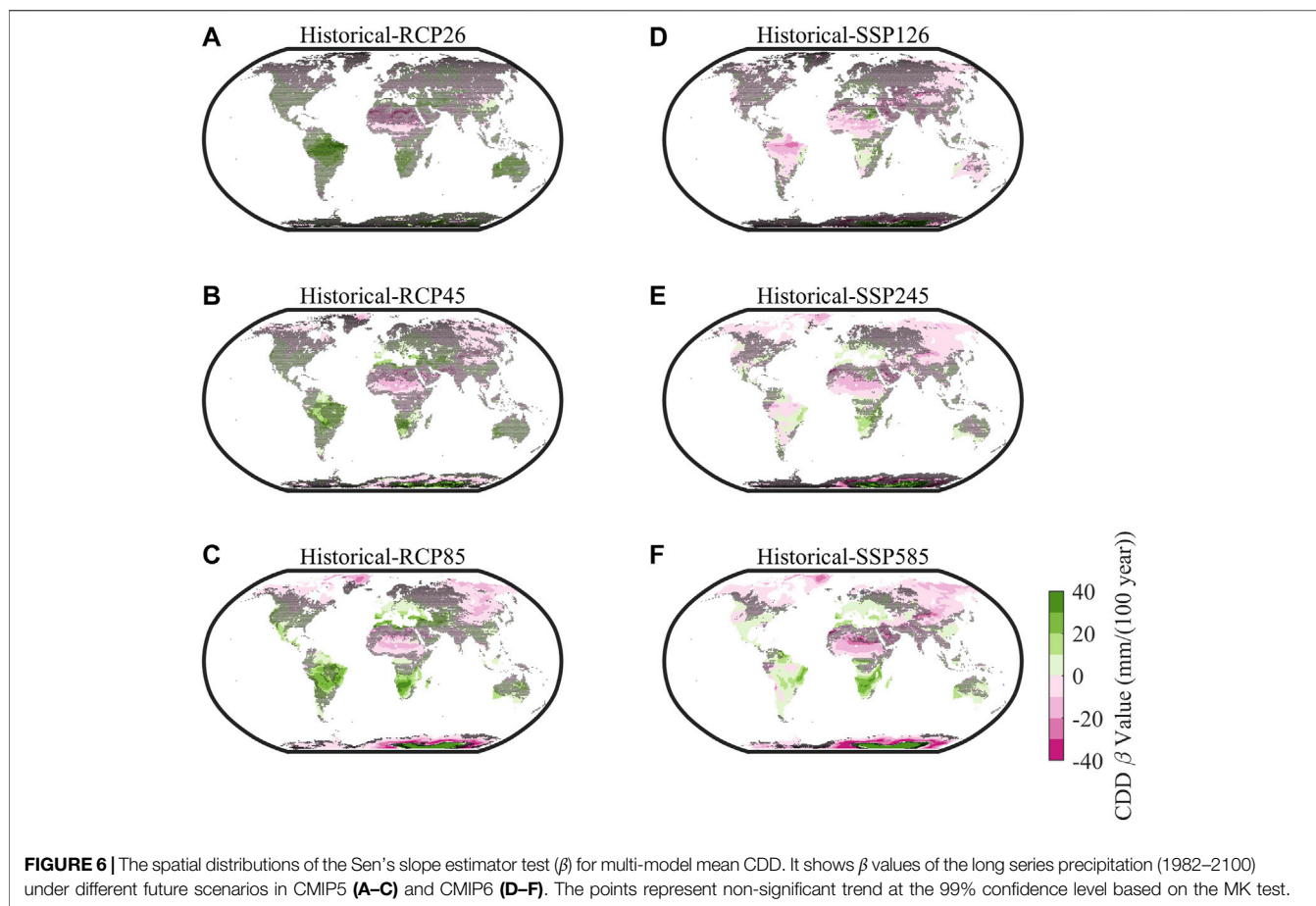
significant decreasing trend mainly in several regions of NAS and EAS, WAF and EAF, and SAH. Under the three future scenarios, the area with the decreasing trend is larger than the area with an increasing trend, indicating that the probability of drought shows a decreasing trend in the future. 2) Comparing the results of CMIP5 and CMIP6 from the historical period to and the future, the CDD of the two projects is different in some regions. For example, some areas in AMZ and SSA show an increasing trend for CMIP5 and a decreasing trend in CMIP6; 3) similar spatial patterns are shown for CMIP5/CMIP6 future scenarios. In addition, there are larger areas showing stronger decrease trend for CDD from SSP126 (RCP26) scenario to SSP585 (RCP85) scenario.

Changing Rate Comparison of Precipitation in the Future Over the History

Calculates the multi-model changing rates in each $1^\circ \times 1^\circ$ land grid in the near-future (2028–2060) and far-future periods (2068–2100) with respect to the historical period (1982–2014) in CMIP5 and CMIP6. Box plots for PRCPTOT multi-model mean changing rates over 21 regions (Table 3) in the near-future and far-future periods are shown in Figures 7, 8, respectively. Box

plots for R95pTOT are shown in Figures 9, 10. Box plots for CDD are shown in Figures 11, 12.

Figures 7, 8 shows that for PRCPTOT, the agreement between CMIP5 and CMIP6 is good because there are few regions where the changing rates simulated by CMIP5 and CMIP6 diverge. Under SSP126 (RCP26) and SSP245 (RCP45), the far future changes slightly compared with the near future. However, for SSP585 (RCP85), the far future changes significantly compared to the near future. PRCPTOT will inevitably increase in the future under the three scenarios, especially in ALA, GRL, NEU, EAF, SAH, EAS, TIB, NAS. However, not all regions of the world will show an increase in PRCPTOT. Compared with CMIP5, CMIP6 changing rate does not show an overall increase or decrease trend. In some regions, the changing rate of CMIP6 is greater; in some regions, the changing rate of CMIP6 is less. In the SSA and WNA regions, the changing rate of CMIP6 in the future under various scenarios is always greater than that of CMIP5; in the ALA, WAF, SAF regions, the changing rate of CMIP6 in the future under various scenarios is always less than that of CMIP5. The difference between land-grids is relatively large in WAF and SAH; the difference between land-grids is also large in GRL under SSP585(RCP85). In EAS, the difference between land-



grids of CMIP6 is always significantly smaller than that of CMIP5 under various scenarios.

