



Reef Habitat Type and Spatial Extent as Interacting Controls on Platform-Scale Carbonate Budgets

Chris T. Perry*, Kyle M. Morgan and Robert T. Yarlett

Geography, College of Life and Environmental Sciences, University of Exeter, Exeter, United Kingdom

A coral reefs carbonate budget strongly influences reef structural complexity and net reef growth potential, and thus is increasingly recognized as a key “health” metric. Despite this, understanding of habitat specific budget states, how these scale across reef platforms, and our ability to quantify both framework and sediment production values remains limited. Here, we use *in-situ* census data from an atoll rim reef platform in the central Maldives to quantify rates of both reef framework and sediment production and loss within different platform habitats, and then combine these data with high-resolution habitat maps to quantify contributions to platform wide carbonate budgets. The net reef framework budget for the entire platform is extremely low (0.12 G, where G = Kg CaCO₃ m⁻² year⁻¹), with a very high proportion (143,745 kg or 65.1%) of total framework production generated within the platform margin reef zones, despite these comprising only ~8% of platform area. Net platform-scale sediment budgets are higher (1.04 G), but most is produced in the reef and platform margin hardground habitats, of which ~80% derives from parrotfish bioerosion. Significant quantities of new sediment (up to ~1 G derived from the calcareous green algae *Halimeda*) are produced only in one habitat. All lagoonal habitats have negative or neutral net carbonate budgets. These data demonstrate the marked inter-habitat differences in reef carbonate budgets that occur across reef platforms, and the major dampening effect on overall platform scale budgets when rates are factored for habitat type and size. Furthermore, the data highlights the disproportionately important role that relatively small areas of reef habitat can have on the maintenance of net positive platform scale budgets. Because of the intrinsic link between carbonate production rates and reef-associated landform development and maintenance, these findings also have implications for understanding reef-associated landform stability. In this context the reef island at this site has been highly mobile over the last ~40 years, and we hypothesize that such instability may be being exacerbated by the measured low overall rates of framework and sediment generation.

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*Correspondence:

Chris T. Perry
c.perry@exeter.ac.uk

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INTRODUCTION

Coral cover loss, reef structural complexity declines, transitions in coral community composition, and declines in reef-associated species abundance, have been fundamental aspects of the on-going degradation of coral reefs globally (Hoey et al., 2016). These changes have had, and are continuing to have, profound impacts on many of the ecosystem goods and services that reefs provide to

society (Pendleton et al., 2016). Of equally pressing concern however are the impacts that these ecological declines are now having on the processes of carbonate production and bioerosion on reefs (Perry et al., 2014a,b). These changes are especially important because they control the ability of reefs to both maintain, and add to, their framework structures, and which in turn influences their capacity to track rising sea levels (Perry et al., 2013b, 2015b; Saunders et al., 2016). Understanding the magnitude of reef carbonate budget changes, and how they have changed within and between reef-building regions has thus been a focus of considerable recent research interest. This has encompassed assessments of both reef framework carbonate production (Perry et al., 2014a; Pratchett et al., 2015) and bioerosion rates (Perry et al., 2014a; Weinstein et al., 2014), and considerations of how rates of both may respond to future climate change drivers (e.g., Fang et al., 2013; Barkley et al., 2015; DeCarlo et al., 2015; Januchowski-Hartley et al., 2017; Schönberg et al., 2017). In addition, there has been considerable effort aimed at improving the understanding of how net reef carbonate budget states vary between reefs and how they may respond to ecological change, these being based on either census (e.g., Stearn et al., 1977; Scoffin et al., 1980; Hubbard et al., 1990; Eakin, 1996; Mallela and Perry, 2007; Perry et al., 2012, 2013b) or hydrochemical (e.g., Smith and Kinsey, 1976; Gattuso et al., 1996; Andersson and Gledhill, 2013; Shaw et al., 2016) *in-situ* measurements. In some cases a combination of both methods have been applied (e.g., Courtney et al., 2016) and up-scaling, based on remotely sensed imagery, applied to derive reef-wide production rates (e.g., Andréfouët and Payri, 2001; Moses et al., 2009; Hamylton et al., 2013b). In a few cases this budget work has also been extended to considerations of future reef growth potential as a function of contemporary budget states (Perry et al., 2015b; Perry and Morgan, 2017), and to the modeling of carbonate budgets under future change scenarios (Alvarez-Filip et al., 2013; Kennedy et al., 2013). Finally, there has been some recent work aimed at using census-based approaches to quantify rates of reef-scale carbonate sediment generation, using habitat specific census-data to explore the linkages between reefs and adjacent reef islands (Perry et al., 2015a; Morgan and Kench, 2016a), and remotely sensed imagery to up-scale production rates for specific sediment producing taxa (e.g., Doo et al., 2017).

Whilst this body of research provides an improved understanding of both site specific and intra-regional scale carbonate budget states, and of the key biological agents of inorganic carbonate production and erosion on reefs, there has been a tendency in much of this work to focus only on assessments of the most highly productive reef and coral-dominated zones (see Yamano et al., 2000; Harney and Fletcher, 2003; Hart and Kench, 2007 for useful exceptions). There are, of course, many valid reasons for doing this, most especially in settings where reefs form semi-continuous linear, or circum-reef platform, structures, and thus play an especially critical role as wave protecting structures (Ferrario et al., 2014), and as key sites of sediment supply (Perry et al., 2015a). However, in many settings reef development can be far more discontinuous. Indeed, in the case of individual reef platforms or coastal shelf environments, the coral-dominated reefal habitats may actually comprise a relatively small proportion of the overall

contemporary platform surface. Instead, complex mosaics of hardgrounds, sand- or rubble-dominated, and seagrass habitats can occupy a greater proportional area of the platform or shelf surface and, as earlier studies on reef carbonate production rates suggested, these lower productivity zones can result in significantly reduced reef-scale carbonate budgets (Smith and Kinsey, 1976; Buddemeier and Smith, 1988; Kinsey and Hopley, 1991). In order to improve our understanding of net inorganic carbonate cycling rates on reefs, to derive more accurate measures of platform scale carbonate budgets, and to better understand how carbonate budgets may influence the development and stability of proximal reef islands or beaches, requires as many of the following as possible to be factored for: (1) rates of both reef framework production and loss based on habitat specific ground-truthed data; (2) measures of sediment production and dissolution rates (see for example recent work by Cyronak et al., 2013; Eyre et al., 2014; Andersson, 2015 which demonstrates the budgetary relevance of the latter); and (3) to account for how both may vary between habitats and, where relevant, how rates may vary with energy (Hamilton et al., 2013a) or depth (Hustan, 1985). Quantifying both the framework and sediment carbonate components of reef budgets is especially desirable in a changing global environmental context, since it is realistic to assume that the processes driving the framework and sediment generation parts of budgets will respond in non-uniform ways to climate change and direct anthropogenic factors.

Here we explore these ideas in the context of an atoll rim reef platform in the central Maldives (Vavvaru, Lhaviyani Atoll), and demonstrate how habitat-specific census-based approaches, combined with remotely-sensed and ground-truthed habitat mapping, can be used to refine estimates of reef scale carbonate productivity. Specifically, we combine recently adapted census-based methodologies to provide quantitative estimates of: (1) net reef framework budgets, based on habitat specific measures of framework carbonate production and bioerosion; and (2) net reef sediment budgets, based on quantification of the sediment derived from both direct sources and from the bioerosional reworking of reef framework, and factored to account for off-reef sediment flushing and dissolution. We calculate both parts of this whole reef carbonate budget approach within each of eight delineated reef-lagoon habitats around Vavvaru, and factor for habitat extent in subsequent assessments of the total net reef carbonate budget (G , where $G = \text{kg CaCO}_3 \text{ m}^{-2} \text{ year}^{-1}$). We then use the resultant data to explore spatial (between-habitat) variations in both framework and sediment production and loss rates, and consider the implications for platform-scale budgets and for the reef-associated landforms that these platforms support.

MATERIALS AND METHODS

Field Setting and Delineation of Geo-Ecological Zones

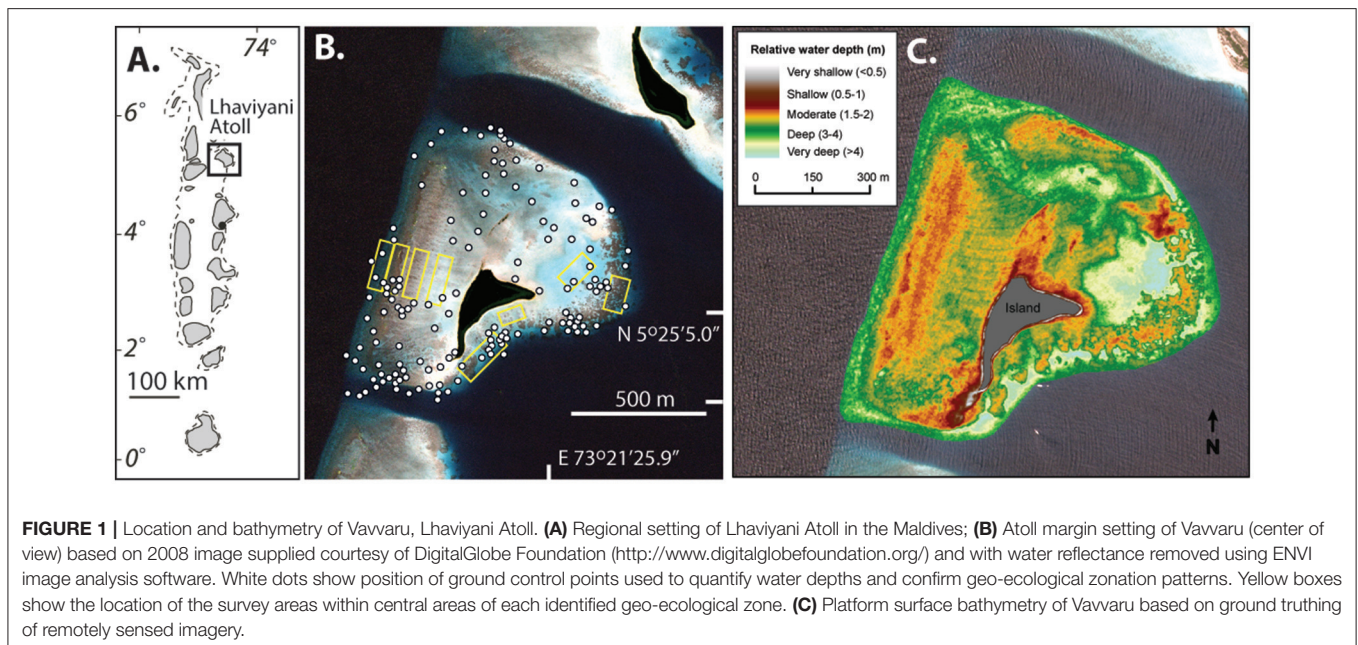
Vavvaru (N 5°25'5.0"; E 073°21'14.0") is located along the western side of Lhaviyani Atoll, in the northern-central region of the Maldives (Figure 1A). Vavvaru comprises a shallow sub-tidal

