



Identifying Drivers of Seasonality in Lena River Biogeochemistry and Dissolved Organic Matter Fluxes

Bennet Juhls^{1*}, Colin A. Stedmon², Anne Morgenstern³, Hanno Meyer³, Jens Hölemann⁴, Birgit Heim³, Vasily Povazhnyi⁵ and Pier P. Overduin³

¹ Department of Earth Sciences, Institute for Space Sciences, Freie Universität Berlin, Berlin, Germany, ² National Institute of Aquatic Resources, Technical University of Denmark, Copenhagen, Denmark, ³ Helmholtz Centre for Polar and Marine Research, Alfred Wegener Institute, Potsdam, Germany, ⁴ Helmholtz Centre for Polar and Marine Research, Alfred Wegener Institute, Bremerhaven, Germany, ⁵ Otto Schmidt Laboratory for Polar and Marine Research, Arctic and Antarctic Research Institute, Saint Petersburg, Russia

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*Correspondence:

Bennet Juhls
bjuhls@www.fu-berlin.de

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Warming air temperatures, shifting hydrological regimes and accelerating permafrost thaw in the catchments of the Arctic rivers is affecting their biogeochemistry. Arctic river monitoring is necessary to observe changes in the mobilization of dissolved organic matter (DOM) from permafrost. The Lena River is the second largest Arctic river and 71% of its catchment is continuous permafrost. Biogeochemical parameters, including temperature, electrical conductivity (EC), stable water isotopes, dissolved organic carbon (DOC) and absorption by colored dissolved organic matter (a_{CDOM}) have been measured as part of a new high-frequency sampling program in the central Lena River Delta. The results show strong seasonal variations of all biogeochemical parameters that generally follow seasonal patterns of the hydrograph. Optical indices of DOM indicate a trend of decreasing aromaticity and molecular weight from spring to winter. High-frequency sampling improved our estimated annual fluvial flux of annual dissolved organic carbon flux (6.79 Tg C). EC and stable isotope data were used to distinguish three different source water types which explain most of the seasonal variation in the biogeochemistry of the Lena River. These water types match signatures of (1) melt water, (2) rain water, and (3) subsurface water. Melt water and rain water accounted for 84% of the discharge flux and 86% of the DOC flux. The optical properties of melt water DOM were characteristic of fresh organic matter. In contrast, the optical properties of DOM in subsurface water revealed lower aromaticity and lower molecular weights, which indicate a shift toward an older organic matter source mobilized from deeper soil horizons or permafrost deposits. The first year of this new sampling program sets a new baseline for flux calculations of dissolved matter and has enabled the identification and characterization of water types that drive the seasonality of the Lena River water properties.

Keywords: Lena River, Arctic, DOC, CDOM, optical indices, stable water isotopes

INTRODUCTION

The current warming of the Siberian Arctic (Richter-Menge et al., 2019) is causing intense changes in atmospheric forcing, precipitation, subsurface water storage and runoff from rivers to the Arctic Ocean (Yang et al., 2002; Velicogna et al., 2012; Bintanja and Selten, 2014; Niederdrenk et al., 2016). A number of studies report an increase of runoff from the Eurasian Arctic Rivers (Peterson, 2002; McClelland et al., 2006; Shiklomanov and Lammers, 2009). Of all Arctic Rivers, the Lena River is the second largest in annual discharge. Its annual discharge increased by 15.6% since 1936 (Ye et al., 2003; Shiklomanov, 2010), which may primarily be due to a winter discharge increase by 93% (Tananaev et al., 2016). Furthermore, increasing Lena River temperature (Yang et al., 2002; Liu et al., 2005) leads to more intense thermal erosion along the river shores and the coast of the Laptev Sea (Aré, 1988; Bareiss and Görden, 2005), warms the surface waters of the central Laptev Sea, and accelerates thawing of newly formed subsea permafrost (Angelopoulos et al., 2019). A change of the river discharge, as well as changes of temperature and biogeochemical properties of river water, can strongly affect the physical state of the Arctic Ocean and marine ecosystems, especially in coastal waters. Increased export of dark, organic-rich river water leads to stronger absorption of sunlight and an increased heat flux, which can, in turn, contribute to sea ice decline (Pegau, 2002; Hill, 2008).

The carbon reservoirs of Arctic river watersheds are currently in transition. Thawing permafrost is releasing previously stored carbon (Schuur et al., 2015; Plaza et al., 2019; Turetsky et al., 2019) and increasing seasonal growth of vegetation (Schuur et al., 2009; Keuper et al., 2012). Both of these processes will increase the particulate and dissolved organic matter flux into rivers and the fluvial flux to coastal waters (Frey and Smith, 2005). Hydrological flow paths will shift toward increased groundwater flow and affect the source and composition of dissolved organic matter (DOM) (Amon et al., 2012). While particulate material is mostly deposited in shelf sea sediments to become mineralised or buried (Charkin et al., 2011; Wegner et al., 2013), dissolved organic matter is exported offshore into the open Arctic Ocean (i.e., Juhls et al., 2019) and approximately 50% is thought to be exported onwards to the Atlantic (Granskog et al., 2012). Most of the DOM transported by the Lena River originates from leaching of surface soils of dominantly boreal forest vegetation (Amon et al., 2012; Kaiser et al., 2017). However, the effect of ongoing permafrost thaw on hydrological pathways will affect DOM fluxes (Freeman et al., 2001; Frey and Smith, 2005). In addition to the amount of DOM, its composition, source and age in the Lena River are expected to change (Amon et al., 2012; Walker et al., 2013; Mann et al., 2016; Wild et al., 2019).

Initiation of the pan-Arctic River sampling programs PARTNERS (2003–2007), Student Partners (2004–2009) and ArcticGRO (since 2009), have provided invaluable insights into the quantitative and qualitative properties of dissolved and particulate matter exported into the Arctic Ocean (Raymond et al., 2007; Holmes et al., 2008; Stedmon et al., 2011; Amon et al., 2012; Mann et al., 2016). Results of PARTNERS and ArcticGRO

(Raymond et al., 2007) show that the Lena River contributes 20–29% of the total circumpolar fluvial dissolved organic carbon (DOC) export to the Arctic Ocean. However, reported estimates of annual Lena River DOC fluxes span a wide range between 3.6 Tg C (Gordeev and Kravchishina, 2009) and 7.67 Tg C (Kicklighter et al., 2013). This is likely because the calculation of annual flux is based on a few water samples per year, which are usually taken during open water season (Raymond et al., 2007; Stedmon et al., 2011; Holmes et al., 2012). Such estimates are susceptible to systematic biases: limited sampling can miss seasonal peaks of DOC concentration and lead to subsequent underestimation of the DOC flux (Jollymore et al., 2012). Furthermore, samples from the Lena River are mostly taken in Zhigansk, which is located ~800 km upstream of the river mouth (Figure 1). Little is known about DOM transformation, mineralization and release to the atmosphere on the way from Zhigansk to the river mouth (Amon et al., 2012). In order to identify a trend in DOC flux, it is crucial to reduce uncertainty in the calculated annual flux to below the interannual variability. While most research attention has focused on potential long-term trends of river DOC export (Kicklighter et al., 2013; Tank et al., 2016), seasonal variations of ice break-up and freeze-up timing and the associated discharge, material load and biogeochemistry are also affected by the changing climate and require research attention. Earlier spring ice break-up and later freeze-up in fall result in longer open water and shorter winter flow periods and will have an impact on the annual organic matter flux. Monitoring river water biogeochemistry may provide insights into the progress of this change, since the river water chemistry, measured near its mouth, integrates the changes occurring in the catchment as a whole. DOM fluxes and composition can be expected to change and this will also ultimately influence the fate of terrestrial carbon in the Arctic, including that carbon stored in shelf sediments, mineralised in the Arctic, and exported to the north Atlantic Ocean.

To accurately capture short-term variability and understand how changing climate influences in-river processes that affect organic matter quality and its function in the fluvial ecosystem, it is necessary to monitor water constituents with a high temporal resolution and throughout the whole season (including both open water and ice-covered periods). To characterize fluxes from land to sea, this should be done as close as possible to the river mouth. Such sampling, carried out over multiple years, can enable prediction of future responses of river flux to projected change. The Research Station Samoylov Island provides an opportunity to meet these criteria in the central Lena River Delta as it serves as the basis for regular, frequent and year-round sampling for major biogeochemical parameters. The goal of this study is to better understand and decipher seasonal variations of hydrochemical characteristics and organic matter and its optical properties of the Lena River. We aim to identify how changing water sources drive seasonal changes in fluvial biogeochemistry by putting such a sampling program into place over 1 year at the Research Station Samoylov Island. With this, we provide a basis for future trend analyses and remote sensing studies that may be used to upscale observations from the hydrographic point scale.

