



Agricultural Greenhouse Gas Fluxes Under Different Cover Crop Systems

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Cultivated lands that support high productivity have the potential to produce a large amount of GHG emissions, including carbon dioxide (CO₂), nitrous oxide (N₂O), and methane (CH₄). Intensive land management practices can stimulate CO₂, N₂O, and CH₄ emissions from the soil. Cover crop establishment is considered as one of the sustainable land management strategies under warm and humid environmental conditions. To better understand how the incorporation of cover crops affect three major GHGs, we compared trace gas fluxes in a no-till maize field over the whole growing season in 2018 in a no cover crop (Tr) system and three cover crop systems: crimson clover (CC), cereal rye (CR), and living mulch (LM) using white clover. In 2019, we further explored potential differences in the three GHGs between in-row (IR) and between-row (BWR) of maize for LM and Tr systems during the early growing season. Measurements were taken using a cavity ring-down spectroscopy gas analyzer in Watkinsville, GA. In 2018, the highest CO₂ flux (7.00 μmol m⁻² s⁻¹) was observed from BWR of maize for LM. The maximum N₂O flux observed in LM on June 20th in 2018 was when soil N increase rate was the largest. Soils served as sinks for CH₄ and Tr system served as the smallest CH₄ sink compared to the other three cover crop systems. For N₂O, the highest fluxes were observed from the TrIR plot (4.13 μmol m⁻² hr⁻¹) in 2019 with the greatest N inputs. In 2019, we observed a smaller CH₄ sink in TrIR (-0.13 μmol m⁻² hr⁻¹) compared to TrBWR (-0.67 μmol m⁻² hr⁻¹) due potentially to greater NH₄⁺ inhibition effects on CH₄ consumption from greater N fertilizer inputs. The net carbon equivalent (CE) from May 23rd to Aug 16th in 2018, taking into account the three GHG fluxes, soil carbon content, and fertilizer, irrigation, and herbicide application, were 32–97, 35–101, 63–139, and 40–106 kg ha⁻¹ yr⁻¹ for CC, CR, LM, and Tr, respectively. LM had the lowest net CE after removing white clover respiration (-16–60 kg ha⁻¹ yr⁻¹). Our results show that implementing different types of cover crop systems and especially the LM system have some potential to mitigate climate change.

Keywords: climate change, soil greenhouse gas (GHG), cover crop, crimson clover, cereal rye, living mulch system

INTRODUCTION

Agriculture is essential for producing food and sustaining livelihoods, but it is also a large source of greenhouse gas (GHG) emissions, contributing to climate change. These environmental impacts are usually associated with the specific agricultural practice of mono-cropping, whereby a single crop is grown on a certain amount of land year after year, with heavy dependence on fertilizers and

pesticides. 80% of the 1.5 billion hectares of arable land globally are devoted to mono-cropping. Another practice that has recently gained attention is to plant cover crops after harvest to prevent soil erosion, accumulate organic carbon (C) in soil, and to suppress weeds (Blanco-Canqui et al., 2015; Hanrahan and Mahl, 2018). Intercropping is yet another agricultural practice whereby more than one crop is grown on the same land area at the same time (Dyer et al., 2012). Intercropping and using cover crops are considered as Climate-Smart Agriculture (CSA) management practices (Paustian et al., 2016). CSA objectives are to: (1) sustainably increase agricultural productivity and incomes; (2) adapt and build resilience to climate change; and (3) reduce GHG emissions and increase C sequestration from soils (FAO, 2013). While the agricultural impacts on the environment are now widely recognized, we still lack the comprehensive understanding of how different agricultural practices affect the soil, atmosphere, and the ecosystems.

How the incorporation of cover crops could influence the mitigation of climate change is also still not well-understood. Studies suggest that cover crops with N-fixing capability (leguminous cover crop) could provide available N to cash crops without additional N inputs (Schomberg et al., 2006; Turner et al., 2016; Andrews et al., 2018). Even cover crops without having the ability to fix N (non-leguminous cover crop) is considered to scavenge the excessive nitrate in the soil and thereby reduce leaching and nitrous oxide (N₂O) emissions from denitrification (Smith and Tiedje, 1979; Jarecki et al., 2009). However, studies found that increases in N₂O after incorporating cover crops sometimes outweigh the mitigation effects from increased SOC, because mineral N fixed by legume cover crop can stimulate N₂O emissions under denitrification (Mitchell et al., 2013; Basche et al., 2014). Moreover, a meta-analysis showed that increased N₂O emissions cannot fully offset the enhanced SOC storage after adopting cover crops, and the mitigation potential from cover crops are highly site-specific (Guenet et al., 2021). Increased carbon (C) inputs from cover crop biomass can also enhance carbon dioxide (CO₂) emissions. Living mulch (LM) is a recently emerging intercrop system, using leguminous cover crops maintained throughout the whole growing season (Zemenchik et al., 2000; Andrews et al., 2018). There are also conflicting findings as to whether LM would increase or decrease N₂O emissions (Gomes et al., 2009; Turner et al., 2016; Peters et al., 2020).

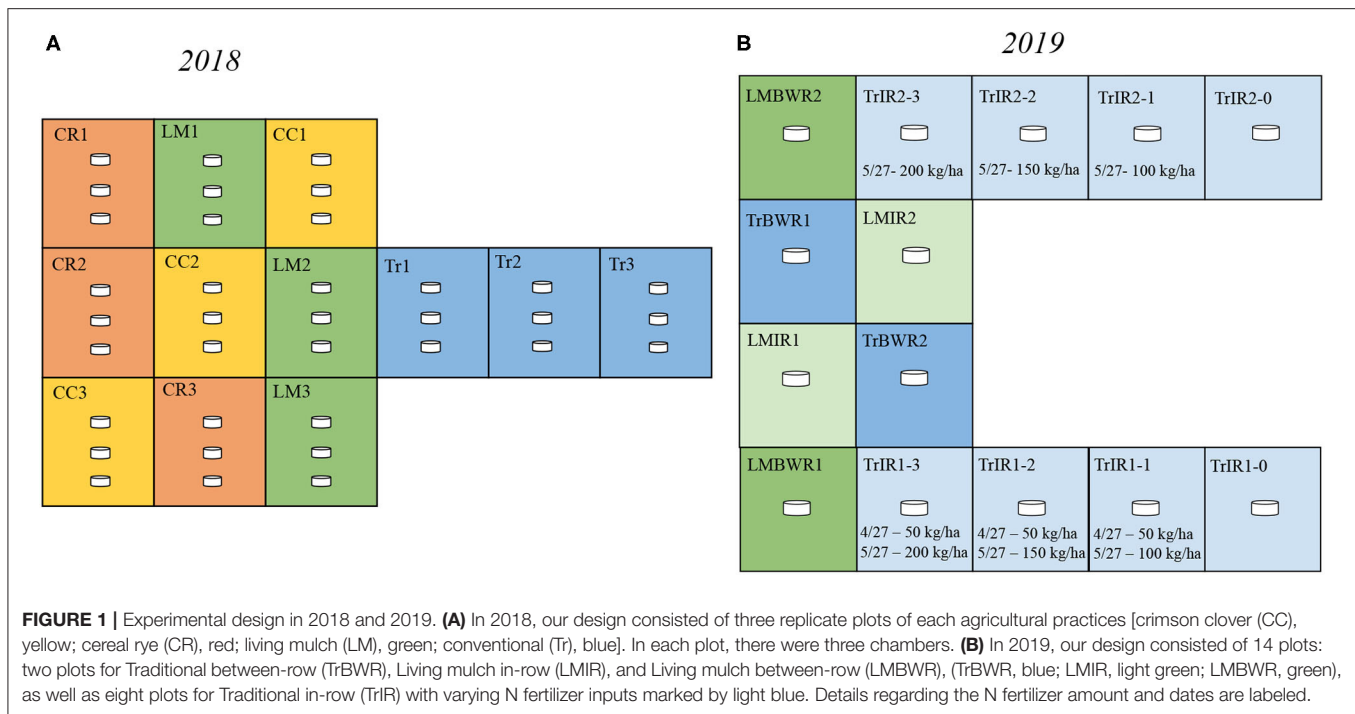
In addition to the impacts on CO₂ and N₂O emissions, there is no simple characterization on how the inputs of cover crop residue will impact CH₄ emissions or uptake capacity. Kim et al. (2012) found that higher CH₄ emissions were observed in cover crops with high C/N ratio due to their higher labile C content, which stimulated CH₄ emissions under anaerobic conditions (Lu et al., 2000; Le Mer and Roger, 2001). Conversely, Boeckx et al. (1996) observed that residue with low C/N led to high CH₄ emissions due to elevated amount of NH₄⁺ and NO₂⁻, which have strong inhibition effects on CH₄ uptake. Meanwhile, previous studies revealed that N fertilizer addition, such as urea, usually exhibits inhibitory effects on CH₄ consumption (Conrad and Rothfuss, 1991; Bronson and

Mosier, 1994; Dunfield and Knowles, 1995). So far, however, it is relatively unknown how N pool in the soil and C/N ratio in cover crop will impact CH₄ uptake in soils (Sanz-Cobena et al., 2014; Guardia et al., 2016; Kaye and Quemada, 2017).