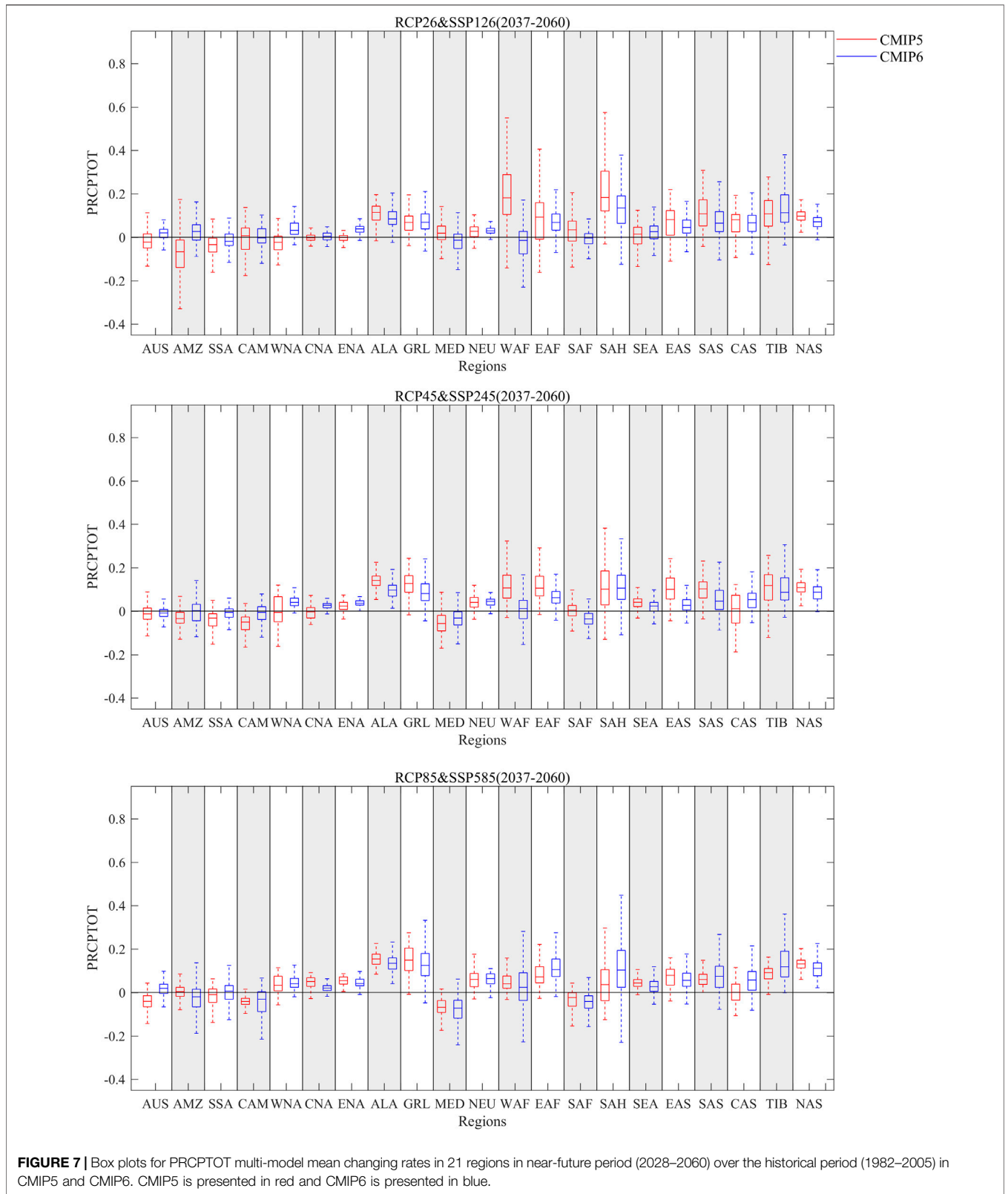
Figures 9, 10 shows that for R95pTOT, the trends of CMIP6 and CMIP5 are similar. As the radiative forcing level increases (from RCP26/SSP126, and RCP45/SSP245 to RCP85/SSP585), the overall changing rate of R95pTOT indicators increases, and the number of regions with a changing rate that is greater than 0 increases for both projects. Under the SSP585 (RCP85) scenario, R95pTOT in all regions increase. For SSP126 (RCP26) and SSP245 (RCP45), the far future changes slightly compared with the near future. However, for SSP585 (RCP85), the far future changes significantly compared to the near future. In the SSA region, the changing rate of CMIP6 in the future under various scenarios is always greater than that of CMIP5; in the region ALA, EAS, NAS regions, the changing rate of CMIP6 in the future under various scenarios is always less than that of CMIP5. The difference between land-grids is relatively large in WAF, EAF, and SAH. In the AMZ, WNA, ALA, WAF, SAH regions, the difference between land-grids of CMIP6 is always bigger than that of CMIP5 under various scenarios.

