



Quantitative Analysis of the Effects of an Earthquake on Rainfall Thresholds for Triggering Debris-Flow Events

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The impacts of destructive earthquakes on rainfall thresholds for triggering the debris flows have not yet been well investigated, due to lacks of data. In this study, we have collected the debris-flow records from the Wenchuan, Lushan, and Jiuzhaigou earthquake-affected areas in Sichuan Province, China. By using a meteorological dataset with 3 h and 0.1° resolutions, the dimensionless effective rainfall and rainfall intensity-duration relationships were calculated as the possible thresholds for triggering the debris flows. The pre- and post-seismic thresholds were compared to evaluate the impacts of the various intensities of earthquakes. Our results indicate that the post-quake thresholds are much smaller than the pre-seismic ones. The dimensionless effective rainfall shows the impacts of the Wenchuan, Lushan, and Jiuzhaigou earthquakes to be ca. 26, 27, and 16%, respectively. The Wenchuan earthquake has the most significant effect on lowering the rainfall intensity-duration curve. Rainfall threshold changes related to the moment magnitude and focal depth are discussed as well. Generally, this work may lead to an improved post-quake debris-flow warning strategy especially in sparsely instrumented regions.

Keywords: debris flows, seismic effects, regional rainfall threshold, early warning, disaster prevention

INTRODUCTION

Massive earthquake-induced loose deposits on slopes or channels can be mobilized by heavy rainfalls and evolve into debris flows. The increased unconsolidated sediment reduces the rainfall threshold for the debris-flow occurrence and significantly augments the magnitude of post-quake debris flows (Chen and Petley, 2005; Ma et al., 2013). The post-quake hazards can remain highly frequent for 5–10 years, even longer as 20 years (Cui et al., 2011). Quantifying the impact of strong earthquakes on the rainfall threshold is essential for post-quake debris-flow early warning (Zhang and Zhang, 2017).

Numerous studies have reported the rainfall threshold variation after strong earthquakes. As an example, the mean value of the accumulated rainfall decreased from 867 mm for 67 debris flows in 1996 to 146 mm for 197 debris flows in 2001 in the intensive-ground-motion areas of the 1999 Chi-Chi earthquake in Taiwan (Liu et al., 2008). Similarly, the maximum effective cumulative precipitation decreased from 950 to 200 mm in the Mt. Ninety-Nine area of Taiwan during the first year after the earthquake (Shieh et al., 2009). After the 2008 Wenchuan earthquake, the accumulated rainfall threshold decreased by 15–22% while the critical value for the hourly rainfall intensity decreased by 25–32% (Tang et al., 2009). Likewise, a sudden drop in the rainfall threshold

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FIGURE 1 | Location of study areas in Sichuan Province, China. The ranges, elevations, main faults, debris-flow records (filled dot), epicenters (pentagram), moment magnitude (Mw), and mean annual rainfall (contour with 100 mm interval) are shown for the WEA (Chinese seismic intensity \geq VII), LEA (Chinese seismic intensity \geq VII), and JEA (Chinese seismic intensity \geq VII), respectively. The location of the Chi-Chi earthquake in Taiwan Province is also marked in the subplot.

was found in the Gaochuan River basin near the Wenchuan earthquake's epicenter (Li et al., 2016). The rainfall intensityduration (I-D) thresholds normalized by the mean annual maximum hourly rainfall, indicate a reduction of at least 30% compared to the pre-quake level (Ma et al., 2017). Yu et al. (2014) found that the ratio of the thresholds before and after the earthquake decreased with increasing peak ground acceleration. However, due to data limitation, the impacts of destructive earthquakes on rainfall thresholds have not been well investigated quantitatively.

The recent three severe earthquakes occurred in Sichuan Province, Southwest China, namely the Wenchuan earthquake in 2008, the Lushan earthquake in 2013 and the Jiuzhaigou earthquake in 2017 have different magnitudes (**Figure 1**), which provides valuable cases to quantitatively examine the influence of the earthquake magnitude on the rainfall thresholds. We have collected debris-flow records in the Wenchuan earthquake-affected area (WEA), Lushan earthquake-affected area (LEA), and Jiuzhaigou earthquakeaffected area (JEA) to examine how much an earthquake can affect the rainfall threshold for triggering debris flows. Two kinds of rainfall thresholds, including the dimensionless effective rainfall and the rainfall I-D curve, were calculated based on the China Meteorological Forcing Dataset from National Tibetan Plateau Data Center to show the seismic effects.

