



Impacts of more frequent droughts on a relict low-altitude *Pinus uncinata* stand in the French Alps

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Cold microclimatic conditions provide exceptional microhabitats to *Pinus uncinata* stands occurring at abnormally low altitudes in seven relict stands of the northern French Alps. Here, *P. uncinata* is located at the lower bounds of its ecological limits and therefore expected to provide a sensitive indicator of climate change processes. We used dendrochronological analysis to study the growth patterns of closely spaced chronologies across an elevational transect and compare a relict low-altitude to a *P. uncinata* stand located at the alpine treeline. Two detrending procedures are used to reveal high- and low-frequencies embedded in annually resolved ring-width series. Growth response of *P. uncinata* to instrumental temperature and precipitation data is investigated by means of moving response function analyses. Results show an increase in the sensitivity of tree-ring widths to drought during previous summer in both stands. At the treeline stand, an increasing response to fall temperature is observed whereby fall temperature and radial tree growth increased in two synchronous steps around ~1930 and from ~1980-present. At the low-altitude stand, *P. uncinata* appears more drought sensitive and exhibits a sharp growth decline since the mid-1980s, coinciding with increasing summer temperatures. Growth divergence between the two stands can be observed since the mid-1980s. We argue that the positive growth trend at the high-altitude stand is due to increasing fall temperatures which would favor the formation of metabolic reserves in conjunction with atmospheric CO₂ enrichment that in turn would facilitate improved water use efficiency. At the relict low-altitude stand, in contrast, it seems that improved water use efficiency cannot compensate for the increase in summer temperatures.

Keywords: relict pine stand, tree-growth decline, climate change, dendroclimatology, response functions, French Alps

INTRODUCTION

In the Northern French Alps, Mountain pines (*Pinus uncinata* Mill. ex Mirb.) are usually at their optimum in the subalpine belt between 1600 and 2300 m a.s.l. (Rolland and Schueller, 1995), but seven stands have been reported to grow at abnormally low elevations within a mountain belt of the Rhone Alpes and at altitudes comprised between 1100 and 1500 m a.s.l. The existence of these exceptional stands in the Devoluy (Combe Obscure; Bertinelli et al., 1993; Bois du Sappey; Cécillon et al., 2010) and Grande Chartreuse (La Plagne cirque; Cassagne et al., 2008) massifs was associated with the circulation of cold air in scree slopes, which may locally create periglacial pedoclimatic conditions (Cécillon et al., 2010). The microclimate of cold scree slopes allows for exceptional microhabitats to develop with very thick and raw humus (Cassagne et al., 2008), forming ideal niches for various cold living plants (Bertinelli et al., 1993) and periglacial

relict species of non-vascular plants (Asta et al., 2001), beetles (Molenda, 1996), spiders and mites (Zacharda et al., 2005). These microhabitats, which are older than the surrounding biological matrix, act as refugia by supporting relict communities unable to survive elsewhere in the landscape. In these refugia, *P. uncinata* is located at the lower elevation bounds of its ecological limits and can therefore be expected to be a sensitive indicator of climate change processes (e.g., Oberhuber et al., 2008). Furthermore, in these open, low-production forests both competition between species and forestry operations are very limited. As a consequence, we though hypothesized that these low-altitude *P. uncinata* stands may bear an important and “pure” imprint of accelerated warming observed in the Alpine region in the last three decades (Böhm et al., 2010; Corona et al., 2010a,b).

This study presents results from a relict low-altitude *P. uncinata* stand situated on a north-facing slope void of periglacial

edaphic conditions, and where growth dynamics are expected to be reflective of changes in climatic conditions. Based on extensive dendrochronological analyses, we compare growth patterns of closely spaced *P. uncinata* chronologies across an elevational transect from a low-altitude (LAS, 1150 m asl) to a high altitude stand (HAS, 1900 m asl) located at treeline. In particular, we (i) identify high- to low-frequencies embedded in annually resolved ring-width series; (ii) use response function analyses to investigate growth responses in *P. uncinata* to instrumental temperature and precipitation records; (iii) compare ring-width chronologies from LAS and HAS with monthly mean temperature and precipitation data; and (iv) discuss growth-climate responses as well as climate related growth shifts.

STUDY SITES

The two sites investigated are located within the Parc National des Ecrins (45°02'N, 6°8'E) where the most easterly of the low-altitude *P. uncinata* stands of the northern French Alps have been identified (Figures 1A,B). The stands are located on the north-facing Liassic limestone-schist slopes of the Grande Aiguille and the Petite Aiguille (Oisans massif, Isère). The stands, formed almost exclusively of *P. uncinata*, occupy an area of ~130 ha located between 1150 and 1900 m a.s.l., with roughly one hundred individuals located at abnormally low altitudes. The low-altitude stand has been described in the literature as a cold subalpine enclave in a montane ecosystem dominated by beech–fir (*Fagus–Abies*) and Norway spruce (*Picea abies* (L.) Karst.) forests (Godron and Salomez, 1995). The cold climatic conditions associated with these north-facing, steep slopes are conducive to low-altitude colonization by mesocryophilic *P. uncinata* (Rameau et al., 1993; Petitcolas, 1998). Ground flora is characterized by an important moss cover, species typical for raw humus soils (*Rhododendron ferrugineum* L., *Vaccinium vitis-idaea* L., *V. myrtillus* L.) and accompanied by subalpine calciphilic species (*Dryas octopetala* L., *Saxifraga oppositifolia* L. or *Arctostaphylos uva-ursi* (L.) Spreng.). The stands grow on relatively shallow soils, predominantly protorendzina, i.e., rendzic and lithic leptosols with moder humus according to the FAO classification system (IUSS Working Group WRB, 2006). They consist of unconsolidated, coarse-textured materials with low water holding capacity. Distinct soil horizons are hardly ever developed.