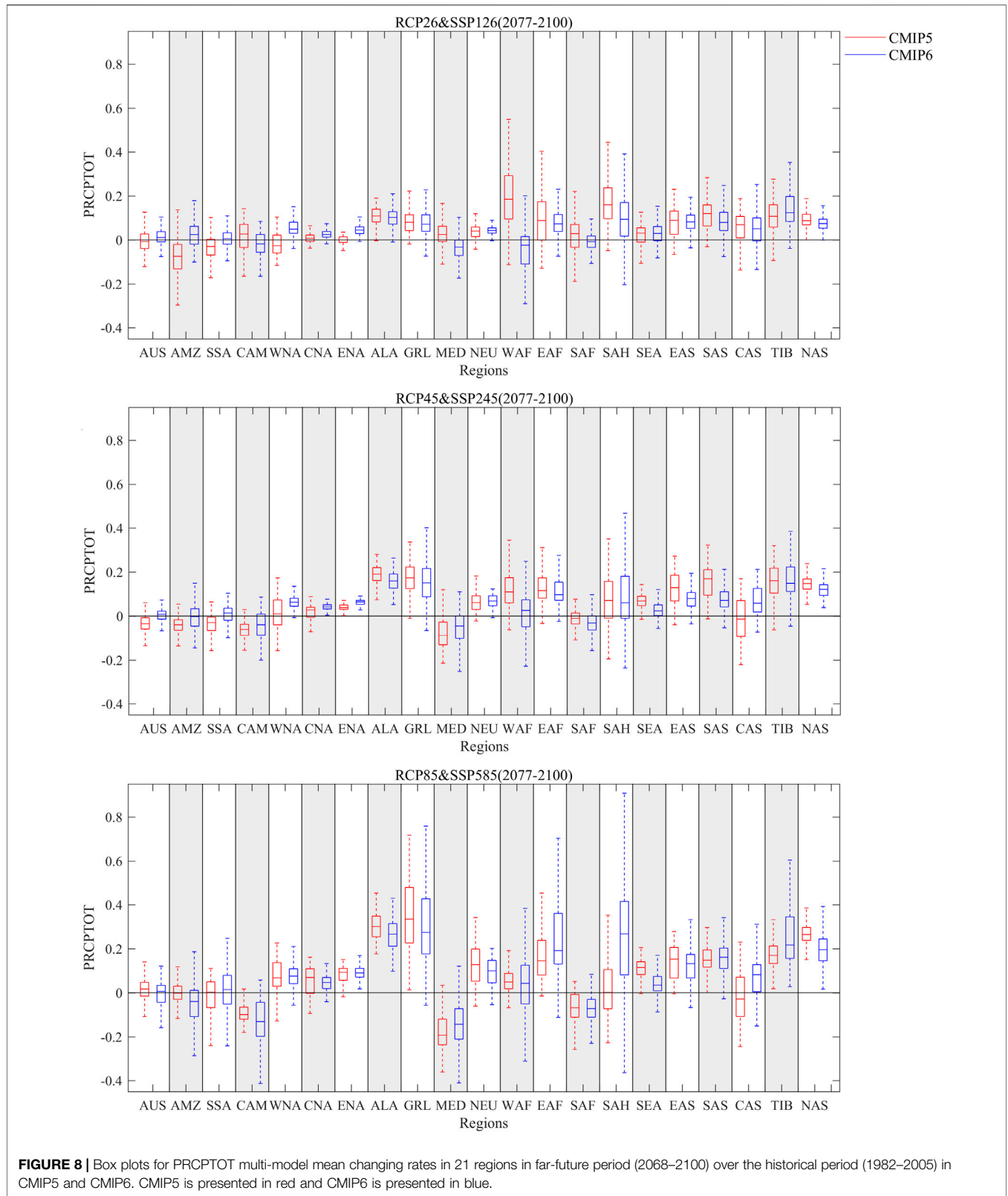
Figures 11, 12 shows that for CDD, the changing rates between the 21 regions differ significantly. Nearly half of the regions experience a future decline in CDD, and other regions

show a future increase in CDD. Under SSP126 (RCP26) and SSP245 (RCP45), the far future changes slightly compared with the near future. However, under SSP585 (RCP85), the far future changes significantly compared to the near future. In almost half of the regions, the changing rate of CMIP6 is less than that of CMIP5 under all scenarios, namely AUS, AMZ, SSA, CAM, WNA, CNA, ENA, CAS, TIB. The difference between land-grids is relatively large in the AMZ and SEA regions. In the WNA, EAF, SAH, EAS, TIB regions, the difference between land-grids of CMIP6 is always significantly smaller than that of CMIP5 under various scenarios. In the GRL, NAS regions, the difference between land-grids of CMIP6 is always significantly bigger than that of CMIP5 under various scenarios.

Comparison of Model Uncertainty in CMIP5 and CMIP6 Precipitation

Figures 13–15 show the data series of land-average PRCPTOT, R95pTOT, and CDD from 1982 to 2100, respectively. The orange and purple solid lines represent the multi-model mean results of CMIP5 and CMIP6, respectively. The upper and lower limits of the orange and purple uncertainty intervals represent the 75 and 25% quantiles, respectively.