platform with a small ($\sim 300 \times 150$ m) island located within the south-eastern area of the platform (**Figure 1B**). The climate of the Maldives is defined by two monsoon periods marked by strong wind reversals, with southwest to northwest winds dominant from April to November (mean wind speed 5.0 ms^{-1}), and northeast-east winds dominant from November to March (mean wind speed 4.8 ms^{-1}) (Kench and Brander, 2006). The region is rarely affected by cyclones and so these seasonal process regime shifts are the main influences on cross- and off-reef platform sediment movement (Morgan and Kench, 2014). Our fieldwork was undertaken prior to the annual seasonal wind regime shift in February 2015. Prior to field census studies, and to provide the data necessary to inform our sampling strategy, we initially mapped the bathymetry and habitat zonation of the platform. Reef bathymetry (see **Figure 1C**) was estimated from multispectral Quickbird satellite imagery of Western Lhaviyani Atoll taken on 09/07/2008 (image provided by DigitalGlobe Foundation; <http://www.digitalglobe.com/>) using the SPEAR Relative Water Depth tool in ENVI v 5.4 with bathymetry calibration data collected in the field ($n = 146$ point depths that were reduced to mean sea level). At the same time the major sub-, and inter-tidal geo-ecological zones around Vavvaru were assessed by swimming cross platform transects and with reference to available Google Earth imagery (using the 2006, 2008, and 2011 images). Based on this imagery and the spot site verifications we identified eight main sub-tidal habitats/zones (**Figure 2A**) defined by different benthic substrate characteristics (percentage cover of consolidated reef, coral rubble, sediment and hard pavement), and the continuity of reef development (**Figures 2B–G; Table 1**). To determine the area occupied by each of these distinctive habitats/zones (m^2 and as a % of total platform area) we generated a habitat map (**Figure 2A**) in ERDAS IMAGINE 2015 using the same 2008 Quickbird imagery provided by DigitalGlobe Foundation. A subset of

the image around Vavvaru Island was orthorectified and pan-sharpened prior to conducting a supervised maximum likelihood classification. Classes corresponding to each of the previously delineated habitat types were defined by creating a signature file using areas of known habitat type, with the same 146 ground truth points used to conduct an accuracy assessment of the classification (overall accuracy was 77.2% compared to these ground truth points). The classified image was imported into ArcMap, and the number of pixels assigned to each class (habitat) extracted. The number of pixels was multiplied by the sensor resolution (0.6 m in both x and y axes) to derive total habitat areas. To calculate island movement over time (see **Figure 6**), the position of the island shoreline (defined as the edge of island vegetation) was digitized from the 2008 satellite image and from an earlier 1969 aerial photograph. The 1969 aerial image was geo-referenced in ArcMap v.10.1 using hard/stable structures (e.g., beachrock, coral heads) visible on the reef as ground control points. The extent of shoreline movement (m) was calculated as the total horizontal change between the consecutive time-series and a rate of shoreline movement (m year^{-1}) was estimated based on the total distance of change relative to the time interval between the images.

Quantifying Reef Framework Carbonate Production and Bioerosion

Benthic survey methodologies were used within each delineated reef platform habitat to determine: (1) general substrate characteristics (e.g., proportion of coral, sand and rubble cover, and substrate rugosity); (2) rates of reef framework carbonate production and bioerosion; and (3) carbonate sediment production. To quantify gross reef framework production and erosion and thus to determine net carbonate framework production rates (G , where $G = \text{kg CaCO}_3 \text{ m}^{-2} \text{ year}^{-1}$) we used an adapted version of the *ReefBudget* methodology of Perry et al.



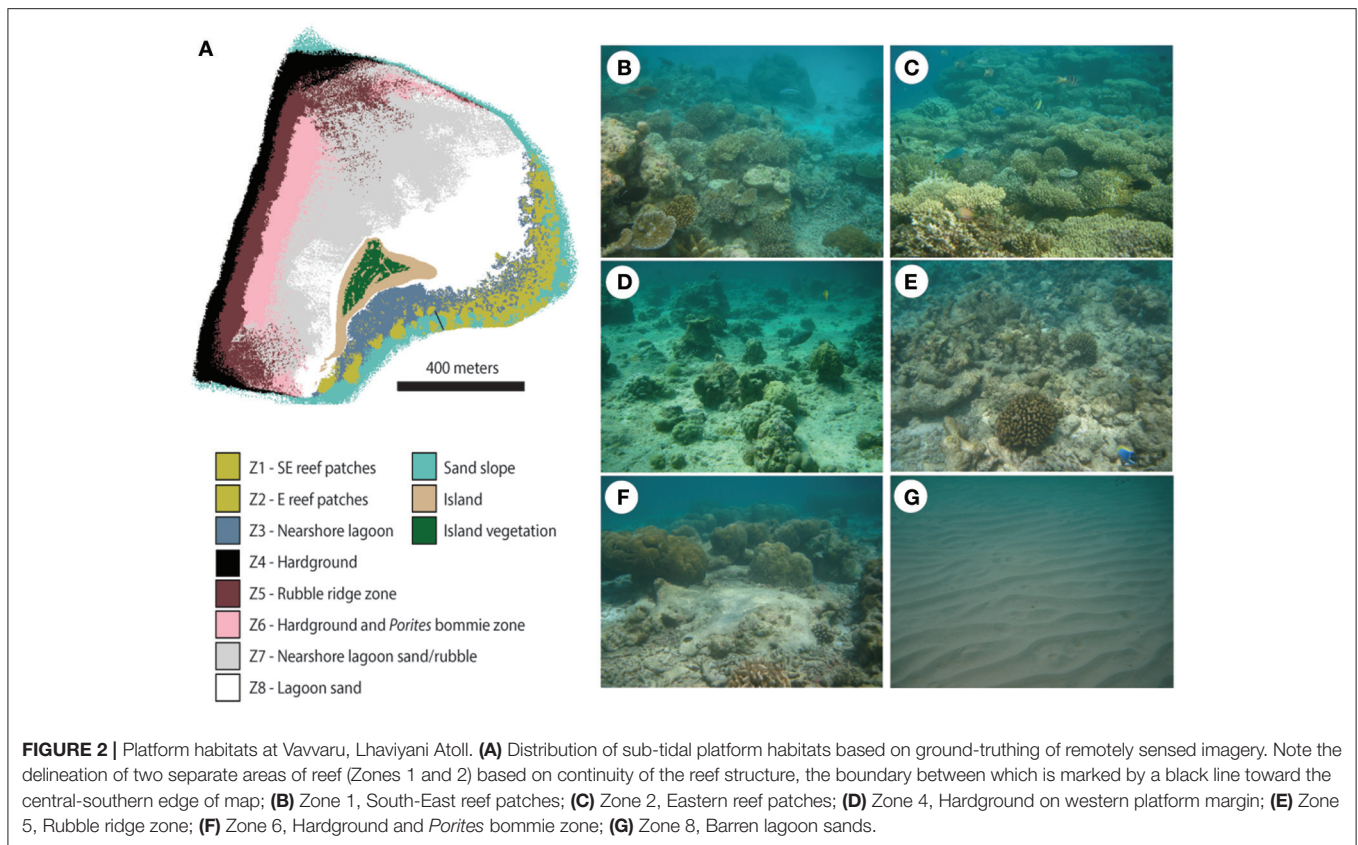


FIGURE 2 | Platform habitats at Vavvaru, Lhaviyani Atoll. **(A)** Distribution of sub-tidal platform habitats based on ground-truthing of remotely sensed imagery. Note the delineation of two separate areas of reef (Zones 1 and 2) based on continuity of the reef structure, the boundary between which is marked by a black line toward the central-southern edge of map; **(B)** Zone 1, South-East reef patches; **(C)** Zone 2, Eastern reef patches; **(D)** Zone 4, Hardground on western platform margin; **(E)** Zone 5, Rubble ridge zone; **(F)** Zone 6, Hardground and *Porites* bommie zone; **(G)** Zone 8, Barren lagoon sands.

TABLE 1 | Benthic characteristics of the eight delineated sub-tidal habitats at Vavvaru, Lhaviyani Atoll.