MATERIALS AND METHODS

Study Area

The study area is located at the eastern edge of the Tibetan Plateau, which has been affected by the three earthquakes (**Figure 1**). The Mw 7.9 Wenchuan earthquake (the "5.12" earthquake) ruptured along the Longmenshan thrust belt on May 12, 2008 (Xu et al., 2009). The shock-induced collapses and landslides produced approximately 2.6×10^9 m³ of loose material (Parker et al., 2011), indicating that debris flows may occur frequently following the earthquake (Cui et al., 2011). Huang and Li (2009) found that there were 8,627 geohazards, including 3,627 landslides, 2,383 slope collapses, and 837 debris flows in the 42 earthquake-hit regions of Sichuan Province. The Mw 6.6 Lushan earthquake

Count

58

12

	WEA		LEA		JEA				
	Period	Count	Period	Count	Period				
Pre-seismic	2000-2007	21	2000-2012	88	2010-2016				
Post-seismic	2008-2010	246	2013-2017	21	2017–2018				

TABLE 1 | Debris-flow numbers and the time range in each area.

occurred in the Longmen Shan fold-and-thrust belt, Sichuan Province, China, on April 20, 2013, causing 196 deaths and 13,484 injuries (Li et al., 2014). The earthquake has reduced hillslope stability and moderately increased the possibility of debris flows occurrence (Cui et al., 2014). On August 8, 2017, the Mw 6.5 earthquake with a focal depth of 20 km hit the Jiuzhaigou World Natural and Cultural Heritage site and its surroundings (Hong et al., 2018). The epicenter was at 33.2° N, 103.82° E, approximately 5 km west of the Jiuzhaigou core scenic area (Hu et al., 2019). 1883 earthquake-induced geological hazards were identified by remote sensing within the 840 km² area affected by the earthquake (Fan et al., 2018).

Considering the sufficient earthquake intensity and available debris-flow records for comparisons, we choose the WEA with the Chinese seismic intensity \geq VIII, the LEA with the intensity \geq VI, and the JEA with the intensity \geq VI as regions of interest. The moist monsoon dominates the regional climate, providing water conditions for debris flows (Cui et al., 2014). In addition, the Mw 7.6 Chi-Chi Earthquake in Taiwan is also used for discussion.

Data Sources

Collection of Debris-Flow Data

The debris flows mentioned in this study limit to the torrential flow of a mixture of water, mud, and debris in the gully or channel of a small catchment, even on a hill slope. The time span of the collected debris-flow events to the earthquakes does not exceed 15 years. Data of some debris-flow events when nearzero rainfall was recorded from a gridded rainfall dataset are excluded (more details in 2.2.2). Then, 267 debris-flow records for 2000-2010 in the WEA, 109 for 2000-2017 in the LEA, and 70 for 2010-2018 in JEA were collected through field survey and data gathering from Sichuan Provincial Natural Resources Department (Figure 1; Table 1). In the WEA, the debris flows that happened on the same day with the earthquake are classified as post-earthquake events, which can be used to show extremely post-seismic impacts on the rainfall conditions. Unlike debris flows in the WEA, we collected no co-seismic debris-flow records in the LEA and JEA.

Gridded Rainfall Data

The gridded near-surface meteorological dataset was developed for evaluating land surface processes in China (He et al., 2020). The dataset was made through the fusion of remote sensing products (e.g., Tropical Rainfall Measuring Mission 3B42 v7), a reanalysis dataset (e.g., Global Land Data Assimilation System and Modern-Era Retrospective Analysis), and *in-situ* observations at weather stations (e.g., daily data from approximately 700 stations of China Meteorological Administration and sub-daily data from the National Oceanic and Atmospheric Administration's National Centers for Environmental Information). The time of the data ranges from January 1979 up to December 2018 with a temporal resolution of 3 h and a spatial resolution of 0.1°. For each 0.1° grid cell, a rainfall record with Greenwich Mean Time (GMT) stands for a mean value of rainfall intensities (mm/h) taken every three hours. So, each grid cell represents 8 records on each day. The data have been adjusted to match the Beijing time, i.e., 8 h ahead of the GMT.

