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Neutral winds from mesosphere to thermosphere—past, present, and future outlook

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The Earth's upper atmosphere (85–550 km) is the nearest region of geospace and is highly dynamic in nature. Neutral winds impact a large portion of the dynamics in this region. They play a critical role in determining the state of the ionosphere-thermosphere system at almost all latitudes and altitudes. Their influences range from wave breaking/dissipation in the mesosphere and lower thermosphere to global redistribution of energy and momentum deposited at high latitudes by the magnetosphere. Despite their known importance, global geospace neutral winds have remained one of the least sampled state parameters of the Earth's upper atmosphere and are still poorly characterized even after multiple decades of observations. This paper presents an overview of historical neutral wind measurements and the critical need for their global height-resolved measurements. Some satellite missions are still operational and deliver valuable information on the contribution of neutral winds in global atmospheric dynamics. However, many significant gaps remain in their global monitoring, and our current understanding of the drivers of neutral winds is incomplete. We discuss the challenges posed by these measurement gaps in understanding geospace physics and weather. Further, we propose some wind observation solutions, including the simultaneous operations of upcoming NASA DYNAMIC and GDC missions as well as support for the development of ground-based observing methodologies, that will lead to fundamental advances in geospace science and address humanity's emerging space needs.

KEYWORDS

neutral wind, upper atmosphere dynamics, thermosphere, ionosphere, ignorosphere, geospace, near earth space, wind dynamics

1 Introduction and motivation

Earth's upper atmosphere (here defined as 85–550 km) is the nearest region of geospace (i.e., the near-earth space environment). It surrounds and protects the planet. It is a weakly ionized plasma whose ionized and neutral components are typically

considered separately as the ionosphere and thermosphere, respectively. The region is highly dynamic in nature and is driven by solar and magnetospheric forcing from above (often impulsively) (e.g., Lu et al., 2016), and continuously by meteorological forcing from the lower atmosphere (e.g., Liu, 2016; Sassi et al., 2019). The weather of this region directly impacts the performance of ground-based and space-based communication and navigation technologies. In addition, a wide variety of low-Earth orbit weather, communications, global navigation, and other operational satellites on which humanity relies on a daily basis, reside in this region of space. Any changes in this atmospheric region can directly impact the orbital trajectories of these satellites (and of the orbital debris that also permeates this region) via atmospheric drag. Given our increasing dependence on space-based technologies, scientific understanding and prediction of upper atmospheric weather and dynamics is of critical importance.

It is well known that neutral winds impact a large portion of the dynamics within the upper atmosphere and act as a coupling agent between the neutrals and plasma in the ionospherethermosphere (I-T) system (e.g., Wang et al., 2021). They are the primary regulators and redistributors of the mass, momentum, and energy within the I-T system and hence play a major role in determining the state of the geospace weather at all latitudes. Neutral winds in the upper atmosphere are driven primarily by a time-varying interplay among heating-induced pressure gradients (caused by absorption of solar ultraviolet irradiance, Joule heating, particle precipitation, and other frictional heating sources), momentum transfer between ion and neutrals, inertial forces (Coriolis and centrifugal), wave/tidal forcing from below, and internal small-scale instabilities. The neutral wind dynamics have a wide spectrum of influences on geospace weather from equatorial to polar latitudes and at all altitudes in mesosphere and thermosphere. Some examples from their wide spectrum of influences are:

- At low latitudes of the E-region, they generate wind dynamo electric fields and push the ionospheric plasma upward and downward along the magnetic field lines, changing the chemical equilibrium of I-T species (e.g., Rishbeth, 1972; Kelly, 1989; Immel et al., 2021).
- In the mesosphere and lower thermosphere (MLT) region, winds impact the energy and momentum deposited by upward propagating waves and tides (e.g., Garcia and Solomon, 1985; Smith, 1996; Forbes, 2007; England, 2012; Yiğit et al., 2016; Sassi et al., 2018).
- At high latitudes, winds feed back into the ionospheric convection and modify Joule heating and other energy deposition processes (e.g., Killeen, 1987; Richmond et al., 2003; Wang et al., 2021).