	Zone 1—SE reef patches	Zone 2—E reef patches	Zone 3—Nearshore (East side)	Zone 4—Hardground zone	Zone 5—Rubble zone	Zone 6— <i>Porites</i> bommie	Zone 7—Nearshore (W side)	Zone 8—Lagoon sand (N side)
Platform habitat area (m ²)	14,551	51,633	54,464	68,769	96,379	80,756	241,250	184,374
Proportion of platform area (%)	1.84	6.52	6.88	8.68	12.17	10.19	30.45	23.27
Reef framework as proportion of zone	0.92	0.97	0.14	0.87	0.95	0.79	0.55	0.00
Sediment as proportion of zone	0.08	0.03	0.86	0.13	0.05	0.21	0.45	1.00
Coral cover (%) ± 1 s.d	23.28 (4.33)	26.72 (3.15)	1.42 (1.52)	18.81 (8.13)	7.89 (2.91)	12.45 (2.54)	0.00 (0.00)	0.00 (0.00)
Rugosity ± 1 s.d	2.23 (0.41)	1.98 (0.14)	1.09 (0.06)	1.34 (0.08)	1.66 (0.23)	1.33 (0.08)	1.06 (0.02)	1.04 (0.02)

(2012) that we have previously applied at other central Indian Ocean, including Maldivian reef sites (Perry et al., 2015a; Perry and Morgan, 2017). Within each habitat we established five 10 m transects. Using the line as a guide, we then measured the distance within each linear 1 m covered by each category of benthic cover. All overhangs, vertical surfaces and horizontal surfaces below the line were surveyed (i.e., if the guide line crossed over a table coral, the upper and lower surfaces of the coral, plus the benthos under the canopy, were recorded). The following groups were recorded: scleractinian corals to the genera and morphological level (e.g., *Acropora* branching, *Porites* massive etc.); crustose coralline algae (CCA) including

CCA below macroalgal or soft coral cover; turf algae; fleshy macroalgae; non-encrusting calcareous algae (e.g., *Halimeda* spp., articulated coralline algae); sediment; bare substrate (e.g., limestone pavement); sediment; rubble; and other benthic organisms. Substrate rugosity was calculated as the total reef surface measured divided by linear distance. We then used the morphology and size of individual coral colonies in combination with genera specific skeletal density (g cm⁻³) and linear growth rates (cm year⁻¹) across each transect to estimate carbonate production rates in kg CaCO₃ m⁻² year⁻¹ (where m² refers to the planar surface of the reef). These were based on mean regional growth rates and densities for each coral genera (see

Perry and Morgan, 2017) for a summary of the rates used). These data were then combined with geometric transformations based on colony morphology to give a growth rate for each colony for the area under the transect line (taking a transect line width of 1 cm).

To determine rates of framework bioerosion by grazing invertebrates (parrotfish and urchins) we undertook two sets of measurements. To measure urchin bioerosion we surveyed urchin species abundance, the main agents of echinoid bioerosion on Indian Ocean reefs belong to the family Diademidae (*Diadema* spp., and *Echinostrephus* spp.), and the genera *Echinometra*, *Echinostrephus*, and *Eucidaris*, and their test size within 1 m wide belt transects along the main benthic transect lines at each site. We then used published test size—erosion rate relationships for different Indo-Pacific urchin species to calculate the erosion rate ($\text{kg urchin}^{-1} \text{ year}^{-1}$) for each individual urchin as follows: *Diademidae* = $0.000001x^{3.4192}$, *Echinometra* = $0.0004x^{1.9786}$, and for all others bioeroding urchins = $0.0001x^{2.323}$ (see Januchowski-Hartley et al., 2017). To calculate bioerosion by urchins in $\text{kg m}^{-2} \text{ year}^{-1}$, the erosion rates of all individual urchins within each transect can then be summed, and divided by the surface areas within each transect. To measure parrotfish erosion, we surveyed parrotfish numbers along eight replicate 30 m transects within each reef zone, with the numbers of parrotfish observed in the area 2 m either side of each transect recorded. Biomass of individual fish was then calculated using estimated length data and length-weight relationships and multiplied by abundance of the species or family of fish (see Perry et al., 2015b for details). To calculate parrotfish bioerosion rates by each individual fish we then used a model based on total length and life phase to predict the bite rates (bites h^{-1}) based on published data for that species, or for similar sized species with the same feeding functional group.

To estimate rates of endolithic bioerosion we utilized published rates of total macro- and microbioerosion measured at Indo-Pacific sites, alongside a census of substrate available for bioerosion from the benthic transects. This comprises of all dead carbonate substrate available to bioeroding sponges, including that covered by macroalgae or algal turf (see Perry et al., 2015b for further details of calculations). We also factored for background off-reef export of sediment and framework from relevant platform margin habitats, using per unit area rates measured from similar platform margin habitats (Morgan and Kench, 2014). On this basis we used a framework rubble export rate of $0.02 \text{ kg m}^{-2} \text{ year}^{-1}$ within the two reef zones and for the platform margin hardground zone, and a sand export rate of $0.15 \text{ kg m}^{-2} \text{ year}^{-1}$ within the two reef zones, the nearshore (east side lagoon zone) and the lagoonal zone on the northern side as these are all areas where some off-reef sediment export is likely to occur.

Sediment Production and Loss

To quantify rates of carbonate sediment generation within each platform zone/habitat we took account of both the sediment generated as a by-product of framework bioerosion, as well as new sediment generated directly by reef associated and benthic

taxa. In terms of sediment generated as a bioerosional by-product we assumed that all of the substrate grazed by parrotfish and urchins is then excreted as sediment—this assumption being based on the fact that there is no quantitative data showing dissolution of this material prior to excretion in either group. For macroendolithic (sponge) bioerosion we assumed that all of the material is converted to sediment, less a proportion that is dissolved during the chamber excavation process. Here we used a loss factor of 20% of the bioerosion rate as an average from all species where this has been reported (data in Nava and Carballo, 2008 and references therein). Microendolithic bioerosion is excluded in terms of sediment generation as the process is entirely chemical in nature and does not produce a sedimentary by-product.

Details of the methods used to measure direct sediment generation are provided in Perry et al. (2015a), but in brief we undertook a census of the following within an area $0.5 \text{ m} \times 0.5 \text{ m}$ either side of each of $3 \times 10 \text{ m}$ survey transect lines: 1. The volume of each *Halimeda* spp. plant based on *in situ* measures of the height, and maximum and minimum widths of the plant, with a carbonate sediment production rate than being determined based on established plant volume: segment count relationships and turnover rates for the main species present (*Halimeda micronesica* and *Halimeda macrophysa*) as determined at other sites in the Maldives (Perry et al., 2016); 2. The number of individual plants of other calcifying green algae (e.g., *Penicillus* sp.)—none were observed at Vavvaru; 3. The number of individual plants of articulated red coralline algae (e.g., *Amphiroa* sp.)—none were observed at Vavvaru; and 4. The number of epifaunal gastropods—with a carbonate production rate per specimen estimated from the rate data in Bosence (1989). In addition, we undertook separate analyses to estimate production rates by benthic foraminifera, infaunal bivalves, and seagrass epiphytes. Benthic foraminifera production rates were estimated based on counts of test abundance in the sediments ($n = 3$ samples per reef zone) and then converted to a production rate based on the relationships in Langer et al. (1997). Based on that methodology we used a conversion ratio of $6 \text{ g m}^{-2} \text{ year}^{-1}$ for every percent of foraminiferal skeletal component determined from reef sediments within reef and rubble dominated areas, and $1.2 \text{ g m}^{-2} \text{ year}^{-1}$ for lagoonal and sediment dominated habitats. Rates per zone were then factored for proportional sediment cover. The abundance of infaunal molluscs was measured by collecting three bulk sediment samples (area $\sim 10 \times 10 \text{ cm}$, and to a depth of 10 cm where possible) at equidistant points along each transect and sieving these to isolate living specimens. No living bivalves were recovered. Our surveys were also designed to account for the density of seagrass, which can produce large amounts of epiphytic carbonate (Nelsen and Ginsburg, 1986; Perry and Beavington-Penney, 2005), although no seagrass was present at Vavvaru. In our calculations we also factored for the effects of post-depositional sediment dissolution, which has recently been shown to be a quantitatively significant process in terms of reef carbonate sediment budgets (Cyronak et al., 2013; Eyre et al., 2014; Andersson, 2015). To this end we applied a rate of $0.15 \text{ kg CaCO}_3 \text{ m}^{-2} \text{ year}^{-1}$ of sediment dissolution occurring within each zone (based on data in Cyronak et al., 2013) and

factored this to account for the proportional cover of sediment per m^{-2} in each zone.

RESULTS

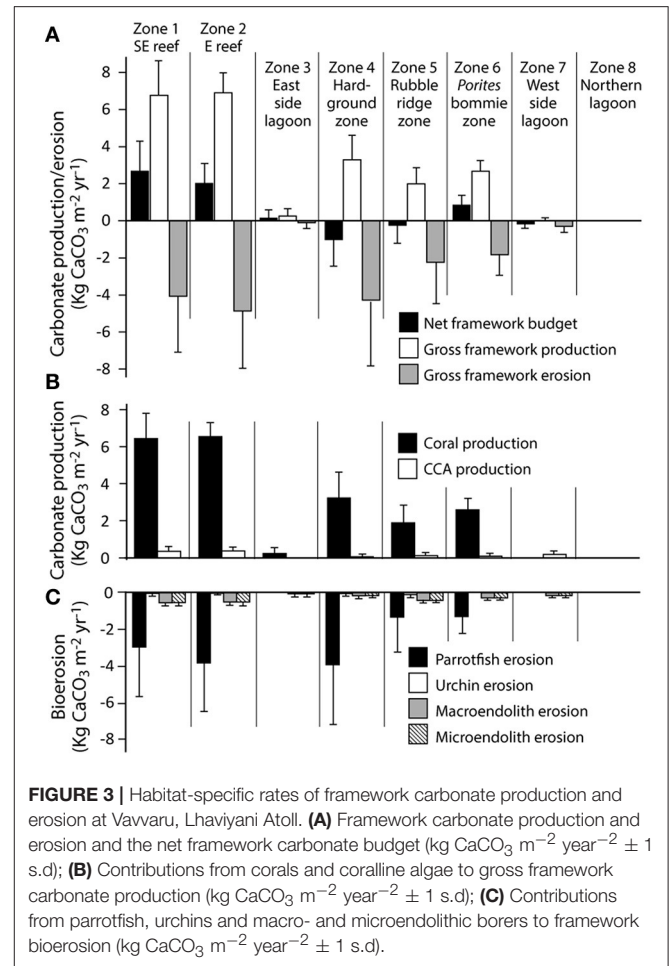
Framework Carbonate Budgets

Habitat scale net reef framework budgets (the balance between carbonate production and biological erosion) vary markedly between habitats around Vavvaru (range: 2.71 to -0.99 G; **Figure 3A**). Highest net rates (2.05 and 2.71 G, respectively; **Figure 3A**, **Table 2**) were measured in the two platform margin reef zones, zones defined by high rates of substrate bioerosion (4.06 and 4.85 G; **Figure 3A**; **Table 2**), but even higher rates of carbonate production (6.77 and 6.90 G, respectively; **Figure 3A**; **Table 2**). Corals are the dominant drivers of the high gross carbonate production rates measured, contributing 94.9 and 94.6%, respectively to calculated carbonate production in each reef zone. The most abundant coral genera in these reef habitats are branching and corymbose morphology *Acropora* spp., branching *Pocillopora* spp., and massive *Porites* sp., and these genera collectively accounted for 62 and 80% of the coral cover in each reef zone. They also contribute to between 80 and 88% of calculated coral carbonate production in the two reef habitats, with massive *Porites* sp. dominating production (46%) in reef habitat Z1 (SE reef patches), and branching and corymbose morphology *Acropora* spp. dominating production (86%) in reef habitat Z2 (Eastern reef patches).