Climate at the study site is reflective of continental conditions observed in the intra-alpine zone of the French Alps. Mean annual precipitation (1951–2007) at the Saint-Christophe-en-Oisans (44°56'N, 6°11', 1570 m) weather station located 3 km east of the study site is 974 ± 186 mm year⁻¹ with a rainfall minimum in summer (mainly July and August) and a rainfall maximum in autumn. Mean annual temperature is 5.9°C with 148 days year⁻¹ of freezing.

MATERIALS AND METHODS

FIELD COLLECTION AND SAMPLE PREPARATION

P. uncinata trees were sampled at LAS (1150 m asl) and HAS (i.e., treeline, 1900 m asl) where slope steepness ranges between 15° and 25°. A total of 72 trees (36 at LAS and 36 at HAS, Figure 2A) with absence of major stem or crown anomalies (due to lightning, wind or snow breakage) was sampled and two increment cores

extracted per tree from opposing radii (144 increment cores). Ring widths were measured to the nearest 0.01 mm using a digital LINTAB positioning table connected to a Leica stereomicroscope and TSAPWin Scientific software (Rinntech, 2009).

CHRONOLOGY DEVELOPMENT AND STATISTICS

Prior to detrending, all ring-width series were screened for missing rings and dating errors using COFECHA (Cook and Holmes, 1984). Comparison of the LAS and HAS chronologies was performed using a recent version of ARSTAN (Cook, 1985) and included Mean Sensitivity (MS) to represent the mean of the relative deviations between consecutive rings for a given tree (Fritts, 1976) and to assess high-frequency variations in the chronologies. In addition, first-order serial autocorrelation (AC) was used to detect the persistence retained in the chronologies.

Expressed Population Signal (EPS) quantifies the degree to which the constructed chronology portrays hypothetically perfect conditions (Wigley et al., 1984). We used an EPS threshold of 0.85 to determine the reliability of our chronologies. Two conceptually different standardization methods were used for this study, namely (i) a hybrid “double-detrending” which is more accurate as far as high-frequency variations concerned (Cook and Peters, 1981), as well as (ii) an Adaptive Regional Growth Curve standardization (ARGC, Nicault et al., 2010) where long-term trends in chronologies are preserved. Indices from the first detrending were used for the response function analyses and relate to the high-frequency growth variations. After smoothing both the tree-ring and meteorological series with 10- and 20-year low pass filters, the indices from the second detrending were visually compared with (multi-) decadal climate variations.

For the “double-detrending,” an initial negative exponential or linear regression was applied for each raw measurement series, followed by a fitting of a 30-year cubic smoothing spline with 50% frequency response cut-off (Cook and Peters, 1981). Then, for each stand, the growth indexes were averaged by year using a bi-weighted robust mean to develop a stand chronology which represented the common high frequency variation of the individual series (Cook, 1987).

The ARGC standardization (Nicault et al., 2010) was employed to preserve potential low-frequency climate variability. The ARGC approach is in fact a refinement of the Regional Curve Standardization (RCS; Briffa et al., 1992) and based on the adaptation of the RCS to each tree by taking into account the differences in cambial age and mean productivity which are functions of competition and soil (Rathgeber et al., 2000). Practically, tree-ring series are standardized into relative tree-ring indices by dividing each measured ring by its expected value estimated from a growth trend (Equation 1). An Artificial Neural Network (ANN; Nicault et al., 2010) was then used to estimate the ARGC for each tree from its cambial age, as well as initial and maximum productivity (Equation 2). In the present case we used a feed-forward network trained with an error back-propagation learning algorithm (Rumelhart et al., 1986) and took account of productivity (i.e., indexed series) on the basis of juvenile growth characteristics as follows:

$$It = Ct/Yt \quad (1)$$

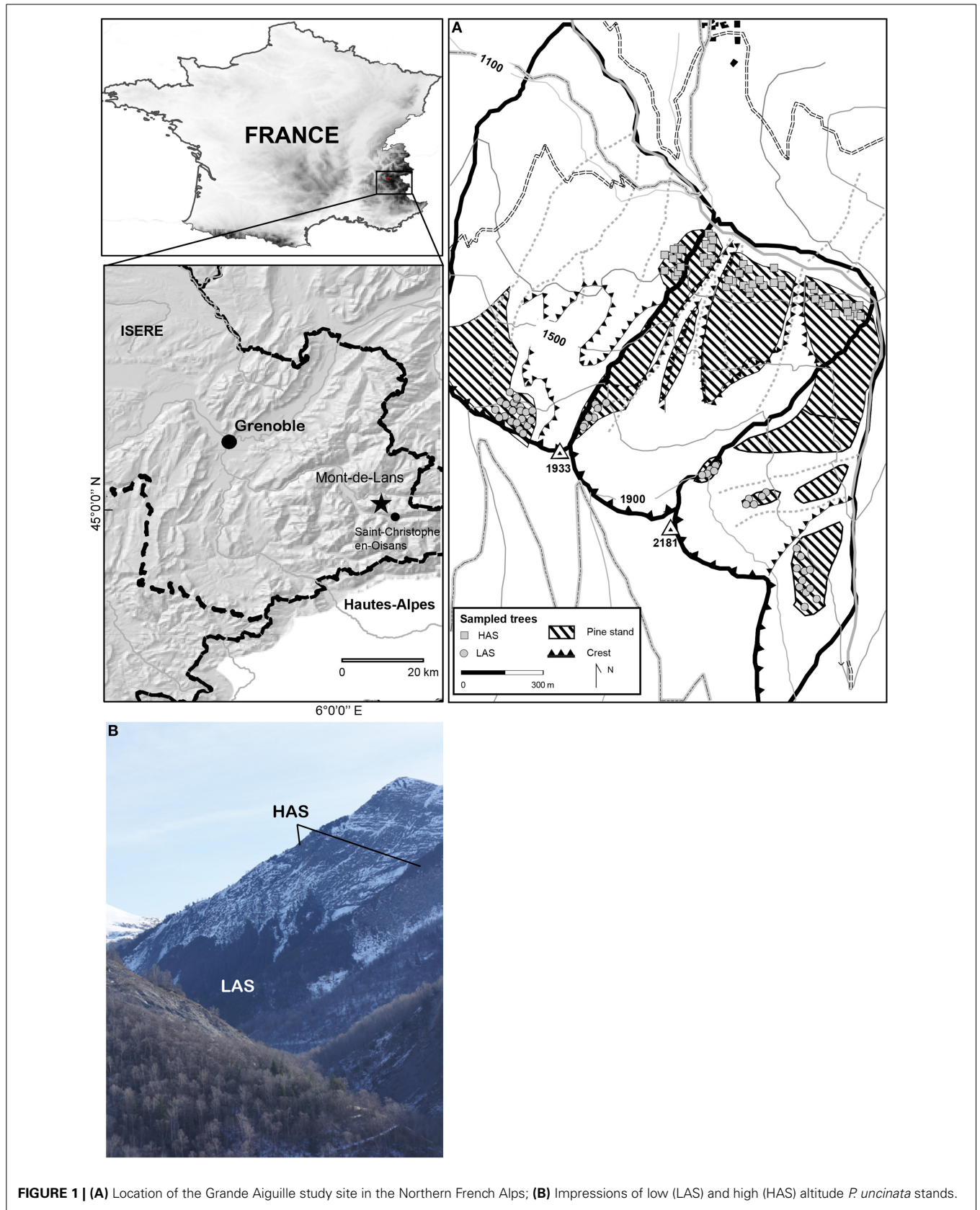
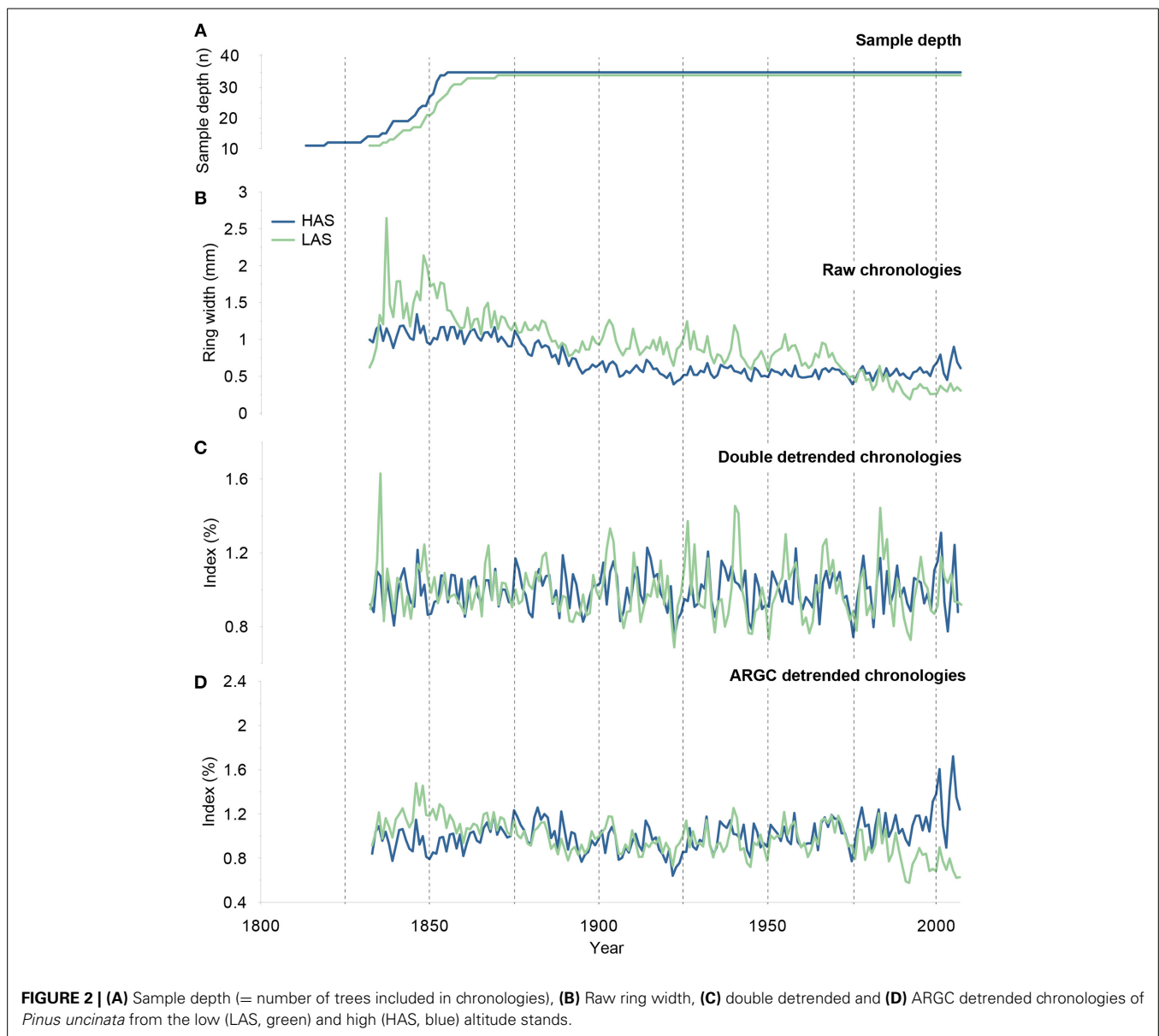