Based on the data, the mean annual rainfall of 1979–2018 is ca. 950, 1,360, and 660 mm for the WEA, the LEA, and the JEA, respectively. According to the time ranges in **Table 1**, the mean annual rainfall of 937 mm after 2008 in the WEA is 7% higher than before 2008. However, the number of debris flows per year has increased ca. 30 times compared with the period before the Wenchuan earthquake. The decrease in the number of debris flows in both the LEA and the JEA cannot be directly linked with the change in the mean annual rainfall, suggesting that more short-term rainfall conditions should be discussed.

Regional Rainfall Threshold Dimensionless Effective Rainfall

Effective rainfall is considered as a function of the regressive rainfall index (Li et al., 2011; Zhuang et al., 2015). We use the ratio of effective rainfall to annual rainfall as the dimensionless effective rainfall (DER) to eliminate the impacts of different annual climate. The equation for calculating DER can be written as:

$$DER = \frac{\sum_{i=1}^{n} K^{i} P_{i} + P_{0}}{P_{annual}}$$
(1)

where P_0 is the total rainfall on the day of debris-flow occurrence; P_i is the accumulative daily rainfall (mm) on the *i*th day before the occurrence; K^i indicating K to the power of *i* is the decay factor of precipitation on the *i*th day before the occurrence; and P_{annual} is the total rainfall for each year. The study areas belong to the Hengduan Mountains in the southwest region of China, therefore, the K is empirically set to 0.84 (Liu et al., 2020). n is set to 3, 7, 11, and 15, respectively, for comparison. The median of DER in each earthquake-affected area is regarded as a proxy of regional rainfall threshold.

Rainfall Intensity-Duration Relationship

The so-called rainfall intensity-duration (I-D) relationship was proposed firstly by Caine (1980). It is a fairly common type of debris-flow threshold (Guzzetti et al., 2007). The original expression of the relationship is of discrepant dimension as the following:



$$I = \alpha D^{-\beta} \tag{2}$$

where the term *I* denotes the mean hourly rainfall intensity (mm/h) of a continuous rainfall process that triggers a debrisflow event; *D* is the rainfall duration (h) for triggering a debris-flow event; and α and β are coefficients. Then, when the curve is drawn in a double logarithmic coordinate system, the α can reflect the intercept, while β can reflect the slope of the straight line. The rainfall conditions above the curve are more likely to trigger debris flows. Although some studies emphasize the physical interpretation of parameters in I-D curve, it is now restricted to some specific kind of debris flows such as runoff-generated debris flows with enough *in-situ* observations (Berti et al., 2020).

To eliminate the impacts of different annual climates before and after an earthquake, the ratio (I_{new}, h^{-1}) of mean hourly rainfall intensity (mm/h) to annual rainfall (mm) is used to replace the term *I* in the **Eq. 2** as follows:

$$I_{new} = \alpha_{new} D^{-\beta_{new}}$$
 2a

where α_{new} and β_{new} are new coefficients. It is noteworthy that only the "day" information of debris-flow occurrence is available. It is assumed that a continue rainfall process begins if when averaged hourly rainfall intensity of grid cells covering the debrisflow gully is greater than 0.05 mm/h. Thus, the I_{new} and D of such rainfall event were used to describe the rainfall process for triggering debris flows.

Impacts of the Earthquake on the Thresholds

The impact of each earthquake on the thresholds is determined by the following equation:

$$Impact = \frac{\left(Thre_{pre} - Thre_{post}\right)}{Thre_{pre}} \times 100\%$$
(3)

where *Thre*_{pre} denotes the median DER or the parameters of I-D curve before the earthquake; and *Thre*_{post} indicates those after the earthquake. For the DER, the positive value of *Impact* indicates the decrease of the regional rainfall threshold. For the parameters of I-D curve, the positive impact on α indicates a decrease of the scaling constant after an earthquake, α_{new} of I_{new} at D = 1 in the power-law relationship indicates that post-earthquake rainfall intensity is smaller or greater than the pre-earthquake intensity. The positive impact on β_{new} indicates a decrease of the shape slope, which should be combined with α_{new} change for threshold analysis.