In contrast, very low or net negative framework budgets (range: 0.85 to -0.99 G) define the lagoon and the western margin hardground and rubble-dominated habitats (**Figure 3A**, **Table 2**). In some cases, such as the sand-dominated lagoon habitats, this reflects the fact that there is little or no coral cover, but in other cases (e.g., the hardground and rubble zones on the western side of the platform) these net negative framework budgets are a function of higher (relative to gross production) rates of bioerosion (**Figures 3B,C**; **Table 2**). Calculated framework bioerosion rates range from 4.85 to 0.30 G across all habitats (**Table 2**). The highest rates (4.85 to 4.06 G) were measured in the reef and hardground habitats (**Figure 3C**; **Table 2**), and across all habitats around Vavvaru parrotfish were responsible for between 60 and 90% of this calculated framework bioerosion, the exceptions being within the sediment-dominated nearshore and lagoonal habitat zones where overall rates of bioerosion were in any case very low (<0.5 G; **Table 2**). Urchins, and macro- and micro-endolithic borers make relatively small contributions to total bioerosion rates within any of the habitats (**Figure 3C**).

Carbonate Sediment Budgets

Rates of calculated gross sediment production also differ markedly between habitats around Vavvaru, ranging from 5.18 to 0.00 G (**Figure 4**; **Table 3**). Highest rates were calculated with the reef habitats Z1 and Z2 (zones supporting extensive reef framework), and within the hardground habitat (3.46, 5.18, and 4.09 G, respectively; **Figure 4**; **Table 3**). Very substantial proportions of these values derive from the sediment generated from the bioerosional activities of parrotfish which excrete

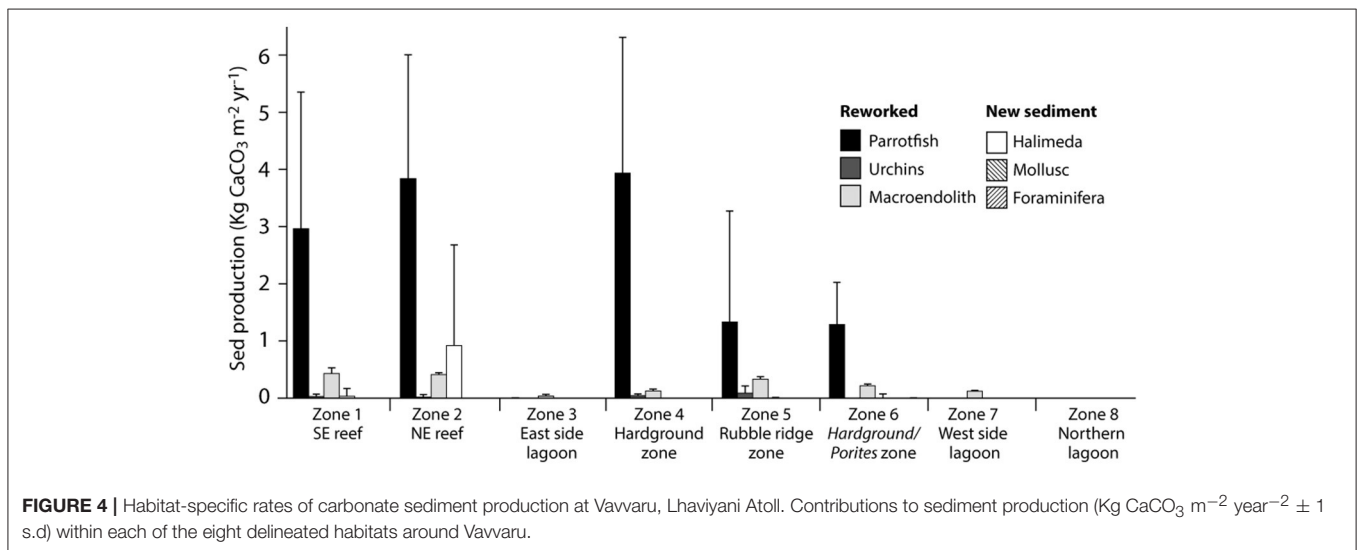


large amounts of carbonate sand as a by-product of substrate grazing (Bellwood, 1996). Rates of 3.42, 4.26, and 4.09 G, respectively, were calculated within each of these habitats (**Table 3**), and accounted for $>70\%$ of total estimated sediment generation (**Table 3**). Lower sediment generation rates (1.75 and 1.51 G) were calculated for the rubble ridge habitat (Z5) and the hardground/*Porites* sp. bommie habitat (Z6), although again most of this (between 75 and 85%) derives from parrotfish framework erosion (**Table 3**). Substantially lower rates of sediment production (<0.2 G) were calculated within each of the sediment-dominated lagoonal habitats (Zones 3, 7, and 8). Very little new sediment is produced by bivalves and gastropods (which were rare in our surveys), or by benthic foraminifera, which do not constitute more than 3% of the sedimentary constituents in any habitat. *Halimeda* is identified as the only other important sediment producer, although was only observed in four habitats (the two reef habitats, the rubble zone and the *Porites* bommie zone) and calculated production rates are only of budgetary relevance with reef habitat 2 (0.92 G; **Table 3**). Thus, it is reasonable to state that most of the sediment being generated on the Vavvaru platform as a whole is derived from the reworking (mainly by parrotfish bioerosion) of reef framework and other hard carbonate substrate.

TABLE 2 | Summary of key budget metrics associated with reef framework production and bioerosion, in each of the eight delineated sub-tidal habitats at Vavaru, Lhaviyani Atoll.

	Zone 1—SE reef patches	Zone 2—E reef patches	Zone 3—Nearshore (East side)	Zone 4—Hardground zone	Zone 5—Rubble ridge zone	Zone 6—Harground/ <i>Porites</i> bommie	Zone 7—Nearshore (W side)	Zone 8—Lagoon sand (N side)
Framework Production (G)								
Coral production (G)	6.428	6.530	0.243	3.234	1.878	2.586	0.000	0.000
CCA and encruster production (G)	0.344	0.372	0.003	0.054	0.116	0.090	0.179	0.000
Gross framework production (G)	6.772	6.902	0.246	3.288	1.994	2.667	0.179	0.000
Framework Erosion (G)								
Parrotfish erosion (G)	2.964	3.840	0.001	3.930	1.332	1.280	0.000	0.000
Urchin erosion (G)	0.030	0.020	0.000	0.040	0.090	0.000	0.000	0.000
Macroendolith bioerosion (G)	0.539	0.500	0.048	0.156	0.413	0.273	0.153	0.000
Microendolith bioerosion (G)	0.533	0.495	0.047	0.154	0.408	0.269	0.151	0.000
Gross framework erosion (G)	4.066	4.855	0.096	4.280	2.243	1.822	0.304	0.000
Net framework production (G)	2.706	2.047	0.150	-0.992	-0.249	0.854	-0.125	0.000
Off-reef coral rubble export rate (G) from relevant habitats ^a	0.020	0.020	N/a	0.020	N/a	N/a	N/a	N/a
Proportion of framework production converted to sediment (%)	52.17	63.17	19.92	125.49	92.03	58.03	85.47	N/a
Total net framework (G) accounting for rubble export	2.686	2.027	0.150	-1.012	-0.249	0.854	-0.125	0.000

^aBased on data in Morgan and Kench (2014). Shaded and bold lines show sub-totals.



DISCUSSION

The carbonate budget of a coral reef strongly influences the development and maintenance of its structural complexity, which is a key control on ecosystem service provisioning, and its long term net growth potential (Kinsey and Hopley, 1991; Perry et al., 2008). In this context, the early pioneering work undertaken on reef carbonate budgets reported that very high rates (in the range 10–12 G) of gross carbonate production may define those shallow water reefs having high branching coral cover (Smith and Kinsey, 1976; Kinsey, 1981; Kinsey and Hopley, 1991), and that

this may equate to potential vertical growth rates of around 7 mm year^{-1} (Kinsey and Hopley, 1991). Lower, but still significant rates (around 4 G) were reported to define reef flat habitats, whilst rates of around 0.5 G were reported for sand and rubble dominated reef habitats (Smith and Kinsey, 1976; Kinsey and Hopley, 1991).

Most of these early calculations were underpinned by novel hydrochemical based methodologies and these significantly advanced our capacities to start quantifying the carbonate budgets of different reef types. Indeed, the above, and other reported rates (e.g., Chave et al., 1972), were used as the basis

TABLE 3 | Summary of key budget metrics associated with sediment production and loss, in each of the eight delineated sub-tidal habitats at Vavvaru, Lhaviyani Atoll.