Based on the complexities of how the incorporation of different cover crops may influence three major GHGs, Peters et al. (2020) conducted some experiments in 2016 and 2017 to explore these effects. From 2016 to 2017, trace gas fluxes (CO₂ and N₂O) were measured in between-row (BWR) of maize in crimson clover (CC), cereal rye (CR), LM using white clover, and no cover crop (Tr) systems. CO₂ was also measured in-row (IR) of maize among those four treatments. What is meant by BWR or IR is whether the measurement was taken between rows of maize (BWR) or directly in the planted rows of maize (IR). In their study, researchers used a portable infrared gas analyzer (IRGA, 6400XT, Li-Cor) for CO₂ and gas chromatography with mass spectrometry (GC-MS) for CO₂ and N₂O. Peters et al. (2020) found that LM showed both highest CO₂ and N₂O fluxes in both 2016 and 2017 compared to other treatments. Moreover, significantly higher CO₂ fluxes were observed in BWR of maize, where clover was present, than IR of maize in the LM system in 2016. However, they were unable to detect CH₄ fluxes due to the detection limit of GC-MS.

In addition to the ecological benefits and mitigation potential provided by cover crops, improving adaptive capacity through the adoption of cover crops could also alleviate the impacts of climate change. Adaptive capacity means that the ability of the agricultural system to maintain crop yields while minimizing nutrient losses when facing climate extremes (Kaye and Quemada, 2017). Concerns regarding lower yields when implementing cover crops compared to conventional cropping may result in a lower rate of adoption (Li et al., 2019). However, an 8-year cover crop study conducted by Olson et al. (2010) reported that average corn and soybean yields were similar comparing with and without cover crops. Complexities regarding the impacts of cover crop on yields also originate from local spatial heterogeneity.

To better understand how the incorporation of different agricultural practices affect three major GHGs and grain yields, we continued our study in 2018 and 2019 but with *in-situ* measurements, using a cavity ring-down spectroscopy gas analyzer (G2508, Picarro). Our first primary research question was: how do GHG fluxes and grain yields differ among CC, CR, LM, and Tr? Second, under the same technique (LM or Tr), are soil GHG fluxes in IR of maize significantly different from those in BWR? In this study, we compared trace gas fluxes (CO₂, N₂O, and CH₄) and grain yields in three cover crop systems (CC, CR, and LM) and a no cover crop system (Tr) in a no-till maize field over the whole growing season in 2018. In 2019, in order to further explore potential differences for all three major GHGs between IR and BWR of maize, we focused on these measurements for LM and Tr systems and conducted more intensive daily measurements surrounding the two fertilization periods in the early growing season.



METHODS

Site Description and Experimental Design

The study site was located at the West Unit of the University of Georgia's J. Phil Campbell Sr. Resource and Education Center in Watkinsville, GA, USA. More details regarding the location of the research station can be found in Peters et al. (2020). This site has been established for agricultural research since 1937 (Melancon, 2014). The classification of the soil is Cecil sandy loam analyzed by soil profile pedology near the research plots. Our experiment in 2018 consisted of twelve 6.1×7.3 m plots in no-till maize fields, three each for four treatments: (1) with dead crimson clover (CC) cover crop; (2) with dead cereal rye (CR) cover crop; (3) living much (LM) system, with white clover intercrop; and (4) traditional (no cover crop) system (Tr) with bare soil (**Figure 1**). Within each plot, three chambers were placed to account for soil heterogeneity. The surface area and volume of the chamber were 0.0182 m^2 and $2.92 \times 10^{-3} \text{ m}^3$, respectively. Each cycle of sample collection consisted of taking CO_2 , N_2O , and CH_4 measurements from nine chambers (three chambers per plot from three plots) in BWR of maize over 81 min. We measured the accumulation of trace gases from nine chambers in one cycle using a multiplexer, sampling each chamber for 90 s, and rotating over the nine chambers six times in a cycle. Four cycles were typically used to collect measurements from all chambers on a sampling day. We quantified soil GHG flux weekly over the whole growing season in 2018. The last day of measurement (September 13) was taken after the maize was harvested. We used this day's observation as a baseline value and removed it from significance tests for GHG fluxes. In 2018, CC, CR, and Tr received the same amount of N inputs at the rate of 56 kg ha^{-1} on April 20. The

second fertilizer application took place on May 18 and CC, CR, and Tr received N inputs at the rate of 168, 224, and 224 kg ha^{-1} , respectively. LM did not receive any N inputs.

In 2019, we conducted an intensive field campaign observing daily soil GHG fluxes at an early growing season, with a focus around the two fertilization periods. We measured the accumulation of the same gases under Tr and LM systems, with a focus on comparing the GHG emissions from IR and BWR of maize. We included the following: (a) traditional in row (TrIR); (b) traditional between row (TrBWR); (c) living mulch in row (LMIR); and (d) living mulch between row (LMBWR) (**Figure 1**). Our experimental setup in 2019 consisted of fourteen plots in no-till maize fields, two each for TrBWR, LMIR, and LMBWR, and remaining eight plots for TrIR. Within each plot, one chamber was established to sample CO_2 , N_2O , and CH_4 fluxes. Each cycle of sample collection (84 min total) consisted of taking measurements from seven chambers (four chambers from TrIR and one chamber per plot from three treatments: TrBWR, LMIR, and LMBWR). We measured the accumulation of trace gases from seven chambers in one cycle using a multiplexer. Each chamber was sampled for 60 s, followed by sampling the ambient air through a filter for another 60 s before sampling the next chamber. This sequence was repeated six times in a cycle for all seven chambers. We typically used two cycles to collect measurements from all chambers on a sampling day.

LMIR received N fertilizer at the rate of 35 kg ha^{-1} on April 27, while both TrBWR and LMBWR did not receive any. For TrIR, eight plots were established to explore GHG flux variations depending on different fertilization amounts and times. TrIR was separated into TrIR1 and TrIR2 with different fertilization rates. In TrIR1, all four paired plots were fertilized twice during our

TABLE 1 | Field events timeline.

2018						
Treatments	Corn sowing	First fertilization (20 April)	Second fertilization (18 May)	Irrigation	Corn harvest	Herbicide application
CC	20 April	56 kg ha ⁻¹	168 kg ha ⁻¹	11 June, 22	22 August	6 April –0.5 kg ha ⁻¹
CR	20 April	56 kg ha ⁻¹	224 kg ha ⁻¹	June, 6 July and		15 March–1 kg ha ⁻¹
LM	20 April	N/A	N/A	31 July –20 mm		6 April–1.5 kg ha ⁻¹
Tr	20 April	56 kg ha ⁻¹	224 kg ha ⁻¹	water application		N/A
2019						
Treatments	Corn sowing	First fertilization (27 April)	Second fertilization (27 May)	Irrigation		
TrIR	27 April	50 kg ha ⁻¹ (TrIR1)	ranging from 0 to 200 kg ha ⁻¹ (TrIR1 and TrIR2), see more details in Figure 1B	3 May–20 mm water application		
TrBWR	27 April	N/A	N/A			
LMIR	27 April	35 kg ha ⁻¹	N/A			
LMBWR	27 April	N/A	N/A			

measurement campaign. On April 27, all four plots (TrIR1-0, TrIR1-1, TrIR1-2, and TrIR1-3) received N fertilizer at the rate of 50 kg ha⁻¹. In addition, four plots in TrIR1 also received fertilizer on May 27, ranging from 0 to 200 kg ha⁻¹ (TrIR1-0: 0 kg ha⁻¹; TrIR1-1: 100 kg ha⁻¹; TrIR1-2: 150 kg ha⁻¹; TrIR1-3: 200 kg ha⁻¹). In TrIR2, all four plots only received fertilizer once on May 27, and the rate ranged from 0 to 200 kg ha⁻¹ (TrIR2-0: 0 kg ha⁻¹; TrIR2-1: 100 kg ha⁻¹; TrIR2-2: 150 kg ha⁻¹; TrIR2-3: 200 kg ha⁻¹). Irrigation events occurred on May 3, when 20 mm water was applied to all four treatments to ensure that the soil water content was above 40%. Summarized field events timeline and the stages of corn development can be found in **Table 1** and **Supplementary Table 1**, respectively.