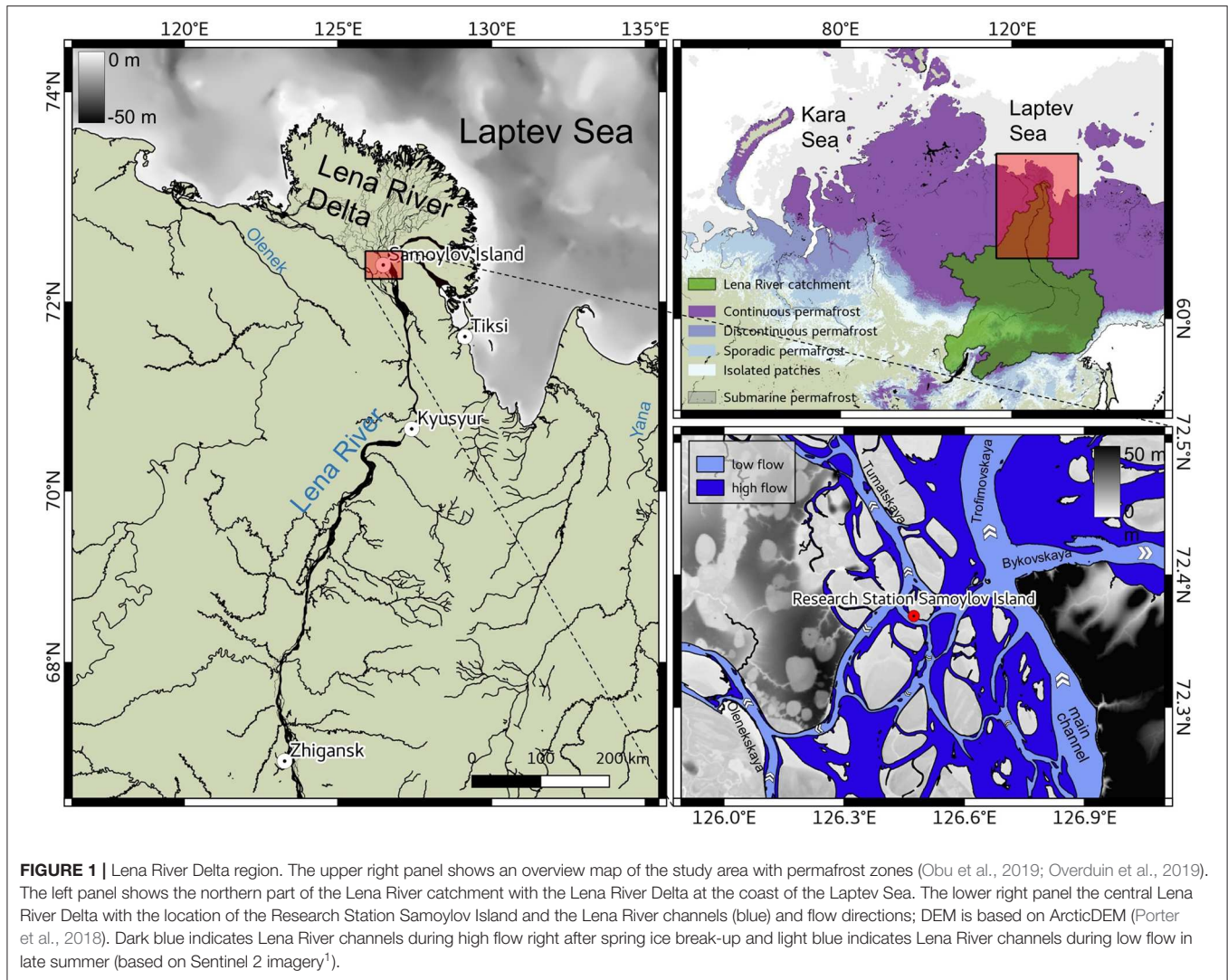


FIGURE 1 | Lena River Delta region. The upper right panel shows an overview map of the study area with permafrost zones (Obu et al., 2019; Overduin et al., 2019). The left panel shows the northern part of the Lena River catchment with the Lena River Delta at the coast of the Laptev Sea. The lower right panel shows the central Lena River Delta with the location of the Research Station Samoylov Island and the Lena River channels (blue) and flow directions; DEM is based on ArcticDEM (Porter et al., 2018). Dark blue indicates Lena River channels during high flow right after spring ice break-up and light blue indicates Lena River channels during low flow in late summer (based on Sentinel 2 imagery¹).

MATERIALS AND METHODS

Study Area and Sample Collection

The Lena watershed area ($2.61 \times 10^6 \text{ km}^2$) extends ~2,400 km from north to south and is underlain by all types of permafrost: continuous (70.5%), discontinuous (10.6%), sporadic (6.0%), and isolated (7.3%). 5.6% of the watershed, on the northern side of the Lake Baikal, is free of permafrost [numbers calculated using permafrost zones by Obu et al. (2019)]. The catchment is dominantly covered by forest (72.1%) and shrubland (12.5%) (Amon et al., 2012).

We collected water samples from the river surface in the center of the Olenekskaya Channel near Samoylov Island using a pre-rinsed HDPE 1 L bottle (in summer) or a UWITEC 1 L water sampler (under ice) for 1 year beginning in April 2018 (Figure 1). The island is located in the central Lena River Delta and hosts a permanently staffed research station (Research Station Samoylov Island), which has been operating since 2013. Additional

sample sets with higher temporal resolution, from different years and from Lena River ice, are listed in Table 1. During the open water period (28 May 2018 to 22 October 2018), water was sampled from a small boat, and during ice-covered period (until 27 May 2018 and from 23 October 2018), through a hole drilled through the river ice. Two samples during the river ice break-up as well as four samples during the ice freeze-up in October were taken from the shore due to the inaccessibility of a more centered location on the river channel. Water samples were subsampled, filtered and conserved at the research station directly after sampling.

Sampling started on 20 April 2018 at a frequency of ~4 days and is ongoing. For this study, we use a dataset of almost one complete year (until 6 April 2019). For each sample, *in situ* temperature and electrical conductivity (EC) were measured using a hand-held conductivity meter (WTW COND 340I, accuracy $\pm 0.5\%$). For all samples, a series of biogeochemical parameters were analyzed (Table 1). The EC was additionally re-measured on each sample in the lab using a hand-held

¹Copernicus Open Access, Hub. Available online at: <https://scihub.copernicus.eu/>.

TABLE 1 | Sampling period, frequency, sample type and measured parameters of datasets used in this study.

Dataset	Location	Period	Type	Frequency	No. of samples	Parameters
This study	Samoylov Island	20 April 2018–06 April 2019	Water samples	~ every 4 days	75	DOC, a_{CDOM} , stable H ₂ O isotopes, cations & anions, EC
This study	Samoylov Island	4 July 2018–15 July 2018	Water samples	2 times per day	23	DOC and a_{CDOM} (using different filter types and filter pore sizes), stable H ₂ O isotopes, EC
This study	Samoylov Island	4 May 2018	Ice core samples	1 time	1 ice core with 57 subsamples	DOC, a_{CDOM} , stable H ₂ O isotopes, cations & anions, EC
Eulenburg et al. (2019)	Samoylov Island	22 May 2014–19 June 2014	Water samples	1 to 2 times per day	43	DOC, a_{CDOM} , stable H ₂ O isotopes, EC
ArcticGRO & Partners; Holmes et al. (2018a,b)	Zhigansk	2002–2019	Water samples	Up to 6 per year	64	DOC, a_{CDOM} , stable H ₂ O isotopes

conductivity meter (WTW Multilab 540, accuracy $\pm 0.5\%$) after transport to Germany. For 11 days, from 4 July 2018 to 15 July 2018, samples were taken at a higher frequency and for additional parameters (Table 1).

Discharge

The discharge of the Lena River is monitored by the Russian Federal Service for Hydrometeorology and Environmental Monitoring (Roshydromet) and data are available at www.r-arcticnet.sr.unh.edu (Shiklomanov, 2010; 1936–2009) and www.arcticgreatrivers.org (Shiklomanov et al., 2018, 1999 to present). All discharge data shown in this study were corrected for the distance difference between the gauge station at Kyusyur and water sampling station at Samoylov Island (~ 220 km) using a Lena River propagation speed estimate (88 km d^{-1}) from Smith and Pavelsky (2008).

Biogeochemical Parameters

For each sample of the Lena River monitoring program, 1 L of surface river water was collected and subsequently subsampled. For DOC and CDOM absorption (a_{CDOM}), the sample was filtered through a $0.45 \mu\text{m}$ cellulose acetate filter which had been rinsed with 20 mL sample water. Over 10 days in July 2018, additional filter type and pore size tests for DOC and a_{CDOM} were carried out (Supplementary Figures 1, 2). DOC samples were filled into a pre-rinsed 20 mL glass vial and acidified with 25 μL HCl Suprapur (10 M) and stored in the dark at 4°C . After transport, DOC samples were analyzed at the Alfred Wegener Institute Helmholtz Center for Polar and Marine Research (AWI), Potsdam, Germany. DOC concentrations were analyzed using high temperature catalytic oxidation (TOC-VCPH, Shimadzu). Three replicate measurements of each sample were averaged. After every ten samples, a blank (Milli-Q water) and a standard was measured. Eight different commercially available certified standards covered a range between 0.49 mg L^{-1} (DWNSVW-15) and 100 mg L^{-1} (Std. US-QC). The results of standards provided an accuracy better than $\pm 5\%$.