FIGURE 1 | (A) Location of the Grande Aiguille study site in the Northern French Alps; **(B)** Impressions of low (LAS) and high (HAS) altitude *P. uncinata* stands.



$$Y_t = F(\alpha(C_t), g(C), G(C)) \quad (2)$$

where Y_t is the tree-ring index, C_t the measured tree-ring width, Y_t the theoretical value of tree ring C_t , $\alpha(C_t)$ the age of tree ring C_t , $g(C)$ the initial growth of tree C , i.e., the average of the first 10 rings, $G(C)$ is the maximum juvenile growth of tree C , i.e., the maximum value reached during its juvenile stage (first 50 year) after smoothing with a 10-year window. As ANN fits each tree, detrending is not sensitive to the occurrence of missing rings at the tree's pith (Nicault et al., 2010). The approach has been demonstrated to be particularly well-suited to heterogeneous Alpine environments (Corona et al., 2010a,b) where rapid variations in soils, microclimate, competition and other factors governing site productivity may induce major divergence in tree-growth rates.

CLIMATE DATA SETS

Long instrumental series of meteorological parameters are largely lacking in the French Alps (Corona et al., 2010b). For the purpose of response function calibration, we therefore used homogenized temperature and precipitation records from the HISTALP dataset (Efthymiadis et al., 2006; Böhm et al., 2010). This database extends back to AD 1760 and covers the Greater Alpine Region (GAR, 4–19°E, 43–49°N, 0–3500 m asl). Data is represented as deviations from the 1961–1990 mean with a latitude-longitude resolution of $0.1^\circ \times 0.1^\circ$ for precipitation and $1^\circ \times 1^\circ$ for temperature. Given the large variability of climate in the Alps, especially for precipitation, we used mean monthly series of temperature and precipitation from the four HISTALP grid points located nearest to the *P. uncinata* stands (6–7°E, 44–45°N for temperature; 6°05'–6°15' E; 44°55'–45°05'N for precipitation). Pearson's correlation analysis with the meteorological series of

Saint-Christophe-en-Oisans (located near the stands) revealed highly significant coefficients for the time interval covered by both series (1951–2007).

CLIMATE GROWTH RELATIONSHIPS

In order to test for the sensitivity of the relict *P. uncinata* stand to climate change, climate–growth relationships have been tested with response function analyses, a form of principal component regression designed to account for collinearity of monthly climate predictors (Fritts, 1976). Bootstrapped climate–growth Response Functions (RF) were computed between double detrended chronologies and monthly climate predictors using the Dendroclim 2002 software (Biondi and Waikul, 2004).

The climatic dataset from the HISTALP database included mean monthly air temperature (°C) and total monthly precipitation (mm) sequenced from May of the year preceding the formation of the tree ring ($t-1$) to September of the year of actual tree-ring formation (t), yielding a set of 34 predictors. To evaluate the stationarity and consistency of responses over time, we computed Moving Response Functions (MRF) by progressively shifting the period of a fixed number of years across time to compute response function coefficients (Biondi and Waikul, 2004). To provide a sufficiently large number of degrees of freedom, the length of the calibration period was double than the number of predictors (34 in this study), so that only periods >68 years were considered. RF and MRF produce a temporal set of coefficients for each monthly predictor, whereby statistical significance at $p < 0.05$ was tested using a bootstrap procedure (Biondi and Waikul, 2004).