- Wind circulation cells in the upper atmosphere (forced by waves and tides as well as thermally driven, shallow layered as well as covering wide altitudinal ranges) influence the seasonal variation of chemical species (e.g., Liu, 2007; Qian et al., 2017; Wang et al., 2022).
- Winds influence the evolution of the ionospheric response during a storm and during subsequent relaxation to quiet conditions (e.g., Crowley et al., 1989; Fuller-Rowell et al., 1996; Strickland et al., 1999). The mass, momentum, and energy deposited from the magnetosphere at high latitudes can be transported over the globe by wind-driven transport, which can lead to significant global thermospheric disturbances (e.g., Lu et al., 2016).
- Along with neutral density, neutral winds contribute directly to atmospheric drag and thus play an important role in precision orbit determination and prediction (e.g., Gaposchkin and Coster, 1988; Gaposchkin, 2003). Their accurate determination will also help in accurately predicting the evolution of debris fields.

One of the exciting new paradigms in upper atmospheric science is the recognition that the meteorology of the lower atmosphere continuously imprints on upwardly propagating tidal components and modulates properties of the thermosphere and the ionosphere (e.g., Lieberman et al., 1994; Oberheide and Forbes, 2008; Forbes et al., 2009; Hagan et al., 2009; Häusler and Luhr, 2009; Oberheide et al., 2015). Tides are also perceived as important vertical coupling agents between the MLT region and the ionosphere (e.g., England et al., 2006; Jin et al., 2012; Pedatella et al., 2012; McDonald et al., 2018). Such driven variability in the entire atmosphere directly impacts winds and can be characterized *via* analysis of wind observations.

The ionosphere is only $\sim 1\%$ as massive as the thermosphere, but it impacts all radiowave propagation through and within the upper atmosphere. Dynamical changes in neutral winds directly or indirectly perturb the ionosphere and hence can disrupt radiowave communications. Despite their known importance, global geospace neutral winds have remained one of the most under-sampled state parameters of the Earth's upper atmosphere and are still poorly characterized even after multiple decades of investigations. Large spatiotemporal gaps exist in historical and contemporary wind observations, especially in the 100-200 km altitude region that has been called the new 'ignorosphere' (Emmert et al., 2021), due to the dearth of data collected there. This paper discusses the past and present state of neutral wind observations, current challenges, and the outlook for the future; it is based on Dhadly et al. (2022) - a white paper submitted to the Decadal Survey on Solar and Space Physics (Heliophysics) 2024-2033.



2 Past and present state of neutral wind observations

Several space-based ITM (Ionosphere-Thermosphere-Mesosphere) missions launched in the past (see Figure 1) have improved our understanding of two-dimensional global horizontal neutral wind flows. The AE-E Neutral Atmosphere Temperature (NATE) instrument (Spencer et al., 1973) measured upper thermospheric in-situ cross track winds at lower latitudes; Wind and Temperature Spectrometer (WATS) and Fabry-Perot Interferometer (FPI) on Dynamic Explorer 2 (DE2 1981-1983) (Spencer et al., 1981) focused on the high latitude I-T interactions using *in-situ* and remote sensing of the region; WIND Imaging Interferometer (WINDII) and High Resolution Doppler Imager (HRDI) on the Upper Atmosphere Research Satellite (UARS 1991-1997) (Hays et al., 1993; Shepherd et al., 1993; 2012) obtained neutral wind profiles using remote sensing equatorward of polar latitudes from the MLT to upper thermospheric altitudes; TIMED Doppler Interferometer (TIDI) on the Thermosphere Ionosphere Mesosphere Energetics and Dynamics (TIMED 2001-present) (Skinner et al., 2003) remotely sensed MLT winds equatorward of auroral latitudes; CHAllenging Minisatellite Payload (CHAMP 2000-2010) (Reigber et al., 2002) and Gravity Field and Steady-State Ocean Circulation Explorer (GOCE 2009-2013) (Doornbos et al., 2013) measured global in-situ cross-track winds in the upper thermosphere; and recently the Michelson Interferometer for Global High-resolution Thermospheric Imaging (MIGHTI 2019-present) (Englert et al., 2017) on the Ionospheric Connection Explorer (ICON) measured neutral winds from the MLT to the upper thermosphere equatorward of middle latitudes. TIDI/TIMED and MIGHTI/ICON are still operational, providing critical pieces of information on the contribution of neutral winds in global atmospheric dynamics from MLT to the upper thermosphere. However, none of these single satellite missions measured global winds synoptically, all have significant altitudinal, latitudinal and local time coverage gaps, and all only resolved one or two components of the 3D vector wind.