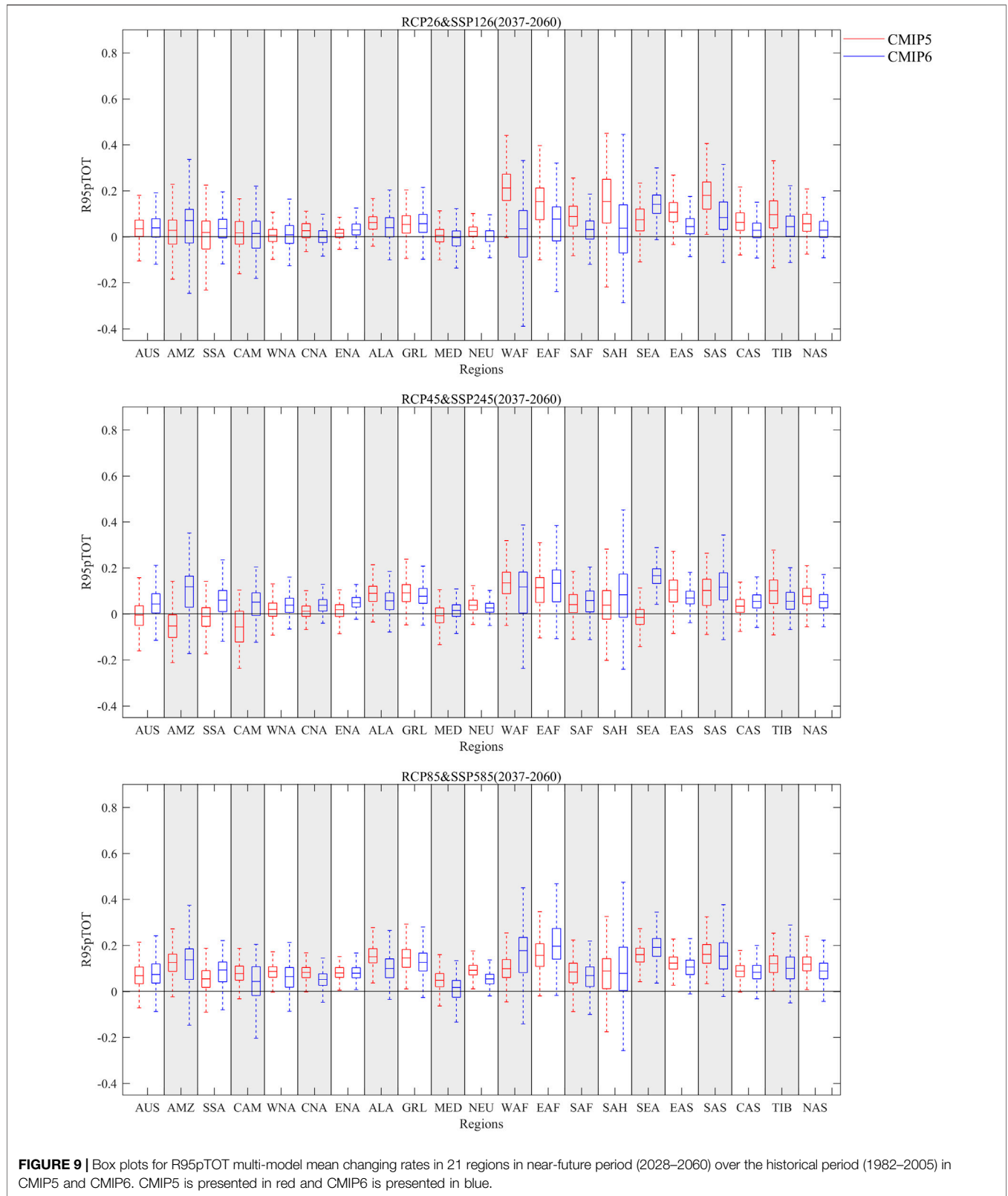
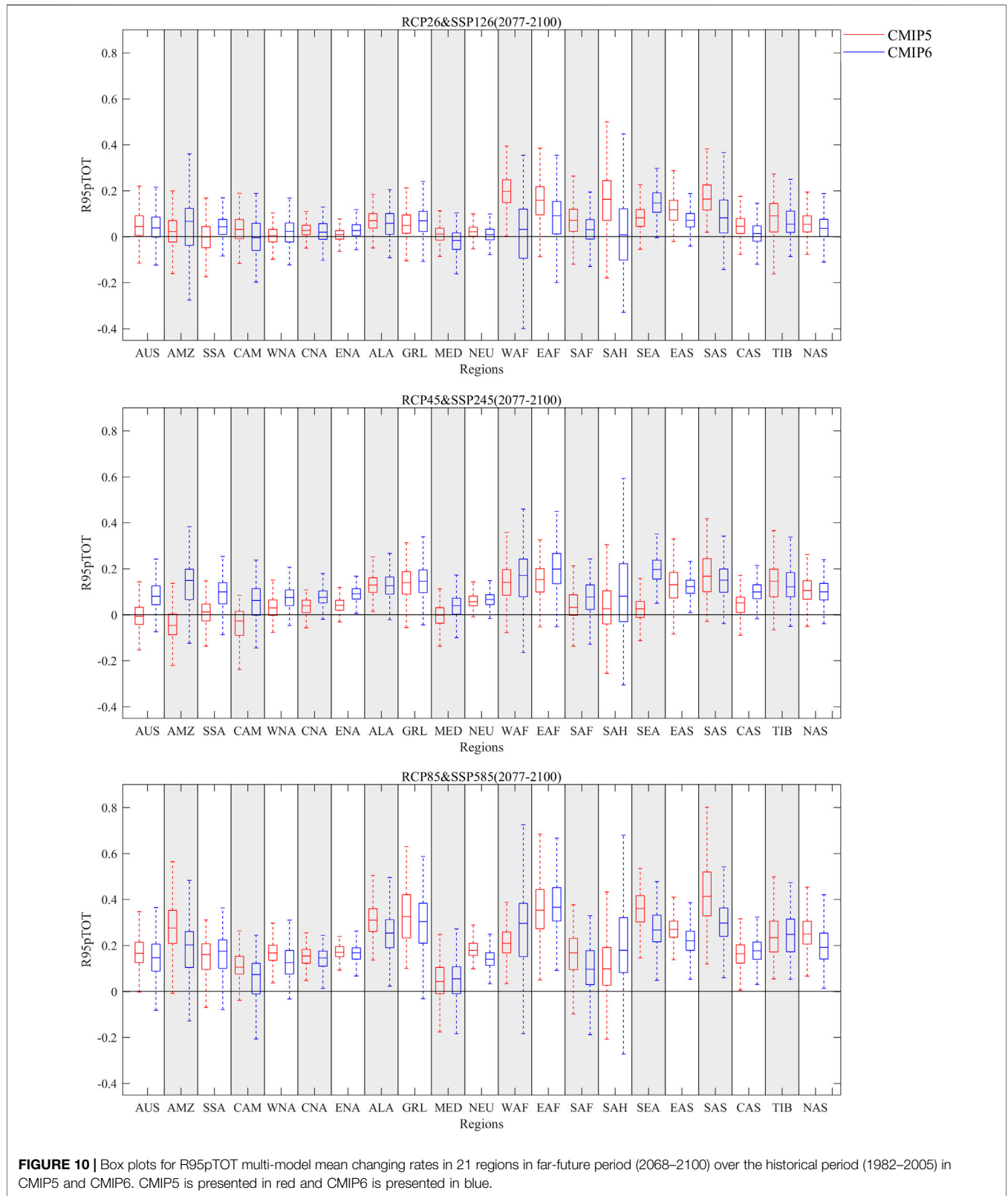