RESULTS AND DISCUSSION

Seismic Impacts on Dimensionless Effective Rainfall Thresholds

In the WEA, the regional rainfall thresholds are ca. 4.2, 5.8, 6.5, and 7.1% of the annual rainfall for effective antecedent rainy-day settings of 3, 7, 11, and 15, respectively, (**Figure 2A**). After the Wenchuan earthquake, the threshold decreased. When excluding the debris-flow samples on the day when the earthquake occurred, the thresholds are ca. 4.0, 5.1, 5.8, and 5.9% of the annual rainfall. When considering the debris-flow records on that day, the thresholds dropped more significantly (**Figure 2B**). The mean reduction of ca. 18% represent the most serious impact of Wenchuan earthquake on the rainfall condition. It suggests that once such earthquake happens, the short-term rainfall threshold could be much lower than that obtained from data in the following several years.

Before the Lushan earthquake, the thresholds are ca. 4.5, 5.4, 6.3, and 7.4% of the annual rainfall. Then, they reduce to 2.4, 3.8, 4.9, and 5.3% after the quake (**Figure 2C**). Similarly, the thresholds in the JEA are ca. 4.5, 6.0, 6.4, and 7.0% before the earthquake (**Figure 2D**). After the shock, they decrease to ca. 4.0, 4.8, and 6.3% for antecedent rainy-day settings of 3, 7, and 11, respectively. However, the threshold unexpectedly increases slightly when the number of antecedent rainy days is set to 15.

In most cases, the earthquake lowered the DER thresholds (**Figure 3**). The regional impacts show a significant variation when n is set to 3. However, when n is set to other values rather than 3, the impacts only show lower variations. Based on the results with n = 7, 11, and 15, the mean impacts (mean level of seismic effects with different n settings) of the Wenchuan (including co-seismic events) and Lushan earthquakes are ca. 26 and 27%, respectively. When excluding the co-seismic samples

("5.12" events), the mean impact of Wenchuan is ca. 13%. When n is set to 11 and 15, the impacts are not significant in the JEA. Maybe the drier climate of Jiuzhaigou in the northern Sichuan province leads to fewer days of DER. When n is equal to 3 or 7, the mean impact of the Jiuzhaigou earthquakes is ca. 16%.

Seismic Impacts on Rainfall Intensity–Duration Relationships

All post-seismic I-D curves are lower than the pre-seismic ones, indicating that triggering post-quake debris flows needs smaller rainfall intensity or shorter duration (**Figure 4**). When drawing the "5.12" events for the Wenchuan earthquake in the picture (the purple crosses in **Figure 4A**), the triggering rainfall level becomes much lower than the post-seismic mean status, indicating extremely low rainfall threshold for co-seismic debris flows. For the WEA, there are more scattered debris-flow records for a shorter duration than a longer duration. Data shows more discrete records after Wenchuan earthquake than before the quake (**Figure 4A**). It may indicate more randomness on the rainfall intensity to trigger debris flows, which is consistent with the random effect of an earthquake on reducing the stability of the surface material (Zhao et al., 2020). Similar results occur at 3 and 9 h durations in both LEA and JLA (**Figure 4B,C**).

According to the **Eq. 3**, impacts of earthquakes on the parameter α_{new} are ca. 45, 24, and 3% in the WEA, the LEA, and the JEA, respectively (**Figure 5**). Impact of the Wenchuan earthquake on the parameter β_{new} is ca. 12%, while those of the Lushan and Jiuzhaigou earthquakes are ca. -34 and ca. -46%. It indicates that, though the impacts of the Lushan and Jiuzhaigou earthquakes on α_{new} are smaller, the increased slopes also made the threshold lower (**Figure 4B,C** and **Figure 5**). However, it should be noted that a little change of β_{new} may not be important



FIGURE 4 | I-D curves for debris flows. The blue and red lines indicate I-D curves for debris flows before and after the Wenchuan (A), Lushan (B), and Jiuzhaigou (C) earthquakes. In the WEA, the pre-quake I-D curve is fitted with 21 debris-flow events from 2000 to 2007, while post-quake one is fitted with 183 events from 2008 to 2010 when excluding the samples on the day of "5.12" earthquake. The red circle denotes debris flows occurring after May 12, 2008, while the purple cross represents the debris flows occurring on that day. All sample points and fitted curves are drawn in the double logarithmic coordinate system.

due to the limited rainfall duration indicated by the observations. Under these circumstances, the changes of the α_{new} could be considered as the more important aspect affected by the earthquakes.