	Zone 1—SE reef patches	Zone 2—E reef patches	Zone 3—Nearshore (East side)	Zone 4—Hardground zone	Zone 5—Rubble ridge zone	Zone 6—Harground/ <i>Porites</i> bommie	Zone 7—Nearshore (W side)	Zone 8—Lagoon sand (N side)
Sediment Production (G)								
<i>Halimeda</i> (G)	0.035	0.921	0.000	0.000	0.003	0.002	0.000	0.000
Gastropods (G)	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Bivalves (G)	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Benthic foraminifera (G) ^a	0.001	0.000	0.002	0.000	0.001	0.004	0.001	0.002
Macrobioerosion-derived sediment ^b	0.431	0.400	0.038	0.125	0.330	0.218	0.122	0.000
Parrotfish (G)	2.964	3.840	0.001	3.934	1.332	1.288	0.000	0.000
Urchins (G)	0.030	0.020	0.000	0.040	0.090	0.000	0.000	0.000
Gross sediment production (G)	3.462	5.181	0.041	4.099	1.757	1.513	0.124	0.002
% from framework bioerosion	98.94	82.22	95.98	99.99	99.76	99.59	99.04	0.00
% contributed from parrotfish	85.61	74.11	2.44	95.98	75.83	85.15	0.00	0.00
Sediment loss from dissolution (G) ^c	0.012	0.005	0.129	0.020	0.008	0.032	0.068	0.150
Sediment export (G) ^d	0.150	0.150	0.150	N/a	N/a	N/a	N/a	0.150
Total net framework-derived sediment (G)	3.425	4.260	0.039	4.099	1.752	1.506	0.122	0.000
Total net new sediment (G)	0.037	0.921	0.002	0.000	0.004	0.006	0.001	0.002
Total net framework and sediment generated, less sediment dissolution and sand/rubble export (G)	5.986	7.054	-0.088	3.067	1.500	2.335	-0.069	-0.298

^aBased on test abundance and then converted to a production rate based on the relationships in Langer et al. (1997) using a conversion ratio of 6 g m⁻² year⁻¹ for every percent within sediments from high productivity reef and rubble dominated areas, and 1.2 g m⁻² year⁻¹ from sediment-dominated habitats, and with rates then factored for proportional sediment cover.

^bAssumes that all macrobioerosional-derived sediment is converted to sediment, less a proportion (using 20% as an average from all species where this has been reported in Nava and Carballo (2008) and references therein) that is dissolved during the chamber excavation process.

^cBased on a rate of 0.15 kg CaCO₃ m⁻² year⁻¹ of sediment dissolution occurring within each sediment-dominated zone (data from Cyronak et al., 2013) and factored to account for the proportional cover of sediment per m⁻² in each zone.

^dBased on a sand export rate of 0.15 kg m⁻² year⁻¹ using data in Morgan and Kench (2014). Shaded and bold lines show sub-totals.

for making some of the first spatial scale assessments of reef-wide carbonate production, both for a range of idealized reef system types (e.g., for the different reef types that define the Great Barrier Reef; Kinsey and Hopley, 1991), and for a range of reef geomorphic types (i.e., atoll, barrier, fringing; Chave et al., 1972). Subsequent reviews of reef carbonate production rates helped further define the optimal rates associated with different coral biogeographic provinces, and the different production rates that define different depth zones within each of these provinces (Vecsei, 2001). This collective body of data has thus provided a framework to support local scale reef production rate estimates (Smith and Kinsey, 1976), to contribute to global scale assessments (Milliman and Droxler, 1995), and more recently to support revised reef- and regional-scale estimates of reef carbonate production based on improved spatial mapping techniques (e.g., Hamylton et al., 2013b; Leon and Woodroffe, 2013).

Such metrics have growing relevance, however, because of the profound changes that are now impacting coral reefs globally, caused both by direct anthropogenic disturbances and by climate change drivers (especially temperature-induced coral bleaching events), and both of which may radically alter reef carbonate

budget states (Eakin, 1996; Edinger et al., 2000; Perry and Morgan, 2017). As a result, there has been a growing interest in using reef budgets as a reef health indicator (e.g., Perry et al., 2008; Mace et al., 2014). This has merit because of the underpinning role that carbonate budgets play in controlling many physical functional aspects of reefs, including their growth potential, their role as generators of coastal sediment, and their role in supporting many economically important reef-associated species. However, to more usefully utilize budget datasets in support of these issues, and to be able to more successfully monitor how contemporary budget states vary both between reefs and between different reef habitats, requires a number of advances to be made. We need to be able to more fully quantify whole reef-scale carbonate budgets that account for the different drivers of both framework and sediment production. We need to be able to better predict how these rates may change as different drivers of carbonate production and erosion change over time (e.g., Kennedy et al., 2013), and we need to better understand how reef-scale rates may vary as a function of the types and extent of different reef habitats.

Here, we have outlined an approach to mapping reef-scale carbonate budgets that allows us to factor for estimates of both

TABLE 4 | Summary of key budget metrics associated with reef framework and sediment budgets (based on data from **Tables 2, 3**) scaled to account for habitat spatial extent (see **Table 1**), in each of the eight delineated sub-tidal habitats at Vavvaru, Lhaviyani Atoll.

	Zone 1—SE reef patches	Zone 2—E reef patches	Zone 3—Nearshore (East side)	Zone 4—Hardground zone	Zone 5—Rubble ridge zone	Zone 6—Harground/ Porites bommie	Zone 7—Nearshore (W side)	Zone 8—Lagoon sand (N side)
TOTAL net framework production within platform zone (kg CaCO ₃ year ⁻¹), accounting for rubble export	39,085	104,660	8,170	-69,594	-23,998	68,966	-30,156	0
TOTAL net sediment budget within platform zone (kg CaCO ₃ year ⁻¹), accounting for sediment dissolution and export (G)	48,020	259,555	-12,960	280,536	168,570	119,606	13,531	-54,958
Total net framework production (kg CaCO ₃ year ⁻¹) across platform	220,880		Total net sediment production (kg CaCO ₃ year ⁻¹) across platform		889,818			
Total net framework loss (kg CaCO ₃ year ⁻¹) across platform	-123,749		Total net sediment loss (kg CaCO ₃ year ⁻¹) across platform		-67,918			
Net framework budget (G) per unit area of total platform	0.123		Net sediment budget (G) per unit area of platform		1.038			

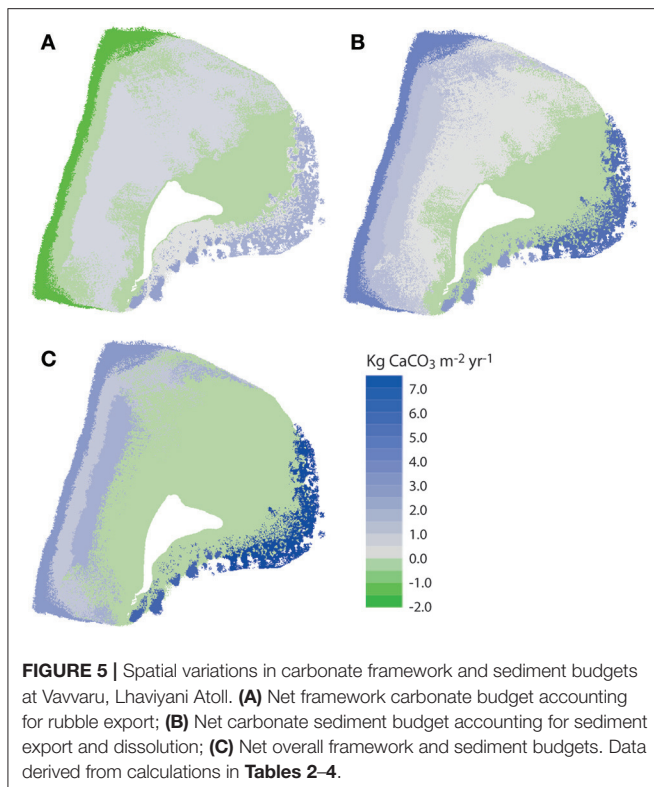
Bold lines show sub-totals.

framework and sediment production and erosion, to identify the different drivers of both, and to assess these processes within different habitat types whose spatial extent has been constrained by a combination of satellite image analysis and ground truthing. Based on the application of this approach at Vavvaru in the Maldives we calculate that the overall platform budget averages 1.15 G. At this site, this results from very small areas of high productivity reef, but much larger areas that are defined either by net negative budgets, or by very low net positive budgets. Highest rates of framework carbonate production occur, not surprisingly, in the two reef habitats, and in the two other habitats where corals are relatively abundant (the shelf-edge hardground habitat and the *Porites* bommie zone; **Tables 2–4**), whilst bioerosion-derived sediment production in these same zones is also responsible for a very high proportion of the sediment generation we calculate. Low rates of framework carbonate production and bioerosion, but also of new sediment generation, define the sand-dominated lagoon habitats around Vavvaru (**Tables 2–4**). Although, not quantified it must thus be assumed that these sand-dominated lagoon zones represent “sinks” for a high proportion of the sediment (new or reworked) that is produced in the more productive reef habitats (e.g., Purdy and Gischler, 2005; Schlager and Purkis, 2013). However, in terms of constructing habitat specific platform scale budgets, rates of framework and sediment export and loss also need to be taken into account. Such rates are not simplistic to measure *in situ* and so, as outlined in the methods, we have utilized published data on rates of framework rubble export and sediment export from other sites in the Maldives, and also factored a generic literature-derived sediment dissolution rate in our calculations, and then applied these to appropriate habitats around Vavvaru (see **Tables 2–3**).