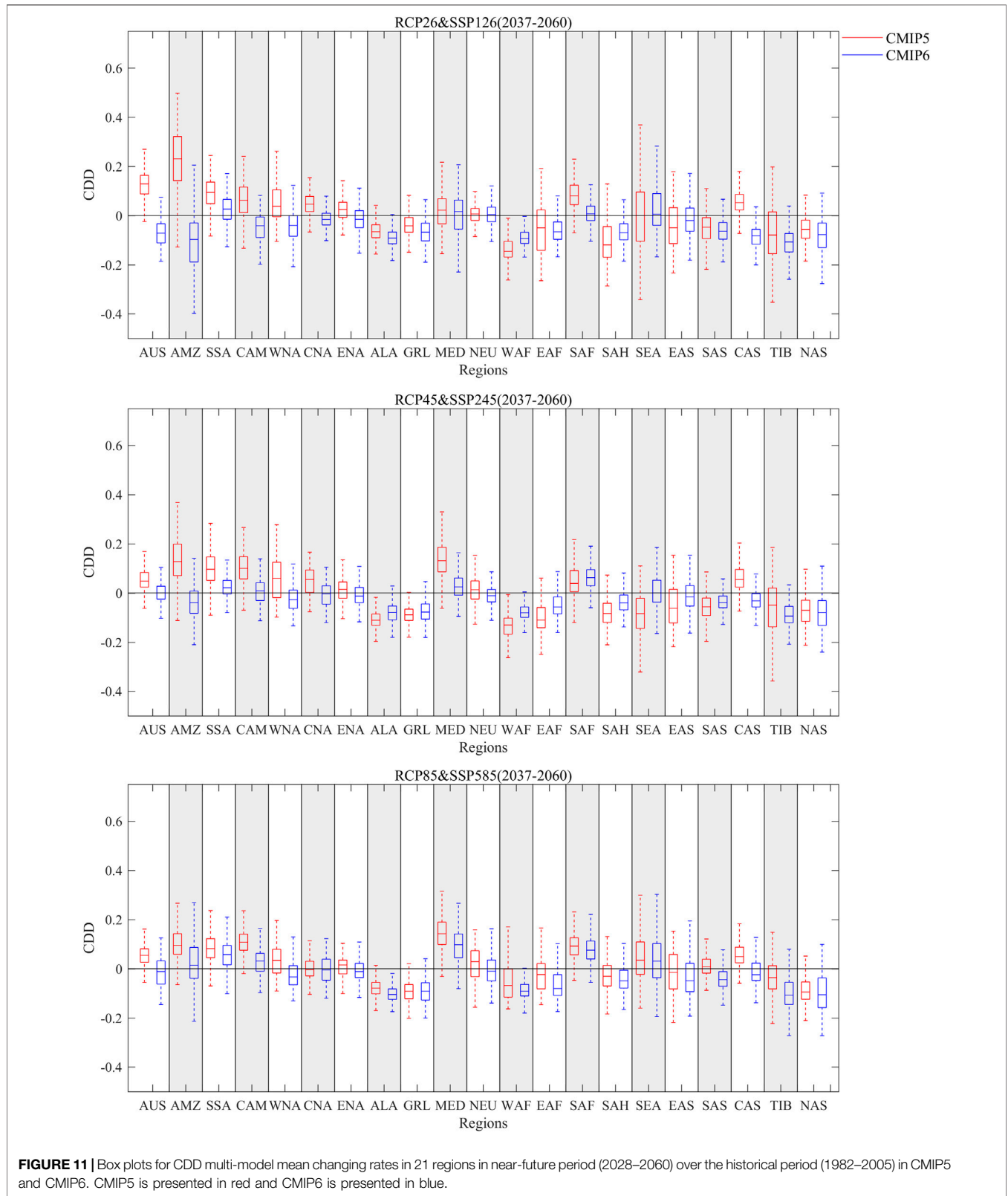
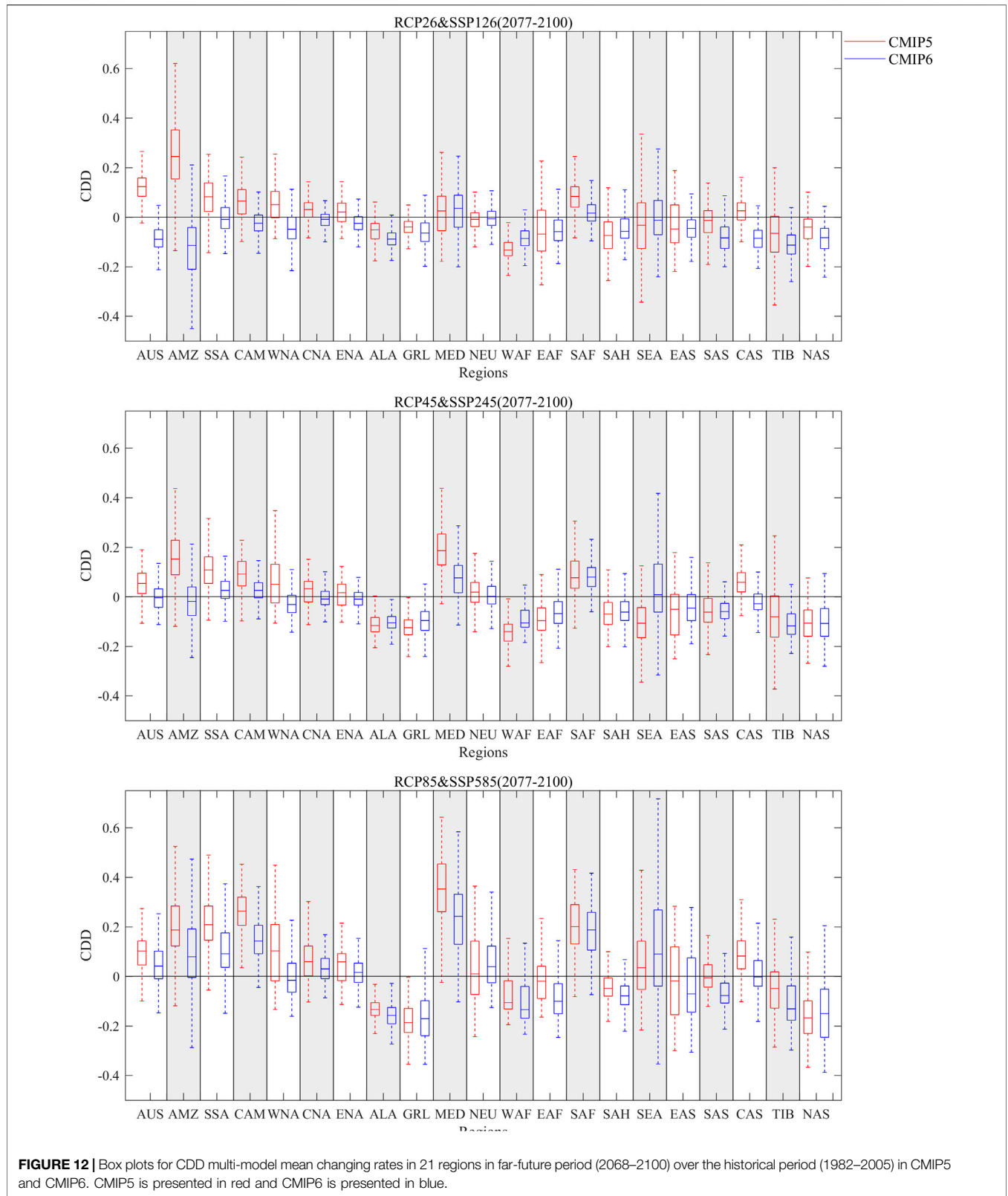