a_{CDOM} samples were collected in pre-rinsed 50 mL amber glass bottles that were stored in the dark at 4°C until analysis. a_{CDOM} was measured at the Otto Schmidt Laboratory for Polar and Marine Research (OSL), Saint Petersburg, Russia using a SPECORD 200 spectrophotometer (Analytik Jena) and at the German Research Center for Geosciences (GFZ), Potsdam, Germany using a LAMBDA 950 UV/Vis (PerkinElmer). The median absorbance (A_λ) of three replicate measurements was used to calculate the $a_{CDOM}(\lambda)$:

$$a_{CDOM}(\lambda) = \frac{2.303 * A_\lambda}{l},$$

where l is the path length (length of cuvette). Fresh Milli-Q water was used as reference. To detect chemical composition and molecular structure of the DOM, two optical indices, SUVA and S_{275–295}, were used. Both indices correlate with aromaticity and molecular weight of bulk DOC (Weishaar et al., 2003; Helms et al., 2008). SUVA ($\text{m}^2 \text{ g C}^{-1}$) was calculated by dividing the decadal absorption A/l (m^{-1}) at 254 nm by DOC concentration (mg L^{-1}). The spectral slope of $a_{CDOM}(\lambda)$ between 275 and 295 nm (S_{275–295}) is an index for photodegradation (Helms et al., 2008). S_{275–295} was calculated by fitting a regression for the wavelength ranges 275–295 nm to the exponential function (Helms et al., 2008):

$$a_{CDOM}(\lambda) = a_{CDOM}(\lambda_0) * e^{-S(\lambda - \lambda_0)},$$

where $a_{CDOM}(\lambda_0)$ is the absorption coefficient at reference wavelength λ_0 and S is the spectral slope of $a_{CDOM}(\lambda)$ for the chosen wavelength range.

Water samples for stable isotopes were collected untreated in 10 mL HDPE vials, sealed tightly, stored in the dark at 4°C . Measurements were conducted at the laboratory facility for stable isotopes at AWI Potsdam using a Finnigan MAT Delta-S mass spectrometer equipped with equilibration units for the online determination of hydrogen and oxygen isotopic composition. The data is given as δD and $\delta^{18}\text{O}$ values, which is the per

mille difference to standard V-SMOW. The deuterium excess (d-excess) is calculated by:

$$d\text{-excess} = \delta D - 8 \cdot \delta^{18}O.$$

The measurement accuracy for hydrogen and oxygen isotopes was better than $\pm 0.8\text{‰}$ and $\pm 0.1\text{‰}$, respectively (Meyer et al., 2000).

Water samples for concentrations of major dissolved components were filtered using a syringe-mounted $0.45 \mu\text{m}$ CA filter and kept cool and dark until analysis. Concentrations of major anions (SO_4^{2-} , Cl^- , Br^- , F^- , NO_3^- , and PO_4^{3-}) were determined using ion chromatography (Thermo-Fischer ICS 2100; Weiss, 2001). Total dissolved elemental concentrations (for Al, Ba, Ca, Fe, K, Mg, Mn, Na, P, Si, and Sr) were measured with inductively coupled plasma optical emission spectroscopy (ICP-OES; Perkin Elmer Optima 8300DV; Boss and Fredeen, 1997).

In addition, one ice core (LE08) was drilled on the river channel at Samoylov Island on 4 May 2018 (Figure 1). The core

was drilled using a Kovacs Mark III and wrapped in HDPE plastic sleeves and stored frozen for transport to AWI Potsdam where further subsampling and analysis were done. The core was retrieved in five pieces with a total length of 144 cm, which was close to the ice thickness measured in the borehole. The ice core was cut into 3 cm slices and analyzed for DOC, $a_{\text{CDOM}}(\lambda)$, stable oxygen isotopes, cations and anions following the methods described above.

Flux Calculations

For visualization, concentrations (C_d) between sampling dates were linearly interpolated. Then, daily fluxes were calculated by multiplying daily concentrations and the daily discharge (Q_d).

$$\text{Flux}_d = C_d \text{ (mg L}^{-1}\text{)} \times Q_d \text{ (m}^3 \text{s}^{-1}\text{)} \times 86400 \text{ s,}$$

Daily fluxes (Flux_d) were summed for individual periods. The annual heat flux was calculated as in Liu et al. (2005), Yang et al. (2013) to enable inter-annual comparisons. Negative river water temperatures were set to 0°C and the specific heat capacity of water was set constant to $4.184 \text{ J g}^{-1} \text{ }^\circ\text{C}^{-1}$.

TABLE 2 | Electrical conductivity (EC) and $\delta^{18}\text{O}$ for water source endmember.

Endmember	EC ($\mu\text{S cm}^{-1}$)	$\delta^{18}\text{O}$ (‰)
EM1	80	-25
EM2	100	-16
EM3	600	-22

Water Source Endmember Calculation

98% of the variability of all measured parameter can be explained by three components (Supplementary Table 1). We distinguished three water source fractions using EC and water

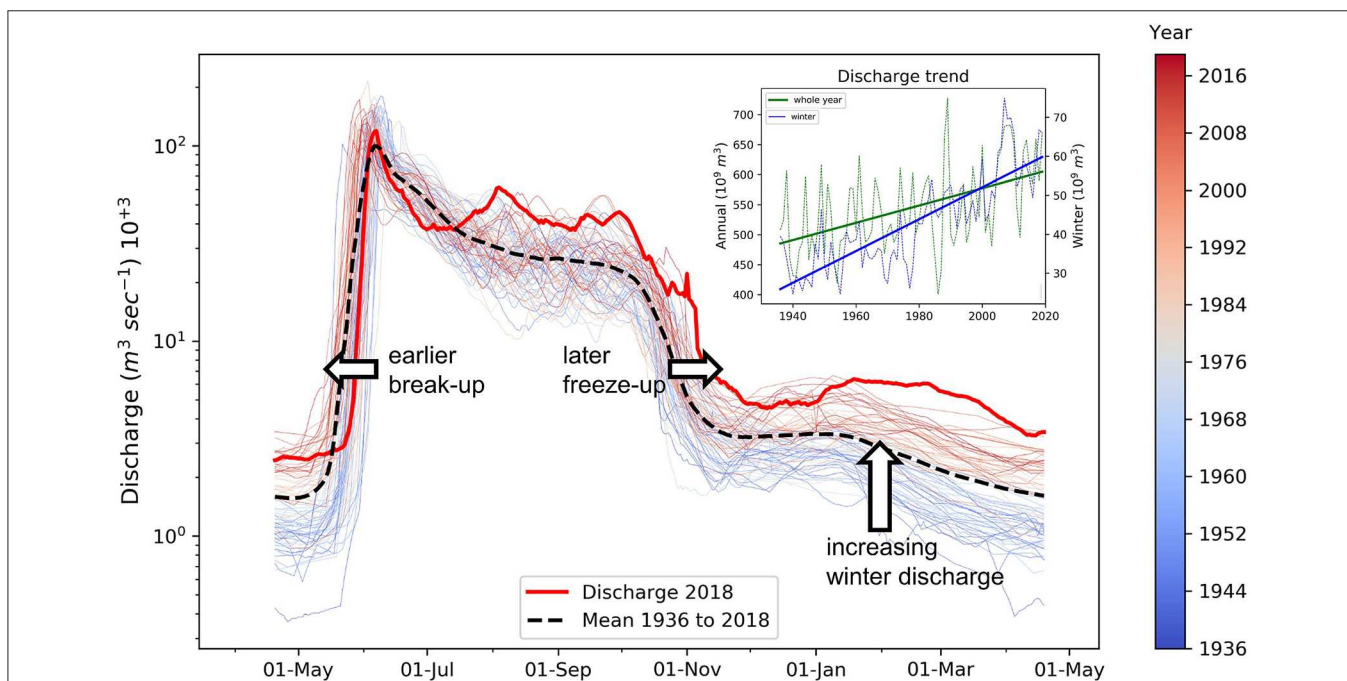


FIGURE 2 | Discharge data from Kyusyur station from 1936 to 2018. Colors from blue to red indicate the year from 1936 to 2018. The red thick solid line shows the year 2018 and the black thick dashed shows the mean of the period from 1936 to 2018. The inset figure shows the discharge flux for the whole year (green) and for November to April (blue). Discharge data: A.I. Shiklomanov (2010) 1939–2008, ArcticGRO 2009–2019.