RESULTS

CHRONOLOGY DESCRIPTIVE STATISTICS

Two ring-width chronologies were developed for each stand (LAS and HAS, **Figure 2B**), whereby most tree-ring series show the typical age-related exponential decrease in ring width, also known as age trend. EPS values meet the signal strength acceptance in both chronologies (>0.85, 0.88 in HAS, 0.89 in LAS) from 1837 to 2007. High first-order autocorrelation (0.83 in HAS and 0.79 in LAS) indicates that radial growth throughout the slope was strongly influenced by conditions in the preceding year. Differences of mean sensitivities are rather small between both stands (0.19 in HAS, 0.18 in LAS).

RESPONSE FUNCTION ANALYSIS

Double detrended ring-width chronologies used for the calculation of RF and MRF are depicted in **Figure 2C**. RF coefficients were computed for the common period 1837–2007. They differ slightly between HAS and LAS (**Figure 3A**) with RF profiles showing that the dominant factors controlling tree-ring widths at HAS and LAS are summer (mainly July and August) temperatures and precipitations in the year preceding tree-ring formation. This influence is particularly acute at LAS. In contrast, below-average May to June temperatures combined with above-average precipitation totals between April and June of the year of growth-ring formation favored radial growth at LAS.

STATIONARITY OF THE MONTHLY CLIMATE RESPONSES OVER TIME

At LAS, MRF analysis revealed that the negative response of tree growth to summer temperatures (JAS) in the year preceding tree-ring formation ($t-1$) as well as positive influence of summer precipitation totals in July and August have been stable since 1837 (**Figures 3B,C**). At HAS, a significant response is observed for $t-1$ July temperature and radial growth, provided that the calibration interval ends before 1980. In addition, positive response is also fairly obvious between $t-1$ August temperatures and radial growth and sharply increased around 1930 (**Figure 3B**).

At LAS, temporally stable key climate variables are represented by t June temperature and April precipitation (**Figures 3B,C**). The strength of the negative response between t May temperature and ring width increased in recent decades and becomes significant provided that the calibration interval ends after the 1940s. Current (t) May, June and July precipitation were the most significant predictors of radial growth at HAS.

Several factors have been identified by the RFs as not being significant over the entire period but still having a periodic influence on tree growth. Among these, $t-1$ November temperature was a significant predictor of radial growth of LAS trees in case that the calibration interval ends after 1980, whereas the regression estimates increased as soon as the calibration period ends after the mid-1930s at HAS. At LAS, the negative relation of radial growth with t August temperatures was even reversed in the last decade.

ANALYSIS OF GROWTH TRENDS

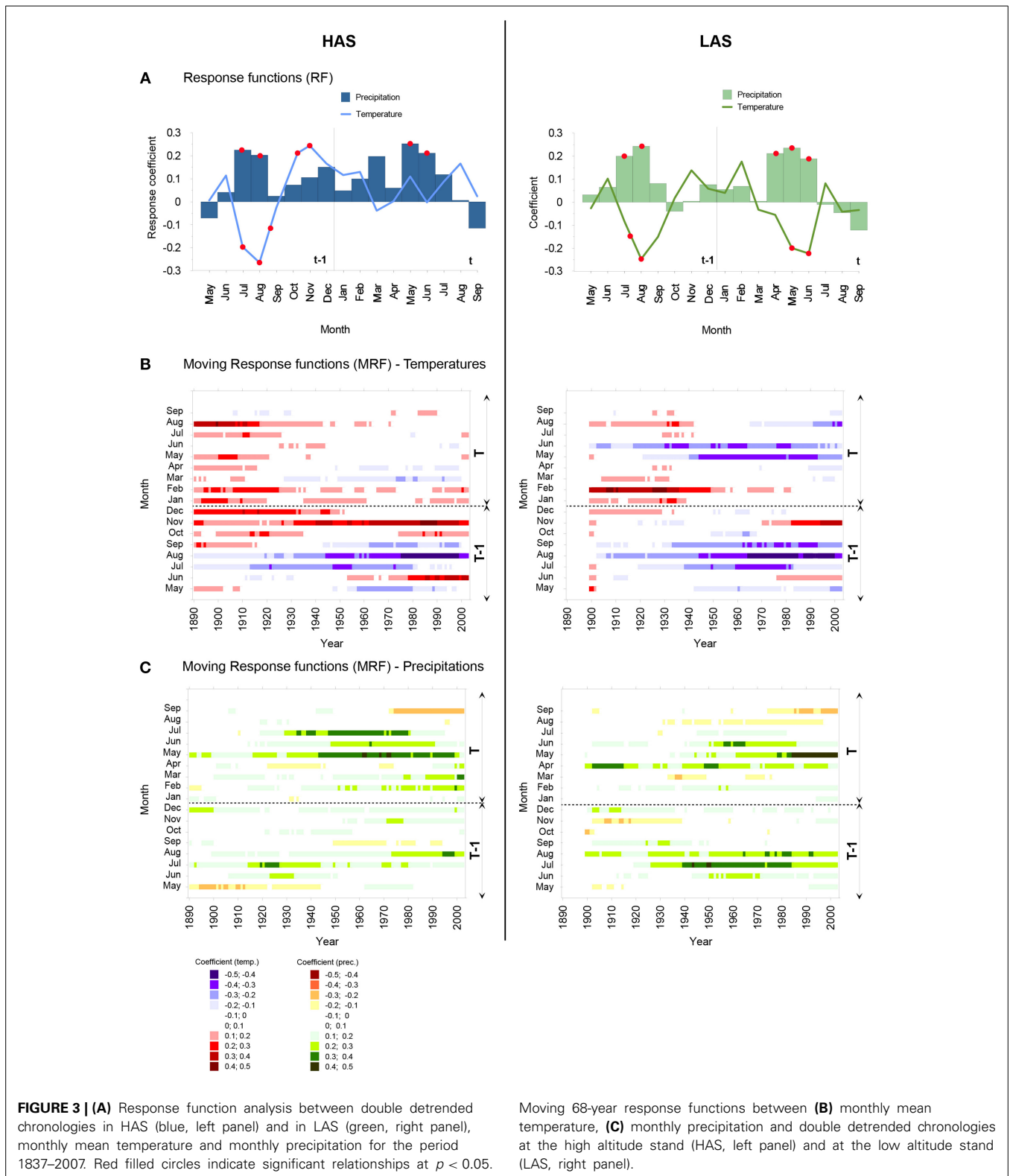
ARGC detrended ring-width chronologies are depicted in **Figure 2D** and are designed to maximize low-frequency information. In **Figure 4A**, ARGC chronologies from HAS are plotted with mean fall (October–November) temperature after a z-transformation and smoothing by a 20-year low-pass filter. We observe that low-frequency variability of temperature and radial tree growth increased in two synchronous steps at ca. 1930 and ca. 1980. Using a 10-year smoothing, HAS chronologies (**Figure 4B**) show a decadal variability in radial tree growth with depressions in 1860, 1895, 1920, 1945, 1975, and 1990 and a clear positive trend since the early 1990s. Decadal ring-width fluctuations properly coincide with rainfall totals during the vegetation period (which locally lasts from May to September, **Figure 4B**).