FIGURE 9 | Box plots for R95pTOT multi-model mean changing rates in 21 regions in near-future period (2028–2060) over the historical period (1982–2005) in CMIP5 and CMIP6. CMIP5 is presented in red and CMIP6 is presented in blue.







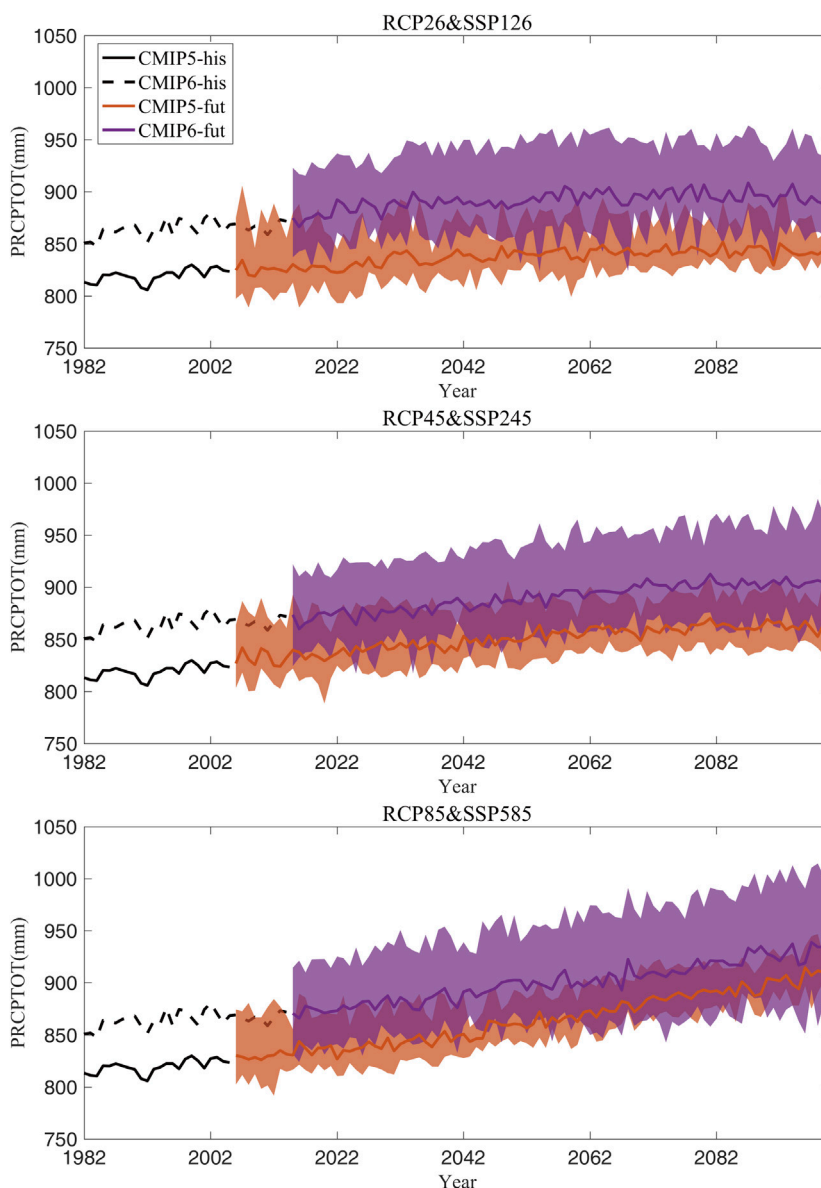


FIGURE 13 | Global land-average multi-model mean PRCPTOT during 1982–2100 in CMIP5 and CMIP6. The shaded region shows the uncertainty range (between the 25 quantiles and the 75 quantiles).