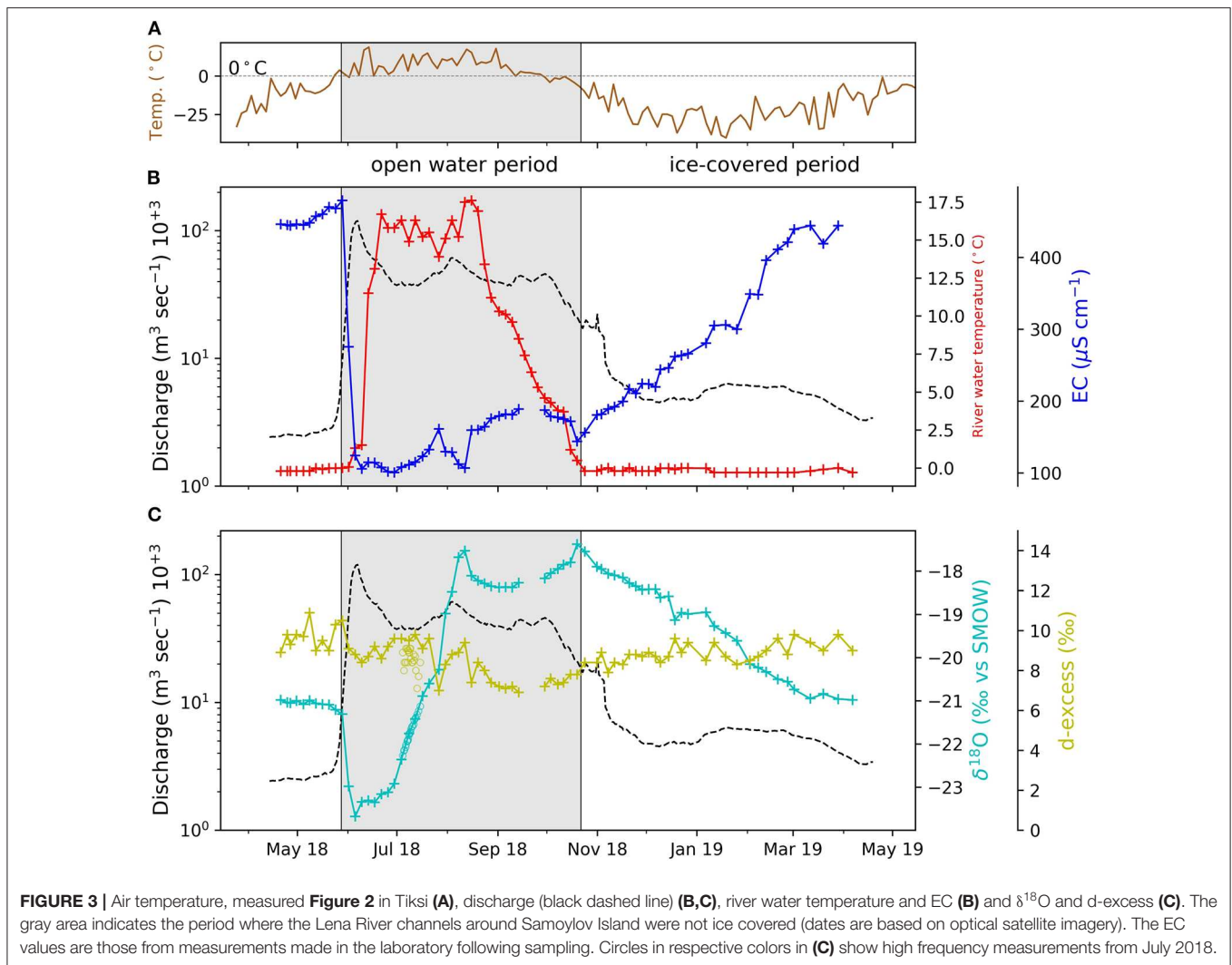


FIGURE 3 | Air temperature, measured **Figure 2** in Tiksi **(A)**, discharge (black dashed line) **(B,C)**, river water temperature and EC **(B)** and $\delta^{18}\text{O}$ and d-excess **(C)**. The gray area indicates the period where the Lena River channels around Samoylov Island were not ice covered (dates are based on optical satellite imagery). The EC values are those from measurements made in the laboratory following sampling. Circles in respective colors in **(C)** show high frequency measurements from July 2018.

$\delta^{18}\text{O}$ by solving the mass balance equation:

$$f_{EM1} + f_{EM2} + f_{EM3} = 1$$

$$f_{EM1}EC_{EM1} + f_{EM2}EC_{EM2} + f_{EM3}EC_{EM3} = EC$$

$$f_{EM1}\delta^{18}\text{O}_{EM1} + f_{EM2}\delta^{18}\text{O}_{EM2} + f_{EM3}\delta^{18}\text{O}_{EM3} = \delta^{18}\text{O}$$

for unknowns, where $f \in [0; 1]$ is the fraction of each water source. We chose endmember for EC and $\delta^{18}\text{O}$ values that encompass the overall variability of observed EC and $\delta^{18}\text{O}$ (**Table 2**).

RESULTS

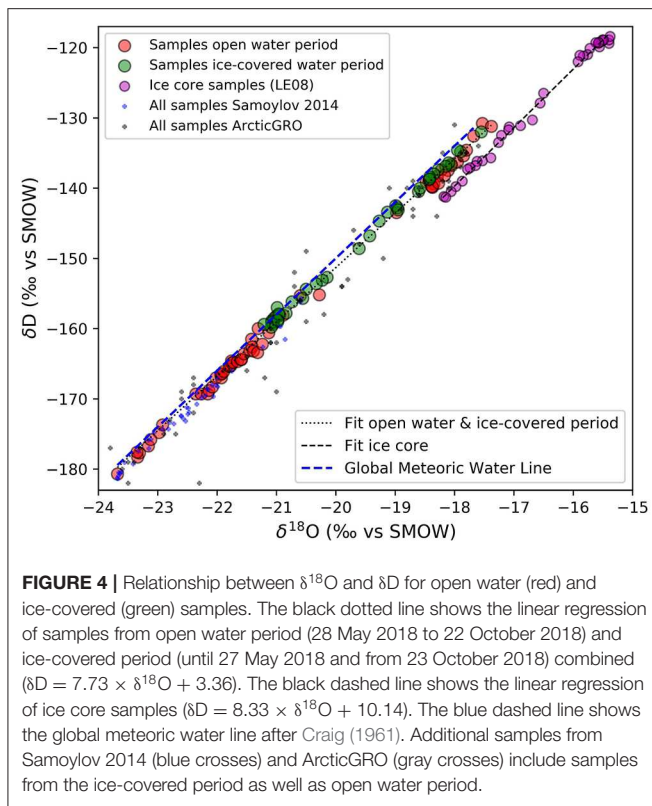
The Lena River discharge shows strong seasonality, with low discharge during the ice-covered winter, a spring peak maximum and a moderate summer discharge (**Figure 2**). In some years, additional discharge peaks in late summer and fall occur. Over the period from 1936 to 2018, the Lena River’s annual discharge increased ($1.44 \times 10^9 \text{ m}^3 \text{ year}^{-1}$), to a large part driven by the increasing winter (November to April) discharge flux (0.41×10^9

$\text{m}^3 \text{ year}^{-1}$; **Figure 2** inset). Furthermore, trends are clearly visible in the colors of **Figure 2**, indicating that river ice breaks up earlier and freezes later in the year.

Putting the year 2018 into the long-term context, the winter (November to April) and late summer (August to October) discharge were significantly higher than their 1938–2018 means. The timing and intensity of the 2018 spring discharge peak was close to the long-term median. Summer 2018 was also characterized by multiple higher discharge events, which rise above the long-term mean. The elevated discharge continues into the first 3 months of 2019 (dashed line in **Figures 3B,C**).

Lena River Temperature and Chemistry

During the period when the Lena River is ice-covered, the mean river water temperature below the ice is slightly below 0°C (-0.2 to -0.01°C ; **Figure 3A**). During the first days of spring peak discharge, the temperature only marginally increases to 1.5°C . About 2 weeks after the start of spring peak discharge, the temperature rises within 1 week to a relatively stable summer level around 16°C . With decreasing air temperatures in fall,



the river temperature gradually decreases and reaches 0°C in mid-October when the Lena River starts to freeze. The highest river water temperature (17.6°C) was recorded on 15. August 2018. The EC drops almost simultaneously with the onset of the spring peak discharge from its annual maximum of $490 \mu\text{S}/\text{cm}$ to its annual minimum of $99 \mu\text{S}/\text{cm}$. The EC is generally low in summer, and gradually increases during the ice-covered period. The summer discharge peaks coincide with decreases in the EC.

A strong seasonality in river water $\delta^{18}\text{O}$ values was observed, ranging between -17.4‰ and -23.7‰ in the sampling period. During late winter, when the Lena River ice thickness reached its maximum (March–April), $\delta^{18}\text{O}$ was stable at around -21‰ . The lowest $\delta^{18}\text{O}$ values (-23.7‰) occurred during the spring freshet in the first days of June, coinciding with the highest discharge. The highest $\delta^{18}\text{O}$ (-17.4‰) occurred shortly before freeze-up at the end of October. Similarly high $\delta^{18}\text{O}$ (-17.5‰) occurred during the mid-summer discharge peak. With decreasing discharge and freezing of the Lena River and its catchment, $\delta^{18}\text{O}$ gradually decreased toward -21‰ . d-excess values were highest ($+10.9\text{‰}$) during late winter and lowest ($+6.9\text{‰}$) in late summer. The d-excess of the measurements in July 2018 were about 0.5‰ lower than long-term values. Additional measurements of $\delta^{18}\text{O}$ in July 2018 agreed with the long-term measurements.

For all samples in the Lena River dataset, we observed a very strong relationship ($r^2 = 0.996$) between $\delta^{18}\text{O}$ and δD (Figure 4). The slope of the linear regression is lower (7.73) than the Global Meteoric Water Line (slope GMWL = 8.0), but almost identical

to the Local Meteoric Water Line (the slope of the LMWL is 7.6 for 381 event-based samples, and 7.7 for 41 monthly means; Spors, 2018). Lena River ice core samples (LE08) showed overall higher $\delta^{18}\text{O}$ and δD than water samples and the regression line showed a clear offset compared to river water samples (Figure 4). The slope of the linear regression of ice core samples (8.32) was higher than that of Lena River water. Lena River water samples from 2014 were similarly strongly correlated ($r^2 = 0.987$) and had a similar slope (7.6) compared to the Lena River samples from 2018. Both slopes, from 2014 and 2018, are lower, compared to the GMWL. Ice core samples had higher mean $\delta^{18}\text{O}$ and δD values, but with a shift to lower δD for similar $\delta^{18}\text{O}$.