At LAS, depressions in radial growth are observed in 1895, 1935, 1945, 1960, and 1990. A negative trend is clearly visible since the 1960 (**Figure 4C**) and low ring-width values concur with below-average July to September precipitation in 1945, 1975, and 1990. In addition, negative coincidence is observed between July to September (JAS) temperature and the LAS ring-width oscillations since the 1940s (**Figure 4D**). Particularly, the growth decline observed in LAS smoothed tree-ring chronology since the 1980s coincides with positive JAS temperature anomalies.

DISCUSSION

RELATIONSHIP BETWEEN CLIMATE AND RADIAL GROWTH

In our study, growth of *P. uncinata* was above all limited by drought in the summer preceding tree-ring formation ($t-1$) and by precipitation in May and June of the year of radial growth (t) at both the low- and high-elevation stands (**Figure 3A**). The predominance of meteorological conditions of the previous growing



season on current radial growth has been reported for other *P. uncinata* stands in the French Alps (Rolland and Florence-Schueller, 1998) and the Spanish Pyrenees (Gutiérrez, 1991; Tardif et al., 2003). Limited radial growth occurring after dry $t-1$

summers likely reflects restricted root growth and/or the depletion of stored reserves (Petitcolas, 1998). Similarly, the positive response of radial growth of *P. uncinata* to spring precipitation has been reported in other subalpine belts of the French