Soil Sampling and GHG Flux Measurements

We used the G2508 Picarro concentration analyzer to measure the accumulation of trace gases in the chambers. The working principle of Picarro G2508 is based on cavity ring-down spectroscopy (CRDS). CRDS technology utilizes the beam that enters into the ring-down cavity formed by two or more high-reflectivity mirrors (Picarro G2508, Santa Clara, CA, USA, 2021). Three mirrors are used in the Picarro concentration analyzer to sustain the continuous traveling wave. Trace gases such as CO₂, N₂O, and CH₄ have their unique absorption spectrum within the near-infrared range. By measuring the absorption intensity under this wavelength, it can determine the concentration of a specific gas (Picarro G2508, Santa Clara, CA, USA, 2021). CRDS extends the effective path length for absorbing up to several kilometers, and the sensitivity of gas concentration can reach parts per billion level in a few seconds (Picarro G2508, Santa Clara, CA, USA, 2021). With the CRDS technique, Picarro has the capability

to measure CO₂, N₂O, CH₄, NH₃, and H₂O simultaneously and ensures the data collection under high temporal resolution (Picarro G2508, Santa Clara, CA, USA, 2021). We used water-corrected trace gas mixing ratios for analysis.

Other studies have found a relationship between soil GHG fluxes and other variables, such as soil volumetric water content, soil temperature, and total amount of soil C and N (Franzluebbers, 2005; Steenwerth and Belina, 2008; Camarotto et al., 2018). CS 625 reflectometers (Campbell Scientific) were used to measure soil moisture and temperature data. The CS 625 reflectometers were placed at 0–15 cm soil depth in the center of the two center rows of each plot. The length of the rods is 30 cm and the rods were installed at an angle of 45° from the surface. Soil moisture and temperature data were collected at a 10-min interval to calculate hourly averages. Total soil C and N amounts were measured by combusting the soil at 1350°C under aerobic condition to convert C and N to CO₂ and N₂ in a weekly-manner in May and June and monthly in July and August. Clover biomass and corresponding soil N content were also measured in all LM plots in 2018 to investigate the effects of clover residue decomposition on N₂O emission spikes at the late growing season. Meteorological data were obtained from a weather station located near experimental plots. Surface pressure and atmospheric temperature were used to calculate soil GHG flux.

Maize Harvest and Yield

We harvested the yield of maize on Aug 22nd for CC, CR, and LM in 2018. After the maize harvest, we also removed the corn residue from all the plots to ensure the re-seeding of CC and CR and similar soil preparation in each system. In LM system, we also applied pendimethalin at the rate of 1 kg ha⁻¹ after the removal of corn residue for controlling the weed in the winter and allowing

the better reestablishment of the white clover during the late summer to the next spring. We harvested maize ears by husking manually from a 3-m segment of the center two maize rows in each plot. We weighed the shelled grain, which was removed by using a mechanical corn sheller after drying the harvested ears at 60°C. The maize yield was calculated by adjusting the moisture at 15%.

Data Analysis

Trace gas flux calculation was carried out using Interactive Data Language (IDL). The mixing ratio ($\mu\text{mol mol}^{-1}$) of trace gases was first converted, following the Ideal Gas Law (1), where P is surface pressure (atm), R is a gas constant 8.205×10^{-5} ($\text{m}^3 \text{ atm mol}^{-1} \text{ K}^{-1}$), and T is air temperature (°C). Trace gas fluxes ($\mu\text{mol m}^{-2} \text{ s}^{-1}$ or $\mu\text{mol m}^{-2} \text{ hr}^{-1}$) were then calculated by the following equation using the change in a gas mixing ratio over a specific time period (t), where t denotes the time period (s or h), V represents a chamber volume [2.92×10^{-3} (m^3)], and A represents chamber surface area [0.0182 (m^2)], (Collier et al., 2014).

$$\text{Flux} = \text{Mixing ratio}/t \times \frac{P \times V}{R * (273 + T) * A} \quad (1)$$

Since CO_2 mixing ratio is greater in magnitude, the flux unit of CO_2 is calculated in ($\mu\text{mol m}^{-2} \text{ s}^{-1}$), while fluxes of N_2O and CH_4 are calculated in ($\mu\text{mol m}^{-2} \text{ hr}^{-1}$). Positive surface flux is an indication that soils emit the gas to the atmosphere. When the surface flux is negative, soils serve as a sink, uptaking gas from the atmosphere. The quality of trace gas fluxes data was checked by assessing the linearity of the mixing ratio increase inside the closed chamber. Only measurements with $R^2 \geq 0.7$ were retained and used in the subsequent statistical analysis (Peters et al., 2020). In 2018, 343, 222, and 323 out of total 343 observations were kept for statistical analysis for CO_2 , N_2O , and CH_4 , respectively. In 2019, 200 out of total 215 observations were kept for statistical analysis for CO_2 , N_2O , and CH_4 fluxes.

Statistics

All the statistical analyses in this study were performed using R 3.6.1 (R Core Team, 2019). Prior to statistical analysis, normality of all fluxes was assessed using Shapiro-Wilks test (Das and Imon, 2016). Mean CO_2 flux throughout the whole measuring period in 2018 was transformed by taking the power of $-1/2$ and mean CO_2 flux separated in both early and late growing season in 2018 was transformed by taking the power of $-1/3$ to meet the normality, as log transformation failed to meet the normality test for both years. ANOVA and Tukey's pairwise comparison was further implemented for transformed CO_2 flux to investigate which specific agricultural practices were statistically significantly different in their means. Welch t -test was carried out for CO_2 flux in 2019, as well as for N_2O and CH_4 fluxes due to the failure of meeting normality assumption for all the attempted transformations. The means of CO_2 , N_2O and CH_4 fluxes in 2019 were compared to explore whether there were significant differences between two sets of data. GHG fluxes among the four practices were also compared between different time periods. In

2018, we also analyzed soil GHG fluxes from four practices in early and late growing seasons separately. In 2019, we similarly analyzed GHG fluxes in the first and the second fertilization periods separately.

Net carbon equivalent (CE) was calculated and compared among four practices (CC, CR, LM, and Tr) from May 23rd to Aug 16th in 2018, including CO_2 , N_2O and CH_4 fluxes generated in the field, soil labile carbon content, as well as CE estimates due to the consumption of fertilizer, irrigation and herbicides (Lal, 2004). For three major GHGs, net CE was calculated by first converting three major GHGs' units from ($\mu\text{mol m}^{-2} \text{ s}^{-1}$ or $\mu\text{mol m}^{-2} \text{ hr}^{-1}$) to ($\text{kg ha}^{-1} \text{ yr}^{-1}$). After that, we used global warming potential (GWP) for a 100-year time horizon to standardize three major GHGs' different climate impacts. N_2O and CH_4 have 273, and 28 times greater GWP compared to CO_2 , respectively (Smith et al., 2021). In addition, land management practices for agriculture also lead to GHG emissions into the atmosphere. Lal (2004) summarized a range of estimates of CE for different farm operations: 0.9–1.8 kg CE per kg N fertilizer use, 6.3 kg CE per kg herbicides use, and 31–227 kg CE for applying 20 cm water irrigation. We calculated net CO_2 eq for N fertilizer and herbicides use (kg) based on the area of our study plots. Each agricultural practice was assumed to occupy an area of 133.59 m^2 . We did not include the time period from April 22nd to May 18th into calculating the CE because of missing N_2O and CH_4 fluxes data which were removed based upon failure of meeting $R^2 \geq 0.7$. To remove the potential impacts of white clover respiration on observed CO_2 flux in the LM system, we further subtracted the estimated white clover respiration in LM. We followed the methodology of Peters et al. (2020) and assumed a similar growth rate of white clover as in 2017.

RESULTS

GHG Fluxes in 2018

We calculated fluxes of three major GHGs (CO_2 , N_2O , and CH_4) by conducting measurements weekly in BWR of maize from the four different agricultural practices during the maize growing season in 2018 (Figure 2). Similar to the findings in Peters et al. (2020), BWR measurements from LM were observed to produce statistically higher CO_2 fluxes compared to CC, CR, and Tr, all at $p < 0.001$ level. In this study, the other two cover crop systems (CC and CR) also emitted significantly higher CO_2 fluxes than Tr (Table 2), which was not observed in Peters et al. (2020). Between leguminous (CC) and non-leguminous (CR) cover crops, CO_2 produced were not statistically significantly different (Table 2).