For PRCPTOT and R95pTOT, both CMIP5 and CMIP6 show a significant increasing trend from the historical period to the end of the 21st century. There are similar fluctuations and trends under the corresponding future scenarios (RCP26 and SSP126; RCP45 and SSP245; RCP85 and SSP585) for CMIP5 and CMIP6. As the level of radiative forcing increases (RCP26/SSP126---RCP45/SSP245----RCP85/SSP585), the increasing trend becomes more apparent. The increase in PRCPTOT is greater than that in R95pTOT. The multi-mode mean PRCPTOT and R95pTOT of CMIP6 are significantly higher than those of CMIP5. The model uncertainty range of PRCPTOT and R95pTOT in CMIP6 is larger than that in CMIP5, with a small overlap area (Figures 13, 14).

However, CDD shows a slight decreasing trend from the historical period to the end of the 21st century in both CMIP5 and CMIP6. The multi-mode mean CDD of CMIP6 is slightly higher than that of CMIP5. The model uncertainty range of CDD in CMIP6 is smaller than that in CMIP5 (Figure 15).

CONCLUSION

- (1) CMIP6 has no overall advantage over CMIP5 in the simulation of PRCPTOT. The performance of CMIP6 of

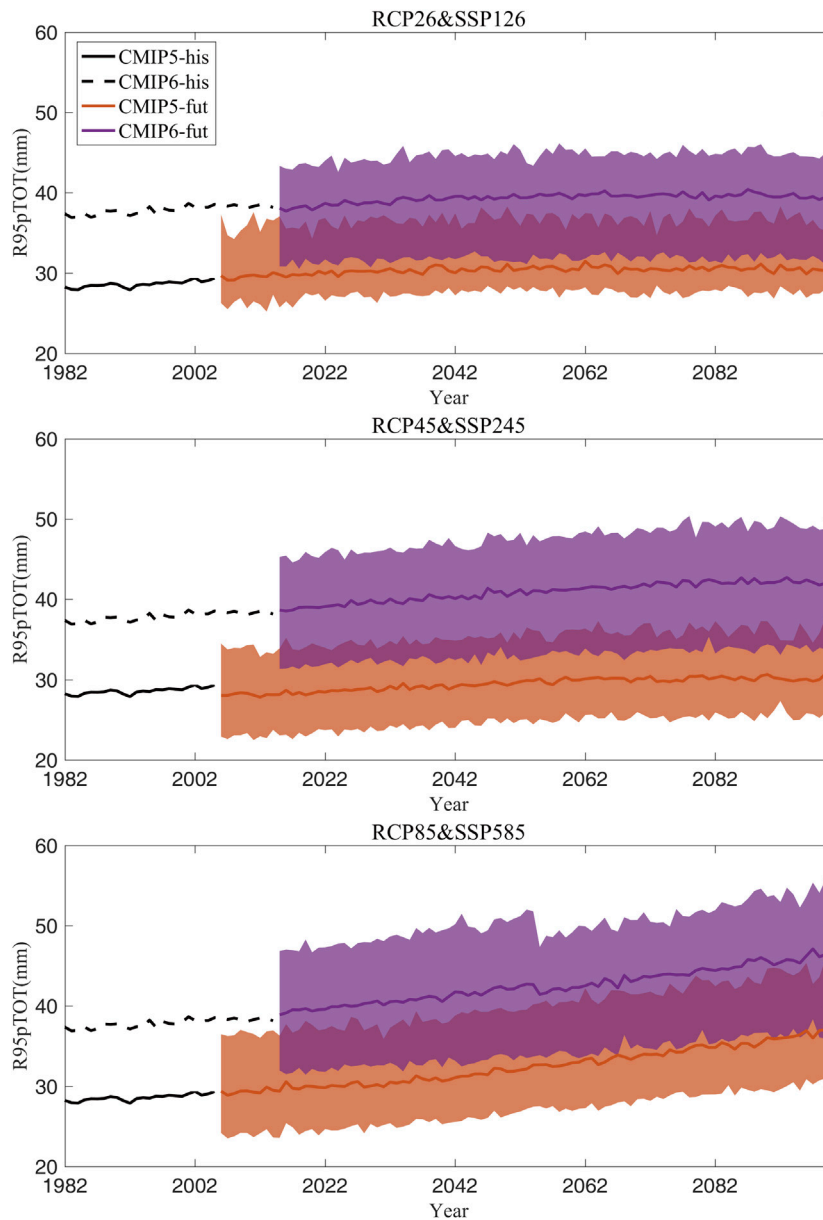


FIGURE 14 | Global land-average multi-model mean R95pTOT during 1982–2100 in CMIP5 and CMIP6. The shaded region shows the uncertainty range (between the 25 quantiles and the 75 quantiles).

PRCPTOT increases or decreases regionally. For the simulation of R95pTOT, the overall effect of CMIP6 is weaker than that of CMIP5, especially in WAF. Same as PRCPTOT, CMIP6 has no advantage over CMIP5 in CDD simulation. The areas with large differences of CDD between CMIP5 and CMIP6 are mainly concentrated at low latitudes.

(2) The MK trend test method and Sen's slope estimator test is used to analyze the trends of long series precipitation. Overall, there are more areas with significant trend based on MK test in CMIP6 compared with that of CMIP5. The β values of PRCPTOT in CMIP5 and CMIP6 show roughly similar spatial distributions in most land areas. There are, however, some exceptions, such as the

fact that there is an opposite trend in some areas of AMZ and WAF for CMIP5 and CMIP6. The differences between the R95pTOT trends mainly noticeable in South America and WAF. The CDD trend of the two projects is different in some regions. For example, some areas in AMZ and SSA show an increasing trend for CMIP5 and a decreasing trend in CMIP6.