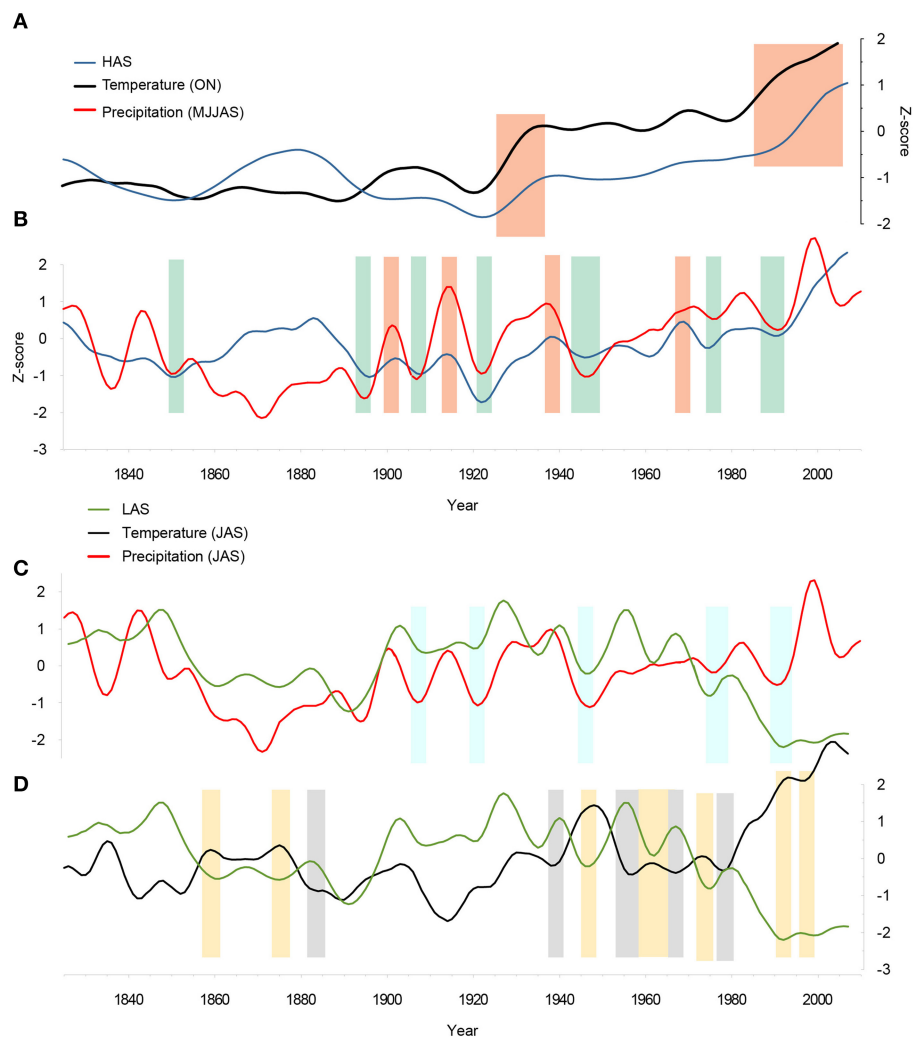


FIGURE 4 | Comparison of low-pass filtered ARGc detrended chronologies and key climate variables for the period 1837–2007: (A) 20-year low-pass filtered HAS chronology and Fall (ON) temperature; (B) 10-year low-pass filtered HAS chronology and May to September precipitations; (C) 10-year low-pass filter LAS chronology, summer (JJA) temperatures and (D) precipitations. Both the climate series and the tree-ring chronologies were transformed to z scores over the 1837–2007

common period. Orange (green) bars emphasize synchronous periods of growth increase (decrease) in tree rings. Based on visual comparison, blue bars emphasize periods of growth decrease coinciding with negative JAS precipitation anomalies. Yellow bars indicate periods of growth decline coinciding with positive JAS temperature anomalies. Gray bars indicate periods of tree-growth increase synchronous with negative JAS temperature anomalies.

Alps, (Rolland and Schueller, 1996; Petitcolas, 1998) and for *P. sylvestris* and *P. cembra* stands growing in dry inner alpine valleys (Oberhuber et al., 1998, 2008; Rigling et al., 2002; Pfeifer et al., 2005). At our sites, growth limitations are probably favored by the presence of shallow and poorly developed soils on schists with low water-holding capacity.

The growth of individuals at HAS is positively correlated with temperatures in $t-1$ autumns. These findings are in agreement with results of Rolland and Schueller (1996) and Petitcolas (1998) for north-facing *P. uncinata* stands of the Briançon area where low temperatures in $t-1$ autumn are likely to limit the formation of metabolic reserves and to consequently affect radial growth in year t (Rolland and Schueller, 1995). Similar observations have been reported for *P. cembra* by Pfeifer et al. (2005) and Gruber

et al. (2010). As photosynthesis can occur until late October (i.e., after the termination of growth processes) under favorable conditions (Tranquillini, 1959), sugar can be stored and will be available at the beginning of the next growing season (Wardle, 1971). As a consequence, warm October and November temperatures will facilitate carbon storage, promote mycorrhizal root growth by maintaining soils above freezing and favor maturation of needles, shoots and buds against detrimental effects of winter stress (Oberhuber et al., 2008).

At LAS, *P. uncinata* trees appear to be drought sensitive with high temperature at the beginning of the growing period also being a limiting factor for cambial activity. This specific relationship was not previously observed at alpine treeline sites (Rolland and Schueller, 1996) and contradicts results of previous

studies realized in the French Alps and the Pyrenees (Rolland and Schueller, 1995; Tardif et al., 2003) for which May temperature was positively correlated with tree-ring width.

TEMPORAL EVOLUTION OF MONTHLY CLIMATE RESPONSES

A temporal change of months significantly correlated with radial growth was found for the LAS and HAS stands since AD 1837. We observe that negative (positive) relationship with $t-1$ August ($t-1$ November) temperature has increased since 1837, therefore suggesting that carry-over effects (such as photosynthetic gain, or the storage of assimilates and water from the $t-1$ growing season) would have impacted increasingly current-year (t) radial growth over the course of the 20th century (Kozłowski and Pallardy, 1997). In a similar way, increasingly positive relationships between growth and precipitation at the beginning of the growing period (May) seem to be closely affected by increasing spring-summer temperatures controlling soil moisture balance (Martínez-Vilalta et al., 2008).

At LAS, the increasing negative relationship with mean temperature in May indicates exacerbated moisture stress at the onset of the growing period (Figures 3B,C). In that sense, results are in agreement with studies from drought-stressed populations of other coniferous species in the Mediterranean (Martin-Benito et al., 2010) and North America (Biondi, 2000).