The concentrations of most of the major cations and anions correlated with the seasonal variability in EC (Supplementary Figure 3). Whereas concentrations of all ions increased during the ice-covered period, the concentration of K behaved differently, with lower values during January and February 2019. During the ice-covered period, Na and Cl^- dominated, while in summer and during the spring discharge peak, Ca and Cl^- and SO_4^{2-} dominated. Whereas concentrations for some species regained over 50% of their late winter concentration during summer (F^- , Si), others did not increase until mid-winter (Na, Ba, Cl^- , Br^-). The end of the open water period was marked by a small decrease in concentration that lasted for almost 2 weeks for all species. Si concentration was hardly affected by precipitation events during the open water period, in contrast to all other measured concentration (for example, the precipitation peak centered at August 20, 2018). Annual fluxes of major ions are presented in Table 3. Note that some concentrations fell below analytical detection limits and this prevented a meaningful flux calculation.

Dissolved Organic Carbon (DOC) and Colored Dissolved Organic Matter (CDOM)

Dissolved organic matter in the Lena River varied with the hydrograph, with DOC concentrations ranging from 4.9 to 18.2 mg L^{-1} and $a_{\text{CDOM}}(254)$ ranging from 40.8 to 150.92 m^{-1} (Figure 5A). Lowest values were observed at the end of winter, right before or with the onset of the spring ice break-up, whereas highest values occurred only a few days later during the spring ice break-up. In summer (July to October) both DOC and $a_{\text{CDOM}}(254)$ returned to values near pre-break-up [$\sim 6\text{--}7 \text{ mg L}^{-1}$ DOC and ~ 44 to $50 \text{ m}^{-1} a_{\text{CDOM}}(254)$] but then increased during the mid-summer discharge peak from 6.2 to 10.2 mg L^{-1} and from 45.3 to $85.6 \text{ m}^{-1} a_{\text{CDOM}}(254)$. High frequency sampling during July 2018 agreed closely with the overall pattern from the annual sampling (circles in Figure 5).

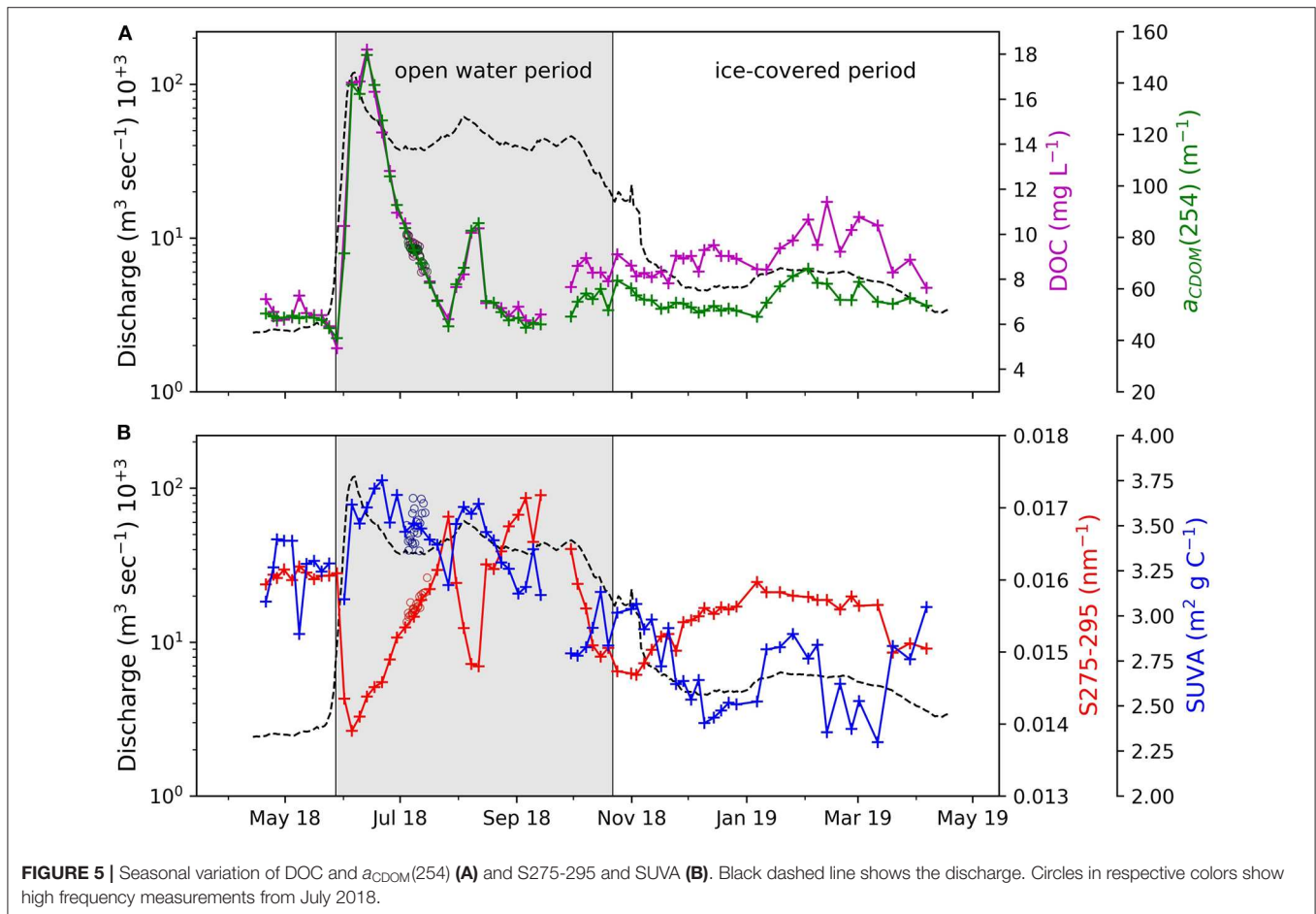
In winter, DOC and $a_{\text{CDOM}}(254)$ increased in response to comparatively minor fluctuations in discharge, with winter maxima of 11.4 mg L^{-1} and 67.8 m^{-1} , respectively. DOC and $a_{\text{CDOM}}(254)$ in the river ice had means of 1.04 mg L^{-1} DOC and $0.96 \text{ m}^{-1} a_{\text{CDOM}}(254)$, with no clear systematic vertical trend down core (Supplementary Figure 4).

Two periods with clear correlations between DOC and $a_{\text{CDOM}}(254)$ were apparent: one for the open water period (28 May 2018 to 22 Oct. 2018), and another for the rest of the

TABLE 3 | Annual mean values and annual fluxes of Lena River biogeochemical parameter.

	Discharge	Temperature	EC	Heat flux	DOC	$a_{CDOM}(254)$
Annual mean	21173.8 m ³ s ⁻¹	4.74°C	250.6 μS cm ⁻¹	6.95 × 10 ¹⁶ J d ⁻¹	8.85 mg L ⁻¹	62.8 m ⁻¹
Annual flux	690.2 × 10 ⁹ m ³	-	-	25.3 EJ	6.79 Tg	52.1 × 10 ¹² m ²

	Ca	K	Mg	Na	Si	Sr	Fl	Cl	SO ₄
Annual mean	22.5 mg L ⁻¹	0.83 mg L ⁻¹	6.69 mg L ⁻¹	18.0 mg L ⁻¹	3.02 mg L ⁻¹	0.19 mg L ⁻¹	0.12 mg L ⁻¹	26.5 mg L ⁻¹	19.2 mg L ⁻¹
Annual flux	11.8 Tg	0.45 Tg	3.29 Tg	6.95 Tg	1.59 Tg	0.09 Tg	0.07 Tg	10.2 Tg	8.76 Tg