The net effect of these calculations i.e., accounting for net framework and sediment generation, less sediment dissolution and sand/rubble export as appropriate, is to exacerbate patterns of inter-habitat heterogeneity. Resultant net total carbonate budgets range from 5.99 and 7.05 G within the two reef habitats, to -0.07 to -0.29 G within the sand-dominated nearshore and

lagoonal habitats (**Figure 5C**). Rates thus differ by more than one order of magnitude between habitats. The implications for the wider platform system are, however, much more profound when account is made of the spatial extent of the different habitats. Most noticeable is that the two habitats with the highest framework, sediment and overall net budgets (reef habitats 1 and 2; **Figures 5A–C**) and which thus drive most of the carbonate production on the platform actually occupy the smallest proportion of the reef platform (1.84 and 6.52%, respectively; **Tables 2–4**). Conversely, the habitats that occupy the largest spatial areas, the western nearshore habitat (Zone 7; 30.4% of platform area) and the large lagoonal area on the northern side of the platform (Zone 8; 23.3% of platform area) have amongst the lowest net budget rates (both being net negative; -0.07 and -0.29 G, respectively; **Figure 5C, Tables 2–4**). If these individual habitat rates for framework and sediment production are then multiplied by the total habitat area and summed they generate estimates of the respective total rates of framework and sediment production and loss for the entire platform. On this basis we calculate that 220,880 kg of framework CaCO₃ year⁻¹ is produced across the entire platform, against a framework breakdown and export figure of 123,749 kg CaCO₃ year⁻¹. This equates to a net platform scale reef framework budget of only 0.12 G (**Table 4**). Sediment production is calculated to total 889,818 kg CaCO₃ year⁻¹, and sediment dissolution and export 67,918 kg CaCO₃ year⁻¹. This equates to a net platform scale budget of 1.04 G, and thus a combined whole platform net carbonate budget of only 1.15 G (**Table 4**).

As evident from these figures, and an implication of the classic early carbonate budget studies outlined above, is that a focus on only the high productivity, high coral cover reef zones on reefs will thus generate much higher budget estimates than are representative of most reef systems as a whole. Indeed, our data suggests that there can be close to an order of magnitude difference between the budgets of the platform margin reef habitats and the much lower budgets that define the overall reef



system, a distinction that is again in line with the ideas presented in many of the early studies which compared reef, reef flat and lagoon habitat rates (see Kinsey, 1981; Kinsey and Hopley, 1991). However, upscaling habitat specific rates in the past often proved problematic due to field mapping constraints, but advances in satellite image resolution (such as utilized here) are making such assessments increasingly accurate (e.g., Andréfouët and Payri, 2001; Moses et al., 2009; Hamylton et al., 2013b). Such data can thus be combined with appropriate field-derived budget estimates to develop not only much improved estimates of reef-scale carbonate production, but also key insights into the most quantitatively significant species that drive the resultant patterns. This, in turn, has the potential to feed into modeling efforts aimed at understanding the implications of different species transitions, caused by environmental change, on reef carbonate production and erosion rates (see Perry et al., 2011 for conceptual models of such changes). Such approaches have recently been applied to project reef framework budget transitions for a hypothetical Caribbean reef under different climate change and management intervention scenarios (Kennedy et al., 2013), but increasingly there is the potential to undertake similar work for the wider carbonate producing system.

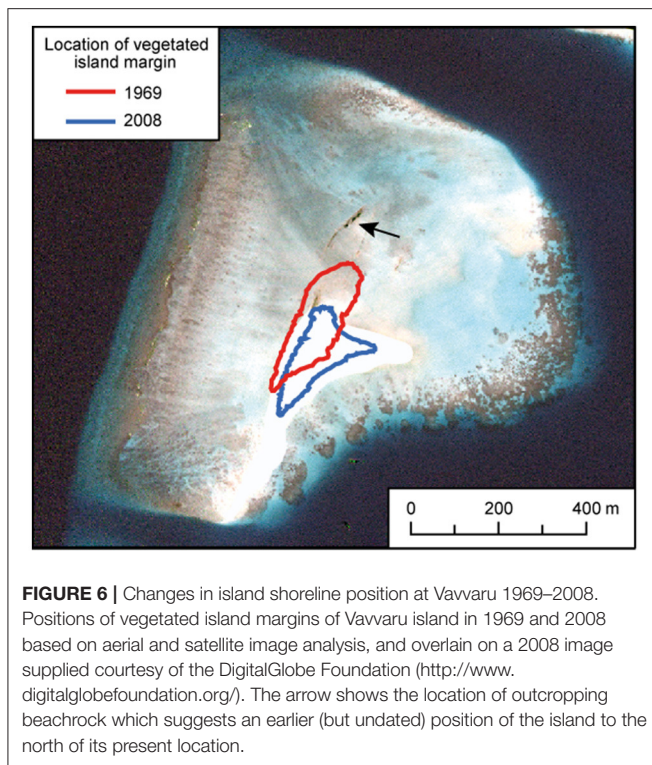
At Vavvaru the high coral carbonate production rates measured within the reefal zones along the platform margin were driven by high branching *Acropora* spp. cover, whilst parrotfish were identified as the major framework bioeroders and thus as major generators of carbonate sand. Parrotfish are well known to produce very large quantities of sediment as a by-product of their substrate grazing (Morgan and Kench, 2016b), and indeed

have been identified as the major sand generators in atoll interior faro settings elsewhere in the Maldives (Perry et al., 2015a). Maintenance of both thriving shallow water coral communities, even when occupying a relatively small area of the wider reef system, and of the associated parrotfish populations, are thus critical to maintaining the framework and sediment budgets of this system in net positive states.

A key issue that arises from the above observations are the impacts of the overall platform-scale carbonate budget on the stability, or lack thereof, of platform top reef islands. The formation and maintenance of these islands depends on there being a net positive supply of sediment from the generation sites to the nodal points of island formation (Gourlay, 1988; Mandlier and Kench, 2012). Given that the overall sediment budget at Vavvaru was calculated to be low (only 1.03 G), and thus that our best estimates suggest that only a little more sediment is produced overall within the platform than is lost, it would seem likely that the platform island sediment system at this site probably already exists close to a budget related stability threshold. In this context it is especially significant to note that Vavvaru Island has been highly mobile on its platform surface over the last ~45 years, the island shoreline migrating 129 m between 1969 and 2008 (a mean rate of 3.3 m year⁻¹). Whilst other factors, such as island size (e.g., Perry et al., 2013a) may contribute to this high mobility state, it is reasonable to hypothesize that the low rates of sediment supply as suggested by the current sediment budget study of the platform may be limiting island expansion and reducing the potential for the island position to stabilize.

Such issues become increasingly significant in the context of on-going reef degradation which can drive both progressive (Edinger et al., 2000) and rapid declines (Eakin, 1996) in reef carbonate budgets. Of immediate and pressing concern in a Maldivian context are the impacts of the major bleaching of mid-2016 caused by the very strong El Niño conditions. At sites in the southern Maldives recent work has reported on the very major declines in coral carbonate production rates and the widespread transition to net negative budget states that occurred at those sites (Perry and Morgan, 2017). Assuming similar transitions have occurred at Vavvaru, and reports from mid-2016 do indicate that bleaching of a similar magnitude was widespread across the Maldives (IUCN, 2016), it is entirely reasonable to hypothesize that carbonate production rates at Vavvaru will also have declined markedly since our work was undertaken in 2015. This would be especially significant because, as outlined above, these narrow coral-dominated zones make a disproportionately important contribution to maintaining the budget of the wider platform in an albeit low, but still net positive, budget state. Indeed, if similar bleaching-driven transitions have occurred at Vavvaru it is realistic to assume that the budget of the overall reef platform may now have become net negative.

Assuming the reef carbonate budget at Vavvaru has followed similar trajectories to sites in the southern Maldives (Perry and Morgan, 2017), there may thus be a number of implications for the island both positive and negative, and that will influence future island stability over different timescales. An initial major impact, would be for a likely overall transition in the framework budget to a net negative state. This may have only a minor



impact on the stability of the island itself in the short-term, but if low coral cover states persist then the growth potential of the reef areas will be diminished, limiting any wave protecting functions that these discrete areas of reef provide. At bleaching impacted sites in the southern Maldives, however, an additional consequence of coral die-off was an initial marked increase in parrotfish bioerosion (Perry and Morgan, 2017). If this has been repeated at Vavvaru the short-term budget impacts may actually be a spike in the generation of parrotfish derived sediment, which as the major sand producer on these reefs may result in a short-term additional pulse of sediment to the island. However, long-term persistence of low coral cover states combined with progressive bioerosion pressure would act to progressively reduce the structural complexity of the reefs, and which provides essential shelter and food availability (e.g., van Rooij et al., 1996). Thus, any short-term enhancement of the sediment budget may be offset by a progressive reduction in parrotfish biomass. Thus, one can hypothesize that future benthic community transitions that impact upon different aspects of carbonate production and erosion, may have rather profound implications for the stability of reef-associated landforms at this, and other similar, sites. Quantifying and mapping the rates and sources of carbonate production and erosion at reef-wide scales thus provides us with tools to consider these landform stability issues in the context of overall budget states and of habitat diversity and extent.

CONCLUSIONS

Early studies of reef carbonate budgets provided not only novel estimates of reef production rates, but also made the critical point

that large areas of most reef complexes i.e., encompassing reef flats and lagoon areas, are likely to have much lower carbonate production rates. It was thus suggested that overall reef-wide carbonate budgets are likely to be much lower, perhaps averaging only round 1.5 G. In other words the average reef system might have an approximately one order of magnitude lower production rate overall compared to that measured in the most productive (shallow, coral-dominated) zones. Our findings support this idea but extend this knowledge by developing an approach that allows us to better quantify which carbonate producing and eroding taxa drive these patterns across reef complexes, and how rates of both reef framework and carbonate sediment production and loss vary between habitats. This is an approach that has value for supporting efforts to up-scale estimates of inorganic carbonate production and cycling to reef-wide and, indeed, regional scales, and as outlined above is an issue with relevance beyond simply calculating reef-wide carbonate budget metrics. These processes of carbonate production, cycling, and loss are likely to have implications for the development and stability of proximal reef-associated landforms, including beaches and, in this example, reef islands. The approach we outline thus has potential to allow us to better explore the complex nature of geo-ecological linkages in reef systems and along reef-fronted coastlines, and to consider the budgetary implications on the form and development of reef-associated landforms. Such approaches are of course only as good as the empirical data that drive the production and erosion estimates, and whilst these are improving for many key aspects of framework and sediment production, there is a pressing need for additional data on coral calcification rates from many regions and for improved approaches to calculating sediment generation rates (see Perry et al., 2016 for such approaches in the content of the carbonate producing algae *Halimeda*). However, especially poorly constrained at present are data on rates of framework breakdown and export from reefs, and of sediment losses, both physical and geochemical. As recent studies have shown (Cyronak et al., 2013; Eyre et al., 2014; Morgan and Kench, 2014; Andersson, 2015) both areas are likely to be of significant budgetary relevance and urgently deserve additional attention. These data constraints aside, the current study highlights the marked differences that can exist in terms of total framework and sediment production between habitats, and the major dampening effect on the net carbonate budget at the platform scale when rates are scaled to account for habitat type and size. The net impact is that whilst individual reef systems may exhibit discrete areas of high coral cover and high associated carbonate production, platform scale budgets can be significantly lower, to the extent that any loss of coral cover and associated reductions in grazer-driven bioerosion in the main reef zones would be sufficient to tip the overall platform budget into a net negative state, with consequent implications for platform top island stability.