Discussion Rainfall Threshold Change

The impact of the Wenchuan earthquake is reported to be ca. 20–30% (Tang et al., 2009; Ma et al., 2017). Compared with these studies, debris-flow records from more timely and detailed investigations after earthquakes have been collected, especially for the Wenchuan earthquake, meaning more debris flows with smaller rainfall have been included for statistical analysis. In addition, we supplemented the data analysis on the Lushan earthquake in 2013 and the recent Jiuzhaigou earthquake in 2017, and further explored the impacts of the earthquake on regional rainfall threshold. The impacts of the Lushan and Jiuzhaigou earthquakes on I-D rainfall thresholds are smaller, which may be due to the fact that moderate magnitude earthquakes (Mw 6–7) produce less landslide-associated loose materials (Stoffel et al., 2014).

Lin et al. (2004) concludes that the triggering thresholds such as cumulative antecedent rainfall and hourly rainfall intensity significantly decreased by 33% after the Chi-Chi Earthquake compared to the values before the earthquake. Shieh et al. (2009) surveyed the Wushihkeng watershed and the Mt. Ninety-Nine watershed in the Chi-Chi earthquake area. The maximum effective cumulative rainfall threshold was about 50% of that before the earthquake. Similar outcomes can be found in the Chen-Yu-Lan watershed (Chen, 2011). The 30–50% change caused by the Chi-Chi earthquake is indeed similar to that by the Wenchuan earthquake.

Implication for Post-quake Debris-Flow Warning

As shown in **Table 2**, the moment magnitude of the Wenchuan earthquake reached 7.9, much higher than the others. Although the change in the median DER is close to that influenced by the Lushan earthquake, the α_{new} of the I-D curve in the WEA has been changed by 45%, which is nearly double in the LEA. With a similar moment magnitude, different focal depths affect the results. The change in the median DER in the LEA with a 13 km focal depth is 11% larger than the JEA with 20 km focal depth.

A warning strategy, including the DER and I-D curves can be considered by using similar settings of methods and data as in this study. The DER threshold should be the initial focus. The antecedent rainfall can be obtained from up-to-date meteorological data, such as near-real-time precipitation produced by the land data assimilation system (CLDAS) from the China Meteorological Data Service Center (http://data.cma. cn/), satellite data, and in-situ observation. The daily rainfall data is obtained from the weather forecast. When the earthquake magnitude is ca. 6.5 and the focal depth is ca. 20 km, the change in the rainfall threshold condition can be estimated by referring to ca. 15% in the JEA. When the magnitude is greater than 6.5, but focal depth is less than 15 km, we can refer to a threshold change of 26% in the LEA and WEA. When the DER exceeds the median DER, we may think the rainfall can reach a warning level at the first stage. However, earthquake-affected areas responded differently to the statistical approaches (Table 2), which may be attributed to the heterogeneity of rainfall for triggering debris flows under different surface conditions. Thus, the hourly I-D



FIGURE 5 | Impacts of each earthquake on the I-D curve parameters for debris flows. The positive impact on α_{new} indicates the decrease of the scaling constant. The positive impact on β_{new} means the decrease of the shape slope.

TABLE 2 Earthquake attributes and their impacts on regional rainfall thresholds.								
Earthquake	Mw	Focal depth (km)	Impacts					
			Median of DER (%)	α _{new} (%)				
Wenchuan	7.9	14	26.05	45.00				
Lushan	6.6 6.5	13 20	26.74 15.56	24.00 2.86				
Jiuzhaigou	6.5	20	15.56	2.86				

curve can be used. The degree of I-D curve lowering after a strong shock are referred in **Table 2**. When the mean hourly rainfall intensity is above the curve, a regional warning should be issued.