We measured higher overall N_2O fluxes from CR ($2.07 \mu\text{mol m}^{-2} \text{ hr}^{-1}$) and N_2O flux from Tr increased on May 23rd after receiving fertilizer inputs (Figure 2). Greater fluxes were also observed in LM on June 20th (Figure 2), which was most likely due to the elevated soil N added from the decomposing white clover. In 2018, N_2O fluxes from the four practices were not statistically significantly different from each other (Supplementary Table 2). This contradicts the finding from

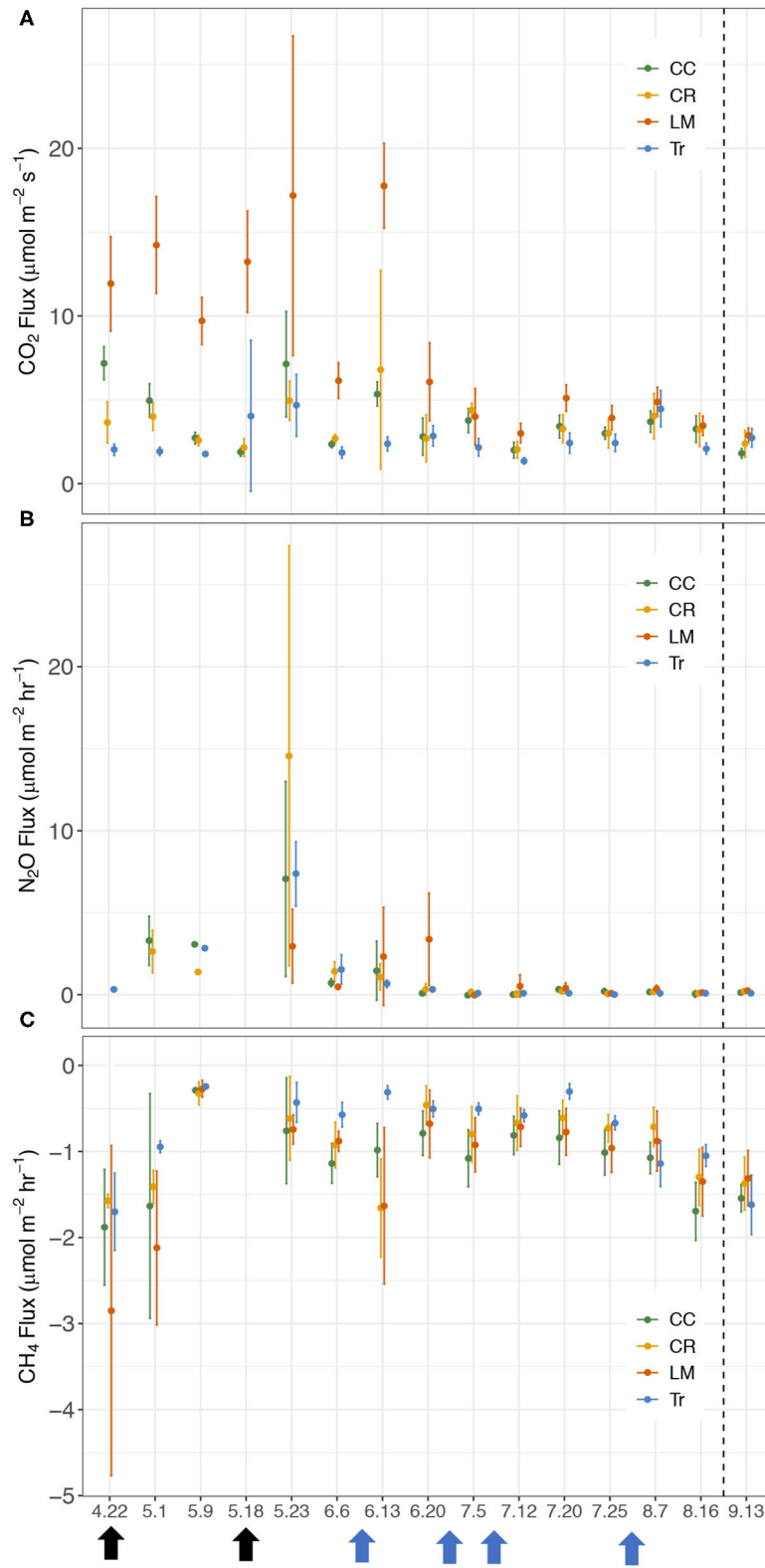


FIGURE 2 | Temporal changes of daily average GHG fluxes between the maize rows in 2018 growing season for crimson clover (CC), cereal rye (CR), white clover living mulch (LM), and conventional (Tr) practices. **(A)** CO₂ ($\mu\text{mol m}^{-2} \text{s}^{-1}$), **(B)** N₂O ($\mu\text{mol m}^{-2} \text{hr}^{-1}$), **(C)** CH₄ ($\mu\text{mol m}^{-2} \text{hr}^{-1}$). Black arrows represent two fertilizer applications (4/22 and 5/18) and blue arrows represent four irrigation events (6/11, 6/22, 7/6, and 7/31).

TABLE 2 | Mean transformed CO₂ flux ($n = 342$) group comparisons from four treatments in 2018.

Groups	Difference ($\mu\text{mol m}^{-2} \text{s}^{-1}$)	95% CI lower bound	95% CI upper bound	p -value
2018				
CR-CC	0.015	-0.036	0.067	0.86
LM-CC	-0.13	-0.18	-0.079	$<1.00 \times 10^{-7***}$
Tr-CC	0.089	0.04	0.14	$2.34 \times 10^{-5***}$
LM-CR	-0.14	-0.19	-0.095	$<1.00 \times 10^{-7***}$
Tr-CR	0.073	0.024	0.12	$7.88 \times 10^{-4} **$
Tr-LM	0.22	0.17	0.27	$<1.00 \times 10^{-7***}$

*** $p < 0.001$.

Peters et al. (2020), where LM was found to emit significantly higher N₂O fluxes than CC, CR, and Tr in both 2016 and 2017.

Soils in all four systems exhibited as CH₄ sinks (Figure 2). Tr showed a significantly lower CH₄ uptake rate ($-0.65 \mu\text{mol m}^{-2} \text{hr}^{-1}$) compared to the other three cover crop systems (Tr-CC: $p < 0.001$; Tr-CR: $p < 0.01$; Tr-LM: $p < 0.001$; Supplementary Table 3). CC ($-1.08 \mu\text{mol m}^{-2} \text{hr}^{-1}$) and LM ($-1.00 \mu\text{mol m}^{-2} \text{hr}^{-1}$) served as larger sinks compared to other treatments and the two did not differ significantly ($p = 0.43$). Comparing CH₄ uptake capacity between leguminous (CC) and non-leguminous cover crop (CR), CC served as a significantly larger CH₄ sink compared to CR, with the difference of $0.25 \mu\text{mol m}^{-2} \text{hr}^{-1}$ ($p < 0.01$; Supplementary Table 3).

Both CO₂ and N₂O fluxes were much lower in mid-July and August compared to the earlier period of the growing season (Supplementary Figure 1). On the other hand, soils served as a larger CH₄ sink in the latter half of the growing season than in earlier period (Supplementary Figure 1). We thus separated the whole growing season into early (April 22nd–July 5th) and late (July 12th–August 16th) periods to explore GHG flux variations among the four treatments better. During the early period, the average CO₂ fluxes in CC, CR, LM, and Tr were 4.07 ± 2.28 , 3.67 ± 2.03 , 9.93 ± 6.33 , and $2.71 \pm 1.54 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively (Table 3; Supplementary Figure 1). In the late period, the average CO₂ fluxes in CC, CR, LM, and Tr were 2.99 ± 0.83 , 3.02 ± 1.08 , 3.93 ± 1.02 , and $2.56 \pm 1.21 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively (Table 3). As found during the whole growing season, the differences between LM and each of the three other agricultural treatments were statistically significantly different in both periods (early: all $p < 0.001$; late: LM-CC: $p < 0.01$; LM-CR: $p < 0.01$; LM-Tr: $p < 0.001$; Table 4). Moreover, both CC and CR also emitted higher CO₂ fluxes than Tr in both periods, as found during the whole growing season (Table 2).