AUTHOR CONTRIBUTIONS

CP conceived the research. CP, KM, and RY contributed to field data collection and analysis, and to the writing of the paper.

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REFERENCES

- Alvarez-Filip, L., Carricart-Ganivet, J. P., Horta-Puga, G., and Iglesias-Prieto, R. (2013). Shifts in coral-assemblage composition do not ensure persistence of reef functionality. *Sci. Reports* 3:3486. doi: 10.1038/srep03486
- Andersson, A. (2015). A fundamental paradigm for coral reef carbonate sediment dissolution. *Front. Mar. Sci.* 2:52. doi: 10.3389/fmars.2015.00052
- Andersson, A. J., and Gledhill, D. (2013). Ocean acidification and coral reefs: effects on breakdown, dissolution, and net ecosystem calcification. *Ann. Rev. Mar. Sci.* 5, 321–348. doi: 10.1146/annurev-marine-121211-172241
- Andréfouët, S., and Payri, C. (2001). Scaling-up carbon and carbonate metabolism in coral reefs using in situ and remote sensing data. *Coral Reefs* 19, 259–269. doi: 10.1007/s003380000117
- Barkley, H. C., Cohen, A. L., Golbuu, Y., Starczak, V. R., DeCarlo, T. M., and Shamberger, K. E. F. (2015). Changes in coral reef communities across a natural gradient in seawater pH. *Sci. Adv.* 1:e1500328. doi: 10.1126/sciadv.1500328
- Bellwood, D. R. (1996). Production and reworking of sediment by parrotfishes (family Scaridae) on the Great Barrier Reef, Australia. *Mar. Biol.* 125, 795–800.
- Bosence, D. (1989). Biogenic carbonate production in Florida Bay. *Bull. Mar. Sci.* 44, 419–433.
- Buddemeier, R. W., and Smith, S. V. (1988). Coral reef growth in an era of rapidly rising sea level: predictions and suggestions for long-term research. *Coral Reefs* 7, 51–56. doi: 10.1007/BF00301982
- Chave, K. E., Smith, S. V., and Roy, K. J. (1972). Carbonate production by coral reefs. *Mar. Geol.* 12, 123–140. doi: 10.1016/0025-3227(72)90024-2
- Courtney, T. A., Andersson, A. J., Bates, N. R., Collins, A., Cyronak, T., de Putron, S. J., et al. (2016). Comparing Chemistry and Census-Based Estimates of Net Ecosystem Calcification on a Rim Reef in Bermuda. *Front. Mar. Sci.* 3:181. doi: 10.3389/fmars.2016.00181
- Cyronak, T., Santos, I. R., and Eyre, B. D. (2013). Permeable coral reef sediment dissolution driven by elevated pCO₂ and pore water advection. *Geophys. Res. Lett.* 40, 4876–4881. doi: 10.1002/grl.50948
- DeCarlo, T. M., Cohen, A. L., Barkley, H. C., Cobban, Q., Young, C., Shamberger, K. E., et al. (2015). Coral macrobioerosion is accelerated by ocean acidification and nutrients. *Geology* 43, 7–10. doi: 10.1130/G36147.1
- Doo, S., Hamylton, S. M., Byrne, M., and Finfer, J. (2017). Spatial and temporal variation in reef-scale carbonate storage of large benthic foraminifera: a case study on One Tree Reef. *Coral Reefs* 36, 293–303. doi: 10.1007/s00338-016-1506-0
- Eakin, C. (1996). Where have all the carbonates gone? A model comparison of calcium carbonate budgets before and after the 1982–1983 El Niño at Uva Island in the eastern Pacific. *Coral Reefs* 15, 109–119.
- Edinger, E. N., Limmon, G. V., Jompa, J., Widjatmoko, W., Heikoop, J. M., and Risk, M. J. (2000). Normal coral growth rates on dying reefs: are coral growth rates good indicators of reef health? *Mar. Poll. Bull.* 40, 606–617. doi: 10.1016/S0025-326X(99)00237-4
- Eyre, B. D., Andersson, A. J., and Cyronak, T. (2014). Benthic coral reef calcium carbonate dissolution in an acidifying ocean. *Nat. Clim. Change* 4, 969–976. doi: 10.1038/nclimate2380
- Fang, J. H. K., Mello-Athayde, M. A., Schoenberg, C. H. L., Kline, D. I., Hoegh-Guldberg, O., and Dove, S. (2013). Sponge biomass and bioerosion rates increase under ocean warming and acidification. *Glob. Change Biol.* 19, 3581–3591. doi: 10.1111/gcb.12334
- Ferrario, F., Beck, M. W., Storlazzi, C. D., Micheli, F., Shepard, C. C., and Airolidi, L. (2014). The effectiveness of coral reefs for coastal hazard risk reduction and adaptation. *Nat. Comms.* 5:3794. doi: 10.1038/ncomms4794
- Gattuso, J. P., Pichon, M., Delesalle, B., Canon, C., and Frankignoulle, M. (1996). Carbon fluxes in coral reefs. I. Lagrangian measurement of community metabolism and resulting air-sea CO₂ disequilibrium. *Mar. Ecol. Prog. Ser.* 145, 109–121. doi: 10.3354/meps145109
- Gourlay, M. R. (1988). “Coral cays: products of wave action and geological processes in a biogenic environment,” in *Proceedings of the 6th International Coral Reef Symposium*, Vol. 2 (Townsville, QLD), 491–496.
- Hamylton, S., Pescud, A., Leon, J. X., and Callaghan, D. P. (2013a). A Geospatial assessment of the relationship between reef flat community calcium carbonate production and wave energy. *Coral Reefs* 32, 1025–1039. doi: 10.1007/s00338-013-1074-5
- Hamylton, S., Silverman, J., and Shaw, E. (2013b). The use of remote sensing to scale up measures of carbonate production on reef systems: a comparison of hydrochemical and census-based estimation methods. *Int. J. Remote Sensing* 34, 6451–6465. doi: 10.1080/01431161.2013.800654
- Harney, J. N., and Fletcher, C. H. III. (2003). A budget of carbonate framework and sediment production, Kailua Bay, Oahu, Hawaii. *J. Sediment. Res.* 73, 856–868. doi: 10.1306/051503730856
- Hart, D. E., and Kench, P. S. (2007). Carbonate production of an emergent reef platform, Warraber Island, Torres Strait, Australia. *Coral Reefs* 26, 53–68. doi: 10.1007/s00338-006-0168-8
- Hoey, A. S., Howells, E., Johansen, J. L., Hobbs, J.-P. A., Messmer, V., McCowan, D. M., et al. (2016). Recent advances in understanding the effects of climate change on coral reefs. *Diversity* 8:12. doi: 10.3390/d8020012
- Hubbard, D. K., Miller, A. I., and Scaturo, D. (1990). Production and cycling of calcium carbonate in a shelf-edge reef system (St. Croix, U.S. Virgin Islands): applications to the nature of reef systems in the fossil record. *J. Sediment. Res.* 60, 335–360. doi: 10.1306/212F9197-2B24-11D7-8648000102C1865D
- Hustan, M. (1985). Variation in coral growth rates with depth at Discovery Bay, Jamaica. *Coral Reefs* 4, 19–25.
- IUCN (2016). *Maldives Coral Reefs under Stress from Climate Change: Research Survey Reveals Over 60% of Corals Bleached*. Available online at: <https://www.iucn.org/news/maldives-coral-reefs-under-stress-climate-change-research-survey-reveals-over-60-corals>. International Union for Conservation of Nature (Accessed on: 01-09-16)
- Januchowski-Hartley, F. A., Graham, N. A. J., Wilson, S. K., Jennings, S., and Perry, C. T. (2017). Predicting divergent reef carbonate budget trajectories following a major climatic disturbance event. *Proc. R. Soc. B* 284:20162533. doi: 10.1098/rspb.2016.2533
- Kench, P. S., and Brander, R. W. (2006). Response of reef island shorelines to seasonal climate oscillations: South Maalhosmadulu atoll, Maldives. *J. Geophys. Res.* 111:F01001. doi: 10.1029/2005JF000323
- Kennedy, E. V., Perry, C. T., Halloran, P. R., Fine, M., Carricart-Ganivet, J. P., Iglesias-Prieto, R., et al. (2013). Avoiding coral reef functional collapse requires combined local and global action. *Curr. Biol.* 23, 912–918. doi: 10.1016/j.cub.2013.04.020
- Kinsey, D. W. (1981). “The Pacific/Atlantic reef growth controversy,” in *Proceedings of the 4th International Coral Reef Symposium*, Vol. 1 (Manila), 493–498.
- Kinsey, D. W., and Hopley, D. (1991). The significance of coral reefs as global carbon sink-response to greenhouse. *Palaeogeog. Palaeoclim. Palaeoecol.* 89, 363–377.
- Langer, M. R., Silk, M. T., and Lipps, J. H. (1997). Global ocean carbonate and carbon dioxide production: the role of foraminifera. *J. Foramin. Res.* 27, 271–277.
- Leon, J. X., and Woodroffe, C. D. (2013). Morphological characterisation of reef types in Torres Strait and an assessment of their carbonate production. *Mar. Geol.* 338, 64–75. doi: 10.1016/j.margeo.2012.12.009
- Mace, G. M., Reyers, B., Alkemade, R., Biggs, R., Chapin, S. F. III., Cornell, S. E., et al. (2014). Approaches to defining a planetary boundary for biodiversity. *Glob. Env. Change* 28, 289–297. doi: 10.1016/j.gloenvcha.2014.07.009