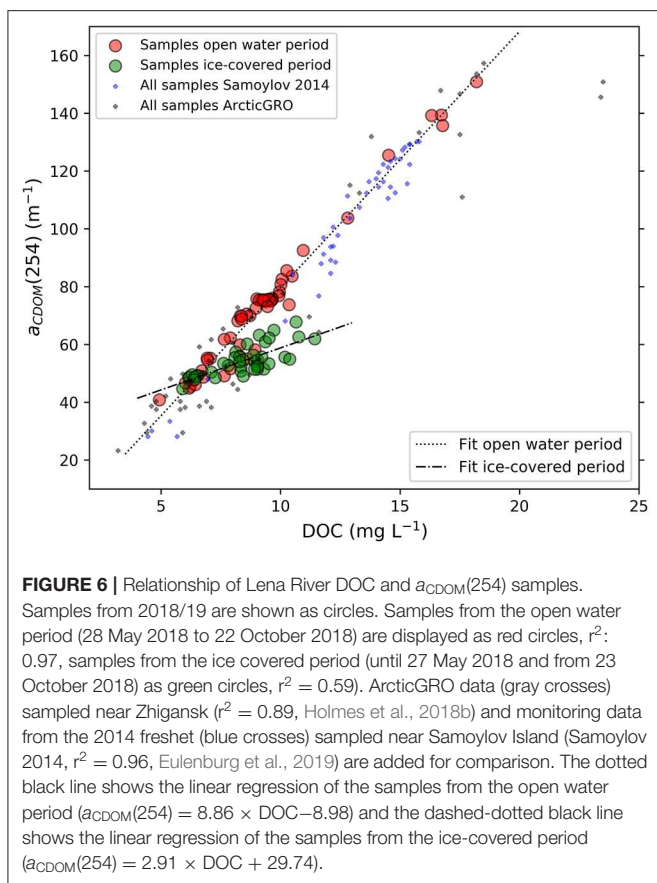
time series, with ice cover. This separation indicates qualitative differences in the river DOM (Figure 6). The slope for the $a_{CDOM}(254)$ vs. DOC regression was significantly steeper for the open water period (8.86 m² g⁻¹) while the ice covered period had a low slope (2.91 m² g⁻¹) and greater intercept. These seasonal qualitative changes were also reflected in S275-295 and SUVA. The lowest slopes (0.0139 nm⁻¹) were measured during the spring ice break-up and the highest during late summer (0.172 nm⁻¹). In winter, during the ice-covered period, S275-295 varied between 0.0147 and 0.0162 nm⁻¹. The SUVA peaked during the spring ice break-up (3.75 m² g C⁻¹) and during the early summer discharge peak (3.62 m² g C⁻¹) and reached a minimum in late winter (2.3 m² g C⁻¹). Beginning at spring ice break-up, SUVA

decreased until late winter. In April 2019, almost 2 months before the return of the spring ice break-up, SUVA increased from 2.2 to 3.1 m² g C⁻¹.

DOC Flux

We calculated an annual DOC flux of 6.79 Tg C and a discharge flux of 690.2 × 10⁹ m³ for the considered period of 1 year starting 20 April 2019, during which 2.78 Tg C (41% of annual flux) were transported in spring, 3.26 Tg C (48% of annual flux) in summer and 0.75 Tg C in winter (11% of annual flux) (Supplementary Figure 5).

Next, we compare the DOC concentration, $a_{CDOM}(254)$ and calculated fluxes of the spring period in 2014 (Figure 7A) with



the hydrologically aligned similar period in 2018 (Figure 7B). In 2014, the spring ice break-up discharge peak occurred 6 days earlier than in 2018. Thus, we compare period of similar length in 2018 (29 days) which is shifted 6 days later. In 2014, the discharge flux was 23.3% higher than in 2018, whereas the DOC flux was only 9% higher. This was due to the maximum DOC concentrations being slightly lower in 2014. Both years showed a similar pattern of decreasing DOM concentrations right before the start of the discharge peak followed by a rapid increase of DOM. Right after the discharge reached its maximum, it decreased again rapidly, whereas the DOM concentration remained high and decreased slowly.

DISCUSSION

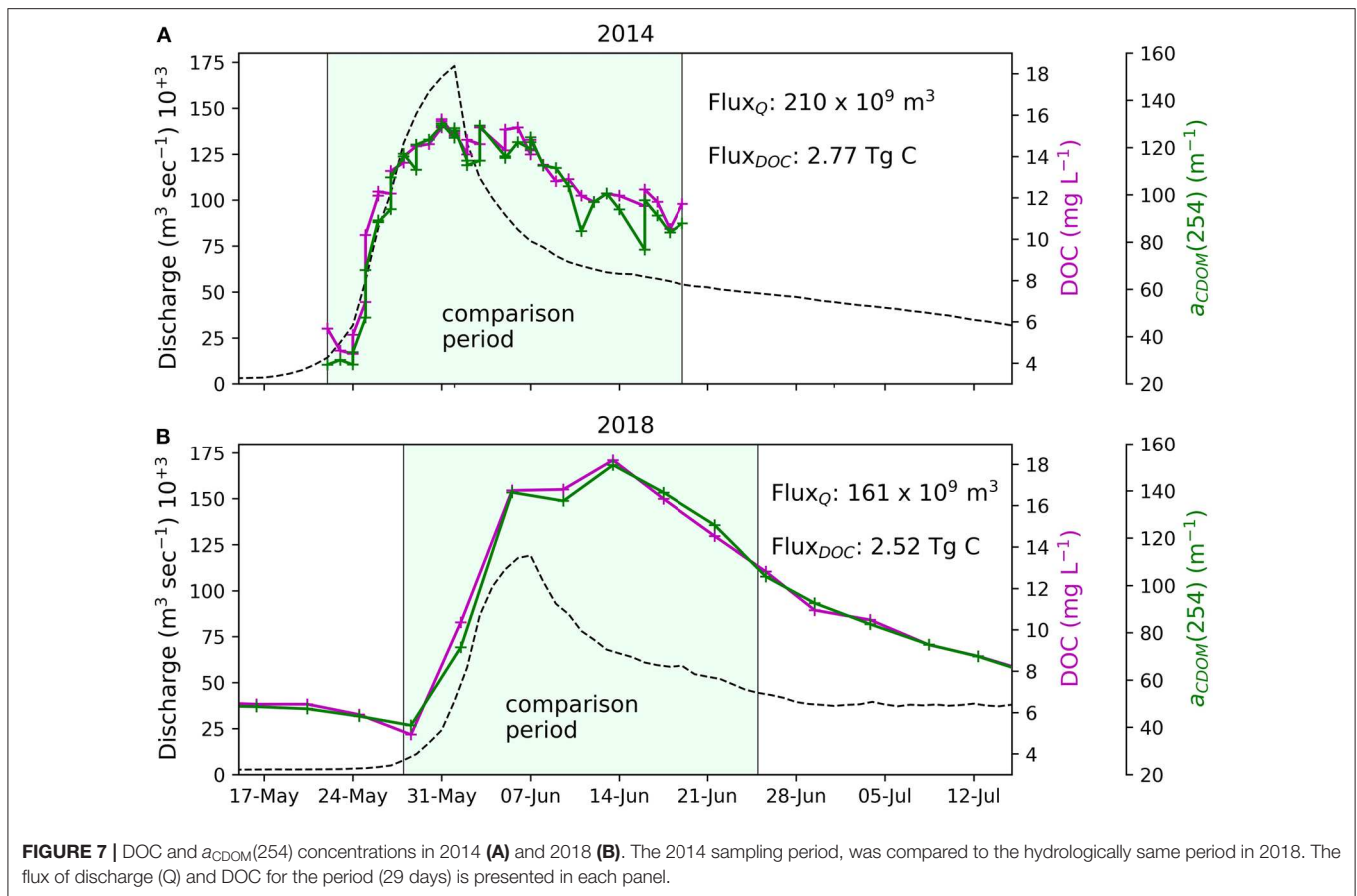
Drivers of Seasonal Variability in Hydrochemistry and DOM

The Lena River is characterized by a strong seasonality in discharge and its water biogeochemistry and the hydrograph can generally be divided into three periods: the spring ice break-up, summer to fall, and winter. Hence, we compare and contrast these three hydrologically and biogeochemically distinct periods and speculate about the shifts in the provenance of the water and its dissolved load.

(1) **Spring Ice Break-up** is the period when air temperatures are consistently positive in most of the catchment and major parts of the winter-accumulated snow and ice melt. Snowmelt does not occur simultaneously across the whole catchment, but begins in the south and moves northward. About 4 days after the ice break-up of the Lena River, the annual discharge maximum ($\sim 120\,000\text{ m}^3\text{ s}^{-1}$) is reached. River water temperatures during the first days of the spring discharge peak remain around 0°C , resulting in an export of large volumes of cold water to the Laptev Sea. Simultaneously with the discharge maximum, the water source endmember EM1 peaks (80% on 5 June) and remains the dominant water source until mid-July (Figure 8). We identified this endmember as melt water due to the sudden drop of EC (Figure 3B), major ions (Supplementary Figure 3) as well as $\delta^{18}\text{O}$ values (Figure 3C). Chosen endmember values for EM1 with low EC ($80\ \mu\text{S cm}^{-1}$) and low $\delta^{18}\text{O}$ (-25‰) are characteristic for snow melt water and agree well with reported values (Sugimoto and Maximov, 2012; Spors, 2018; Bonne et al., 2020). With regards to the DOM results it is clear that this pulse has two portions associated with it. The initial input of snow meltwater has low DOM and likely reflects snow not in contact with soil. This results in an initial decrease in DOM in the river which lasts on the order of a week, captured by the high temporal resolution of sampling. This interpretation is supported by the fact that the qualitative measures of DOM (SUVA and S275-295) do not change, essentially reflecting dilution of river water. This is then followed by the discharge maximum, where DOC reaches highest annual concentrations (18.2 mg L^{-1}). In only 2 months (between 2. June 2018 and 2 August 2018), 53.2% (3.62 Tg C) of the total annual DOC flux (6.79 Tg C) is exported to the Laptev Sea. By applying water source fractions, we estimate that 43.3% (2.9 Tg C) of the annual DOC flux is transported with melt water into the river (Figure 9). Moreover, melt water accounts for 35% ($242 \times 10^9\text{ m}^3$) of the total annual discharge flux ($690.18 \times 10^9\text{ m}^3$). This volume agrees well with the accumulated snow volume in the Lena River catchment. Using a mean (1988 to 2000) winter peak value of 120 mm as the snow water equivalent in the catchment (Yang et al., 2007), we derive a snow to water equivalent volume of $295.2 \times 10^9\text{ m}^3$.