In addition to the space-based observations which provide complete longitudinal sampling of neutral wind dynamics, ground-based Fabry-Perot Interferometer (FPIs) (Hernandez and Roble, 1976a; Hernandez and Roble, 1976b; Smith et al., 1988; Smith et al., 1989; Makela et al., 2012; Meriwether, 2006) and Scanning Doppler Imaging FPIs (SDIs) (Conde and Smith, 1995; Conde et al., 2018) have contributed extensively to our current understanding of geographically local neutral wind dynamics. They have been operational at multiple sites around the globe since the 1970s. Balloon borne FPI High altitude Interferometer WIND experiment (HIWIND) provided both day and night thermospheric winds in 2011 and 2018 (Wu et al., 2012; 2019). HIWIND daytime capability is significant for understanding thermosphere and ionosphere interaction during the daylight times as most ground-based FPI and SDI techniques are limited to observing during the nighttime. Further, sporadic sounding rocket chemical release wind measurements carried out at different sites around the globe, incoherent scatter radars (ISRs), meteor radars, and lidars have contributed significantly to lower thermospheric wind science.

For more detailed summaries of historical and current neutral wind observations, see, e.g., Dhadly et al. (2017); Dhadly et al. (2018); Dhadly et al. (2019); Drob et al. (2008); Drob et al. (2015); Emmert et al. (2008). Despite the historical space-based and ground-based observational efforts, there still exist wide gaps in the archive of wind measurements, especially in the ignorosphere (see **Figure 2**) and in the longitudinal coverage at high latitudes, such that the global evolution of neutral wind patterns in response to the solar, magnetospheric, and terrestrial drivers is still poorly understood even after combining four decades of wind measurements.

3 Current challenges

Given the important role they play in determining the state of geospace weather, thermospheric winds have remained a subject of interest over the past several decades, but the Heliophysics community still does not have extensive, contemporaneous, and complete databases to close on many fundamental questions. Prior to the UARS and TIMED missions, wind measurements in the thermosphere were obtained primarily from sparsely located ground-based FPIs/SDIs. With exceedingly few exceptions (e.g., Gerrard and Meriwether, 2011), all FPI and SDI ground-based neutral wind measurements are nighttime only because these techniques are overwhelmed during the daytime by Rayleigh-scattered light from the lower atmosphere. In addition, they provide wind measurements at fixed altitudes in E- (~100-130 km) and F-regions (~240 km). Thus, the overwhelming majority of daytime MLT and upper thermospheric neutral wind measurements are from spacebased instruments, meteor radars, and incoherent scatter radars (ISRs). There have also been rocket-based experiments (e.g., Edwards et al., 1963) that have provided temporally limited measurements of the daytime atmosphere, but by their nature they are unable to provide the continuous global-scale coverage necessary to advance understanding of global dynamics. The ground-based instruments provide limited spatial coverage but at high temporal cadence and some with good local time coverage. These local measurements are significant for investigation of ion/neutral interactions when combined with other observations. Space-based instruments, of course, provide wide geographical coverage, but with limited local time coverage.

Satellites have provided measurements of thermospheric winds at a broad range of thermospheric altitudes. However, the remote sensing of nighttime ignorosphere is very limited (Figure 2). In addition, the ignorosphere is too low to be routinely probed in-situ and too high to be probed with radars, lidars, and balloons. Thus, it has remained extremely under-sampled and poorly understood compared to the other regions. As an example, even diurnal variations of neutral winds in this region are largely undocumented. Unfortunately, the ignorosphere strongly influences the upper thermosphere (>200 km altitude); it is the region where the majority of the large amplitude upward propagating tides and waves excited in the lower atmosphere deposit their energy and momentum (due to tidal dissipation and wave breaking) and hence disrupt the neutral dynamics primarily shaped by solar heating, Coriolis force, and ion drag. In addition, most of the solar and magnetospheric impacts (such as large scale traveling atmospheric and ionospheric disturbances from the auroral zone) propagating from the topside are absorbed in this region.

Tides and waves propagating upward from the lower atmosphere (modulated by the meteorology of the lower atmospheric weather) add significant short-term variability in winds and the I-T system which is currently unexplained. For example, major components of the tidal spectrum (migrating tides such as DW1 and SW2 and non-migrating tides such as DE3) originate from the lower atmosphere. In order to estimate their short-term variability in winds, ground-based wind sensors equally spread in longitudes are required, which is hard to achieve due to limited landmass. Single spacecraft missions launched in the past provide only limited local time coverage. Thus, neither the space-based nor ground-based sampling required to understand short-term variability propagating up from the lower atmosphere is currently available. Without sampling multiple local times, especially in the ignorosphere, it is extremely difficult to characterize the vertical coupling in the atmosphere and parse out the exact role of terrestrial weather on the upper atmosphere.