(3) Analyzing the changing rates of the 21 regions in the near-future period and far-future period over the historical period, it can be found that in most regions, the simulated changing rates of CMIP6 have no obvious change direction compared with CMIP5. However, in some regions, CMIP6 has a steady upward or decreasing trend compared to CMIP5. In the SSA

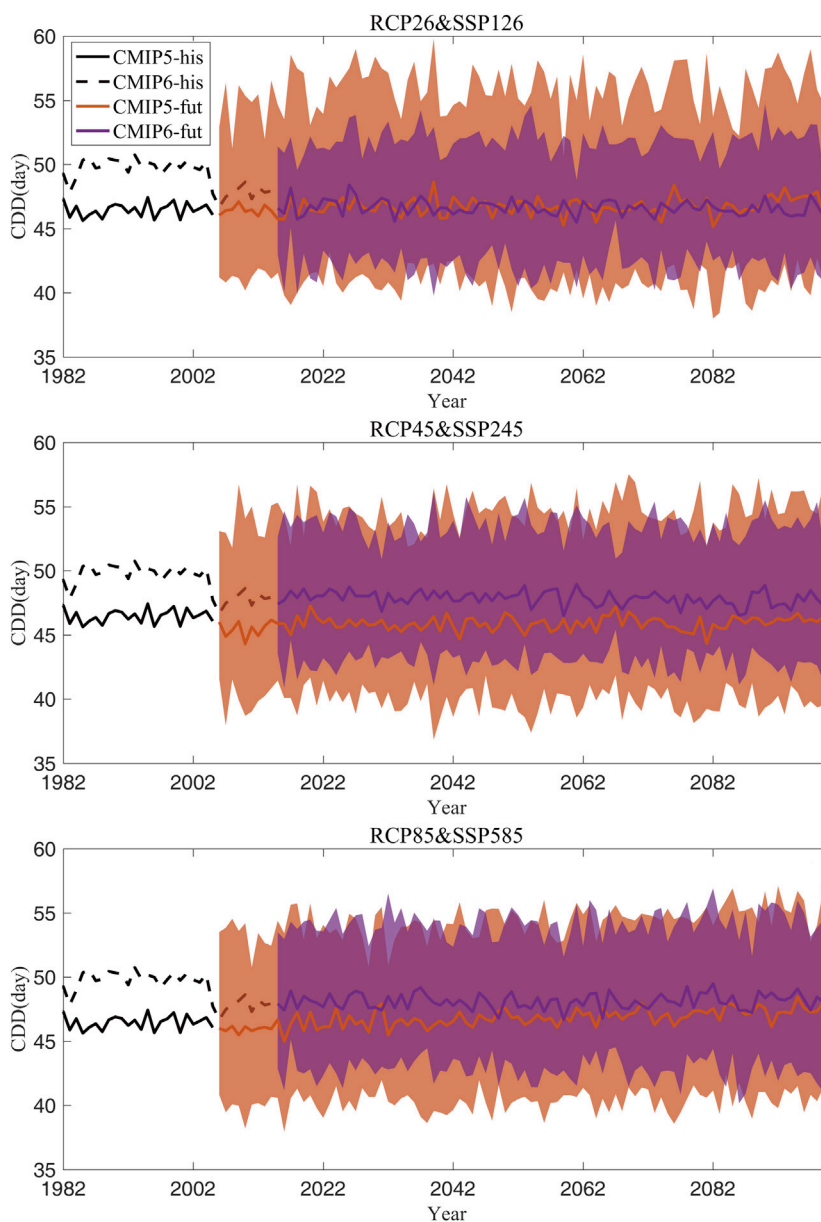


FIGURE 15 | Global land-average multi-model mean CDD during 1982–2100 in CMIP5 and CMIP6. The shaded region shows the uncertainty range (between the 25 quantiles and the 75 quantiles).

and WNA regions, the PRCPTOT changing rate of CMIP6 in the future under various scenarios is always greater than that of CMIP5; in the ALA, WAF, SAF regions, the PRCPTOT changing rate of CMIP6 in the future under various scenarios is always less than that of CMIP5. In the SSA region, the R95pTOT changing rate of CMIP6 in the future under various scenarios is always greater than that of CMIP5; in the ALA, EAS, NAS regions, the R95pTOT changing rate of CMIP6 in the future under various scenarios is always less than that of CMIP5. In almost half of the regions, the changing rate of CMIP6 is less than that of CMIP5 under all scenarios, namely AUS, AMZ, SSA, CAM, WNA, CAN, ENA, CAS, TIB.

(4) For PRCPTOT and R95pTOT, the uncertainty of CMIP6 is larger than that of CMIP5; for CDD, the uncertainty of CMIP6 is smaller than that of CMIP5.

DATA AVAILABILITY STATEMENT

Publicly available datasets were analyzed in this study. This data can be found here GPCP: <https://climatedataguide.ucar.edu/climate-data/gpcp-global-precipitation-climatology-centre> CMIP5: <https://esgf-node.llnl.gov/search/cmip5/> CMIP6: <https://esgf-node.llnl.gov/search/cmip6>.

AUTHOR CONTRIBUTIONS

JL, RH, and HC carried out the study design. YZ and TZ download the CMIP5 CMIP6, and GPCC data. JL, RH, TZ, and YZ carried out the calculation. HC was responsible for supervision, project administration and funding acquisition. JL and RH checked the data processing progress. RH and HC reviewed and edited the manuscript after JL wrote the original draft. All authors contributed to the article and approved the final manuscript.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2021.687976/full#supplementary-material>

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Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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