Due to the lack of meteorological in-situ observations, advanced satellite-based rainfall data can be used as an effective alternative for near real-time analysis, which is widely adopted for disaster research (Jia et al., 2020). The Global Precipitation Measurement (GPM) mission, led by the National Aeronautics and Space Administration and the Japanese Aerospace Exploration Agency has been the latest generation of rainfall products since March 2014 (Hou et al., 2013). With the resolution of 0.1° and 0.5 h, the up-to-date GPM products may be suitable for near real-time analysis based on long-term historical records, which have been expanded back to start-ups (in 2000) by calibration (Huffman et al., 2019). We use the gridded rainfall dataset to check the reliability of the new GPM data for future application in debris-flow monitoring and warning. Correlation coefficients of 0.67, 0.53, and 0.56 are obtained for Jiuzhaigou, Wenchuan, and Lushan areas, respectively, (Figure 6), which indicates that the reliability needs to be further improved, as also confirmed by Liu et al. (2019). Though any satellite-based method will cause inevitable uncertainties (Carr et al., 2015), a combination with the methods used in this study could be an important supplement for debrisflow warning in ungauged areas with earthquakes.



Uncertainty and Limitation

The 40 years gridded rainfall data can provide enough information for this study, but the 3 h resolution and GMT format could cause uncertainty for determining the accurate time of debris-flow occurrence. Furthermore, even if eliminating the effects of fluctuant annual rainfall, the rainfall thresholds cannot avoid abnormal distribution of rainfall within certain year. The statistics in small regions (specified high earthquake-intensity areas) depend on the local information and temporal variability (Nikolopoulos et al., 2017; Gariano et al., 2020). However, most debris-flow disasters always occur in places where there is insufficient or ineffective instrumental observation. **Figure 4** shows that effects of the Lushan and Jiuzhaigou earthquake in the curves are not very obvious as that of Wenchuan earthquake. I-D scatters obtained by other study also support the point (Guo et al., 2021). Guo et al. (2021) pointed that debris flows after Lushan earthquake are triggered by shorter-duration rainfall than in most other regions, including the Wenchuan earthquake area, which may magnify the uncertainty from the rainfall data. In addition, some regional I-D threshold is difficult to be obtained, due to the limited data which cannot meet demand for a rigorous fitting. That is why we emphasize more on the change in the I-D relationships rather than the accurate threshold.

Van Asch et al. (2014) made a link between the rainfall thresholds and sediment material delivered by seismic landslides retained in gullies. Thousands of landslides triggered by the Lushan earthquake have become the material source of debris flows (Tang et al., 2015). But the erosion of channel deposits by overland runoff after Lushan earthquake was the dominant mechanism of debris-flow formation in the LEA (Guo et al., 2021). Compared with the WEA, the debris-flow volume in the LEA is smaller, which may be due to the smaller earthquake magnitude and the wetter climate for vegetation recovery. In addition, free placement of wastes produced in post-quake reconstruction can be initiated more easily to form debris flows. The complex spatial heterogeneity of natural and anthropogenic factors makes great differences of critical rainfall conditions for debris flows in different regions. Even so, it is still difficult to isolate any other factors for discussing the earthquake effects on rainfall for triggering the debris flows.

CONCLUSION

We collected the debris-flow records in the Wenchuan, Lushan, and Jiuzhaigou earthquake-affected areas in Sichuan Province, China. Based on the gridded meteorological dataset with 3 h and 0.1° resolutions, the DER and I-D thresholds were calculated. Then, comparison of the pre and post-seismic rainfall thresholds shows the impacts of the earthquakes with different magnitudes on the rainfall thresholds. Our results demonstrate that the post-quake thresholds are much smaller than the pre-seismic thresholds. The mean impacts of the Wenchuan, Lushan, and Jiuzhaigou earthquakes are ca. 26, 27, and 16%, respectively, with respect to the DER. For the Wenchuan earthquake we found a significantly lowered I-D curve.

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The impacts of the Wenchuan and Chi-Chi earthquakes are comparable. In general, this work is useful for calibrating and improving the rainfall thresholds of post-quake debris-flow warning.

DATA AVAILABILITY STATEMENT

Because of tough field survey and requirements of the cooperative institutions, the debris-flow records should be requested upon approval of the corresponding author.

AUTHOR CONTRIBUTIONS

SL and KH analyzed the results, and wrote the manuscript. SL, QZ, and DT documented all the data. KH, SZ, and XH reviewed and enhanced the paper.

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Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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