Apart from the ignorosphere, even comprehensive spacetime sampling in the regions that can be easily monitored was never obtained in the past. This is because of the multidimensional dependence of neutral wind circulation. A survey showing the geographical sparseness of the winds due to their multi-dimensionality even after fusing 4 decades of data is presented in Figure 3. Some major data gaps that currently exist in neutral wind observations impart large uncertainties in wind characterizations and are known to introduce numerical artifacts in fitting techniques. Some geographical regions are more densely sampled by FPIs and SDIs than others, which can skew statistical analyses or statistical fitting algorithms (e.g., Dhadly et al., 2019) when included with sparsely sampled regions. In addition, optical and insitu wind sensors operate independently of each other, possess different geometries, different spatial and solar coverage, and may contain large mutual biases (e.g., Dhadly et al., 2019; Dhadly et al., 2021). Fusion of these diverse data in a selfconsistent manner to characterize winds is a formidable challenge. Although various investigations have pieced together the patchy historical and current neutral wind observations to build a comprehensive picture of neutral wind dynamics (e.g., Drob et al., 2008; Drob et al., 2015; Emmert et al., 2008; Dhadly et al., 2017; Dhadly et al., 2018; Dhadly et al., 2019), it is still not enough to understand the large-scale evolution of neutral winds in shape and magnitude in response to all the solar and magnetospheric drivers, continuous terrestrial forcing, seasons, and interhemispheric magnetic asymmetry.

It is worth mentioning that DE-2 was the first ITM mission to measure drivers and impacts of neutral winds simultaneously and provided a tantalizing glimpse of the strongly coupled high latitude I-T system by simultaneously measuring important ionospheric and thermospheric state parameters. Unfortunately, none of the ITM missions after DE-2 measured the drivers and impacts simultaneously especially at high latitudes. The most extreme space weather occurs at these latitudes, and the Heliophysics community does not yet have a complete picture of neutral winds and their drivers.

The basic morphology of horizontal neutral winds and its dominant drivers are known to some extent, thanks to past ITM missions; however, the dynamics of the vertical winds is largely unknown. None of the past ITM missions after DE-2 measured vertical winds. Our current observational understanding of



vertical neutral wind dynamics is either based on nighttime localscale studies by FPIs/SDIs or DE-2. Previous missions such as CHAMP and GOCE measured anomalies in neutral densities in the upper atmosphere. The physical understanding of the thermosphere suggests that they likely caused by neutral winddriven vertical transport. However, due to a nearly complete lack of vertical wind measurements, their role has yet to be delineated.

Recent advances in physics-based models (e.g., TIEGCM, TIMEGCM, GITM, WACCM-X, CITP-e, etc) and empirical models (e.g., HWM14, HL-TWiM, etc - see reference list) have pushed the boundaries of our current understanding and have shown promising results in capturing the intricacy of the full I-T system. However, due to the lack of appropriate observations, their advances are still to be tested, especially at high latitudes where statistical wind behaviors diverge significantly from physics-based simulations (e.g., Wu et al., 2012; Wu et al., 2015; Liuzzo et al., 2015). If realistic forecasting of space weather is ever to be achieved, the I-T coupling mechanisms need to be understood and validated using observations.

4 Future outlook and recommendations

Given our ever-increasing reliance on space-based technologies, and recognizing the imapcts of neutral winds on I-T dynamics, accurate knowledge of the global neutral wind system, its drivers, and continuous global monitoring is crucial for advancing space weather research and forecasting capabilities needed for space assets. As discussed earlier, the biggest obstacles are current data gaps and observational challenges. The groundbased platforms provide limited geographical coverage (local view) but for extended periods of time and measure the diurnal cycle without any temporal averaging. In contrast, space-based platforms provide wide geographical coverage (synoptic view), but their local time coverage is limited. To overcome the disadvantage of each platform, an effective solution is to use a combination of space-based platforms with the ground-based networks. Thus, investments are needed to develop space-based as well as longitudinally and latitudinally distributed arrays of ground-based sensors. The Super Dual Auroral Radar Network (SuperDARN), which synoptically measures ion velocity fields at high latitudes, has demonstrated the value of ground-based networks to both basic I-T science and space weather research (e.g., Chisham et al., 2007; Nishitani et al., 2019). It's success derives in large part from international participation by the SuperDARN community.