The highest N₂O fluxes were measured from the CR system ($3.76 \mu\text{mol m}^{-2} \text{hr}^{-1}$) during the early growing season and from the LM system ($0.25 \mu\text{mol m}^{-2} \text{hr}^{-1}$) in the late growing seasons. In the early growing season, we did not observe statistically significant difference in N₂O fluxes among four practices (Supplementary Table 4). However, in the late growing season, N₂O fluxes in LM were significantly higher than those in CR and Tr (LM-CR: $p < 0.05$; LM-Tr: $p < 0.01$; Supplementary Table 4). Differences in N₂O fluxes between

TABLE 3 | Average and standard deviation of trace gas fluxes (CO₂: $\mu\text{mol m}^{-2} \text{s}^{-1}$, N₂O and CH₄: $\mu\text{mol m}^{-2} \text{hr}^{-1}$) for four systems throughout the whole measurement period (CO₂: $n = 342$; N₂O: $n = 222$; CH₄: $n = 323$) and both early (CO₂: $n = 181$; N₂O: $n = 129$; CH₄: $n = 163$) and late (CO₂: $n = 161$; N₂O: $n = 93$; CH₄: $n = 160$) growing seasons in 2018.

Treatment	CO ₂	N ₂ O	CH ₄
The whole measurement period (April 22–Aug 16)			
CC	3.55 ± 1.81	1.42 ± 3.15	-1.08 ± 0.51
CR	3.35 ± 1.66	2.07 ± 5.99	-0.83 ± 0.44
LM	7.00 ± 5.47	1.24 ± 1.99	-1.00 ± 0.62
Tr	2.64 ± 1.40	1.33 ± 2.49	-0.65 ± 0.36
Early growing season (April 22–July 5)			
CC	4.07 ± 2.28	2.12 ± 3.76	-1.05 ± 0.59
CR	3.67 ± 2.03	3.76 ± 7.85	-0.83 ± 0.51
LM	9.93 ± 6.33	2.22 ± 2.46	-1.06 ± 0.81
Tr	2.71 ± 1.54	2.06 ± 2.89	-0.56 ± 0.35
Late growing season (July 12–August 16)			
CC	2.99 ± 0.83	0.16 ± 0.19	-1.11 ± 0.43
CR	3.02 ± 1.08	0.11 ± 0.12	-0.83 ± 0.36
LM	3.93 ± 1.02	0.25 ± 0.27	-0.95 ± 0.38
Tr	2.56 ± 1.21	0.08 ± 0.09	-0.75 ± 0.34

TABLE 4 | Mean transformed CO₂ flux group comparisons from four treatments in early ($n = 181$) and late ($n = 161$) growing seasons in 2018.

Groups	Difference ($\mu\text{mol m}^{-2} \text{s}^{-1}$)	95% CI lower bound	95% CI upper bound	p -value
2018 Early growing season (April 22–July 5)				
CR-CC	0.02	-0.043	0.081	0.86
LM-CC	-0.16	-0.22	-0.096	$<1.0 \times 10^{-6***}$
Tr-CC	0.084	0.026	0.14	$1.21 \times 10^{-3**}$
LM-CR	-0.18	-0.24	-0.11	$<1.0 \times 10^{-7***}$
Tr-CR	0.065	0.007	0.12	$2.12 \times 10^{-2*}$
Tr-LM	0.24	0.18	0.30	$<1.0 \times 10^{-7***}$
2018 Late growing season (July 12–Aug 16)				
CR-CC	-0.0032	-0.100	0.094	0.99
LM-CC	0.14	0.042	0.24	$1.46 \times 10^{-3**}$
Tr-CC	-0.088	-0.18	0.0065	5.81×10^{-2}
LM-CR	0.14	0.045	0.24	$1.07 \times 10^{-3**}$
Tr-CR	-0.088	-0.18	0.005	7.23×10^{-2}
Tr-LM	-0.23	-0.32	-0.14	$1.00 \times 10^{-7***}$

* $p < 0.05$, ** $p < 0.01$, *** $p < 0.001$.

leguminous cover crop (CC) and non-leguminous cover crop (CR) did not differ significantly in either period (early: $p = 0.32$; late: $p = 0.35$; Supplementary Table 4).

Soils in all four treatments functioned as CH₄ sinks. In the early growing season, Tr showed lower CH₄ uptake rate compared to the other three cover crop systems (CC-Tr: $p < 0.001$; CR-Tr: $p < 0.01$; LM-Tr: $p < 0.001$; Supplementary Table 5). In the late growing season, however, statistically significantly lower CH₄ uptake rate in Tr was only

observed in comparison with the N-fixing cover crop systems—CC and LM (**Supplementary Table 5**). Compared to CC and CR, soils in LM were not a larger CH₄ sink in either of the early or late growing seasons (**Supplementary Table 5**). Between leguminous cover crop (CC) and non-leguminous cover crop (CR), a larger CH₄ sink was observed in CC compared with CR only in the late growing season (early: CC-CR: $p = 0.09$; late: CC-CR: $p < 0.01$; **Supplementary Table 5**).

From the measurements of the average maize yields in 2018, Hill et al. (2021) reported that similar yields were found for CC (14.73 Mg ha⁻¹) and CR (16.06 Mg ha⁻¹), while statistically significantly lower yields under LM (12.78 Mg ha⁻¹). Achieving both soil health and increased yields remains a challenge. In order to better assess the mitigation potential of different cover crop systems compared to the conventional agricultural system, we also compared four agricultural practices with a net CE. The net CE from CC, CR, LM, and Tr were 32–97, 35–101, 63–139, and 40–106 kg ha⁻¹ yr⁻¹, respectively. After removing the estimated white clover respiration, the net CE from LM was -160–60 kg ha⁻¹ yr⁻¹.

GHG Fluxes in 2019

We conducted an intensive daily measurement campaign surrounding the two fertilization dates (April 27th–May 4th and May 23rd–June 1st) during the early growing season in 2019. Daily GHG fluxes were measured for a week in both BWR of maize and IR of maize from Tr and LM plots (**Figure 3**). On the first fertilization date (April 27th), LMIR received N fertilizer inputs at the rate of 35 kg ha⁻¹, and three chambers in TrIR1 (TrIR1-1, TrIR1-2, TrIR1-3) also received N fertilizer at the rate of 50 kg ha⁻¹. On the second fertilization period (May 27th), six chambers in TrIR received N fertilizer, ranging from 0 to 200 kg ha⁻¹. LMBWR and TrBWR did not receive any N addition in either period. Higher CO₂ and N₂O fluxes were observed in the second measurement period compared to the first (**Supplementary Figure 2**). At the same time, the magnitude of the CH₄ sink was also larger in the second period than the first (**Supplementary Figure 2**). We thus first compared GHG fluxes during the whole measurement period for TrIR, TrBWR, LMIR, and LMBWR, then separated the measurements in 2019 into the first measurement period (April 27th–May 4th) and second (May 23rd–June 1st) to explore how GHG fluxes generated in BWR of maize and IR of maize from Tr and LM plots varied in the two measurement periods.

The highest CO₂ fluxes were observed from LMBWR (8.31 μmol m⁻² s⁻¹; **Table 5**), similar to what was reported in Peters et al. (2020) for years 2016 and 2017. Peters et al. (2020) found that in 2016, higher CO₂ flux in LM compared to Tr was only observed in BWR of maize, where clover was present but not within rows. However, in 2019, higher CO₂ fluxes were observed in both LMIR and LMBWR as compared to TrIR and TrBWR, respectively (both $p < 0.001$; **Table 6**). All four (TrIR, TrBWR, LMIR, and LMBWR) showed increased CO₂ fluxes during the second fertilization period than the first, especially for LMBWR (**Figure 3**). The CO₂ difference between LMIR and LMBWR was greater after the second fertilization period than the first (first:

LMBWR-LMIR: 1.87 μmol m⁻² s⁻¹; second: LMBWR-LMIR: 6.26 μmol m⁻² s⁻¹; **Table 7**).

Throughout the whole measurement period in 2019, we observed the highest average N₂O fluxes from TrIR (4.13 μmol m⁻² hr⁻¹; **Table 5**). Peters et al. (2020) observed higher N₂O flux in LM compared to Tr in 2016 and 2017. Unlike the previous study, we did not observe the highest N₂O fluxes from LMBWR in 2019. The average N₂O fluxes generated in both IR and BWR of Tr were higher in the second fertilization period than those in the first, while the opposite was true for the LM system. Around the first fertilization period, the highest average N₂O flux was observed in LMIR (2.94 μmol m⁻² hr⁻¹), which was however only slightly greater than those in TrBWR ($p < 0.05$; **Supplementary Table 6**). Comparing N₂O fluxes between IR and BWR, higher N₂O fluxes were found in TrIR than in TrBWR in the first fertilization period ($p < 0.001$; **Supplementary Table 6**). In the second fertilization period, we observed significantly higher N₂O fluxes in TrIR (5.93 μmol m⁻² hr⁻¹) compared to TrBWR and LMIR (TrIR-TrBWR: $p < 0.05$; TrIR-LMIR: $p < 0.05$; **Supplementary Table 6**). In LM, IR and BWR measurements did not show significant differences in N₂O fluxes generated in neither first nor second fertilization periods (first: LMIR-LMBWR: $p = 0.69$; second: LMIR-LMBWR: $p = 0.50$; **Supplementary Table 6**).