RISING TEMPERATURES AS A DRIVER OF DIVERGENT GROWTH

At HAS, radial growth of *P. uncinata* exhibits successive stages of increase starting around 1920 and 1980 synchronous with increases in October and November temperature (Figure 4A). Comparable results were reported for *P. uncinata* in the French Alps (Rolland et al., 1998) and for *P. cembra* in the central Swiss Alps (Vittoz et al., 2008), where increased radial growth could be successfully correlated with mild autumn temperatures since the 1980s.

Conversely, the *P. uncinata* trees sampled at LAS exhibit a strong negative growth trend since the mid-1980s. The sharp decline observed in radial growth since the mid-1980s is likely reflective of increasing summer temperatures. Results from LAS are thus in line with other studies reporting growth decline in other *Pinus* sp. for various sites in the Mediterranean (see Martin-Benito et al., 2010 for a review), more xeric treeline locations in boreal forests (Lloyd and Fastie, 2002) or within dry inneralpine valleys where altered temperatures have been held responsible for increased mortality of susceptible or less competitive species (Bigler et al., 2006).

POSSIBLE EFFECTS OF OTHER ENVIRONMENTAL FACTORS ON TREE GROWTH DIVERGENCES BETWEEN SITES

Given the poor productivity of stands growing close to or at the treeline, difficult access to the study sites as well as the sylvo-pastoral activity practiced in the region since the Middle Ages (Allix, 1929), forestry operations such as selective harvesting (Rolland et al., 1998) can be excluded as a driver of changes in growth in our 19th century relict *P. uncinata* stand. As a consequence, two drivers may be involved in explaining divergent tree growth between HAS and LAS from 1980 onwards, namely

enrichment of atmospheric CO₂ and/or a deposition of N from air pollution by NO₂.

Increased atmospheric CO₂ concentrations have been suggested as a possible driver of changes. The CO₂ fertilization effect would result in increased water-use efficiency (WUE) mainly due to partial closure of stomata and to a rise of the temperature optimum of photosynthesis (Broadmeadow and Randle, 2002; Keenan et al., 2013). Such a CO₂-induced increase of WUE is associated with greater net primary productivity and positive growth trends and at high elevation sites (e.g., LaMarche et al., 1984; Keenan et al., 2013). It is more evident at sites with greater stress and more severely limited resources (Martínez-Vilalta et al., 2008), as is the case at HAS and other *Pinus* stands in the Alps and the Mediterranean (Leal et al., 2008; Martínez-Vilalta et al., 2008). However, recent studies also point to the fact that temperature-induced drought may limit C uptake (Barber et al., 2000). Lower C uptake during dry summers may cancel out increased uptake in spring (Angert et al., 2005). At LAS, we can thus hypothesize that an improvement of WUE (if any) cannot compensate for the increase in summer temperatures as shown for other species such as *P. sylvestris* (Martínez-Vilalta et al., 2008) and *F. sylvatica* (Jump et al., 2006) at their southern distribution ranges.

As a result of the proximity of our study sites to a major city and a main highway, possible effects of N deposition and its interactions with increased CO₂ concentrations should be considered as well when addressing drivers of growth in *P. uncinata*, as increased N uptake from soil could allow N-deficient forests to respond to elevated CO₂. Tree growth might be particularly sensitive to such ecological effects because (i) alpine habitats are known to be strongly N-limited (Bowman, 1992) and as (ii) the benefit from N fertilization has been shown to be higher in nutrient-limited soils (Hattenschwiler et al., 1996; Huang et al., 2007). Both sites used in this study have mainly poor, shallow soils and N deposition has possibly contributed, along with the CO₂ increase, to boost tree growth at HAS, but fails to explain the divergence observed between the stands since the 1980s.

CONCLUSION

This study is among the first to detail the dendroclimatic response of a low altitude *P. uncinata* relict forest stand located at the lower bounds of its altitudinal limits (Pawłowsky, 1970). Our results demonstrate that tree growth has been driven increasingly by summer droughts, in particular and much more severely during the last ~30 year. Our findings provide evidence of divergence in growth between relict low (LAS) and the high-altitude (HAS) *P. uncinata* stands since the 1980s. At HAS, response function suggest that increasing autumn temperatures have favored radial growth presumably as a result of longer growing season (Penuelas et al., 2009). Response functions also suggest that the pronounced and continuous growth decline at LAS seems to be driven by higher summer temperatures and increased soil water stress in summer. The relict stand at LAS can be seen as a model case for the evolution of forests in the subalpine belt as well as for dynamics of trees located at the lower bounds of their ecological range.

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- Conflict of Interest Statement:** The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Received: 16 May 2014; accepted: 28 November 2014; published online: 20 January 2015.

Citation: Corona C, Lopez-Saez J, Stoffel M, Rovéra G, Edouard J-L and Guibal F (2015) Impacts of more frequent droughts on a relict low-altitude *Pinus uncinata* stand in the French Alps. *Front. Ecol. Evol.* 2:82. doi: 10.3389/fevo.2014.00082

This article was submitted to *Paleoecology*, a section of the journal *Frontiers in Ecology and Evolution*.

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