As discussed earlier, the ignorosphere is the most undersampled region of the thermosphere. Newer technologies such as THz limb sounder (Wu et al., 2016; Yee et al., 2021) and Doppler Wind and Temperature Sounder (DWTS) (McHugh et al., 2014) have the potential to measure winds along with temperature and density at all local times in the ignorosphere when flown on a space-based mission. For details, please refer to the papers cited above. In addition, optical remote sensing technologies for this region have a strong heritage, at least for daytime conditions. Miniaturized versions of some spacebased remote sensing sensors are in development or have been developed and undergoing demonstrations, such as INDI (Interferometer for Neutral-thermosphere Dynamics Imaging (Harlander et al., 2020) and mini-MIGHTI (Harlander et al., 2019; Harlander and Englert, 2020). NASA has recently invested in the ICON/MIGHTI, which is currently monitoring the state of lower latitude neutral winds from MLT to the upper thermosphere from low Earth orbit. However, it is a singlespacecraft mission monitoring only lower latitudes and does not allow us to fully investigate the short-term variability (for example shorter than a month for ICON and 60 days for TIMED) entering the I-T system from the lower atmosphere. A dedicated fleet of multiple identical instruments in polar orbit spread longitudinally would provide an unprecedented sampling of global wind circulation covering a wide range of latitudes/longitudes that is necessary for fundamental advances in our understanding of neutral wind dynamics and its drivers.

The 2013 Decadal Survey on Solar and Space Physics recommended a dedicated platform to address the atmospheric vertical coupling from the MLT to the upper thermosphere, and to study I-T dynamics as a global system using a constellation of satellites in polar orbit spread longitudinally. To fulfill these recommendations, the Decadal Survey endorsed Dynamical Neutral Atmosphere-Ionosphere Coupling (DYNAMIC—a two satellite mission https://soma.larc.nasa.gov/stp/dynamic/ index.html) and Geospace Dynamic Constellation (GDC—a six to eight satellite constellation mission—https://lws.larc.nasa.gov/ gdc/). The GDC mission, set to launch in 2027 (as per



NASA Program Element Appendix on https://lws.larc.nasa.gov/ gdc/), will provide a quasi-global coverage of in-situ upper thermospheric winds (~400 km), initially on local scales and finally on large scales when all the GDC spacecrafts are equally spread in local time and longitudes. DYNAMIC will measure neutral winds at least four local solar times per day (6 h apart if spaced equally in longitude) from the MLT to the upper thermospheric altitudes. DYNAMIC is also being seriously considered by NASA. Flying DYNAMIC and GDC together would amplify the scientific return of each mission - DYNAMIC focusing on the I-T system below 250 km and GDC in-situ above 300 km altitude. These constellations will enable science that goes well beyond what was possible with previous ITM missions. Their combined observations would allow the Heliophysics community to resolve top-down and bottom-up forcing mechanisms which is a difficult observational challenge and is impossible to address without simultaneous multipoint observations by constellation-type missions. In addition to filling data gaps, the data obtained from them would allow the characterization of the multidimensional nature of global neutral winds, which in turn would provide a convenient and reliable reference for the validation and tuning of general circulation models. Figure 4 illustrates the power of having constellation type missions. It shows a survey performed using NASA provided GDC Design Reference Mission (DRM) ephemeris. Assuming the solar cycle 25 after the year 2027 will follow a similar decline as the solar cycle 23 after the year 2004, similar geomagnetic conditions as in 2004-2006 are expected to occur in 2027-2029 (GDC prime mission). For the survey, the geomagnetic conditions of 2004-2006 are utilized as a proxy for 2027-2029. Thus, over the course of the GDC prime mission, it alone will provide much better high latitude coverage of winds than obtained by combining all the available past wind data from all the instruments (as shown in **Figure 4**).

Further, FPIs/SDIs are well known for their extensive contributions to our current understanding of neutral wind dynamics, but they provide only nighttime wind measurements (with some exceptions-e.g., Gerrard and Meriwether (2011)). Past experiments have suggested that daytime FPI/SDI observations are feasible; however, they require further development and testing. Successful implementation of groundbased wind instruments that can measure winds at all local times and their employment in longitudinally and latitudinally distributed arrays with international cooperation can provide continuous monitoring of E- and F-region neutral winds. Nevertheless, missions like DYNAMIC are needed to measure vertical profile of neutral winds to understand the lower and upper atmospheric vertical coupling. Thus, a combined approach of using ground-based array and space-based sensor wind measurements would be a huge leap forward for the Heliophysics community.

Author contributions

Primary author—MD All other authors contributed by either editing or helping with data analysis.

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