We observed a statistically significantly smaller CH₄ sink in TrIR (-0.13 μmol m⁻² hr⁻¹; **Table 5**) compared to TrBWR and LMBWR in 2019 (both $p < 0.001$; **Supplementary Table 7**). A statistically significantly larger CH₄ sink was observed in TrBWR compared with TrIR in both first and second fertilization periods (early: TrBWR-TrIR: $p < 0.001$; late: $p < 0.05$; **Supplementary Table 8**). However, LMIR and LMBWR only showed statistically significant CH₄ flux difference in the first fertilization period (LMIR-LMBWR: -0.33 μmol m⁻² hr⁻¹, $p < 0.05$; **Supplementary Table 8**).

DISCUSSION

CO₂ and N₂O Fluxes

We observed statistically significantly higher CO₂ fluxes in both LMIR and LMBWR compared with TrIR and TrBWR, respectively. Our findings differ from a previous study, where higher CO₂ fluxes were only found in LMBWR compared to TrBWR (Peters et al., 2020). There are two potential reasons explaining why we also observed higher CO₂ fluxes in LMIR compared to both TrIR and TrBWR. The first reason is that we conducted our experiment in an intensive daily manner with *in-situ* measurement, which made it possible to observe the difference better in 2019. The second potential reason is that the soil C and N amount changed in BWR of maize as well as IR of maize after 3-year incorporation of LM in the soil. For example, from observing measurements of soil C and N amount, in 2019, soils in LMIR had a larger C but a lower N during the early growing season when we conducted our measurements (**Figure 4**). Although white clover is present only in BWR, it is clear that soils in IR of maize were also impacted from the clover presence as the LM soil C content was greater than that in Tr system in 2019.

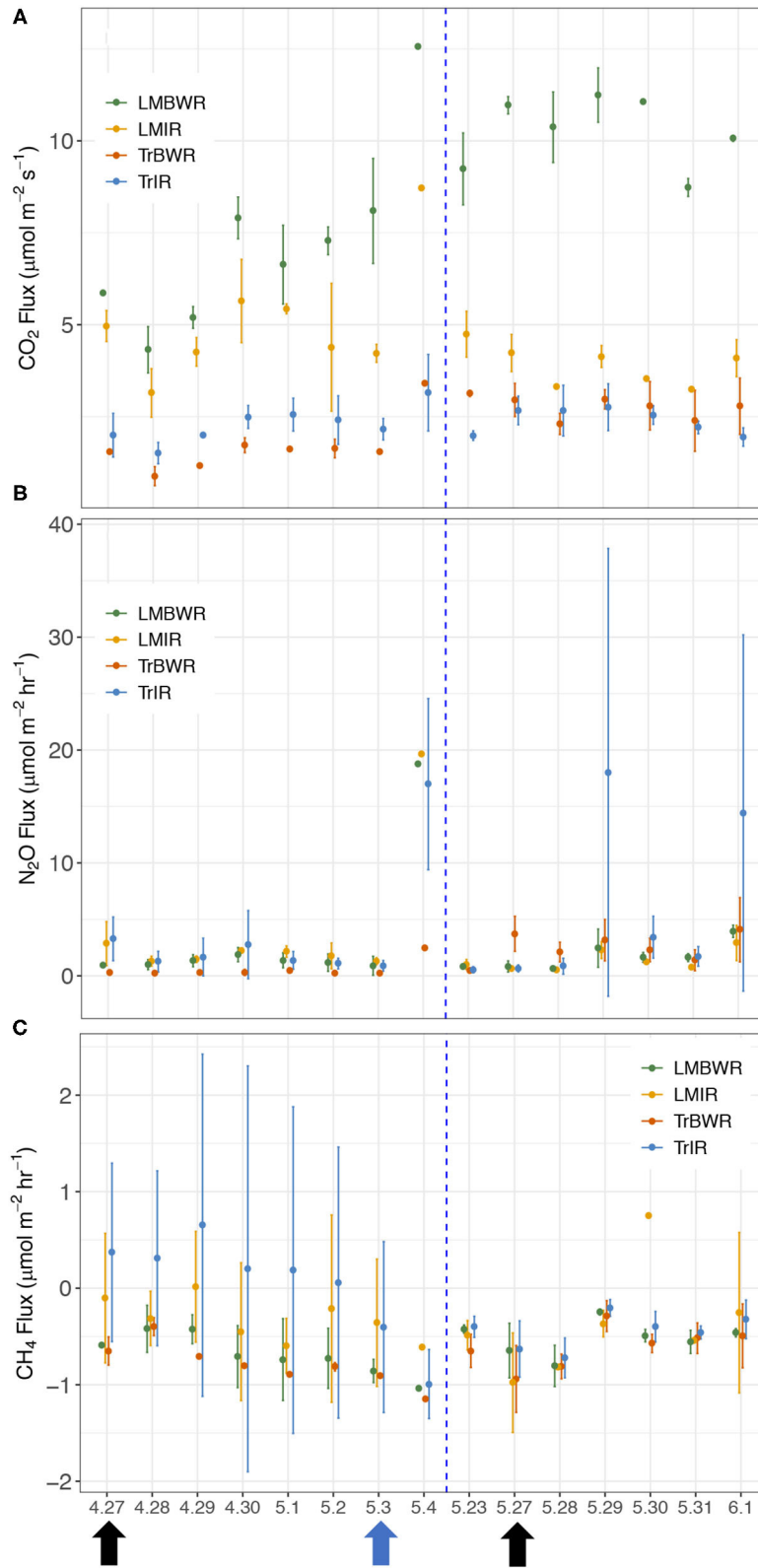


FIGURE 3 | Temporal changes of daily average GHG fluxes between the maize rows in 2019 during early growing season in tradition in row (TrIR), traditional between the row (TrBWR), living mulch in row (LMIR), and living mulch between the row (LMBWR). **(A)** CO₂ ($\mu\text{mol m}^{-2} \text{s}^{-1}$), **(B)** N₂O ($\mu\text{mol m}^{-2} \text{hr}^{-1}$), **(C)** CH₄ ($\mu\text{mol m}^{-2} \text{hr}^{-1}$). Black arrows represent twice fertilizer application (4/27 and 5/27), and blue arrow represents one irrigation event (5/3) during the measurement period.

TABLE 5 | Average and standard deviation of trace gas fluxes (CO_2 : $\mu\text{mol m}^{-2} \text{s}^{-1}$, N_2O and CH_4 : $\mu\text{mol m}^{-2} \text{hr}^{-1}$) for four measurements throughout the whole measurement period ($n = 200$ for all trace gas species) and both first ($n = 107$ for all trace gas species) and second ($n = 93$ for all trace gas species) fertilization periods in 2019.

Treatment	CO_2	N_2O	CH_4
The whole measuring time (April 27–June 1)			
TrIR	2.25 ± 0.59	4.13 ± 8.62	-0.13 ± 1.00
TrBWR	2.23 ± 0.81	1.54 ± 1.64	-0.67 ± 0.25
LMIR	4.42 ± 1.23	2.28 ± 3.54	-0.38 ± 0.51
LMBWR	8.31 ± 2.53	2.03 ± 3.30	-0.58 ± 0.25
The first fertilization period (April 27–May 23)			
TrIR	2.14 ± 0.64	2.62 ± 4.34	0.13 ± 1.29
TrBWR	1.61 ± 0.65	0.46 ± 0.64	-0.75 ± 0.22
LMIR	4.75 ± 1.48	2.94 ± 4.54	-0.31 ± 0.48
LMBWR	6.62 ± 2.23	2.31 ± 4.42	-0.64 ± 0.28
The second fertilization period (May 27–June 1)			
TrIR	2.38 ± 0.49	5.93 ± 11.64	-0.43 ± 0.23
TrBWR	2.76 ± 0.50	2.46 ± 1.68	-0.61 ± 0.26
LMIR	3.98 ± 0.60	1.40 ± 1.07	-0.47 ± 0.55
LMBWR	10.24 ± 1.02	1.71 ± 1.25	-0.52 ± 0.20

TABLE 6 | Mean CO_2 flux Welch t -test ($n = 200$) between select two measurements in 2019.