With the discharge peak and high DOM concentrations, the quality of the organic matter also changes. Optical indices of DOM (SUVA and S275-295) were comparable to reported values (Walker et al., 2013). The changes occurring during the freshet indicate an organic matter source with high molecular weight and aromaticity (Figure 5B) (Weishaar et al., 2003; Helms et al., 2008) which is likely young DOM (Amon et al., 2012). During this period, the frozen ground beneath the snowpack during melt limits infiltration, confining most flow to a thin desiccated organic layer and to overland or snowpack flow. The DOM thus likely originates from modern plant litter, which is washed into the Lena River by the rapid and extensive snowmelt (Amon et al., 2012). Very low S275-295 during the spring ice break-up also suggests the input of fresh DOM.

(2) **Open Water Period in Summer and Fall** is the period during which Lena River water temperatures reach the annual maximum. In this period, the discharge decreases to about half of the spring peak values ($40\,000$ to $60\,000\text{ m}^3\text{ s}^{-1}$). The

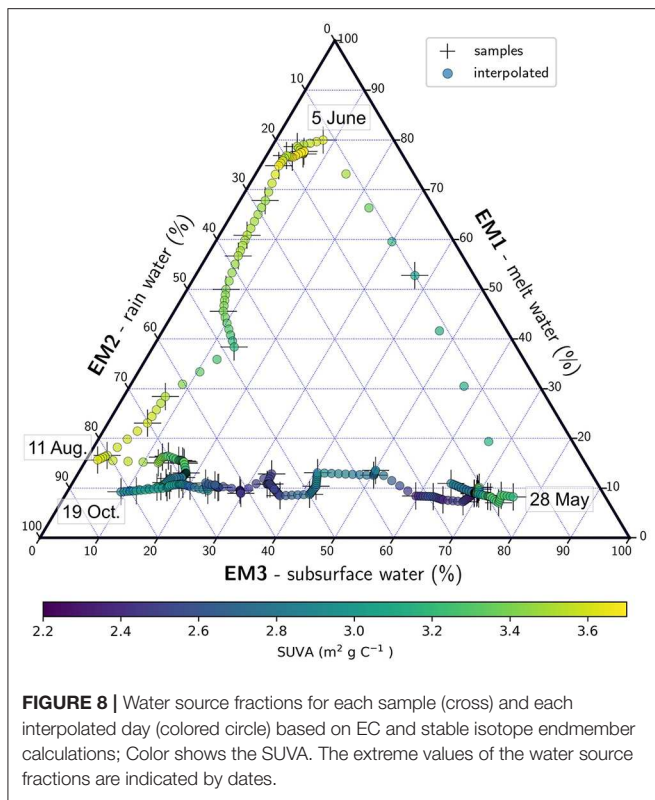


fraction of EM1 (melt-water) decreases and another water source endmember—EM2—becomes dominant. f_{EM2} peaks during the discharge events in summer (82.3% on 11 August and 81.6% on 19 October) (Figure 8). At the same time, $\delta^{18}O$ reaches its highest annual values and the EC is dropping. Such high $\delta^{18}O$ and low EC suggest that this end member is meteoric in origin. For these rain-induced discharge maxima, $\delta^{18}O$ and d-excess co-vary. This is likely related to less evaporative fractionation (and thus a higher d-excess), when more moisture is transported to the catchment. On the other hand, the highest evaporative fractionation is observed during the late summer (lowest d-excess). When discharge in summer is low, DOM concentrations are likewise low ($\sim 6 \text{ mg L}^{-1}$). During these periods, little DOM is mobilized and transported into the Lena River. In total, rain water accounts for 49.3% ($340 \times 10^9 \text{ m}^3$) of annual discharge. 42.6% (2.9 Tg C) of DOC is transported by rain water into the Lena River (Figure 9).

During the summer period the quality of the DOM changes (lower SUVA, higher S275-295). This could reflect the progress of soil thawing and a deepening active layer. Percolation and the leaching of older and more degraded DOM from deeper in the soil would explain both observations. Changes in SUVA can be linked to changes in source and age of DOM (Stedmon et al., 2011; O'Donnell et al., 2014; Coch et al., 2019) and can indicate the intensity of permafrost degradation within the

catchment (Abbott et al., 2014). Furthermore, direct relationships are found between dissolved organic $\delta^{14}C$ and SUVA (Neff et al., 2006; Butman et al., 2012; O'Donnell et al., 2014). Annual maxima in S275-295 during low discharge in summer points toward a higher degree of (photo)-degradation in the river, which can be a result of longer transport time from its source to the sampling station and an exposure of DOM to photodegradation.

There are peaks in DOM concentrations during the distinct rain-induced discharge peaks in summer and fall (Figure 2). During these periods, that are likely caused by precipitation events over a large area, more DOM is mobilized from the catchment and transported into the Lena River. The general trend toward older and more degraded DOM is interrupted for those events and DOM similar to that transported by the spring snow melt enters the river. However, during the second strong rain event with peak values of rain water fraction (19 Oct.), higher SUVA values compared to the first summer rain event (11 Aug.) point toward a smaller fraction of young organic material. We suggest that reservoirs of young organic material on the land surface of the Lena River catchment are continuously washed out throughout the summer and thus higher relative fractions of older DOM from deeper soil horizons or permafrost become visible in the water samples. Additionally, rain events cause higher water levels in the Lena River that can re-connect isolated



waterbodies and flush lakes, ponds and wetlands. In this way, DOM from phytoplankton production and bleached DOM from standing water bodies can be introduced to the Lena River.

Generally, during the open water period a trend toward lower SUVA and higher S275-295 values is present and continues into the early winter along with a decreasing fraction of rain water. Thus, the amount of fresh and young DOM transported with rain water from the Lena River catchment decreases.

(3) **Winter** is the period when the Lena River is entirely covered with ice and the discharge is low ($< 6,000 \text{ m}^3 \text{ s}^{-1}$) compared to the other seasons and to the annual mean of $21,885 \text{ m}^3 \text{ s}^{-1}$. During this period, most precipitation in the Lena River catchment accumulates as snow on land and the water level is gradually lowered. The remaining rain water is removed from the hydrological system of the Lena River catchment and the rain water fraction decreases gradually during the first half of the winter (**Figure 8**).

In early March, $\delta^{18}\text{O}$ and EC reach a stable level. At this time, the third endmember (EM3) becomes the dominant water source (76.1% on 28 May) (**Figure 8**). We suggest that this water source endmember represents subsurface water such as groundwater, soil and pore water (Abbott et al., 2014), which are the only significant natural water sources in winter, when air temperatures are permanently below 0°C . Subsurface water represents 15.7% ($105 \times 10^9 \text{ m}^3$) of the annual Lena River water flux. 14.1% (1 Tg C) of the DOC flux was transported with this water during our year of observation (**Figure 9**).

Optical DOM characteristics in winter continue the trend of summer and fall toward older and more degraded DOM until

January 2019 when DOM concentrations and SUVA values begin to increase, coincident with a small increase in discharge. During low discharge in winter, a substantial part of the basin outlet discharge originates from the Vilui reservoir (Tananaev et al., 2016), which is regulated by a dam (Viluyskaya HPS) constructed in the 1970s. Changes of DOM concentration and composition may be caused by the higher relative fraction of reservoir water that is released through the Vilui dam. In reservoir water all water sources are pooled throughout the year, leading to DOC concentrations elevated above the otherwise low winter levels. The major increase in long term winter discharge of the Lena River in the late 1970s (**Figure 2**) coincides with the completion of Vilui dam construction and its reaching full capacity (Ye et al., 2003; Tananaev et al., 2016). In late winter, before the spring ice break-up, SUVA increases which likely coincides with first input of fresh DOM from the southern most parts of the catchment, where temperatures begin to rise above 0°C during the day.

Data from the ice core samples suggest that only low concentrations of DOM are incorporated into the river ice. Once the river ice is exported to the Laptev Sea shelf after spring ice break-up, it tends to dilute the shelf waters rather than act as a source of high DOM (Amon et al., 2012). However, patches of high sediment load on the Lena River ice (Fedorova et al., 2015) and the erosion of shorelines during ice jams (van Huissteden, 2020) contribute to the particulate organic matter flux to the Laptev Sea. There is likely a contribution to DOM from dissolution of the transported POM but this is expected to be quantitatively low in comparison to the river load of DOM.

The results presented here show that high frequency sampling reveals subtle changes in water chemistry reflecting the seasonal changes in the hydrology of the Lena River catchment. Most of the intra-annual variation in the Lena River water biogeochemistry can be captured by the 3-component endmember analysis presented here. However, river water at the basin outlet represents an integrated signal from the entire catchment and geographical and temporal variations of biogeochemical signals within the catchment cannot be detected by sampling.