- Mallela, J., and Perry, C. T. (2007). Calcium carbonate budgets for two coral reefs affected by different terrestrial runoff regimes, Rio Bueno, Jamaica. *Coral Reefs* 26, 129–145. doi: 10.1007/s00338-006-0169-7
- Mandlier, P. G., and Kench, P. S. (2012). Analytical modelling of wave refraction and convergence on coral reef platforms: Implications for island formation and stability. *Geomorph* 159, 84–92. doi: 10.1016/j.geomorph.2012.03.007
- Milliman, J. D., and Drozler, A. W. (1995). Calcium carbonate sedimentation in the global ocean: linkages between the neritic and pelagic environments. *Oceanography* 8, 92–94.
- Morgan, K. M., and Kench, P. S. (2014). A detrital sediment budget of a Maldivian reef platform. *Geomorph* 222, 122–131. doi: 10.1016/j.geomorph.2014.02.013
- Morgan, K. M., and Kench, P. S. (2016a). Reef to island sediment connections on a Maldivian carbonate platform: using benthic ecology and biosedimentary depositional facies to examine island-building potential. *Earth Surface Proc. Land* 41, 1815–1825. doi: 10.1002/esp.3946
- Morgan, K. M., and Kench, P. S. (2016b). Parrotfish erosion underpins reef growth, sand talus development and island building in the Maldives. *Sed. Geol.* 341, 50–57. doi: 10.1016/j.sedgeo.2016.05.011
- Moses, C. S., Andréfouët, S., Kraneburg, C. J., and Muller-Karger, F. E. (2009). Regional Estimates of reef carbonate dynamics and productivity using Landsat 7 ETM+ and potential impacts from ocean acidification. *Mar. Ecol. Prog. Ser.* 380, 103–115. doi: 10.3354/meps07920
- Nava, H., and Carballo, J. L. (2008). Chemical and mechanical bioerosion of boring sponges from Mexican Pacific coral reefs. *J. Exp. Mar. Biol. Ecol.* 211, 2827–2831. doi: 10.1242/jeb.019216
- Nelsen, J. E., and Ginsburg, R. N. (1986). Calcium carbonate production by epibionts on *Thalassia* in Florida Bay. *J. Sed. Pet.* 56, 622–628.
- Pendleton, L., Comte, A., Langdon, C., Ekstrom, J. A., Cooley, S. R., and Suatoni, L. (2016). Coral reefs and people in a High-CO₂ world: where can science make a difference to people? *PLoS ONE* 11:e0164699. doi: 10.1371/journal.pone.0164699
- Perry, C. T., and Beavington-Penney, S. J. (2005). Epiphytic calcium carbonate production and facies development within sub-tropical seagrass beds, Inhaca Island, Mozambique. *Sed. Geol.* 174, 161–176. doi: 10.1016/j.sedgeo.2004.12.003
- Perry, C. T., Edinger, E. N., Kench, P. S., Mumby, P. J., Murphy, G., Steneck, R. S., et al. (2012). Estimating rates of biologically driven coral reef framework production and erosion: a new census-based carbonate budget methodology and applications to the reefs of Bonaire. *Coral Reefs* 31, 853–868. doi: 10.1007/s00338-012-0901-4
- Perry, C. T., Kench, P. S., O'Leary, M. J., Morgan, K. M., and Januchowski-Hartley, F. (2015a). Linking reef ecology to island-building: Parrotfish identified as major producers of island-building sediment in the Maldives. *Geology* 43, 503–506. doi: 10.1130/G36623.1
- Perry, C. T., Kench, P. S., O'Leary, M., Riegl, B. R., Smithers, S. G., and Yamano, H. (2011). Implications of reef ecosystem change for the stability and maintenance of coral reef islands? *Glob. Change Biol.* 17, 3679–3696. doi: 10.1111/j.1365-2486.2011.02523.x
- Perry, C. T., Kench, P. S., Smithers, S. G., Yamano, H., O'Leary, M., and Guilliver, P. (2013a). Timescales and modes of reef lagoon infilling in the Maldives and controls on the onset of reef island formation. *Geology* 41, 1111–1114. doi: 10.1130/G34690.1
- Perry, C. T., and Morgan, K. M. (2017). Bleaching drives collapse in reef carbonate budgets and reef growth potential on southern Maldives reefs. *Sci. Rep.* 7:40581 doi: 10.1038/srep40581
- Perry, C. T., Morgan, K. M., and Salter, M. A. (2016). Sediment generation by *Halimeda* on atoll interior coral reefs of the southern Maldives: a census-based approach for estimating carbonate production by calcareous green algae. *Sed. Geol.* 346, 17–24. doi: 10.1016/j.sedgeo.2016.10.005
- Perry, C. T., Murphy, G. N., Graham, N. A. J., Wilson, S. K., Januchowski-Hartley, F. A., and East, H. (2015b). Remote coral reefs can sustain high growth potential and may match future sea-level trends. *Sci. Reps.* 5:18289. doi: 10.1038/srep18289
- Perry, C. T., Murphy, G. N., Kench, P. S., Edinger, E. N., Smithers, S. G., Steneck, R. S., et al. (2014a). Changing dynamics of Caribbean reef carbonate budgets: emergence of reef bioeroders as critical controls on present and future reef growth potential. *Proc. R. Soc. B* 281, 2014–2018. doi: 10.1098/rspb.2014.2018
- Perry, C. T., Murphy, G. N., Kench, P. S., Smithers, S. G., Edinger, E. N., Steneck, R. S., et al. (2013b). Caribbean-wide decline in carbonate production threatens coral reef growth. *Nat. Comms.* 4:1402. doi: 10.1038/ncomms2409
- Perry, C. T., Spencer, T., and Kench, P. (2008). Carbonate budgets and reef production states: a geomorphic perspective on the ecological phase-shift concept. *Coral Reefs* 27, 853–866. doi: 10.1007/s00338-008-0418-z
- Perry, C. T., Steneck, R. S., Murphy, G. N., Kench, P. S., Edinger, E. N., Smithers, S. G., et al. (2014b). Regional-scale dominance of non-framework building corals on Caribbean reefs affects carbonate production and future reef growth. *Glob. Change Biol.* 21, 1153–1164. doi: 10.1111/gcb.12792
- Pratchett, M. S., Anderson, K. D., Hoogenboom, M. O., Widman, E., Baird, A. H., Pandolfi, J. M., et al. (2015). Spatial, temporal and taxonomic variation in coral growth – implications for the structure and function of coral reef ecosystems. *Oceanogr. Mar. Biol.* 53, 215–295.
- Purdy, E. G., and Gischler, E. (2005). The transient nature of the empty bucket model of reef sedimentation. *Sed. Geol.* 175, 35–47. doi: 10.1016/j.sedgeo.2005.01.007
- Saunders, M. I., Albert, S., Roelfsema, C. M., Leon, J. X., Woodroffe, C. D., Phinn, S. R., et al. (2016). Tectonic subsidence provides insight into possible coral reef futures under rapid sea-level rise. *Coral Reefs* 35, 155–167. doi: 10.1007/s00338-015-1365-0
- Schlager, W., and Purkis, S. J. (2013). Bucket structure in carbonate accumulations of the Maldives, Chagos and Laccadive archipelagos. *Int. J. Earth Sci.* 102, 2225–2238. doi: 10.1007/s00531-013-0913-5
- Schönberg, C. H., Fang, J. K., Carreiro-Silva, M., Tribollet, A., and Wisshak, M. (2017). Bioerosion: the other ocean acidification problem. *ICES J. Mar. Sci.* 74, 895–925. doi: 10.1093/icesjms/fsw254
- Scoffin, T., Stearn, C. W., Boucher, D., Frydl, P., Hawkins, C. M., Hunter, I. G., et al. (1980). Calcium carbonate budget of a fringing reef on the west coast of Barbados. I. erosion, sediments and internal structure. *Bull. Mar. Sci.* 30, 475–508.
- Shaw, E., Hamylton, S. M., and Phinn, S. (2016). Incorporating benthic community changes into hydrochemical-based projections of coral reef calcium carbonate production under ocean acidification. *Coral Reefs* 35:739. doi: 10.1007/s00338-016-1407-2
- Smith, S. V., and Kinsey, D. W. (1976). Calcium carbonate production, coral reef growth, and sea level change. *Science* 194, 937–939. doi: 10.1126/science.194.4268.937
- Stearn, C. W., Scoffin, T. P., and Martindale, W. (1977). Calcium carbonate budget of a fringing reef on the West Coast of Barbados part I – zonation and productivity. *Bull. Mar. Sci.* 27, 479–510.
- van Rooij, J. M., de Jong, E., Vaandrager, F., and Videler, J. J. (1996). Resource and habitat sharing by the stoplight parrotfish, *Sparisoma viride*, a Caribbean reef herbivore. *Environ. Biol. Fish* 47, 81–91. doi: 10.1007/BF00002381
- Vecsei, A. (2001). Fore-reef carbonate production: development of a regional census-based method and first estimates. *Palaeogeog. Palaeoclim. Palaeoecol.* 145, 185–200. doi: 10.1016/S0031-0182(01)00371-6
- Weinstein, D. K., Smith, T. B., and Klaus, J. S. (2014). Mesophotic bioerosion: variability and structural impact on U.S. Virgin Island deep reefs. *Geomorph* 222, 14–24. doi: 10.1016/j.geomorph.2014.03.005
- Yamano, H., Miyajima, T., and Koike, I. (2000). Importance of foraminifera for the formation and maintenance of a coral sand cay: Green Island, Australia. *Coral Reefs* 19, 51–58. doi: 10.1007/s003380050226

Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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