Groups	Difference ($\mu\text{mol m}^{-2} \text{hr}^{-1}$)	95% lower CI	95% upper CI	t	df	p -value
2019						
TrIR-	0.019	-0.32	0.36	0.11	31.08	0.91
TrBWR						
TrIR-LMIR	-2.17	-2.66	-1.68	-9.09	30.00	$4.00 \times 10^{-10}^{***}$
TrIR-LMBWR	-6.06	-7.01	-5.11	-13.01	29.80	$7.98 \times 10^{-14}^{***}$
TrBWR-LMIR	-2.19	-2.76	-1.62	-7.77	47.05	$5.55 \times 10^{-10}^{***}$
TrBWR-LMBWR	-6.08	-7.07	-5.09	-12.43	35.69	$1.59 \times 10^{-14}^{***}$
LMIR-LMBWR	-3.89	-7.51	-4.93	-2.84	42.60	$2.51 \times 10^{-9}^{***}$

*** $p < 0.001$.

We observed lower average N_2O fluxes throughout our measurement period in LM in 2018 compared to the other three treatments (CC, CR, and Tr), which is different from the previous findings. Peters et al. (2020) observed the highest N_2O fluxes occurred in the LM system in 2016 and 2017. We did observe an N_2O fluxes peak in LM on June 20th, which was likely due to the elevated soil N added as the white clover decomposed. From observing the relationship between the soil N amount and white clover biomass in the soil, the soil N content increased as white clover biomass decreased between the end of May to mid-July in 2018 (Figure 4). At the same time, from June 5th to July 10th is the time when the rate of soil N amount increase is one of the highest under the LM system in 2018 (Figure 4). In 2019, from June 11th to June 24th, the rate of soil N amount increase was

one of the highest when we observed the maximum N_2O flux generated in LM (Figure 4). Therefore, N_2O production in LM might also be associated with the rate of soil N content increase rather than only with the absolute soil N content as we found in 2018.

Greater soil N was observed in Tr than LM from April to the end of May when we conducted measurements (Figure 4). Soil N in LM increased from the end of May to mid-July due to the decomposition of white clover, which corresponded perfectly to the higher N_2O flux observed in late June to early July in LM in 2018 (Figure 2). This is similar to what Peters et al. (2020) observed in 2016 and 2017. However, in 2019, since we focused our measurements in the early growing season during fertilization periods, higher N_2O fluxes induced by the added N from clover were not observed in LM compared to Tr. As discussed earlier in 2018, even though we conducted measurements throughout the whole growing season, we still did not observe the highest N_2O fluxes generated in LM. Despite the lack of N inputs in LM in 2018, the soil N amount in LM was similar to that in Tr from late June to July (Figure 4) and N_2O flux levels in both plots did not differ.

Among eight chambers of TrIR in 2019, chambers that received N inputs twice (TrIR1) tended to emit higher CO_2 flux than their paired chambers, which only received fertilizer once (TrIR2, Supplementary Figure 3; Supplementary Table 9). For N_2O , the chamber with the greatest amount of N inputs emitted the highest N_2O fluxes in the second fertilization period (Supplementary Table 9). Among four chambers in either TrIR1 or TrIR2, chambers that received higher amount of N fertilizer inputs emitted higher N_2O fluxes (Supplementary Table 9). More results describing GHG fluxes among eight chambers in TrIR could be found in the Supplementary Material. Higher N_2O flux in LM from mid-June to July compared to CC and CR in 2018 could also be related to greater mineralizable C in the soil. Mineralizable C is used as a substrate to stimulate N_2O fluxes generated from the denitrification process (Mitchell et al., 2013). From observing the soil C amount measured under LM in 2018, from June 14th to July 10th, the rate of soil C amount increase is one of the highest when we observed the maximum N_2O flux on June 20th (Figure 4). Comparing CO_2 and N_2O fluxes between leguminous (CC) and non-leguminous cover crops (CR), no significant difference was observed between those two from observations in 2018. However, studies suggest that cover crops with N fixing ability tend to increase the amount of soil inorganic C, whereas cover crops with no N fixing ability leads to N immobilization (Schmatz et al., 2020).

CH_4 Fluxes

Soils in cover crop systems were observed as larger CH_4 sinks compared to Tr in 2018. This could result from fewer N inputs under cover crop systems compared to Tr. The addition of N fertilizer could increase the NH_4^+ amount, hence reducing the CH_4 oxidation capacity of the soil, as found in previous study (Boeckx et al., 1996). The similarity in physical properties of NH_4^+ and CH_4 molecules allow them to compete for binding sites in the Methane Monooxygenase (MMO) enzyme. NH_4^+ inhibition effects refer to the fact that NH_4^+ , which serves as a more

TABLE 7 | Mean CO₂ flux Welch *t*-test between certain two measurements among TrIR, TrBWR, LMIR, and LMBWR in in first (*n* = 107) and second (*n* = 93) fertilization periods in 2019.

Groups	Difference ($\mu\text{mol m}^{-2} \text{hr}^{-1}$)	95%lower CI	95% upper CI	<i>t</i>	df	<i>p</i> -value
2019 First fertilization period (April 27–May 4)						
TrIR-TrBWR	0.53	0.09	0.96	2.56	15.26	$2.14 \times 10^{-2*}$
TrIR-LMIR	-2.61	-3.41	-1.81	-6.89	16.44	$3.15 \times 10^{-6***}$
TrIR-LMBWR	-4.48	-5.68	-3.29	-7.96	15.63	$6.97 \times 10^{-7***}$
TrBWR-LMIR	-3.13	-4.00	-2.28	-7.55	21.80	$1.63 \times 10^{-7***}$
TrBWR-LMBWR	-5.01	-6.24	-3.77	-8.51	18.31	$8.84 \times 10^{-8***}$
LMBWR-LMIR	1.87	0.49	3.25	2.79	26.09	$9.62 \times 10^{-3**}$
2019 Second fertilization period (May 23–June 1)						
TrIR-TrBWR	-0.38	-0.69	-0.069	-2.55	20.16	$1.89 \times 10^{-2*}$
TrIR-LMIR	-1.60	-1.99	-1.20	-8.62	14.47	$4.46 \times 10^{-7***}$
TrIR-LMBWR	-7.85	-8.46	-7.26	-27.88	14.59	$4.57 \times 10^{-14***}$
TrBWR-LMIR	-1.22	-1.68	-0.77	-5.60	21.44	$1.34 \times 10^{-5***}$
TrBWR-LMBWR	-7.48	-8.12	-6.84	-24.58	18.79	$9.56 \times 10^{-16***}$
LMBWR-LMIR	6.26	5.58	6.93	19.30	21.44	$4.86 \times 10^{-15***}$

p* < 0.05, *p* < 0.01, ****p* < 0.001.

aggressive competing substrate, can inhibit the CH₄ oxidation process (Gulledge et al., 1997). Among three cover crop systems, slightly lower CH₄ consumption in CR potentially resulted from its higher C/N ratio due to its inability to fix N. CR most likely had a higher soil C amount, which may have stimulated CH₄ fluxes under anaerobic conditions (Le Mer and Roger, 2001). Future studies should include measurements of soil NH₄⁺ and soil C/N ratio to further investigate their impacts on CH₄ uptake ability.

Intensive land management activities could also influence the CH₄ uptake ability. In 2018, decreasing CH₄ uptake occurred from April 22nd to early May, suggesting that soil disturbance event, such as planting maize and increased N inputs from fertilizer during this time affected the role of soil as absorbing CH₄ (Conrad and Rothfuss, 1991; Bronson and Mosier, 1994; Dunfield and Knowles, 1995). In 2019, a higher CH₄ sink from BWR than IR was observed under both LM and Tr systems. This indicates that higher soil N in the IR may have inhibited CH₄ uptake and resulted in a smaller CH₄ sink. However, instead of being the largest CH₄ sink, TRIR2-0, the chamber which did not receive any N inputs, served as the smallest CH₄ sink among the eight chambers in TrIR. Furthermore, chambers that received higher levels of N fertilizers tended to be larger CH₄ sinks than those that received lower N. Additional experiments are needed to better understand what fertilization pattern enhances soil CH₄ sink potential so that mitigation strategy can be considered in agriculture.