Comparison of Reported DOM Values and Fluxes for the Lena River

Overall, our observed ranges of DOC concentrations and $a_{\text{CDOM}}(254)$ agree well with reported values from sampling programs of shorter duration or substantially lower sampling frequency (ArcticGRO and PARTNERS (Holmes et al., 2018a,b); Lena River freshet 2014 (Eulenburg et al., 2019)). ArcticGRO and PARTNERS samples have significantly higher DOC concentrations compared to samples from Samoylov Island in 2014 and 2018. Generally, these samples show a lower r^2 (0.89) of their DOC to $a_{\text{CDOM}}(275)$ regression, compared to the dataset shown in this study ($r^2 = 0.97$). This may indicate either a prominent local influence of anomalous DOM or the impact of analytical differences and different sampling protocols.

A number of studies have reported annual organic carbon fluxes of the Lena River (**Table 4**). Reported values range between 3.6 Tg C (Gordeev and Kravchishina, 2009) and 7.67 Tg C

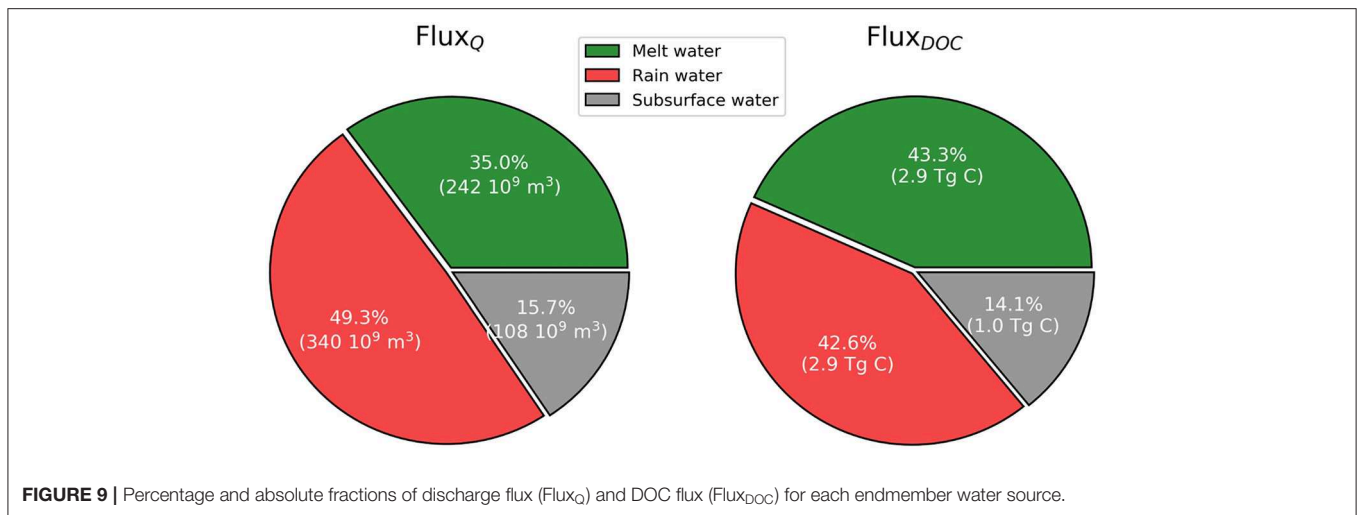


TABLE 4 | Previously published discharge and DOC fluxes and type of sampling and calculation.

References	Type	Year(s)	Annual discharge flux (10 ⁹ m ³)	Annual DOC flux (Tg C)
Opsahl et al. (1999)	<i>In situ</i>	1999	n.a.	4.7
Raymond et al. (2007)	<i>In situ</i> + load model	2004 2005	566 654	5.26 6.39
Gordeev and Kravchishina (2009)	<i>In situ</i>	n.a.	n.a.	3.6
Stedmon et al. (2011)	<i>In situ</i> + load model	2004 & 2005	615.1	7.27
Holmes et al. (2012)	<i>In situ</i> + load model	1999-2008	581	5.68 (varied between 4.1–7.4 Tg C yr ⁻¹)
Kicklighter et al. (2013)	Terrestrial ecosystem model	1990 - 2006	352	7.67
Wild et al. (2019)	<i>In situ</i> + load model	2003 - 2013	n.a.	5.71
This study	<i>In situ</i> , long period, high frequency	2018/19	690.18	6.8

(Kicklighter et al., 2013). Variability in DOC flux estimates is likely not only the result of inter-annual variability, but probably also results from differences in sampling frequency, sampling strategy and methods of calculating flux. All of the reported studies are based on a lower number of samples per year and/or on statistical models that are used to pair discharge measurements with concentration (Table 4). With statistical models, such as LOADEST (Runkel et al., 2004), estimated daily concentrations can be generated, with which seasonal or annual fluxes can be calculated. Although seasonal changes for the relation of discharge to concentration are taken into account in such models, the very low number of samples is critical and simplifying assumptions can cause systematic biases in flux estimates. Our dataset improves sample frequency and reveals that there is no clear relationship between discharge and DOC concentration. A relationship only persists during spring ($r^2 = 0.87$, not shown). The impact of methodological differences for flux calculations becomes clear when comparing reported DOC flux values for 2004 and 2005 from Raymond et al. (2007) and Stedmon et al. (2011), that are based on precisely the same data but differ by 22.3% (1.46 Tg C). Furthermore, the reported fluxes that used models to relate discharge and DOC concentration tacitly assume that variations in DOC concentration are driven

by processes changing the discharge, which is equivalent to a one-component system. Thus, the expected increase of DOC mobilization from degrading and thawing permafrost, which does not significantly affect discharge, cannot be reflected by such models.

The identification of seasonal fluxes by selecting certain time periods (Stedmon et al., 2011; Holmes et al., 2012) throughout different years is a method of addressing this shortcoming but prevents meaningful direct comparisons. The timing of seasonal changes such as the spring ice break-up varies between years and thus influence flux calculations for time periods defined by the calendar. Differences in seasonal fluxes between years (Supplementary Figure 5) thus are more likely to show shifts in seasonal timing than in water or dissolved load provenance.

We compared DOC fluxes during the spring discharge peak of 2 years (2014 and 2018). Although the discharge flux between both years differs by 26.4% (4.9 km³), the difference in DOC flux is not as severe (9.4%). At least for the spring discharge peak, lower discharge is compensated by higher concentrations, which results in a similar flux. This is, however, not the case for summer and fall rain events where high DOC concentrations are triggered by rain that washes organic matter from land surfaces into the Lena River. Thus, recurrence intervals and the

intensity of such rain events dominate inter-annual variability in DOC fluxes.

The high temporal resolution of the sampling program presented in this study improves our ability to capture all seasonal events relevant to organic carbon fluxes, without assuming a relationship between DOC concentration and discharge. Thus, we are confident that our flux calculation can be used as a baseline for future trend analysis of Lena River DOC fluxes and, in particular, for the identification of changing seasonality. Little is known about changes in biogeochemistry between sampling locations further upstream (e.g., Zhigansk) and Samoylov Island. Continuous sampling of both programs (Samoylov sampling and ArcticGRO) will, however, enable future comparative studies.

CONCLUSION

In a warming Arctic, we expect permafrost landscapes to change dramatically. One result will be the remobilization of dissolved matter from permafrost. In particular, the release of dissolved organic carbon is under immense research attention due to its strong potential feedbacks for the climate. By monitoring the biogeochemistry of Arctic river water at the river mouth, we can observe ongoing changes that reflect change at the catchment scale.

In this study, we present 1 year of biogeochemical data from the Lena River, one of the largest Arctic Rivers. We improve on existing studies by sampling at high temporal frequency throughout the whole year. The main drivers that are responsible for the strong seasonality of water discharge and DOC flux were ascribed to three water sources—melt water, rain water and subsurface water. Melt and rain water were found to be the prevailing water sources that transport together 5.8 Tg C dissolved organic matter (85% of annual flux) into the Lena River. Optical DOM indices revealed changing composition and sources of DOM throughout the year. The high resolution sampling also revealed two phases of melt water introduction, with an initial phase of approximately 1 week, during which melt water carries little or no DOM and dilutes river DOM concentrations without altering DOM character, followed by a large pulse of fresher organic matter from the catchment that substantially changes the river DOM characteristics. Future studies including $\delta^{14}\text{C}$ measurements of the age of the DOM will enable direct relation of optical DOM indices to the DOM age and the contribution of permafrost thaw as shown in Neff et al. (2006).

The results of this sampling program provide a baseline for future shifts in seasonal variations as well as inter-annual variation of DOM and the chemistry of the Lena River. The annual flux of 6.8 Tg C was calculated without recourse to load models, which are probably the source of discrepancies between existing estimates. Continuous under-ice sampling revealed that high winter DOM concentrations are probably related to the discharge of reservoir water from the Vilui tributary.

This dataset represents the first year of a planned long-term monitoring program at the Research Station Samoylov Island and provides a reference data set against which future change of this

large integrative system may be measured. Continuous sampling of Arctic River water will facilitate identification of intra and inter-annual trends during ongoing climate change.

DATA AVAILABILITY STATEMENT

The data sets used here are available online (<https://doi.org/10.1594/PANGAEA.913196>, <https://doi.org/10.1594/PANGAEA.913279>).

AUTHOR CONTRIBUTIONS

BJ, PO, and AM initiated the sampling program. BJ and PO designed the study with contributions of all authors. BJ prepared the figures and led the writing of the paper. All authors contributed to the discussion and the editing of the paper.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fenvs.2020.00053/full#supplementary-material>

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