One limitation of our study is the lack of soil microbial data. Including soil microbial measurements would have added insights in understanding GHG fluxes variations that occurred between cover crop systems and the conventional agricultural system. In order to better understand the live clover impacts on GHG emissions, both long-term soil GHG observation and microbial level laboratory studies are needed. From the results

of this study, we observed increased soil CO₂ fluxes also from LMIR, which comes from the accelerated decomposition of native soil due to clover incorporation. Studies have shown that incorporating fresh organic matter such as clover residue could stimulate mineralization of soil organic matter, which is named as priming effect (Bingeman, 1952; Fontaine et al., 2003). Priming effect depends on both the biomass and microbial amounts in the soil (Camarotto et al., 2018). For example, incorporating biomass with high C/N ratio into the soil would accelerate the decomposition by microorganisms to acquire more N in the soil (Camarotto et al., 2018). Meanwhile, studies conducted by Chen et al. (2015) suggest that addition of residue could formulate a more diverse microbial community and lead to stronger priming effects, which potentially results in increased CO₂ fluxes observed under cover crops, as found in this study. Future studies should also investigate the microbial mechanism of priming effects induced by cover crop residues. Moreover, soil parameters were only collected as total C and total N amounts in our study. Future studies should also include the measurements of both particulate and mineral soil organic matter types under different cover crop systems. Different types of soil organic matter measurements could better evaluate the type of organic matter added to the soil by cover crops, their susceptibility to further decomposition and CO₂ and N₂O emissions. A more comprehensive N budget should also be included in future studies to better evaluate the mitigation potential comparison between legume and non-legume cover crops. Reducing GHG emissions from agriculture is complex since the success of realizing mitigation potential through cover crop establishment is site-specific. Spatial heterogeneity acts as barriers to wider adoption of cover crops regarding different local conditions. There is no one-size-fits all solution, and future studies still need to fully evaluate both locally and globally about barriers and opportunities for successful cover crop implementation.

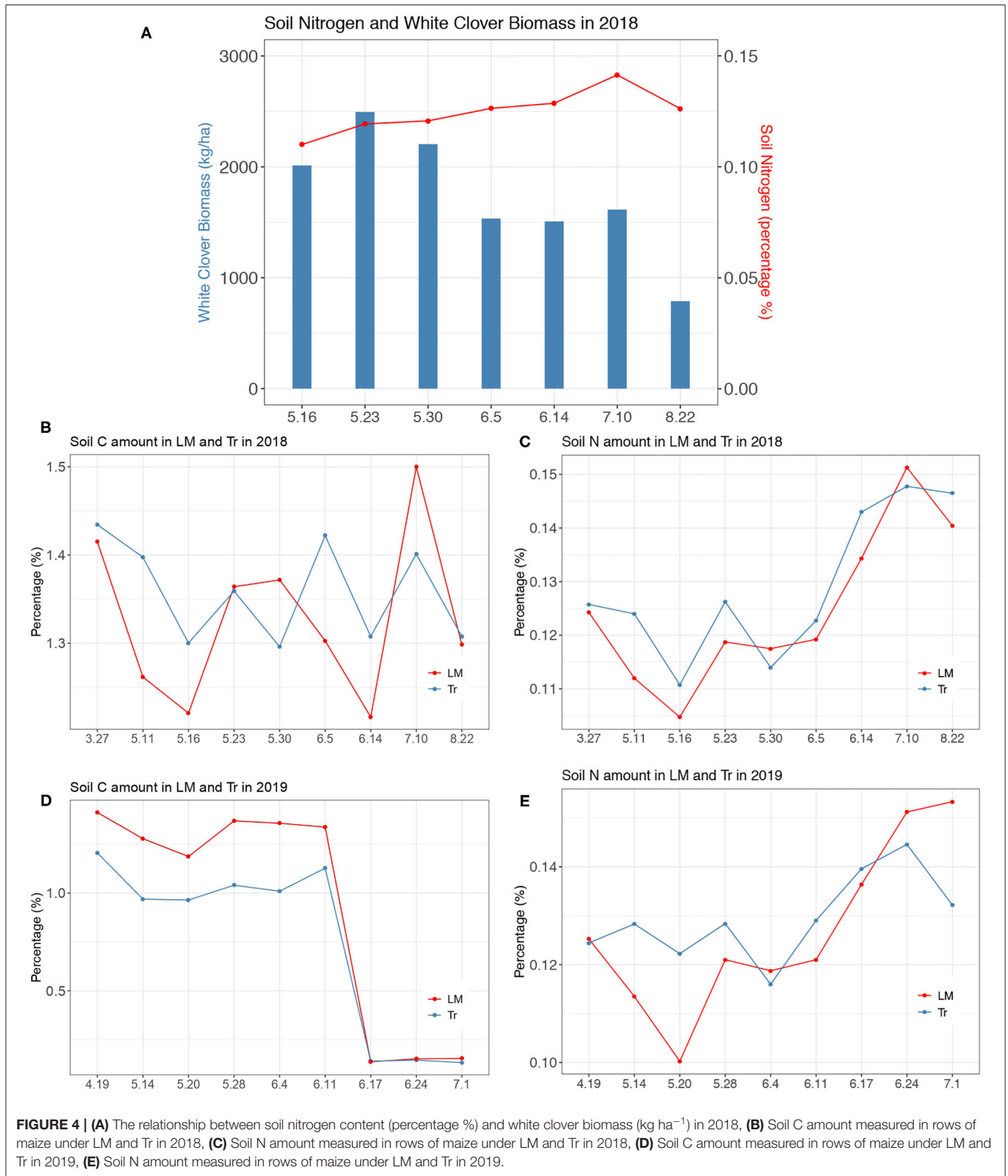


FIGURE 4 | (A) The relationship between soil nitrogen content (percentage %) and white clover biomass (kg ha⁻¹) in 2018, **(B)** Soil C amount measured in rows of maize under LM and Tr in 2018, **(C)** Soil N amount measured in rows of maize under LM and Tr in 2018, **(D)** Soil C amount measured in rows of maize under LM and Tr in 2019, **(E)** Soil N amount measured in rows of maize under LM and Tr in 2019.

Practicality of Cover Crops

Crop yields and associated economic returns typically have a larger impact on the farmers’ decision to implement cover crops than the potential reduction of GHG emissions. Even though

maize yields under LM system were lower than CC and CR in 2018, the evaluation of economic impact of cover crops at a farm level should also include other factors as well. Gabriel et al. (2013) found that incorporating cover crops brought extra

expense in establishment, seed, and killing, although there were potential savings from reduced fertilizer use. Selling the biomass of the cover crops as animal feed could bring additional savings as well. For example, from the case study conducted by Gabriel et al. (2013) in Spain, the extra cost of incorporating vetch as a cover crop was 71.65 € ha⁻¹ but if the biomass of vetch was collected and sold as forage, the extra cost of the cover crops was reduced to 31.63 € ha⁻¹. The spatial heterogeneity also makes the economic analysis of cover crop a challenging endeavor as it also depends on product and commodity prices. A comprehensive economic analysis is essential in order to provide a better reference regarding cover crop practicality with farmers.

Even though we witnessed reduced maize yields under the LM system in 2018, conflicting studies exist regarding the impacts of cover crop on grain yields. A review by Abdalla et al. (2019) reported that grain yields were reduced by 4% under cover crop systems as compared to the control treatment. However, a case study conducted in Spain showed that the yields of maize were two times higher after adopting cover crops than in fallow (Kaye and Quemada, 2017). Grain yields were also found to vary between legume and non-legume cover crops. For example, similar yields of cash crop were found under non-legume cover crop treatment compared to the control treatment. In contrast, Finney et al. (2016) found that adopting legume cover crop could increase the yields by 5–30%. Long-term studies are needed to better explore the impacts of cover crops on yields. It is also possible that the positive effects of the restored soil health on crop yields are more pronounced as the time of mulch incorporation increases.

CONCLUSION

This study finds that in addition to soil and water quality improvement, increased soil C storage, and reduced soil erosion and nitrate leaching, the implementation of cover crops can add another benefit for mitigating climate change, stemming from the larger CH₄ sink potential of the soils under this agricultural practice. When we subtracted a rough estimate of white clover respiration from the LM system, the net CE turned out to be the lowest of all practices. In order to assess climate change mitigation potential from agriculture, it is important to analyze

the three GHG fluxes holistically and understand the impacts of different practices in more detail. The economic returns also need to be quantified, as reduced fertilizer and herbicide costs could be substantial when implementing the LM system at a farm scale. Long-term studies are essential to better assess the practicality of cover crops in terms of crop yields, economic impacts, and climate mitigation potential.

DATA AVAILABILITY STATEMENT

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

AUTHOR CONTRIBUTIONS

ES and AA conceptualized the study and conducted field experiment. ES and NH obtained funding. YW analyzed the data and wrote the draft with ES. Everyone reviewed the manuscript and helped in revision. All authors have given approval to the final version of the manuscript.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fclim.2021.742320/full#supplementary-material>

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